

# **NIST GCR 03-854**

# Guidelines for Evaluating Liquefaction Resistance Using Shear Wave Velocity Measurement and Simplified Procedures



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# **NIST GCR 03-854**

# Guidelines for Evaluating Liquefaction Resistance Using Shear Wave Velocity Measurement and Simplified Procedures

Prepared for

U.S. Department of Commerce Materials and Construction Research Division National Institute of Standards and Technology Gaithersburg, MD 20899-8611

By

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#### ABSTRACT

Predicting the liquefaction resistance of soil is an important step in the engineering design of new and the retrofit of existing structures in earthquake-prone regions. The procedure currently used in the U.S. and throughout much of the world to predict liquefaction resistance is termed the simplified procedure. This simplified procedure was originally developed by H. B. Seed and I. M. Idriss in the late 1960s using blow count from the Standard Penetration Test. Small-strain shear wave velocity measurements provide a promising supplement and in some cases, where only geophysical measurements are possible, may be the only alternative to the This report presents guidelines for evaluating liquefaction penetration-based approach. resistance using shear wave velocity measurements. These guidelines were written in cooperation with industry, researchers and practitioners, and evolved from workshops in 1996 and 1998 as well as review comments received on an earlier draft. The guidelines present a recommended procedure, which follows the general format of the penetration-based simplified procedure. The proposed procedure has been validated through case history data from more than 20 earthquakes and 70 measurement sites in soils ranging from clean fine sand to sandy gravel with cobbles to profiles including silty clay layers. Deterministic liquefaction resistance curves were established by applying a modified relationship between the shear wave velocity and cyclic stress ratio for the constant average cyclic shear strain suggested by R. Dobry. These curves correctly predict moderate to high liquefaction potential for over 95 % of the liquefaction case histories, and are shown to be consistent with the penetration-based curves in sandy soils. From logistic regression and Bayesian models, the recommended deterministic curve is characterized with a probability of liquefaction of about 26 %. To further validate the procedure, additional case histories are needed with all soil types that have and have not liquefied, particularly from deeper deposits (depth > 8 m) and from denser soils (shear wave velocity > 200 m/s) shaken by stronger ground motions (peak ground acceleration > 0.4 g). The guidelines serve as a resource document for practitioners and researchers involved in evaluating soil liquefaction resistance.

KEYWORDS: building technology; earthquakes; in situ measurements; seismic testing; shear wave velocity; soil liquefaction

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#### ACKNOWLEDGMENTS

The NIST program to develop Guidelines for Evaluating Liquefaction Resistance Using Shear Wave Velocity Measurements and Simplified Procedures began in October of 1995. The initial work involved a review of proposed simplified procedures, collection of available case history data, participation in two workshops, and development of a recommended procedure. The first workshop was held on January 4-5, 1996 in Salt Lake City, Utah, and was sponsored by the National Center for Earthquake Engineering Research (NCEER). The second workshop was held on August 14-15, 1998 also in Salt Lake City, and was sponsored by the Multidisciplinary Center for Earthquake Engineering Research (MCEER, formally NCEER) and the National Science Foundation (NSF). Workshop participants included: T. Leslie Youd (Chair), Brigham Young University; Izzat M. Idriss (Co-chair), University of California at Davis; Ronald D. Andrus, formerly with National Institute of Standards and Technology; Ignacio Arango, Bechtel Corporation; John Barneich, Woodward-Clyde Consultants; Gonzalo Castro, GEI Consultants, Inc.; John T. Christian, Consulting Engineer; Ricardo Dobry, Rensselaer Polytechnic Institute; W. D. Liam Finn, University of British Columbia; Leslie F. Harder, Jr., California Department of Water Resources; Mary Ellen Hynes, U.S. Army Corps of Engineers, WES; Kenji Ishihara, Science University of Tokyo, Joseph P. Koester, U.S. Army Corps of Engineers, WES; Sam S. C. Liao, Parsons Brinckerhoff; Faiz Makdisi, Geomatrix Consultants; William F. Marcuson, III, Virginia Tech; Yoshiharu Moriwaki, Woodward-Clyde Consultants; Maurice S. Power, Geomatix Consultants; Peter K. Robertson, University of Alberta: Raymond B. Seed, University of California at Berkeley; and Kenneth H. Stokoe, II. University of Texas at Austin. The workshop participants provided expert review for the initial work.

Draft guidelines were published in a NIST technical report (*NISTIR 6277*, Andrus et al., 1999), based on the results of the initial work. Prior to publishing the draft guidelines, six technical experts were asked to review the guidelines. The six technical experts were: Ricardo Dobry; Mary Ellen Hynes; Izatt M. Idriss; Robert Pyke, Consulting Engineer; Richard D. Woods, University of Michigan; and T. Leslie Youd.

Since the publication of the draft guidelines, efforts have been made to obtain feedback on them from a broader base of practitioners and researchers. This was accomplished by distributing more than 100 copies of the draft guidelines with a request for comments. Written responses were received from the following individuals: John Barneich; Leo Brown for Robert Nigbor, GeoVision; Gonzalo Castro; A. G. Franklin, Consulting Engineer; C. Hsein Juang and Caroline J. Chen, Clemson University; Michael K. Lee and Ken Y. Lum, BC Hydro; Paul W. Mayne and James Schneider, Georgia Institute of Technology; Alan F. Rauch and James Chrisley, University of Texas at Austin; Soheil Nazarian, University of Texas at El Paso; and Zhenming Wang, Oregon Department of Geology and Mineral Industries. Comments were also received from anonymous reviewers of two journal papers summarizing this work. In addition, presentations were made at various professional meetings, including the Transportation Research Board Seventy-Eighth Annual Meeting Workshop on New Approaches to

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Liquefaction Analysis. The guidelines presented herein are based on the draft guidelines, and feedback from the mail reviews, journal paper reviews, and professional meetings. The authors thank the many people who provided comments and suggestions.

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Draft copies of this report were distributed to fifteen technical experts for their final review and comment. The fifteen technical experts were: Donald G. Andersen, CH2M Hill, Gonzalo Castro, John T. Christian, Ricardo Dobry, Mary Ellen Hynes, Izzat M. Idriss, Sam S. C. Liao, James K. Mitchell, Maurice S. Power, Robert Pyke, Glenn J. Rix, Georgia Institute of Technology, Peter K. Robertson, Raymond Seed, T. Leslie Youd, and Richard D. Woods. Comments were received from several of these experts and, to the extent possible, have been incorporated into the report.

The final report was submitted to the National Institute of Standards and Technology on September 28, 2000 for review. Based on subsequent work, some minor errors are corrected in this publication.

Finally, the authors express their thanks to the staff at the National Institute of Standards and Technology. Harry Brooks, Rose Estes, Bonnie Gray and other library staff assisted with the collection of several references cited in this report. Nicholas Carino and Fahim Sadek of the Structures Division reviewed the earlier draft guidelines, and provided many helpful suggestions. William Guthrie of the Statistical Engineering Division shared insights into the logistic regression technique for determining probability. John Gross served as the Technical Information Contact for the final guidelines.

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#### CHAPTER 1

#### INTRODUCTION

#### **1.1 BACKGROUND**

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A major cause of damage from earthquakes is liquefaction-induced ground failure. For example, direct property loss caused by liquefaction during the 1989 Loma Prieta, California earthquake (moment magnitude,  $M_w = 7.0$ ) was over \$100 million (Holzer, 1998). Large indirect property loss by fire almost occurred in 1989 when liquefaction-induced ground deformation ruptured water mains that served the Marina District of San Francisco. Fortunately, the fire in the Marina District at Divisadero and Beach Streets was contained to the three-story apartment building where it ignited. It was also fortunate that the 1989 earthquake did not occur closer to the San Francisco Bay area. The cities of Kobe and Osaka, Japan were not so fortunate. The 1995 Hyogoken-Nanbu earthquake ( $M_w = 6.9$ ) directly struck this metropolitan area, causing over \$100 billion in property damage (Kimura, 1996). A significant portion of the damage in Kobe can be attributed to liquefaction-induced ground deformation. These are just two of many examples of major damage caused by liquefaction-induced ground failure. Predicting soil liquefaction resistance is an important step in the engineering design of new and the retrofit of existing structures in earthquake-prone regions.

The procedure widely used in the United States and throughout much of the world for predicting the liquefaction resistance of soils is termed the *simplified procedure*. This simplified procedure was originally developed by Seed and Idriss (1971) using blow count from the Standard Penetration Test (SPT) correlated with a parameter called *cyclic stress ratio* that represents the seismic loading on the soil. Since 1971, the procedure has been revised and updated (Seed, 1979; Seed and Idriss, 1982; Seed et al., 1983; Seed et al., 1985). Correlations based on the Cone Penetration Test (CPT) and shear wave velocity measurements have also been developed by various investigators. General reviews of the simplified procedure are contained in a report by the National Research Council (1985), a workshop report edited by Youd and Idriss (1997), and a journal paper by Youd et al. (2001).

Small-strain shear wave velocity,  $V_s$ , measurements provide a promising supplement and in some cases, where only geophysical measurements are possible, may be the only alternative to the penetration-based approach. The use of  $V_s$  as an index of liquefaction resistance is soundly based because both  $V_s$  and liquefaction resistance are similarly influenced by many of the same factors (e.g., void ratio, state of stress, stress history, and geologic age).

Some advantages of using  $V_s$  are (Dobry et al., 1981; Seed et al., 1983; Stokoe et al., 1988a; Tokimatsu and Uchida, 1990): (1) Measurements are possible in soils that are hard to sample, such as gravelly soils where penetration tests may be unreliable. (2) Measurements can be performed on small laboratory specimens, allowing direct comparisons between laboratory and field behavior. (3)  $V_s$  is a basic mechanical property of soil materials, directly related to small-strain shear modulus,  $G_{max}$ , by:

$$G_{max} = \rho \, V_S^2 \tag{1.1}$$

where

 $\rho$  = the mass density of soil.

(4)  $G_{max}$ , or  $V_s$ , is in turn a required property in analytical procedures for estimating dynamic shearing strain in soil in earthquake site response and soil-structure interaction analyses. (5)  $V_s$  can be measured by the Spectral-Analysis-of-Surface-Waves (SASW) test method at sites where borings may not be permitted, such as capped landfills, and sites that extend for great distances where rapid evaluation is required, such as lifelines and large building complexes.

Three concerns when using  $V_s$  to evaluate liquefaction resistance are: (1) Measurements are made at small strains, whereas pore-water pressure buildup and liquefaction are medium- to high-strain phenomena (Jamiolkowski and Lo Presti, 1990; Teachavorasinskun et al., 1994; Roy et al., 1996). This concern can be significant for cemented soils, since small-strain measurements are highly sensitive to weak interparticle bonding which is eliminated at medium and high strains. It also can be significant in silty soils above the water table where negative pore water pressures can increase  $V_s$ . (2) No samples are obtained for classification of soils and identification of non-liquefiable soft clayey soils. According to the so-called Chinese criteria, non-liquefiable clayey soils have clay contents (particles smaller than 5  $\mu$ m) > 15 %, liquid limits > 35 %, or moisture contents < 90 % of the liquid limit (Seed and Idriss, 1982). Andrews and Martin (2000) refined this criteria to soils with clay contents (particles smaller than 2  $\mu$ m) ≥ 10 % and liquid limits ≥ 32 % (by Casagrade-type percussion apparatus) for non-liquefiable clayey soils. (3) Thin, low  $V_s$  strata may not be detected if the measurement interval is too large (USBR, 1989; Boulanger et al., 1997). In general, borings should always be a part of the field investigation. Surface geophysical measurements and cone soundings are often conducted first to help select the best locations for borehole sampling and testing. Surface geophysical tests usually involve making measurements at several different locations, and provide general, or average, stratigraphy for sediments beneath the area tested. The ability of surface geophysical methods to resolve a layer at depth depends on the thickness, depth, and continuity of that layer, as well as the test and interpretation procedures employed. Cone soundings provide detailed stratigraphy at each test location for sediments that can be penetrated. The preferred practice when using  $V_s$  measurements to evaluate liquefaction resistance is to drill sufficient boreholes and conduct sufficient tests to detect and delineate thin liquefiable strata, to identify non-liquefiable clay-rich soils, to identify soils above the ground water table that might have lower values of  $V_s$  should the water table rise, and to detect liquefiable weakly cemented soils.

#### **1.2 PURPOSE**

This report presents guidelines for evaluating liquefaction resistance through shear wave velocity measurements. The guidelines are based on an earlier report entitled "Draft Guidelines for Evaluating Liquefaction Resistance Using Shear Wave Velocity Measurements and Simplified Procedures" by Andrus et al. (1999), which evolved from two workshops. The first workshop was held on January 4-5, 1996 in Salt Lake City, and was sponsored by the National Center for Earthquake Engineering Research (NCEER). The second workshop was held on August 14-15, 1998 also in Salt Lake City, and was sponsored by the Multidisciplinary Center for Earthquake Engineering Research (MCEER, formally NCEER) and the National Science Foundation (NSF). These two workshops are herein called the 1996 NCEER Workshop and 1998 MCEER Workshop. The guidelines present a recommended procedure based on the suggestions given at the workshops, as well as review comments received on the draft guidelines. The guidelines provide guidance on selecting site variables and correction factors that are consistent with the shear-wave-based procedure.

From the comments received on the earlier draft guidelines, several improvements have been made to the guidelines. The major improvements include: (1) Much of the background and development information is moved to the appendixes to provide a clearer, more practical description of the recommended procedure. (2) A table is added to compare advantages and disadvantages of the various *in situ*  $V_s$  test methods for liquefaction assessment. (3) A more comprehensive description of the database is given, particularly on the nature of the case histories of gravelly soils. (4) The results of a probability study of the  $V_s$ -based case history data are added to calibrate the recommended liquefaction resistance curve and to compare with the results of probability studies of the SPT-based case history data.

#### **1.3 REPORT OVERVIEW**

Following this introduction, Chapter 2 presents the recommended procedure for evaluating liquefaction resistance using  $V_{S}$ . Chapter 3 illustrates the application of the procedure using two case studies. And Chapter 4 summarizes the recommended procedure and identifies issues that remain to be resolved.

Eight appendixes are included to assist the reader, and to provide information used in the development of the guidelines. Appendix A presents a list of references cited in the guidelines. Appendix B provides a list of Symbols and Notation, and Appendix C provides a Glossary of Terms. Appendix D reviews six proposed  $V_s$ -based liquefaction resistance curves. Appendix E describes the general characteristics of case history data used to develop the recommended liquefaction resistance curves. Appendix F presents the development of the recommended curves. Appendix G considers three probability models for the case history data. Finally, Appendix H presents a summary of the case history data.

#### CHAPTER 2

#### LIQUEFACTION EVALUATION PROCEDURE

This chapter presents guidelines for using the  $V_{s}$ -based liquefaction evaluation procedure originally proposed by Andrus and Stokoe (1997) and subsequently updated in the report by Andrus et al. (1999) and the paper by Andrus and Stokoe (2000). The evaluation procedure follows the general format of the Seed-Idriss simplified procedure, and the general recommendations of the 1996 NCEER Workshop (Youd et al., 1997) and 1998 MCEER Workshop (Youd et al., 2001). It requires the calculation of three parameters: (1) the level of cyclic loading on the soil caused by the earthquake, expressed as a cyclic stress ratio; (2) the stiffness of the soil, expressed as an overburden stress-corrected shear wave velocity; and (3) the resistance of the soil to liquefaction, expressed as a cyclic resistance ratio. Each parameter is discussed below.

#### 2.1 CYCLIC STRESS RATIO (CSR)

The cyclic stress ratio,  $\tau_{av} / \sigma'_v$ , at a particular depth in a level soil deposit can be expressed as (Seed and Idriss, 1971):

$$CSR = \frac{\tau_{av}}{\sigma'_{v}} = 0.65 \left(\frac{a_{\max}}{g}\right) \left(\frac{\sigma_{v}}{\sigma'_{v}}\right) r_{d}$$
(2.1)

where

- $\tau_{av}$  = the average equivalent uniform shear stress caused by the earthquake and is assumed to be 0.65 of the maximum induced stress,
- $a_{max}$  = the peak horizontal ground surface acceleration,

g = the acceleration of gravity,

- $\sigma_v$  = the total vertical (overburden) stress at the depth in question,
- $\sigma'_{v}$  = the initial effective overburden stress at the same depth, and
- $r_d$  = a shear stress reduction coefficient to adjust for the flexibility of the soil profile.

#### 2.1.1 Peak Horizontal Ground Surface Acceleration

Peak horizontal ground surface acceleration is a characteristic of the ground shaking intensity, and is defined as the peak value in a horizontal ground acceleration record that would occur at the site without the influence of excess pore-water pressures or liquefaction that might develop (Youd et al., 2001).

Peak accelerations are commonly estimated using empirical attenuation relationships of  $a_{max}$  as a function of earthquake magnitude, distance from the energy source or surface project of the fault rupture, and local site conditions. Since many published attenuation relationships are based on both peak values obtained from ground motion records for the two horizontal directions (sometimes referred to as the randomly oriented horizontal component), the geometric mean (square root of the product) of the two peak values is used. According to Youd et al. (2001), use of the geometric mean is consistent with the derivation of the SPT-based procedure and is preferred for use in engineering practice. However, use of the larger of the two horizontal peak accelerations would be conservative and is allowable.

Regional or national seismic hazard maps (<u>http://geohazards.cr.usgs.gov/eq/</u>; Frankel et al., 2000) are also often used to estimate peak accelerations. If peak acceleration is estimated from a map, the magnitude and distance information should be obtained from the deaggregated matrices used to develop the map. The value of  $a_{max}$  selected will depend on the target level of risk and compatibility of site conditions. For site conditions not compatible with available probabilistic maps or attenuation relationships, the value of  $a_{max}$  may be corrected based on dynamic site response analyses or site class coefficients given in the latest building codes.

#### 2.1.2 Total and Effective Overburden Stresses

Required in the calculation of  $\sigma_v$  and  $\sigma'_v$  are densities of the various soil layers, as well as characteristics of the ground water. For non-critical projects involving hard-to-sample soils below the ground water table, densities are often estimated from typical values for soils with similar grain size and penetration or velocity characteristics. Fortunately, *CSR* is not very sensitive to density, and reasonable estimates of density yield reasonable results.

The values of  $\sigma'_{\nu}$  and *CSR* are sensitive to the ground water table depth. Other ground water characteristics that may be significant to liquefaction evaluations include seasonal and long-term water level variations, depth of and pressure in artesian zones, and whether the water table is perched or normal.

#### 2.1.3 Stress Reduction Coefficient

Equation 2.1 is based on Newton's second where force is equal to mass times acceleration. The coefficient  $r_d$  is added because the soil column behaves as a deformable body rather than a rigid body.

2.1.3.1 Relationship by Seed and Idriss (1971)—Values of  $r_d$  are commonly estimated from the chart by Seed and Idriss (1971) shown in Fig. 2.1. This chart was determined analytically using a variety of earthquake motions and soil conditions. Average  $r_d$  values given in the chart can be estimated using the following functions (Liao and Whitman, 1986; Robertson and Wride, 1997):

$r_d = 1.0 - 0.00765 z$	for $z \le 9.15$ m	(2.2a)
$r_d = 1.174 - 0.0267 z$	for 9.15 m $< z \le 23$ m	(2.2b)
$r_d = 0.744 - 0.008 z$	for 23 m $< z \le$ 30 m	(2.2c)

where

z = the depth below the ground surface in meters.

Figure 2.1 shows the average  $r_d$  values approximated by Eq. (2.2). These average  $r_d$  values were suggested by the 1996 NCEER Workshop (Youd et al., 1997; 2001) for non-critical projects.

2.1.3.2 Revised Relationship Proposed by Idriss (1998; 1999)—Figure 2.2 presents revised average values of  $r_d$  proposed by Idriss (1998; 1999) for various earthquake magnitudes. The plotted curves are averages of many individual curves derived analytically by Golesorkhi (1989) under the supervision of the late Prof. H. B. Seed. They are defined by the following relationship (after Idriss, 1998; modified for depth in meters):

$$ln(r_d) = \alpha(z) + \beta(z) M_w \tag{2.3}$$

where

$$\alpha(z) = -1.012 - 1.126 \sin\left(\frac{z}{11.7} + 5.133\right)$$
, and (2.4)

$$\beta(z) = 0.106 + 0.118 \sin\left(\frac{z}{11.3} + 5.142\right). \tag{2.5}$$



Fig. 2.1 - Relationship Between Stress Reduction Coefficient and Depth Developed by Seed and Idriss (1971) with Approximate Average Value Lines from Eq. 2.2. (after Youd et al., 1997)



Fig. 2.2 - Relationship Between Average Stress Reduction Coefficient and Depth Proposed by Idriss (1998; 1999) with Average of Range Determined by Seed and Idriss (1971). As shown in Fig. 2.2, the curve defined by Eq. 2.3 for  $M_w = 7.5$  is almost identical to the average of the range published by Seed and Idriss (1971).

The scatter in the individual curves used to determine the average curves shown in Fig. 2.2, as well as Fig. 2.1, is rather large. For example, coefficients determined for a 30 m thick, loose sand deposit and magnitude 5.5 earthquakes exhibit standard deviations of about 0.1 at a depth of 5 m and 0.15 at a depth of 10 m. These standard deviation values would be larger if soil deposits of various thicknesses and stiffnesses are considered. Thus, as an alternative approach, the variation of  $r_d$  with depth may be calculated analytically using site-specific layer thicknesses and stiffnesses.

#### 2.2 STRESS-CORRECTED SHEAR WAVE VELOCITY

The *in situ*  $V_s$  can be measured by several seismic tests including crosshole, downhole, seismic cone penetrometer (SCPT), suspension logger, and Spectral-Analysis-of-Surface-Waves (SASW). Recent reviews of these test methods are given in Woods (1994), Kramer (1996), and Ishihara (1996). ASTM D-4428M-91 provides a standard test method for crosshole seismic testing. Standard test methods do not exist for the other seismic tests. Primary advantages and disadvantages of the various *in situ*  $V_s$  test methods are presented in Table 2.1. The accuracy of each test method can be sensitive to equipment and procedural details, soil conditions, and interpretation techniques.

One important factor influencing  $V_s$  is the state of stress in soil (Hardin and Drnevich, 1972; Seed et al., 1986). Laboratory test results (Roesler, 1979; Stokoe et al., 1985; Belloti et al., 1996) show that the velocity of a propagating shear wave depends equally on principal stresses in the direction of wave propagation and the direction of particle motion. Thus,  $V_s$  measurements made with wave propagation or particle motion in the vertical direction can be related by the following empirical relationship:

$$V_{S} = A \left(\sigma_{\nu}^{\prime}\right)^{m} \left(\sigma_{h}^{\prime}\right)^{m} \tag{2.6}$$

where

A	=	a parameter that depends on the soil structure,
$\sigma'_h$	=	the initial effective horizontal stress at the depth in question, and
т	=	a stress exponent with a value of about 0.125.

Table 2.1 - Comparison of Advantages and Disadvantages of Various In Situ Vs Test Methods for Liquefaction Assessment.

Feature			Measurement Method	states and the second	
	Crosshole	Downhole & Seismic Cone Penetrometer	Suspension Logger	Spectral-Analysis-of Surface-Waves	Surface Reflection/ Refraction
No. of boreholes (or soundings) required	2 or more	1	1	None	None
Quality control and repeatability <sup>1</sup>	Good	Good	Good	Good to fair, complex interpretation technique	Fair; often difficult to distinguish shear wave arrival
Resolution of soil stiffness variability <sup>2</sup>	Good; constant with depth	Good to fair; decreases with depth	Good at depth; poor close to the surface	Good to fair; decreases with depth; provides good global average	Fair to poor; provides coarse global average
Major component of particle motion or wave propagation in vertical direction	Yes, with vertically polarized shear waves	Yes, with test depth greater than distance between borehole and shear beam source	Yes	Yes, with vertical source	Yes, with horizontal source for reflection and vertical source for refraction
Limitations	Possible refraction problems; senses stiffer material at test depth; most expensive test method	Possible refraction problems with shallow layers; wave travel path increases with depth	Fluid-filled hole required; may not work in cased holes and soft soils; high test frequency (500 Hz- 2000 Hz)	Horizontal layering assumed; poor resolution of thin layers and soft material adjacent to stiff layers; no samples recovered	In refraction test, only works for velocity increasing with depth; no samples recovered
Other	Most reliable test method	Penetration data also obtained from seismic cone; detailed layered profile with cone		Well-suited for tomo- graphic imaging large areas and testing difficult to penetrate soils	Well-suited for screening large areas
			Linnention technique	ac for all mothode	

<sup>1</sup>Good quality depends on good equipment and procedural details, and good interpretation <sup>2</sup>Resolution depends on test spacing for all methods.

Following the traditional procedures for correcting SPT blow count and CPT tip resistance, one can correct  $V_s$  to a reference overburden stress by (Sykora, 1987b; Robertson et al., 1992):

$$V_{SI} = V_S C_{VS} = V_S \left(\frac{P_a}{\sigma'_v}\right)^{0.25}$$
(2.7)

where

V <sub>SI</sub>	=	the overburden stress-corrected shear wave velocity,
Cvs	=	a factor to correct measured shear wave velocity for overburden pressure;
Pa	=	a reference stress, 100 kPa or approximately atmospheric pressure, and
$\sigma'_v$	=	the initial effective overburden stress in kPa.

A maximum  $C_{VS}$  value of 1.4 is generally applied to  $V_S$  data at shallow depths, similar to the SPT and CPT procedures. In using Eq. (2.7), it is implicitly assumed that the initial effective horizontal stress,  $\sigma'_h$ , is a constant factor of the initial effective overburden stress (because both  $\sigma'_v$  and  $\sigma'_h$  affect  $V_S$  as shown in Eq. (2.6)). This factor, generally referred to as  $K'_o$ , is assumed to be approximately 0.5 at natural, level-ground sites where liquefaction has occurred or is likely to occur. Also, in applying Eq. (2.7), it is implicitly assumed that  $V_S$  is measured with both the directions of particle motion and wave propagation polarized along principal stress directions and one of those directions is vertical.

Since the direction of wave propagation and the direction of particle motion is different with respect to the stress in the soil for each *in situ* seismic test method, some variations between measured  $V_s$  is expected. These variations are minimized by performing the tests with at least a major component of wave propagation or particle motion in the vertical direction. To have a major component of wave propagation or particle motion in the vertical direction, crosshole tests are conducted with particle motion in the vertical direction, downhole and seismic cone tests are conducted at depths greater than the distance between the shear beam source and the borehole or cone sounding such that wave propagation is in the vertical direction, and SASW tests are conducted with a vertical source.

In soils above the ground water table, particularly silty soils, negative pore pressures increase the effective state of stress and, hence, the value of  $V_s$  measured in seismic tests. This effect should be considered in the estimation of  $\sigma'_{\nu}$  for correcting  $V_s$  to  $V_{sI}$ , and for computing *CSR* using Eq. (2.1).

#### 2.3 CYCLIC RESISTANCE RATIO (CRR)

The value of CSR separating liquefaction and non-liquefaction occurrences for a given  $V_{SI}$ , or corrected penetration resistance, is called the cyclic resistance ratio, CRR. Figure 2.3 presents the  $CRR-V_{SI}$  curves developed by Andrus et al. (1999) for magnitude 7.5 earthquakes and uncemented, Holocene-age soils. The curves are dashed above CRR of about 0.35 to indicate that they are based on limited field performance data, as discussed in Appendix F. The curves do not extend much below 100 m/s, since there are no field data to support extending them to the origin. They are defined by:

$$CRR = MSF \left\{ 0.022 \left( \frac{K_c V_{SI}}{100} \right)^2 + 2.8 \left( \frac{1}{V_{SI}^* - K_c V_{SI}} - \frac{1}{V_{SI}^*} \right) \right\}$$
(2.8)

where

MSF = the magnitude scaling factor to account for the effect of earthquake magnitude,

 $V_{S1}^*$  = the limiting upper value of  $V_{S1}$  for cyclic liquefaction occurrence, and

 $K_c$  = a factor to correct for high  $V_{SI}$  values caused by cementation and aging.

The first (or squared) term in Eq. (2.8) is based on a relationship between *CRR* and  $V_{SI}$  for constant average cyclic strain derived by R. Dobry. The second term is a hyperbola with small value at low values of  $V_{SI}$ , and a very large value as  $V_{SI}$  approaches  $V_{SI}^*$ .

#### 2.3.1 Magnitude Scaling Factor

The magnitude scaling factor is traditionally applied to *CRR*, rather than the cyclic loading parameter *CSR*, and equals 1 for earthquakes with a magnitude of 7.5. For magnitudes other than 7.5, Table 2.2 presents scaling factors developed by various investigators. These magnitude scaling factors were derived from laboratory test results and representative cycles of loading (Seed and Idriss, 1982; Idriss, personal communication to T. L. Youd, 1995; Idriss, 1998; Idriss, 1999), correlations of field performance data and blow count measurements (Ambrasey, 1988; Youd and Noble, 1997), estimates of seismic energy for laboratory and field data (Arango, 1996), and correlations of field performance data and *in situ*  $V_s$  measurements (Andrus and Stokoe, 1997). Figure 2.4 shows a plot of the various magnitude scaling factors along with the range recommended by the NCEER Workshop (Youd et al., 1997; 2001).



Fig. 2.3 - Curves Recommended for Calculation of CRR from  $V_{SI}$  (Andrus et al., 1999).

<u></u>	Magnitude Scaling Factor (MSF)										
Moment Magnitude, M <sub>w</sub>	Seed & Idriss (1982)	Idriss (personal communi cation to T. L. Youd, 1995)	Idriss (1998)	Idriss (1999)	Ambraseys (1988)	Үот < 20	nd & No (1997) P <sub>L</sub> , % < 30	obie < 50	Ага (199	ngo 6)**	Andrus & Stokoe (1997)
(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)	(9)	(10)	(11)	(12)
5.5	1.43	2.20	1.625	1.68	2.86	2.86	3.42	4.44	3.00	2.20	2.8*
6.0	1.32	1.76	1.48	1.48	2.20	1.93	2.35	2.92	2.00	1.65	2.1
6.5	1.19	1.44	1.28	1.30	1.69 ·	1.34	1.66	1.99	1.60	1.40	1.6
7.0	1.08	1.19	1.12	1.14	1.30	1.00	1.20	1.39	1.25	1.10	1.25
7.5	1.00	1.00	0.99	1.00	1.00			1.00	1.00	1.00	1.0
8.0	0.94	0.84	0.88	0.87	0.67			0.73	0.75	0.85	0.8*
8.5	0.89	0.72	0.79	0.76	0.44			0.56			0.65*

Table 2.2 - Magnitude Scaling Factors Obtained by Various Investigators. (modified from Youd and Noble, 1997)

\*Extrapolated from scaling factors for  $M_w = 6, 6.5, 7, \text{ and } 7.5 \text{ using } MSF = (M_w/7.5)^{-3.3}$ .

\*\*Based on equivalent uniform number of cycles and consideration of distant liquefaction sites (Column 10), and energy principles (Column 11).

Although the 1996 NCEER Workshop (Youd et al., 1997) recommended a range of magnitude scaling factors for engineering practice, a consensus has not yet been reached by the workshop participants. At the August 1998 MCEER Workshop, a revised set of magnitude scaling factors and stress reduction coefficients (see Section 2.1.3.2) were proposed by I. M. Idriss. Liao and Lum (1998) present results of a statistical analysis supporting the original Seed and Idriss (1982) factors. The magnitude scaling factors recommended by the 1996 NCEER Workshop and the revised factors proposed by Idriss (1999) are discussed below.

2.3.1.1 Factors Recommended by 1996 NCEER Workshop—The magnitude scaling factors recommended by the 1996 NCEER Workshop (Youd et al., 1997) can be represented by:

$$MSF = \left(\frac{M_w}{7.5}\right)^n \tag{2.9}$$

where

 $M_w$  = moment magnitude, and n = an exponent.



Fig. 2.4 - Magnitude Scaling Factors Derived by Various Investigators with Range Recommended by the 1996 NCEER Workshop (after Youd et al., 1997; Youd and Noble, 1997).
Moment magnitude is the scale most commonly used for engineering applications, and is preferred for liquefaction resistance calculations (Youd et al., 1997). When only other magnitude scales are available, they can be converted to  $M_w$  using the relationship of Heaton et al. (1982) shown in Fig. 2.5.

The lower bound for the range of magnitude scaling factors recommended by the 1996 NCEER Workshop is defined with n = -2.56 (Idriss, personal communication to T. L. Youd, 1995) for earthquakes with magnitudes  $\leq 7.5$ . The upper bound of the recommended range is defined with n = -3.3 (Andrus and Stokoe, 1997) for earthquakes with magnitudes  $\leq 7.5$ . For earthquakes with magnitudes > 7.5, the recommended factors are defined with n = -2.56. Magnitude scaling factors defined by Eq. (2.9) and average  $r_d$  values originally proposed by Seed and Idriss (1971) should be used together when applying Eqs. (2.1) and (2.8).

2.3.1.2 Revised Factors Proposed by Idriss (1999)—The magnitude scaling factors proposed by Idriss (1999) are derived using laboratory data from Yoshimi et al. (1984) and a revised relationship between representative cycles of loading and earthquake magnitude. They are defined by the following equation:

$$MSF = 6.9 \exp\left(\frac{-M_w}{4}\right) - 0.06$$
 for  $M_w > 5.2$  (2.10a)

$$MSF = 1.82$$
 for  $M_w \le 5.2$  (2.10b)

where exp is the constant *e* raised to the power of the number given in the parentheses. Magnitude scaling factors defined by Eq. (2.10) and revised  $r_d$  proposed by Idriss (1999) should be used together when applying Eqs. (2.1) and (2.8).

**2.3.1.3 Recommended Magnitude Scaling Factors**—There is little difference in using magnitude scaling factors and  $r_d$  values recommended by the 1996 NCEER Workshop (Youd et al., 1997) and those proposed by Idriss (1999) for magnitudes near 7.5 and depths less than 11 m (see Appendix F, Sections F.2.2 and F.2.3). At magnitudes near 5.5, the difference is significant. The lower bound for the range of magnitude scaling factors defined by Eq. (2.9) with n = -2.56 is recommended in these guidelines because it provides more conservative assessment than with n = -3.3 for magnitudes less than 7.5. While the magnitude scaling factors defined by Eq. (2.9) with n = -2.56 are less conservative than the factors proposed by Idriss (1999) for magnitudes less than 7.5, the findings of Ambrasey (1988), I. M. Idriss (personal communication to T. L. Youd, 1995), Arango (1996), Youd and Noble (1997), and Andrus and Stokoe (1997; as indicated by the very conservative *CRR-V<sub>S1</sub>* curves shown in Figs. F.13 and F.14) support their use.



Fig. 2.5 - Relationship Between Moment Magnitude and Various Magnitude Scales (after Heaton et al., 1982).

## 2.3.2 Limiting Upper Value of $V_{SI}$ in Sandy Soils

The assumption of a limiting upper value of  $V_{SI}$  is equivalent to the assumption commonly made in the SPT- and CPT-based procedures dealing with clean sands, where liquefaction is considered not possible above a corrected blow count of about 30 (Seed et al., 1985) and a corrected tip resistance of about 160 (Robertson and Wride, 1998). Upper limits for  $V_{SI}$  and penetration resistance are explained by the tendency of dense soils to exhibit dilative behavior at large strains, causing negative pore-water pressures. While it is possible in a dense soil to generate pore-water pressures close to the confining stress if large cyclic strains or many cycles are applied, the amount of water expleied during reconsolidation is dramatically less for dense soils than for loose soils. As explained by Dobry (1989), in dense soils, settlement is insignificant and no sand boils or failure take place because of the small amount of water expelled. This is important because the definition of liquefaction used to classify the field behavior here, as well as in the penetration-based procedures, is based on surface manifestations such as boils and ground cracks.

The case history data above a CSR value of about 0.35 are limited, as discussed in Appendix F. Thus, current estimates of  $V_{S1}^*$  rely, in part, on penetration- $V_S$  correlations and, in part on the case histories. Values of  $V_{S1}^*$  can be estimated from:

$V_{S1}^* = 215 \text{ m/s}$	for sands with $FC \leq 5$ %	(2.11a)
$V_{S1}^* = 215 - 0.5(FC-5)$ m/s	for sands with 5 % $< FC < 35$ %	(2.11b)
$V_{S1}^* = 200 \text{ m/s}$	for sands and silts with $FC \ge 35$ %	(2.11c)

where

FC = average fines content in percent by mass.

Equations (2.8) and (2.11a) provide a CRR value of about 0.6 at  $V_{SI} = 210$  m/s. A  $V_{SI}$  value of 210 m/s is considered equivalent to a corrected blow count of 30 in sands with FC = 5 %, based on penetration- $V_S$  correlations.

# 2.3.3 Limiting Upper Value of V<sub>S1</sub> in Gravelly Soils

Although the  $V_{S1}^*$  values given in Eq. (2.11) were determined for sandy soils, the case history data indicate that these limits also represent reasonable limits for gravelly soils divided into the same categories based on fines content (see Fig. F.7). This might be considered rather surprising based on the penetration- $V_S$  correlations presented in the literature for gravelly soils. For instance, the correlation by Ohta and Goto (1978) suggested a  $V_{SI}$  value of 227 m/s for Holocene gravels at an equivalent  $(N_1)_{60}$  of 30. Similarly, the correlation by Rollins et al. (1998) provided a best-fit  $V_{SI}$  value of 232 m/s for Holocene gravels. On the other hand, all the liquefaction case history data exhibit  $V_{SI}$  values of about 200 m/s or less, suggesting that 230 m/s may be inappropriately high. To investigate further the value of  $V_{S1}^*$  in gravely soils, laboratory studies involving  $V_s$  measurements in gravelly soils were reviewed. Kokusho et al. (1995) clearly showed that the shear wave velocity (or stiffness) of gravelly soils varies greatly and is highly dependent on the particle gradation. Weston (1996) showed similar results for In both cases, the results show that increasing the uniformity coarse sands with gravels. coefficient can significantly increase the shear wave velocity in medium dense to dense gravels. On the other hand, very loose gravely soils, even well-graded gravels, can exhibit shear wave velocities similar to those of loose sands (Kokusho et al., 1995). The case history data presented in Fig. F.7 support the premise that gravely soils that are loose enough to exhibit significant liquefaction effects (boils, ground cracks, etc.) have shear wave velocities similar to loose sands. Hence, the boundaries developed for sandy soils are recommended as preliminary boundaries for gravely soils. However, additional work is clearly needed to understand the relationship between  $V_{SI}$  and liquefaction resistance of gravels.

## 2.3.4 Cementation and Aging Correction Factor

The recommended  $CRR-V_{SI}$  curves shown in Fig. 2.3 are limited to the characteristics of the database used to develop them. The database consists of relatively level ground sites with the following general characteristics: (1) uncemented soils of Holocene age; (2) average depths less than about 10 m; (3) ground water table depths between 0.5 m and 6 m, and (4) all  $V_S$  measurements are from below the water table. Correction factors may be used to extend the curves to site conditions different from the database.

The correction factor  $K_c$  is 1 for areas of uncemented, Holocene-age soils. For Pleistocene-age soils (>10 000 years), average estimates of  $K_c$  range from 0.6 to 0.8 based on the penetration- $V_{SI}$  correlations by Rollins et al. (1998a) and Ohta and Goto (1978), respectively. Figures 2.6 and 2.7 illustrate two methods for estimating the value of  $K_c$  using SPT and CPT test results, respectively. Shown in the figures are correlations for clean sands and silty soils implied by the  $CRR-V_{SI}$  relationship defined by Eq. (2.8) and 1996 NCEER Workshop recommended CRR-penetration relationships (Youd et al., 1997). In the example, the measured values of  $V_{SI}$ ,  $(N_I)_{60}$ ,  $q_{cIN}$ , and fines content are 220 m/s, 8, 55, and 10 %, respectively. The  $V_{SI}$ -penetration correlations. The  $K_c$  value is assumed to be the ratio of the predicted value of  $V_{SI}$ , based on the corrected penetration resistance and fines content, to the measured value of  $V_{SI}$ .



Fig. 2.6 - Suggested Method for Determining the Correction Factor  $K_c$  from  $(N_1)_{60}$ ,  $V_{S1}$ , and Fines Content at a Weakly Cemented Soil Site.



Fig. 2.7 - Suggested Method for Determining the Correction Factor  $K_c$  from  $q_{clN}$ ,  $V_{Sl}$ , and Fines Content at a Weakly Cemented Soil Site.

The method for estimating  $K_c$  described above assumes that the strain level induced during penetration testing is the same strain level causing liquefaction, which may not be true because pore-water pressure buildup to liquefaction can occur at medium strains in several loading cycles (Dobry et al., 1982; Seed et al., 1983). The method also assumes that liquefaction potential, blow count, and cone penetration resistance are <u>not</u> affected by cementation, which may not be a reasonable assumption. Hence, this suggested method should be used cautiously and with engineering judgment.

# 2.4 FACTOR OF SAFETY

A common way to quantify the potential for liquefaction is in terms of a factor of safety. The factor of safety,  $F_s$ , against liquefaction can be defined by:

$$F_S = \frac{CRR}{CSR} \tag{2.12}$$

By convention, liquefaction is predicted to occur when  $F_s \le 1$ . When  $F_s > 1$ , liquefaction is predicted not to occur.

As is the case with the SPT- and CPT-based charts, it is possible that liquefaction could occur outside the region of predicted liquefaction shown in Fig. 2.3. Consequently, the Building Seismic Safety Council (1997, page 158) suggests a factor of safety of 1.2 to 1.5 is appropriate when applying the Seed-Idriss simplified procedure in engineering design. The acceptable value of  $F_s$  for a particular site will depend on several factors, including the type and importance of structure and the potential for ground deformation. Based on V<sub>s</sub>-SPT blow count correlations (see Section F.4) and probability studies (see Section G.2), the recommended V<sub>s</sub>-based procedure is as conservative as the SPT-based procedure outlined by Seed et al. (1985) and updated by the NCEER Workshop (Youd et al., 2001). Thus, the same range of factor of safety is recommended for the V<sub>s</sub>-based method.

### 2.5 PROBABILITY-BASED EVALUATION

Probability of liquefaction,  $P_L$ , is required information for making risk-based design decisions. As discussed in Appendix G, the relationship between  $P_L$  and  $F_S$  for the deterministic procedure described above can be expressed as (Juang et al., 2002; modified from Juang et al., 2001a):

$$P_L = \frac{1}{1 + \left(\frac{F_S}{0.73}\right)^{3.4}}$$
(2.13)

In Eq. (2.13), a  $F_s$  value of 1 corresponds to points on the deterministic curves shown in Fig. 2.3. Thus, on average, the deterministic curves are characterized by a  $P_L$  value of 26 %. This average  $P_L$  value is similar to probability estimates determined for the SPT-based curves (Liao et al., 1988; Youd and Noble, 1997; Juang et al., 2000a). The relationship defined by Eq. (2.13) is plotted in Fig. 2.8, and provides the link between the probabilistic and deterministic methods. By combining Eqs. (2.8), (2.12) and (2.13), one can obtain a family of  $P_L$  curves for risk-based design. The family of  $P_L$  curves for magnitude 7.5 earthquakes and soils with  $FC \leq 5$  % is presented in Fig. 2.9.

It is important to note that Figs. 2.8 and 2.9 are developed assuming  $F_S$  to be a fixed variable, and possible variations in *CRR* and *CSR* are not considered directly. Previous studies by Juang et al. (2000b) and Chen and Juang (2000) concluded that practically the same  $P_L$ - $F_S$  relationship would be obtained even if the uncertainties in *CSR* and *CRR* were incorporated in the formulation for  $P_L$ . Thus, in general, if the variations of *CRR* and *CSR* are not too great, the figures can be used directly without considering the variations (Juang et al., 2001b).

# 2.6 SUMMARY

In this chapter, guidelines are presented for evaluating liquefaction resistance through  $V_s$  measurements using the procedure outlined in Andrus and Stokoe (2000). The procedure can be summarized in the following ten steps:

- 1. From available subsurface data, develop detailed profiles of  $V_s$ , soil type, fines content and, if possible, soil density and penetration resistance. Identify the depth of the ground water table, noting any seasonal fluctuations and artesian pressures.
- 2. Calculate the values of  $\sigma_{\nu}$  and  $\sigma'_{\nu}$  for each measurement depth at which seismic testing has been performed.
- 3. Correct the  $V_S$  measurements to the reference overburden stress of 100 kPa using Eq. (2.7). The correction factor  $C_{VS}$  is limited to a maximum value of 1.4 at shallow depths.



Fig. 2.8 - Suggested Relationship for Selecting  $F_s$  Based on Probability of Liquefaction and the Recommended  $CRR-V_{sl}$  Curves. (Juang et al., 2002)



Fig. 2.9 - Curves Suggested for Probability-Based Evaluation in Clean Soils. Note that  $P_L = 0.26$  Corresponds to the Recommended Deterministic Curve Shown in Fig. 2.3. (Juang et al., 2002)

- 4. Determine the value of  $V_{S1}^*$  for each measurement depth using Eq. (2.11) which is recommended for sandy as well as gravely soils. If the fines content is unknown, assume 215 m/s for  $V_{S1}^*$ .
- 5. Determine the value of  $K_c$ .  $K_c$  can be assumed equal to 1, if the soil to be evaluated is uncemented and less than 10 000 years old. If the soil conditions are unknown and penetration data are not available, assume 0.6 for  $K_c$ .
- 6. Determine the design earthquake magnitude and expected value of  $a_{max}$ .
- 7. Calculate CSR for each measurement depth below the water table using Eq. (2.1). The value of  $r_d$  can be estimated from the average curve originally proposed by Seed and Idriss (1971).
- 8. Calculate CRR for each value of  $V_{SI}$  using Eqs. (2.8) and (2.9) with n = -2.56. It is important to note that Eq. (2.8) is for extreme behavior where boils and ground cracks occur.
- 9. Calculate the value of  $F_s$  for each value of  $V_{sl}$  using Eq. (2.12). By convention, liquefaction is predicted to occur when  $F_s \le 1$ , and not to occur when  $F_s > 1$ .
- 10. Plot the values of  $V_{SI}$ , CSR, CRR and  $F_S$  to visually note how they vary with depth, and how many points fall in the regions of liquefaction and no liquefaction.

The deterministic  $V_s$ -based procedure outlined above is characterized with an average probability of liquefaction of 26 %. In other words, a soil with a calculated  $F_s = 1$  has a 26 % chance of liquefaction occurrence based on the case histories analyzed in this study. As mentioned previously, the  $V_s$ -based procedure is as conservative as the SPT-based procedure by Seed et al. (1985). A factor of safety of 1.2 to 1.5, as suggested by the Building Seismic Safety Council (1997, page 158), is considered appropriate for design of typical buildings using the SPT- and  $V_s$ -based procedures. This range of factor of safety corresponds to a probability of liquefaction of about 8 % (for  $F_s = 1.5$ ) to 16 % (for  $F_s = 1.2$ ). The acceptable value of  $F_s$  for a particular site will depend on several factors, including the type and importance of structure and the potential for ground deformation. For critical structures, a smaller probability of liquefaction might be required. Equation (2.13) provides an important link between  $F_s$  and  $P_L$ , and is suggested for probability liquefaction evaluations. ł

# CHAPTER 3

### **APPLICATION OF THE LIQUEFACTION EVALUATION PROCEDURE**

To illustrate the application of the liquefaction evaluation procedure described in Chapter 2, two sites shaken by the 1989 Loma Prieta, California, earthquake  $(M_w = 7)$  are considered below. The two sites are Treasure Island Fire Station and Marina District Winfield Scott School.

## 3.1 TREASURE ISLAND FIRE STATION

Treasure Island is a man-made island located in the San Francisco Bay along the Bay Bridge between the cities of San Francisco and Oakland. It was constructed in 1936-37 by hydraulic filling behind a perimeter rock dike. The perimeter dike served to contain the hydraulic fill and was raised in sections over the previously placed fill. In 1991, Treasure Island was selected as a national geotechnical experimentation site. Much of the work to date at the Treasure Island national geotechnical experimentation site centers arounds a ground response experiment (de Alba and Faris, 1996) with six accelerometers and eight piezometers operating at various elevations near the fire station.

Extensive field tests have been conducted near the Treasure Island fire station to characterize ground conditions. Figures 3.1 and 3.2 present  $V_s$  and general soil profiles for the site. The  $V_s$  profile shown in Fig. 3.1(a) is from Fuhriman (1993), and was determined by crosshole testing. The  $V_s$  profile shown in Fig. 3.2(a) is based on unpublished SASW test results by The University of Texas at Austin in 1992. From the description by de Alba et al. (1994), the upper 4.5 m of soil consists of silty sand fill, possibly formed by dumping. Between depths of 4.5 m and 12.2 m, the soil consists of silty sand to clayey sand, formed by hydraulic filling. Beneath the hydraulic fill are natural clayey soils. The ground water table lies near the ground surface at a depth of 1.4 m.

During the 1989 Loma Prieta earthquake, a seismograph station at the fire station recorded ground surface accelerations. Unlike recordings at other seismograph stations located on soft-soils in the Bay area, there is a sudden drop in the recorded acceleration at about 15 seconds and small motion afterward (Idriss, 1990). De Alba et al. (1994) attribute this behavior to liquefaction of an underlying sand layer, although no sand boils or ground cracks occurred at the site. The nearest liquefaction effect observed is a sand boil located 100 m from the site (Geometric Consultants, 1990; Bennett, 1994; Power et al., 1998).



of 1.5 m to 14 m).

30



Fig. 3.2 - Application of the Recommended Procedure to the Treasure Island Fire Sation Site, SASW Test Array (Depths of 2 m to 13 m). Figure 3.1 presents the liquefaction evaluation for the crosshole test array B1-B4 and the 1989 Loma Prieta earthquake. Values of  $V_{SI}$  and CSR shown in Figs. 3.1(a) and 3.1(d), respectively, are calculated assuming densities of 1.76 Mg/m<sup>3</sup> above the water table and 1.92 Mg/m<sup>3</sup> below the water table. Based on peak values of 0.16 g and 0.11 g recorded in two horizontal directions at the fire station during the 1989 earthquake (Brady and Shakal, 1994), a geometric mean value of 0.13 g is used to calculate CSR. Stress reduction coefficients are estimated using the average curve by Seed and Idriss (1971) shown in Fig. 2.1.

For the crosshole measurement at a depth of 4.6 m, values of CSR and  $V_{Sl}$  are calculated as follows:

$$CSR = 0.65 \left(\frac{a_{\text{max}}}{g}\right) \left(\frac{\sigma_{\nu}}{\sigma'_{\nu}}\right) r_{d} = 0.65 \left(\frac{0.13g}{g}\right) \left(\frac{84.0}{52.7}\right) 0.97 = 0.131$$
(3.1)

and

$$V_{SI} = V_S \left(\frac{P_a}{\sigma'_v}\right)^{0.25} = 134 \left(\frac{100}{52.7}\right)^{0.25} = 158 \text{ m/s}$$
 (3.2)

Assuming an average fines content of 24 %, from Fig. 3.1(c), and a  $K_c$  value of 1, the values of  $V_{S1}^*$ , CRR, and  $F_S$  are calculated by:

$$V_{S1}^* = 215 - 0.5(FC-5) = 215 - 0.5(24-5) = 206 \text{ m/s}$$
 (3.3)

and

$$CRR = \left\{ a \left( \frac{K_c V_{S1}}{100} \right)^2 + b \left( \frac{1}{V_{S1}^* - K_c V_{S1}} - \frac{1}{V_{S1}^*} \right) \right\} MSF$$

$$= \left\{ 0.022 \left( \frac{158}{100} \right)^2 + 2.8 \left( \frac{1}{206 - 158} - \frac{1}{206} \right) \right\} \left( \frac{7}{7.5} \right)^{-2.56}$$

$$= 0.119$$

$$(3.4)$$

and

$$F_{S} = \frac{CRR}{CSR} = \frac{0.119}{0.131} = 0.91$$
(3.5)

Since the value of  $F_s$  is less than 1, liquefaction is predicted at this depth.

Values of  $F_s$  shown in Fig. 3.1(e) are less than 1 for the depths of 4 m to about 9 m. Between the depths of 4 m and 7 m, the sand contains non-plastic fines and is considered liquefiable. Between the depths of 7 m and 9 m, the soil exhibits plastic characteristics and may be non-liquefiable by the simple clay criteria (see Section 1.1). Thus, the layer most likely to liquefy, or the critical layer, lies between the depths of 4 m and 7 m.

Figure 3.2 presents the liquefaction evaluation for the SASW test array. Locations of  $V_s$  measurements for the SASW test array are assumed at the center of the layer used in forward modeling of surface wave measurements. Values of  $F_s$  shown in Fig. 3.2(e) are less than 1 between the depths of about 3.5 m and 11 m. The lowest values of  $F_s$  in the non-plastic soil is 0.75 at a depth of 5.3 m. This  $F_s$  value is similar to the lowest  $F_s$  value of 0.77 determined from crosshole measurements in the critical layer.

Figures 3.3 and 3.4 present the liquefaction evaluations directly on the recommended liquefaction assessment chart for the crosshole test array and SASW test array, respectively. Plots of this type are particularly useful in comparing the range and distribution of  $V_{SI}$  and CSR values with the case histories used to develop the assessment chart. Based on Fig. F.15, the assessment chart is well supported within the range of the Treasure Island Fire Station data.

Although no sand boils or ground cracks occurred at the fire station during the 1989 earthquake, the prediction of liquefaction agrees with the conclusion stated above that liquefaction of an underlying sand cause the sudden drop in the acceleration time histories recorded at this site (de Alba et al., 1994). A similar sudden drop in the strong ground motion recordings occurred at the Port Island Downhole Array site in Kobe, Japan, during the 1995 Hyogo-ken Nanbu earthquake (Aguirre and Irikura, 1997), where liquefaction and sand boils did occur. It is possible that the 4 m thick layer capping the site, predicted not to liquefy, as shown in Figs. 3.1(e) and 3.2(e), prevented the formation of sand boils at the ground surface (Ishihara, 1985).

## 3.2 MARINA DISTRICT WINFIELD SCOTT SCHOOL

Kayen et al. (1990) conducted downhole seismic tests at the Winfield Scott School in the Marina District of San Francisco. Figures 3.5(a) and 3.5(b) present the  $V_s$  and soil profiles for the site. The  $V_s$  profile was originally determined based on best-fit line segments through travel time measurements plotted versus depth. However, the layering assumed in the best-fit segment method did not seem appropriate for the fill. Figures 3.5(a) presents the  $V_s$  profile for the site determined using the pseudo-interval method (see Appendix E). Figure 3.5(c) presents a profile of fines content that are based on information provided by Kayen et al. (1990). The upper 7.6 m of soil at the site consists of sand with 1 % to 8 % fines. The ground water table lies at a depth of 2.7 m.



Fig. 3.3 - Liquefaction Assessment Chart for Magnitude 7 Earthquakes with Data for the 1989 Loma Prieta Earthquake and the Treasure Island Fire Station Site, Crosshole Test Array B1-B4 (Depths of 1.5 m to 14 m).



Fig. 3.4 - Liquefaction Assessment Chart for Magnitude 7 Earthquakes with Data for the 1989 Loma Prieta Earthquake and the Treasure Island Fire Station Site, SASW Test Array (Depths of 2 m to 13 m).



Fig. 3.5 - Application of the Recommended Procedure to the Marina District School Site (Depths of 3 m to 10 m).

Many structures, pavements, and public works near the school sustained heavy damage during the 1989 Loma Prieta earthquake (Kayen et al., 1990). This damage was due to liquefaction of the sand fill. From maps prepared by Pease and O'Rourke (1995), the site lies on the margin of the 1906 water front and artificial fill where about 40 mm of settlement occurred. Mapped sand boils and ground cracks lie just east of the site. Based on these observations, this site is classified as a liquefaction site during this earthquake.

The Marina District and Treasure Island are located about 82 km from the 1989 surface fault rupture. Assuming a distance of 82 km from the fault rupture, the attenuation relationship by Idriss (1991) for 1989 strong ground motion records from soft-soil sites provides a median value of 0.16 g. This value is slightly higher than the geometric mean value of 0.13 g for the two peak horizontal accelerations recorded at Treasure Island fire station. Thus, a peak horizontal ground surface acceleration of 0.15 g, the average of these two estimates, is assumed in the analysis.

Figure 3.5 presents the liquefaction evaluation for the Marina District School site and the 1989 Loma Prieta earthquake. Values of  $V_{SI}$  and CSR are calculated assuming densities of 1.76 Mg/m<sup>3</sup> above the water table and 1.92 Mg/m<sup>3</sup> below the water table. They are plotted in Fig. 3.5(a) and 3.5(d) at the depths midway between receiver locations. Since the sand is uncemented and less than 10 000 years old, the value of  $K_c$  is 1. Calculated values of  $F_s$  are 0.42, 0.90, and 0.51 at the depths of 3 m, 4 m, and 6.7 m within the sand fill. The silty clay layer beneath the sand fill is non-liquefiable by the simple clay criteria (see Section 1.1). Thus, having the lowest average value of  $F_s$ , the sand fill just below the water table between the depths of 2.7 m and 4.4 m is identified as the critical layer that liquefied. A prediction of liquefaction agrees with the observed field behavior.

Figure 3.6 present the liquefaction evaluation directly on the recommended liquefaction assessment chart for the Marina District. Based on Fig. F.15, the assessment chart is well supported within the range of the Marina District data.



Fig. 3.6 - Liquefaction Assessment Chart for Magnitude 7 Earthquakes with Data for the 1989 Loma Prieta Earthquake and the Marina District School Site (Depths of 3 m to 7 m).

## CHAPTER 4

### SUMMARY AND RECOMMENDATIONS

### 4.1 SUMMARY

Presented in this report are guidelines for evaluating the liquefaction resistance of soils through shear wave velocity,  $V_s$ , measurements. The guidelines are based on an earlier report entitled "Draft Guidelines for Evaluating Liquefaction Resistance Using Shear Wave Velocity Measurements and Simplified Procedures." From comments received on the earlier report, the draft guidelines are updated in this report. The guidelines present a recommended procedure for evaluating soil liquefaction resistance and guidance for its use.

The recommended procedure follows the general format of the simplified penetrationbased procedure originally proposed by Seed and Idriss (1971). Cyclic stress ratios, CSR, are calculated using Eq. (2.1), with the average stress reduction coefficient estimated from Fig. 2.1. Shear wave velocity measurements are corrected for overburden stress using Eq. (2.7). Figure 2.3 presents the recommended evaluation curves for uncemented, Holocene-age soils and magnitude 7.5 earthquakes. These curves are defined by Eq. (2.8) with MSF = 1,  $V_{S1}^* = 200$ m/s to 215 m/s (depending on fines content), and  $K_c = 1$ . Equation (2.8) can be adjusted for other magnitude earthquakes using MSF values defined by Eq. (2.9) with n = -2.56. Corrections for cemented and aged soils are suggested in Section 2.3.4. A ten-step summary of the procedure is given in Section 2.6.

The recommended liquefaction evaluation curves, defined by Eq. (2.8), are based on a modified relationship between overburden stress-corrected shear wave velocity,  $V_{SI}$ , and CSR for constant average cyclic shear strain suggested by R. Dobry. As discussed in Section 2.3 and Appendix F, the quadratic relationship proposed by Dobry is modified so that it is asymptotic to some limiting upper value of  $V_{SI}$ . This limit is related to the tendency of dense granular soils to exhibit dilative behavior at large strains, as well as the fact that dense soils expel dramatically less water during reconsolidation than loose soils. Liquefaction and non-liquefaction case histories from 26 earthquakes and more than 70 measurement sites in soils ranging from clean fine sand to sandy gravel with cobbles to profiles including silty clay layers are analyzed to determine the parameters of Eq. (2.8). Penetration- $V_S$  correlations are also considered. The evaluation curves correctly bound over 95 % of the case histories where liquefaction occurred.

By constructing relationships between  $V_{SI}$  and penetration resistance from the recommended evaluation curves and plotting available *in situ* test data, it is shown that the  $V_{SI}$ -based evaluation curves are generally more conservative than the penetration-based evaluation curves. From logistic regression and Bayesian interpretation techniques (see Appendix G), the recommended curve for clean soils is characterized with an average probability of 26 %.

Caution should be exercised when applying the procedure to sites where conditions are different from the case history data. The case history data used to develop the procedure are limited to relatively level ground sites with the following general characteristics: (1) uncemented soils of Holocene age; (2) average depths less than about 10 m; and (3) ground water table depths between 0.5 m and 6 m. All  $V_S$  measurements are from below the water table. About three-quarters of the case history data are for soils with fines content greater than 5 %. Almost half of the case histories are for earthquakes with magnitudes near 7.

Three concerns when using shear wave velocity as an indicator of liquefaction resistance are (1) its higher sensitivity (when compared with the penetration-based methods) to weak interparticle bonding, (2) the lack of a physical sample for identifying non-liquefiable clayey soils, and (3) not detecting thin liquefiable strata because the test interval is too large. The preferred practice is to drill sufficient boreholes and conduct sufficient other *in situ* tests to detect thin liquefiable strata, identify non-liquefiable clay-rich soils, identify soils above the ground water table that might have lower values of  $V_s$  should the water table rise, and detect liquefiable weakly cemented soils.

# 4.2 FUTURE STUDIES

The following future studies are recommended:

1. Additional well-documented case histories with all types of soil that have and have not liquefied during earthquakes should be compiled, particularly from deeper deposits (depth > 8 m) and from denser soils ( $V_s > 200$  m/s) shaken by stronger ground motions ( $a_{max} > 0.4$  g), to further validate the recommended curves. Also, case histories from lower magnitude earthquakes ( $M_w < 7$ ) may improve estimates of the magnitude scaling factor.

2. Laboratory and field studies should be conducted to further refine estimates of  $V_{S1}^*$ , the limiting value of  $V_{S1}$  for cyclic liquefaction occurrence. For example, careful laboratory studies may identify more clearly the influence of fines content, gravel content, and particle gradation on  $V_{S1}^*$ . Additional careful penetration- $V_S$  correlation studies may also help refine the  $V_{S1}^*$  estimates.

3. Laboratory studies should also be conducted to evaluate the implied assumption observed in Fig. 2.3 that at low values of  $V_{SI}$  (say 100 m/s) liquefaction resistance is independent of fines content.

4. Additional work is needed to evaluate the significance of ignoring soil type and horizontal stress in the overburden correction.

5. Standard test procedures exist only for the crosshole test. Standard test methods should be developed for the other *in situ* seismic tests.

# APPENDIX A

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# **APPENDIX B**

# SYMBOLS AND NOTATION

The following symbols and notation are used in this report:

Α	=	parameter that depends on soil structure;
a	=	parameter related to slope of CRR-V <sub>SI</sub> curve;
$a_1, a_2, a_3$		regression coefficients;
a <sub>max</sub>	=	peak horizontal ground surface acceleration;
$B_{1}, B_{2}$	=	parameters relating $V_{SI}$ and penetration resistance;
b	=	parameter related to slope of CRR-V <sub>S1</sub> curve;
$b_1, b_2, b_3, b_4$	=	regression coefficients;
CRR	=	average cyclic resistance ratio;
$CRR_{tx}$	=	CRR for cyclic triaxial tests;
CRRI	=	CRR corrected for high overburden stress;
CRR <sub>7.5</sub>	=	CRR for magnitude 7.5 earthquakes;
CSR		cyclic stress ratio;
$C_{VS}$	=	factor to correct $V_s$ for overburden pressure;
D50	=	median grain size by mass;
DA		double-amplitude axial strain;
Dr	=	relative density;
exp	=	the constant e raised to the power of a given number;
$F_{1}, F_{2}$	=	age and soil type factors for correlating $V_S$ and $N_j$ ;
FC	=	fines content (particles smaller than 75 $\mu$ m);
$F_{S}$	=	factor of safety;
f	=	high overburden stress exponent;
$f(e_{min})$	=	function of minimum void ratio;
$f(V_{Sl})$	=	function of $V_{SI}$ ;
f(x)	=	function of $x (= F_s)$ ;
$f(\gamma_{av})$	=	function of average peak cyclic shear strain;
$f_L(F_S)$	=	probability density function of calculated $F_s$ for
		liquefaction case histories;
$f_{NL}(F_S)$	=	probability density function of calculated $F_s$ for
		non-liquefaction case histories;

G	=	shear modulus;
$G_{max}$	=	small-strain shear modulus;
$G_N$	=	$G_{max}$ corrected for confining stress and void ratio;
$(G)_{\gamma_{av}}$	=	secant shear modulus at $\gamma_{av}$ ;
g	=	acceleration of gravity;
Kc	=	cementation and aging correction factor;
K <sub>fc</sub>	=	fines content correction factor;
K'o		coefficient of effective lateral earth pressure at rest;
$K_{\sigma}$	=	high overburden stress correction factor;
ln	=	natural logarithm function;
MSF	=	magnitude scaling factor;
$M_w$		earthquake moment magnitude;
m	=	stress exponent;
$N_j$	=	SPT blow count in Japanese practice;
$N_{60}$	=	SPT energy-corrected blow count;
$(N_1)_{60}$	=	SPT energy- and overburden stress-corrected blow count;
n	==	magnitude scaling factor exponent;
Pa		reference overburden stress (= 100 kPa),
$P_L$	=	probability of liquefaction occurrence
r <sub>c</sub>	. =	factor to account for effects of multidirectional shaking;
r <sub>d</sub>	=	shear stress reduction coefficient;
sin		sine function;
Sres	=	residual standard deviation;
$V_S$	=	small-strain shear wave velocity;
$V_{SI}$		overburden stress-corrected V <sub>s</sub> ;
$V_{SI,CS}$		equivalent clean soil value of $V_{SI}$ ;
$V_{S1}^{*}$	=	limiting upper value of $V_{SI}$ for liquefaction occurrence;
$V_{SIm}$	=	mean stress-corrected $V_S$ ;
Z	=	depth;
$\alpha(z)$		function of depth;
$\beta(z)$	=	function of depth;
Yav	=	average peak cyclic shear strain;
μ	=	mean;
ρ	=	mass density of soil;
σ	=	standard deviation;
$\sigma_d$	=	cyclic deviator stress in cyclic triaxial tests;
$\sigma_h$	-	initial effective horizontal confining stress;
$\sigma_m$	=	mean effective confining stress;
σ'。	=	initial effective confining stress in cyclic triaxial tests:

$\sigma_{v}$	=	total vertical (or overburden) stress;
$\sigma'_v$	=	initial effective vertical (or overburden) stress;
$\tau_{av}$	=	average cyclic equivalent uniform shear stress generated by
		earthquake; and
T <sub>max</sub>	=	maximum cyclic shear stress generated by earthquake.

,

# **APPENDIX C**

# **GLOSSARY OF TERMS**

The following definitions apply to this report:

Case History	An earthquake and a test array.
Critical Layer	The layer of non-plastic soil below the ground water table where corrected values of shear wave velocity and penetration are the least, and where cyclic stress ratios are the greatest.
Liquefaction Occurrence	Surface manifestations of excess pore-water pressure at depth, such as sand boils, ground cracks and fissures, and ground settlement.
Moment Magnitude	An earthquake magnitude scale defined in terms of energy.
Overburden Stress- Corrected Shear Wave Velocity	Shear wave velocity measurement corrected to a reference vertical (or overburden) stress of 100 kPa.
Peak Horizontal Ground Surface Acceleration	The peak value in a horizontal ground surface acceleration record that would occur at the site in the absence of liquefaction or excess pore-water pressures.
Shear Wave Velocity	The velocity of a propagating shear wave within a material with either the direction of wave propagation or the direction of particle motion in the vertical direction.
Shear Wave	A body wave with the direction of particle motion transverse to the direction of wave propagation.

Test Array The two boreholes used for crosshole measurements, the borehole and source used for downhole measurements, the cone sounding and source used for seismic cone measurements, the borehole used for suspension logging measurements, or the line of receivers used for Spectral-Analysis-of-Surface-Waves (SASW) measurements.

#### APPENDIX D

## COMPARISON OF Vs-BASED LIQUEFACTION RESISTANCE CURVES

During the past two decades, several studies have been conducted to investigate the relationship between  $V_s$  and liquefaction resistance. These studies involved laboratory tests (Dobry et al., 1981; Dobry et al., 1982; de Alba et al., 1984; Hynes, 1988; Tokimatsu and Uchida, 1990; Tokimatsu et al., 1991a; Rashidian, 1995; Rauch et al., 2000), analytical investigations (Bierschwale and Stokoe, 1984; Stokoe et al., 1988c; Andrus, 1994), penetration- $V_s$  correlations (Seed et al., 1983; Lodge, 1994; Kayabali, 1996; Rollins et al., 1998b; Andrus et al., 1999), or field performance observations (Stokoe and Nazarian, 1985; Robertson et al., 1992; Kayen et al., 1992; Andrus and Stokoe, 1997; Andrus et al., 1999; Juang and Chen, 2000; Andrus and Stokoe, 2000; Juang et al., 2001a). Many of the liquefaction evaluation procedures developed from these studies follow the general format of the Seed-Idriss simplified procedure, where  $V_s$  is corrected to a reference overburden stress and correlated with the cyclic stress, or resistance, ratio.

This appendix reviews seven proposed liquefaction evaluation curves based on CRR and  $V_{SI}$ . The seven CRR- $V_{SI}$  curves are shown in Fig. D.1. Each of the curves is briefly discussed below.

#### D.1 CURVE BY TOKIMATSU AND UCHIDA (1990)

The best-fit curve by Tokimatsu and Uchida (1990) shown in Fig. D.1 was determined from laboratory cyclic triaxial test results for various sands with less than 10 % fines (silt and clay) and 15 cycles of loading. Figure D.2 presents the cyclic triaxial test results. The solid symbols in Fig. D.2 correspond to specimens obtained by the *in situ* freezing technique. The open symbols correspond to specimens reconstituted in the laboratory. Tokimatsu and Uchida defined the cyclic resistance ratio for cyclic triaxial tests,  $CRR_{tx}$ , as the ratio of cyclic deviator stress to initial effective confining stress,  $\sigma_d/2\sigma_o'$  when the double-amplitude (or peak-to-peak)



Fig. D.1 - Comparison of Seven Proposed CRR-V<sub>S1</sub> Curves for Clean Granular Soils.



Fig. D.2 - Relationship Between Liquefaction Resistance and Normalized Shear Modulus for Various Sands with Less than 10 % Fines Determined by Cyclic Triaxial Testing. (modified from Tokimatsu and Uchida, 1990)

axial strain, DA, reaches 5 %. They measured the elastic shear modulus of the specimen at a shear strain of  $10^{-3}$  % just prior to the liquefaction test. This small-strain shear modulus was normalized to correct for the influence of confining pressure and void ratio by:

$$G_{N} = \frac{G_{\max}}{f(e_{\min})(\sigma'_{m})^{2/3}}$$
(D.1)

and

$$f(e_{\min}) = \frac{(2.17 - e_{\min})^2}{1 + e_{\min}}$$
(D.2)

where

$G_N$	=	the normalized shear modulus,
e <sub>min</sub>	=	the minimum void ratio determined by standard test method, and
$\sigma'_m$	=	the mean effective confining stress.

Tokimatsu and Uchida selected an exponent of 2/3 rather than 1/2, as determined by Hardin and Drnevich (1972), because it seemed that a slightly better correlation could be obtained. Values of  $e_{min}$  ranged from 0.61 to 0.91 for the sands tested. The actual values of void ratio in each test were greater than  $e_{min}$ , with values ranging from about 0.65 to about 1.4.

By combining Eqs. (1.1) and (D.1), one obtains the following relationship for converting  $G_N$  to mean stress-corrected  $V_S$ :

$$V_{SIm} = V_S \left(\frac{1}{\sigma'_m}\right)^{0.33} = \left(\frac{G_N f(e_{\min})}{\rho}\right)^{0.5}$$
(D.3)

where

 $V_{SIm}$  = mean stress-corrected  $V_s$ , and  $\sigma'_m$  = the mean effective confining stress in kgf/cm<sup>2</sup> (1 kgf/cm<sup>2</sup> = 98 kPa).

Tokimatsu and Uchida (1990) suggested using 0.65 as an average value of  $e_{min}$  for clean sands.

The overburden stress-corrected  $V_s$  and  $V_{sim}$  can be related by:

$$V_{SIm} = V_{S} \left(\frac{1}{\sigma'_{\nu}}\right)^{0.33} \left(\frac{3}{1+2K'_{o}}\right)^{0.33} \approx V_{SI} \left(\frac{1}{\sigma'_{\nu}}\right)^{0.08} \left(\frac{3}{1+2K'_{o}}\right)^{0.33}$$
(D.4)

where

 $K'_o$  = the coefficient of lateral earth pressure at rest  $(=\sigma'_h/\sigma'_v)$ .

Values of  $V_{SI}$  for the best fit curve by Tokimatsu and Uchida (1990) shown in Fig. D.1 are determined (Andrus et al., 1999; after Tokimatsu et al., 1991a) from Fig. D.2 using Eqs. (D.3) and (D.4), and assuming  $K'_o = 0.5$ ,  $e_{min} = 0.65$ ,  $\sigma'_m = 100$  kPa, and soil density of 1.9 Mg/m<sup>3</sup>.

For converting  $CRR_{tx}$  to an equivalent field CRR, Tokimatsu and Uchida (1990) suggested the following expression originally proposed by Seed (1979):

$$CRR = \frac{(1+2K'_o)}{3} r_c (CRR_{tr})$$
(D.5)

where

r<sub>c</sub>

= a constant to account for the effects of multi-directional shaking with a value between 0.9 and 1.0.

Values of *CRR* for the best fit curve by Tokimatsu and Uchida shown in Fig. D.1 are determined from Fig. D.2 using Eq. (D.5) and assuming  $K'_o = 0.5$  and  $r_c = 0.9$ .

Because the other liquefaction resistance curves shown in Fig. D.1 were drawn to bound liquefaction case histories, the more conservative lower bound curve for the laboratory test results by Tokimatsu and Uchida (1990) also is shown. This curve was drawn (Andrus et al., 1999) from Fig. D.2 following the procedure outlined above.

## D.2 CURVE BY ROBERTSON ET AL. (1992)

The bounding curve by Robertson et al. (1992) was developed using field performance data from primarily sites in Imperial Valley, California, along with data from four other sites, as shown in Fig. D.3. The soil at these sites contained as much as 35 % fines. Robertson et al. corrected  $V_s$  using Eq. (2.7). The shape of their curve was based on the analytical results of



Fig. D.3 - Liquefaction Resistance Curve for Magnitude 7.5 Earthquakes and Case History Data from Robertson et al. (1992).

Bierschwale and Stokoe (1984). They reasoned that the curve should pass close to the Imperial Valley (Wildlife site) data point, since liquefaction did and did not occur at this site during the 1987 Superstition Hills ( $M_w = 6.5$ ) and Elmore Ranch ( $M_w = 6.2$ ) earthquakes, respectively. Robertson et al. used magnitude scaling factors similar to those suggested by Seed and Idriss (1982), Column 2 of Table 2.2, to position their curve for magnitude 7.5 earthquakes.

## D.3 CURVE BY KAYEN ET AL. (1992)

Kayen et al. (1992) studied four sites that did and did not liquefy during the 1989 Loma Prieta, California, earthquake ( $M_w = 7.0$ ). The four sites are: Port of Richmond, Bay Bridge Toll Plaza, Port of Oakland, and Alameda Bay Farm Island South Loop Road. The fines content for soils at these sites ranged from less than 5 % to as much as 57 %. Values of  $V_S$  were measured by the SCPT method and corrected for overburden stress using Eq. (2.7). Figure D.4 presents their data and bounding curve. The curve by Kayen et al. shown in Fig. D.1 was adjusted for magnitude 7.5 earthquakes by assuming a *MSF* of 1.19 (see Column 3 of Table 2.2).

#### D.4 CURVE BY LODGE (1994)

Lodge (1994) considered the same sites that Kayen et al. (1992) studied, as well as other sites shaken by the 1989 Loma Prieta earthquake. The curve by Lodge was developed as follows. First, cyclic stress ratios for the entire soil profile at each site were calculated. Second, available SPT blow counts were corrected for overburden pressure and energy. Soil layers with high and low liquefaction potential were identified with the procedure of Seed et al. (1985). Soil layers with corrected blow count within 3 of the SPT-based curve were eliminated due to uncertainties in the correlation. Third,  $V_s$  measurements from SCPT and crosshole tests were corrected for overburden stress using Eq. (2.7). Fourth, on a "meter by meter" basis, values of  $V_{Sl}$  and cyclic stress ratio were plotted for both layer types, those which were predicted liquefiable and those which were predicted non-liquefiable. Fifth, published data for sites shaken by the 1983 Borah Peak, Idaho, and 1964 Niigata, Japan, earthquakes were added to the plot. Finally, a curve was drawn to include all liquefiable layers, as shown in Fig. D.5. The curve by Lodge shown in Fig. D.1 was adjusted for magnitude 7.5 earthquakes by assuming a *MSF* of 1.19 (see Column 3 of Table 2.2).



Fig. D.4 - Liquefaction Resistance Curve for Magnitude 7 Earthquake and Case History Data from Kayen et al. (1992).



Fig. D.5 - Liquefaction Resistance Curve for Magnitude 7 Earthquakes and Case History Data from Lodge (1994).

#### **D.5 CURVE BY ANDRUS AND STOKOE (1997)**

The curve by Andrus and Stokoe (1997) shown in Fig. D.1 was developed for the proceedings of the 1996 NCEER Workshop (Youd and Idriss, eds., 1997). Several suggestions were offered at, and after, the workshop concerning how site variables should be define, as well as the shape of the boundary curve separating liquefaction and no liquefaction. Following the suggestions and using field performance data from 20 earthquakes and *in situ*  $V_s$  measurements from over 50 sites in soils ranging from clean fine sand to sandy gravel with cobbles to profiles including silty clay layers, Andrus and Stokoe constructed curves for uncemented, Holocene-age soils with various fines content. The values of  $V_s$  were corrected using Eq. (2.7). The curve by Andrus and Stokoe (1997) for fines content  $\leq 5$  % along with the case history data are presented in Fig. D.6. The shape of the curve by Andrus and Stokoe (1997) was based on a modified relationship between  $V_{SI}$  and CSR for constant average cyclic shear strain suggested by R. Dobry.

## **D.5.1** Cyclic Shear Strain and Cyclic Shear Stress

Liquefaction results from the rearranging of soil particles and the tendency for decrease in volume. Experimental and theoretical studies show that decrease in volume is more closely related to cyclic strain than cyclic stress (Silver and Seed, 1971); a threshold cyclic strain exists below which neither rearrangement of soil particles nor decrease in volume take place (Drnevich and Richart, 1970; Youd, 1972; Pyke et al., 1975), and no pore water pressure buildup occurs (Dobry et al., 1981; Seed et al., 1983); and that there is a predictable correlation between cyclic shear strain and pore pressure buildup of saturated soils (Martin et al., 1975; Park and Silver, 1975; Finn and Bhatia, 1981; Dobry et al., 1982; Hynes, 1988). The threshold cyclic strain is limited to a narrow range of variation, ranging from about 0.005 % for gravels to 0.01 % for normally consolidated clean sands and silty sands to 0.03 % for overconsolidated clean sands. In addition, cyclic strain-controlled test results are less affected than stress-controlled tests by factors such as density, confining stress, anisotropic confining stress, fabric and prestaining (Martin et al., 1975; Dobry and Ladd, 1980; Dobry et al., 1982; Hynes, 1988). It should also be noted that the steady state approach to liquefaction evaluation by Poulos et al. (1985) is based These findings confirm the fact that cyclic strain is more on a triggering strain level. fundamentally related to pore pressure buildup than cyclic stress, and are strong arguments in favor of a cyclic strain approach to liquefaction evaluation.



Fig. D.6 - Liquefaction Resistance Curve for Magnitude 7.5 Earthquakes and Uncemented Clean Soils of Holocene Age with Case History Data from Andrus and Stokoe (1997).

Cyclic shear strain and cyclic shear stress can be related by the following equation:

$$\gamma_{av} = \frac{\tau_{av}}{(G)_{\gamma_{av}}} \tag{D.6}$$

where

 $\gamma_{av}$  = the average peak cyclic shear strain during a cyclic stress-controlled test of uniform cyclic shear stress  $\tau_{av}$ , and

 $(G)_{\gamma_{-}}$  = the secant shear modulus at  $\gamma_{av}$  during the same cyclic test.

In the cyclic strain approach proposed by Dobry et al. (1982), the average cyclic shear strain caused by an earthquake is estimated from:

$$\gamma_{av} = 0.65 \frac{a_{\max}}{g} \frac{\sigma_v r_d G_{\max}}{\rho V_S^2(G)_{\gamma_{av}}}$$
(D.7)

Equation (D.7) is obtained by combining Eqs. (1.1), (2.1) and (D.6). The variation of shear modulus with strain is commonly expressed in terms of  $(G)_{\gamma_{av}}/G_{max}$ , called the modulus reduction factor. The modulus reduction factor can be estimated from an experimentally determined correlation. Neither pore pressure buildup nor liquefaction will occur when  $\gamma_{av}$  is less than the threshold strain. When  $\gamma_{av}$  is greater than the threshold strain, then pore pressure buildup can occur. The amount of pore pressure buildup can also be estimated from an experimentally determined correlation.

### **D.5.2** Dobry's Relationship Between CRR and $V_{SI}$

R. Dobry (personal communication to R. D. Andrus, 1996) derived a relationship between  $V_{SI}$  and CSR for constant average cyclic shear strain using Eqs. (1.1) and (D.6). Combining Eqs. (1.1) and (D.6), and dividing both sides by  $\sigma'_{\nu}$  leads to:

$$\frac{\tau_{av}}{\sigma'_{v}} = \gamma_{av} \left(\frac{\rho}{\sigma'_{v}}\right) \frac{(G)_{\gamma_{av}}}{G_{\max}} V_{S}^{2}$$
(D.8)

For an overburden stress of 100 kPa,  $V_S = V_{SI}$  and curves of constant average cyclic strain can be expressed by:

$$CSR = \frac{\tau_{av}}{\sigma'_{v}} = f(\gamma_{av}) (V_{S1})^{2}$$
(D.9)

where

$$f(\gamma_{av}) = \gamma_{av} \left(\frac{\rho}{P_a}\right) \frac{(G)_{\gamma_{av}}}{G_{\max}}$$
(D.10)

Since CSR equals CRR at the point separating liquefaction from no liquefaction, Eq. (D.9) provides an analytical basis for establishing the CRR- $V_{SI}$  curve at low values of  $V_{SI}$  (say  $V_{SI} \le 125$  m/s) and extending them to zero at  $V_{SI} = 0$ .

## D.5.3 Modified CRR-V<sub>SI</sub> Relationship

Andrus and Stokoe (1997) reasoned that the curve separating liquefiable and nonliquefiable soils would become asymptotic to some limiting upper value of  $V_{SI}$ . They modified Eq. (D.9) to:

$$CRR = \left\{ a \left( \frac{V_{S1}}{100} \right)^2 + b \left( \frac{1}{V_{S1}^* - V_{S1}} - \frac{1}{V_{S1}^*} \right) \right\} MSF$$
(D.11)

where

$$V_{S1}^*$$
 = the limiting upper value of  $V_{S1}$  for liquefaction occurrence, and   
a, b = curve fitting parameters.

The first term in Eq. (D.11) is a form of Eq. (D.9), assuming  $f(\gamma_{\alpha\nu})$  is independent of initial effective confining pressure and of pore water pressure buildup. The second term is a hyperbola with a small value at low values of  $V_{SI}$ , and a very large value as  $V_{SI}$  approaches  $V_{SI}^*$ .

The curve by Andrus and Stokoe (1997) shown in Figs. D.1 and D.6 is defined by Eqs. (D.11) and (2.9) with a = 0.03, b = 0.9, n = -3.3, and  $V_{S1}^* = 220$  m/s.

#### D.6 CURVE BY ANDRUS ET AL. (1999)

Since the publication of the 1996 NCEER Workshop proceedings (Youd and Idriss, eds., 1997), the case history data compiled by Andrus and Stokoe (1997) have been revised based on new information, and expanded to include field performance data from 26 earthquakes and more than 70 measurements sites. Also, the 1998 MCEER Workshop was held to discuss developments since the 1996 workshop. From the suggestions given at the second workshop and using the expanded database, the curve proposed by Andrus and Stokoe (1997) was revised in the report by Andrus et al. (1999) and paper by Andrus and Stokoe (2000). The revised curve for uncemented soils with fines content  $\leq 5$  % along with the case history data are shown in Fig. D.7. The development of the revised curve is discussed in Appendix F.

#### **D.7 SUMMARY**

Seven proposed curves relating *CRR* and  $V_{SI}$  were discussed in this Appendix. Many of the differences among the seven curves (see Fig. D.1) can be explained by the following four factors: (1) The best-fit curve by Tokimatsu and Uchida (1990) is more of a median curve, while the other curves bound the liquefaction case history data. (2) Portions of the proposed curves are based on limited data, and the investigator(s) have assumed different levels of conservatism. In particular, the curves by Robertson et al. (1992), Kayen et al. (1992), and Lodge (1994) were based on little or no data above  $V_{SI}$  of 200 m/s, and were conservatively drawn in this region. (3) Methods for selecting some site variables and correction factors are different among investigator(s). (4) Some errors exist in the database by Andrus and Stokoe (1997), and lead to more conservative curve than the updated curve by Andrus et al. (1999) above a  $V_{SI}$  value of 150 m/s. Thus, the *CRR-V<sub>SI</sub>* curve proposed by Andrus et al. (1999) for clean soils is recommended because it was based on the largest, most correct case history data set and procedures recommended by the 1996 NCEER Workshop (Youd et al., 1997).



Fig. D.7 - Revised Liquefaction Resistance Curve for Magnitude 7.5 Earthquakes and Uncemented Clean Soils of Holocene Age with Case History Data from Andrus et al. (1999).

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## APPENDIX E

# **CASE HISTORY DATA AND THEIR CHARACTERISTICS**

Shear wave velocity measurements have been made for field liquefaction studies at many sites during the past twenty years. Table E.1 presents a list of over 70 sites and 26 earthquakes that have been investigated. Of the 26 earthquakes listed, 9 occurred in the United States; and the other 15 in Japan, Taiwan, and China. The field performance information for these earthquakes along with the  $V_s$  measurements provides an important opportunity to determine the relationship between liquefaction resistance and  $V_s$  directly from case histories. A detailed summary of available case history data is presented in Appendix H. This appendix describes the site variables and characteristics of the database.

## E.1 SITE VARIABLES AND DATABASE CHARACTERISTICS

### E.1.1 Earthquake Magnitude

Earthquake magnitudes for the 26 earthquakes listed in Table E.1 range from 5.3 to 8.3, based on the moment magnitude scale. Moment magnitude is the scale most commonly used for engineering applications, and is the preferred scale for liquefaction resistance calculations (Youd et al., 1997). When other magnitude scales are reported by the investigator(s), they are converted to  $M_w$  using the relationship of Heaton et al. (1982) shown in Fig. 2.5.

## E.1.2 Shear Wave Velocity Measurement

At the more than 70 investigation sites listed in Table E.1, shear wave velocity measurements were made with 139 test arrays. A test array is defined in this report as the two boreholes used for crosshole measurements, the borehole and source used for downhole measurements, the cone sounding and source used for seismic cone measurements, the borehole used for suspension logging measurements, or the line of receivers used for Spectral-Analysis-of-Surface-Waves (SASW) measurements. Of the 139 test arrays, 39 are crosshole, 21 downhole, 27 seismic cone, 15 suspension logger, 36 SASW, and one is unknown.

Earthquake	Moment Magnitude	Site	Reference
(1)	(2)	(3)	(4)
1906 San Francisco, California	7.7	Coyote Creek; Salinas River (North, South)	Youd & Hoose (1978); Barrow (1983); Bennett & Tinsley (1995)
1957 Daly City, California	5.3	Marina District (2, 3, 4, 5, School)	Kayen et al. (1990); Tokimatsu et al. (1991b); T. L. Youd (personal communication to R. D. Andrus, 1999)
1964 Niigata, Japan	7.5	Niigata City (A1, C1, C2, Railway Station)	Yoshimi et al. (1984; 1989); Tokimatsu et al. (1991a)
1975 Haicheng, China	7.3	Chemical Fiber; Construction Building; Fishery & Shipbuilding; Glass Fiber; Middle School; Paper Mill	Arulanandan et al. (1986)
1979 Imperial Valley, California 1981 Westmorland, California 1987 Elmore Ranch, California 1987 Superstition Hills, California	6.5 5.9 5.9 6.5	Heber Road (Channel fill, Point bar); Kornbloom; McKim; Radio Tower; Vail Canal; Wildlife	Bennett et al. (1981; 1984); Sykora & Stokoe (1982); Youd & Bennett (1983); Bierschwale & Stokoe (1984); Stokoe & Nazarian (1984); Dobry et al. (1992); Youd & Holzer (1994)
1980 Mid-Chiba, Japan 1985 Chiba-Ibaragi-Kenkyo, Japan	5.9 6.0	Owi Island No. 1	Ishihara et al. (1981; 1987)
1983 Borah Peak, Idaho	6.9	Andersen Bar; Goddard Ranch; Mackay Dam Downstream Toe; North Gravel Bar; Pence Ranch	Youd et al. (1985); Stokoe et al. (1988a); Andrus et al. (1992); Andrus (1994)
1986 Event LSST2, Taiwan Event LSST3, Taiwan Event LSST4, Taiwan Event LSST6, Taiwan Event LSST7, Taiwan Event LSST8, Taiwan Event LSST12, Taiwan Event LSST13, Taiwan Event LSST16, Taiwan	5.3 5.5 6.6 5.4 6.6 6.2 6.2 6.2 6.2 7.6	Lotung LSST Facility	Shen et al. (1991); EPRI (1992)
1987 Chiba-Toho-Oki, Japan	6.5	Sunamachi	Ishihara et al. (1989)

Table E.1 - Earthquakes and Sites Used to Establish Liquefaction Resistance Curves

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Earthquake	Moment Magnitude	Site	Reference
(1)	(2)	(3)	(4)
1989 Loma Prieta, California	7.0	Bay Bridge Toll Plaza, Bay Farm Island (Dike, South Loop Road); Port of Oakland; Port of Richmond	Stokoe et al. (1992); Mitchell et al. (1994)
		Coyote Creek; Salinas River (North, South);	Barrow (1983); M. J. Bennett (personal communication to R. D. Andrus, 1995); Bennett and Tinsley (1995)
		Marina District (2, 3, 4, 5, school)	Kayen et al. (1990); Tokimatsu et al. (1991b)
		Moss Landing (Harbor Office, Sandholdt Road, State Beach)	Boulanger et al. (1995); Boulanger et al. (1997)
		Santa Cruz (SC02, SC03, SC04, SC05, SC13, SC14)	Hryciw (1991); Hryciw et al. (1998)
		Treasure Island Fire Station	Hryciw et al. (1991); Redpath (1991); Gibbs et al. (1992); Furhriman (1993); Andrus (1994); de Alba et al. (1994)
```		Treasure Island Perimeter (Approach to Pier, UM03, UM05, UM06, UM09)	Geomatrix Consultants (1990); Hryciw (1991); R. D. Hryciw (personal communication to R. D. Andrus, 1998); Hryciw et al. (1998); Andrus et al. (1998a, 1998b)
1993 Kushiro-Oki, Japan	8.3	Kushiro Port (2, D)	Iai et al. (1995); S. Iai (personal communication to R. D. Andrus, 1997)
1993 Hokkaido-Nansei-Oki, Japan	8.3	Pension House	Kokusho et al. (1995a, 1995b, 1995c)
		Hakodate Port	S. Iai (personal communication to R. D. Andrus, 1997)
1994 Northridge, California	6.7	Rory Lane	Abdel-Haq & Hryciw (1998)

Table E.1 (cont.) - Earthquakes and Sites Used to Establish Liquefaction Resistance Curves.

Earthquake	Moment Magnitude	Site	Reference
(1)	(2)	(3)	(4)
1995 Hyogo-Ken Nanbu, Japan	6.9	Hanshin Expressway 5 (3, 10, 14, 25, 29); Kobe- Nishinomiya Expressway (3, 17, 23, 28)	Hamada et al. (1995); Hanshin Expressway Public Corporation (1998)
		KNK; Port Island (Downhole Array); SGK	Sato et al. (1996); Shibata et al. (1996)
		Port Island (Common Factory)	Ishihara et al. (1997); Ishihara et al. (1998)
		Kobe Port (7C); Port Island (1C, 2C)	Inatomi et al. (1997); Hamada et al. (1995)
		Kobe Port (LPG Tank Yard)	S. Yasuda (personal communication to R. D. Andrus, 1997)

Table E.1 (cont.) - Earthquakes and Sites Used to Establish Liquefaction Resistance Curves.

Values of  $V_s$  reported by the investigator(s) are used directly. The one exception is for the downhole array located at the Marina District School site in San Francisco, California. A reevaluation of the field data indicates that  $V_s$  values reported for the critical layer at this site are too high. They are recalculated using the pseudo-interval method, as discussed in Section E.2.2. Only the crosshole measurements made with shear waves having particle motion in the vertical direction are used. Crosshole measurements near the critical layer boundary that seem high, and could represent refracted waves, are not included in the average. Some  $V_s$  values are from measurements made before the earthquake, others followed the earthquake. No adjustments are made to compensate for changes in soil density and  $V_s$  due to ground shaking.

#### E.1.3 Measurement Depth.

In situ  $V_s$  measurements may be reported at discrete depths or for continuous intervals, depending on the test method. When velocities are reported for continuous intervals, as is typically the case for downhole, seismic cone, suspension logger and SASW measurements, the depth to the center of each interval is assumed. Thus, if the reported  $V_s$  profile has ten velocity layers, it is assumed that the profile consists of ten "measurements" with depths at the center of each layer.

## E.1.4 Case History

In this report, a case history is defined as a seismic event and a test array. For example, at the Treasure Island Fire Station site, crosshole measurements were made between five different pairs of boreholes, downhole measurements were made by two different investigators, seismic cone measurements were made at one location, and SASW measurements were made along one alignment. Thus, a total of nine case histories are identified for the Fire Station site and the 1989 Loma Prieta, California earthquake. At the Marina District School site, downhole measurements were made at one location. Estimates of ground surface acceleration at this site are available for the 1957 Daly City and 1989 Loma Prieta earthquakes. Thus, two case histories are identified for the Marina District School site. Combining the 26 seismic events and 139 test arrays, a total of 225 case histories are obtained with 149 from the United States, 36 from Taiwan, 34 from Japan, and 6 from China.

The two exceptions to this definition are the Owi Island No. 1 site and the Moss Landing Sandholdt Road UC-4 site where additional subsurface information is available. At Owi Island, pore pressure transducers recorded pore-water pressure buildup for two separate layers. At Moss Landing, inclinometer measurements indicated lateral movement in an upper loose layer and no lateral movement in a lower dense layer. Thus, two case histories are identified for each of these two test arrays.

## **E.1.5 Liquefaction Occurrence**

It is important to realize that the occurrence of liquefaction, in this evaluation, is based on the appearance of surface evidence, such as sand boils, ground cracks and fissures, and ground settlement. Case histories are classified as non-liquefaction when no liquefaction effects were observed. At the Owi Island No. 1, Lotung LSST Facility, Sunamachi, Wildlife (1987 earthquakes), and Port Island sites, the assessment of liquefaction or non-liquefaction occurrence is supported by pore-water pressure measurements. In addition, liquefaction occurrence is assigned (in This Report) to the Treasure Island, California, Fire Station case histories where the strong ground motion records from the 1989 Loma Prieta earthquake exhibit a sudden drop at about 15 seconds and small motion afterward (Idriss, 1990), indicating liquefaction (de Alba et al., 1994). Of the 225 case histories, 99 are liquefaction case histories and 126 are non-liquefaction case histories. Figure E.1 shows the distribution of case histories with earthquake magnitude.



Fig. E.1 - Distribution of Liquefaction and Non-Liquefaction Case Histories by Earthquake Magnitude.

## E.1.6 Critical Layer

The layer of soil most likely to liquefy at a site, or the critical layer, is the layer of nonplastic soil below the ground water table where values of  $V_{SI}$ , as defined in Chapter 2, and penetration resistance are generally the least and cyclic stress ratio relative to  $V_{SI}$  is the greatest. Figure E.2 presents the cumulative relative frequency distributions for the case histories by critical layer thickness and predominate soil type (gravel, or sand and silt). Critical layer thicknesses range from 1 m to as much as 15 m. About 50 % of the case histories have a critical layer thickness less than 3.5 m; 90 % of the case histories have a critical layer thickness for the sand and silt cases

Figure E.3 presents the cumulative relative frequency distributions for the case histories by average  $V_s$  measurement depth in the critical layer and predominate soil type. The average measurement depths are between 2 m and 11 m for nearly all case histories. Over 50 % of the case histories have average measurement depths less than 5.5 m. About 90 % of the case histories have average measurement depths less than 8 m. Overall, the measurement depths for the gravel cases are less than the measurement depths for the sand and silt cases.



Fig. E.2 - Cumulative Relative Frequency of Case History Data by Critical Layer Thickness.



Fig. E.3 - Cumulative Relative Frequency of Case History Data by Average Depth of  $V_s$ Measurements in Critical Layer.

Materials comprising the critical layers range from clean fine sand to sandy gravel with cobbles to profiles including silty clay layers. In Fig. E.4, the distribution of case histories with earthquake magnitude, predominate soil type (gravel, or sand and silt) and average fines content (silt and clay) is presented. Of the 225 case histories, 28 were for sands with fines content (FC)  $\leq 5$  %, 90 for sands with FC = 6 % to 34 %, 71 for sands and silts with FC  $\geq 35$  %, 26 for gravels with FC  $\leq 5$  %, and 10 for gravels with FC = 6 % to 34 %.

About 70 % of the case histories are for natural soils deposits, with many formed by alluvial processes. The other 30 % are for hydraulic or dumped fills. Eight of the fills have been densified by soil improvement techniques.

At least 85 % of the case histories are of Holocene age (< 10 000 years). Although the age of the other 15 % is unknown, they are believed to be also of Holocene age.



Fig. E.4 - Distribution of Case Histories by Earthquake Magnitude, Predominate Soil Type, and Average Fines Content.

# E.1.7 Ground Water Table

Figure E.5 presents the cumulative relative frequency distributions for the case histories by depth to the ground water table and predominate soil type. The ground water table for nearly all case histories lies between depths of 0.5 m and 6 m. Nearly 60 % of the case histories have water table depths less than 2 m. About 90 % of the case histories have water table depths less than 2 m.

Artesian pressures are reported for the Lotung Large-Scale Seismic Test (LSST) Facility site in Taiwan. At this site, the pore-water pressure distribution is assumed to vary linearly from a pressure head of 8.1 m at a depth of 7 m to a pressure head of 1.9 m at a depth of 2 m.

# E.1.8 Total and Effective Overburden Stresses

Values of total and effective overburden stresses are estimated using densities reported by the investigator(s). When no densities are reported, typical values for soils with similar grain size, penetration and velocity characteristics are assumed. In most instances, the assumed densities are  $1.76 \text{ Mg/m}^3$  for soils above the water table and  $1.92 \text{ Mg/m}^3$  for soils below the water table.



Fig. E.5 - Cumulative Relative Frequency of Case History Data by Depth to the Ground Water Table.

#### E.1.9 Average Peak Ground Acceleration

Average values of peak horizontal ground surface acceleration,  $a_{max}$ , are determined by averaging estimates reported by the investigator(s) and estimates made as part of this study using attenuation relationships developed from published ground surface acceleration data. Because many published attenuation relationships are based on both peak values obtained from ground motion records for the two horizontal directions (sometimes referred to as the randomly oriented horizontal component), the geometric mean (square root of the product) of the two peak values is used. Use of the geometric mean is consistent with the development of the SPTbased procedure (Youd et al., 1997; 2001). For the cases in this study, the difference between the geometric mean and arithmetic mean values is generally small, within about 5 %.

## E.1.10 Average Cyclic Stress Ratio

Cyclic stress ratios, CSR, are first calculated for each "measurement" depth within the critical layer using Eq. 2.1 and then averaged. Values of  $r_d$  are estimated using the average relationship developed by Seed and Idriss (1971) shown in Fig. 2.1. These  $r_d$  values are used to follow the traditional format of the SPT- and CPT-based procedures where the magnitude scaling factor is used to account for all effects of earthquake magnitude.

### E.1.11 Average Overburden Stress-Corrected Shear Wave Velocity

Values of  $V_S$  within the critical layer are first corrected for overburden stress using Eq. 2.7 and then averaged. The number of values included in the average range from 1 to 22 (see Appendix H). Values of  $C_{VS}$  used to correct measured shear wave velocities range from 1.4 to 0.9 for most of the data. About 80 % of the case histories have two to seven values in the average. No adjustments are made for possible variations between seismic test methods due to different source-receiver orientations with respect to the stress state in the soil. In the calculations, each site is assumed to be level ground.
### **E.2 SAMPLE CALCULATIONS**

Calculations for two sites shaken by the 1989 Loma Prieta, California earthquake ( $M_w =$  7.0) are presented below to illustrate how values of *CSR* and overburden stress-corrected shear wave velocity,  $V_{Sl}$ , are determined. The two sites are Treasure Island Fire Station and Marina District School.

### **E.2.1** Treasure Island Fire Station

Treasure Island is a man-made island located in the San Francisco Bay along the Bay Bridge between the cities of San Francisco and Oakland. It was constructed in 1936-37 by hydraulic filling behind a perimeter rock dike.

Extensive field tests have been conducted at the fire station on Treasure Island. Figure E.6 presents two  $V_s$  profiles for the site. The  $V_s$  profile determined by crosshole testing is from Fuhriman (1993). The other  $V_s$  profile is based on unpublished SASW test results by The University of Texas at Austin in 1992. Also presented in Fig. E.6 is the soil profile for the site. From the description by de Alba et al. (1994), the upper 4.5 m of soil consists of silty sand fill, possibly formed by dumping. Between depths of 4.5 m and 12.2 m, the soil consists of silty sand fill, soils. The ground water table lies near the ground surface at a depth of 1.4 m. The critical layer is determined to be between depths of 4.5 m and 7 m, where the soil is non-plastic, lies below the water table, and exhibits the lowest values of  $V_{SI}$  relative to the highest values of CSR in the layer (see Fig. 3.1).

During the 1989 Loma Prieta earthquake, a seismograph station at the fire station recorded ground surface accelerations. The peak values in the two horizontal accelerometer records are 0.16 g and 0.11 g (Brady and Shakal, 1994).

Sample calculations for the crosshole and SASW test arrays are summarized in Tables E.2 and E.3, respectively. The data points used in the calculations are shown by the open symbols in Fig. E.6. Total and effective overburden stresses are calculated assuming densities of  $1.76 \text{ Mg/m}^3$  above the water table and  $1.92 \text{ Mg/m}^3$  below the water table. Stress reduction coefficients are estimated using the average curve by Seed and Idriss (1971) shown in Fig. 2.1. The geometric mean of the two peak values observed in the horizontal ground surface acceleration records is 0.13 g. Using these parameters, values of *CSR* and *V*<sub>SI</sub> are calculated for the crosshole measurement at depth of 4.6 m as follows:



Fig. E.6 - Shear Wave Velocity and Soil Profiles for the Treasure Island Fire Station Site.

$$CSR = 0.65 \left(\frac{a_{\text{max}}}{g}\right) \left(\frac{\sigma_{\nu}}{\sigma_{\nu}'}\right) r_d = 0.65 \left(\frac{0.13g}{g}\right) \left(\frac{84.0}{52.7}\right) 0.97 = 0.131 \quad (E.1)$$

and

$$V_{SI} = V_S \left(\frac{P_a}{\sigma'_v}\right)^{0.25} = 134 \left(\frac{100}{52.7}\right)^{0.25} = 158 \text{ m/s}$$
 (E.2)

Representative values of CSR and  $V_{SI}$  used to defined the two case histories are determined by averaging values for each "measurement" depth within the critical layer, as shown in Tables E.2 and E.3.

### **E.2.2 Marina District School**

Kayen et al. (1990) conducted downhole seismic tests at the Winfield Scott School in the Marina District of San Francisco. Figure E.7 presents soil and velocity profiles for the site. The critical layer lies between depths of 2.7 m, the ground water table depth, and 4.3 m, the base of sand fill. The average  $V_s$  profile shown in Fig. E.7 was determined by Kayen et al., and was based on best-fit line segments through travel time measurements plotted versus depth. The second  $V_s$  profile is determined using the pseudo-interval method (This Report), as illustrated in Fig. E.8. Both methods should provide similar average values over the same depth interval. However, the layering assumed for the best-fit line segment method does not seem appropriate for the fill. For this reason, values of  $V_s$  based on the pseudo-interval method are used in this analysis.

As discussed in Section 3.2, the Marina District of San Francisco experienced a peak horizontal ground surface acceleration of about 0.15 g during the 1989 Loma Prieta earthquake.

Sample calculations for the Marina District School site are summarized in Table E.4. The locations of  $V_s$  measurements are assumed midway between receiver positions, as shown in Fig. E.7. Total and effective overburden stresses are estimated assuming densities of 1.76 Mg/m<sup>3</sup> above the water table and 1.92 Mg/m<sup>3</sup> below the water table. The ground water table is at a depth of about 2.7 m. Average values of *CSR* and  $V_{SI}$  defining the case history are determined by averaging values for the two "measurement" depths, as shown in Table E.4

# Table E.2 - Sample Calculations for the Treasure Island Fire Station Site, Crosshole Test ArrayB1- B4, and the 1989 Loma Prieta Earthquake.

Measurement Number	Average Depth, m	Measured Shear Wave Velocity, <i>V</i> s, m/s	Total Overburden Stress <sup>1</sup> , kPa	Effective Overburden Stress <sup>1</sup> , kPa	Stress Reduction Coefficient <sup>2</sup> , $r_d$	Cyclic Stress Ratio <sup>3</sup> , <i>CSR</i>	Overburden Stress- Corrected Shear Wave Velocity, V <sub>S1</sub> , m/s
(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)
1	4.57	134	84.0	52.7	0.97	0.13	158
2	5.49	133	111.3	60.9	0.96	0.14	150
3	6.40	144	118.5	69.2	0.95	0.14	158
Average	5.5	137	101.3	60.9	0.96	0.14	155

<sup>1</sup>Assuming water table at 1.4 m; and material densities are 1.76 Mg/m<sup>3</sup> above the water table

and 1.92 Mg/m<sup>3</sup> below the water table.

<sup>2</sup>Based on average values determined by Seed and Idriss (1971).

<sup>3</sup>Assuming peak horizontal ground surface acceleration is 0.13 g.

Table E.3 -	Sample	Calculations	for the	Treasure	Island	Fire	Station	Site,	SASW	Test A	Аггау,
	and the	1989 Loma	Prieta E	arthquak	e.						

Measurement Number (1)	Average Depth, m (2)	Measured Shear Wave Velocity, Vs, m/s (3)	Total Overburden Stress <sup>1</sup> , kPa (4)	Effective Overburden Stress <sup>1</sup> , kPa (5)	Stress Reduction Coefficient <sup>2</sup> , $r_d$ (6)	Cyclic Stress Ratio <sup>3</sup> , <i>CSR</i> (7)	Overburden Stress- Corrected Shear Wave Velocity, V <sub>S1</sub> , m/s (8)
1	5.34	131	98.4	59.6	0.96	0.14	149
Average	5.3	131	98.4	59.6	0.96	0.14	149

<sup>1</sup>Assuming water table at 1.4 m; and material densities are 1.76 Mg/m<sup>3</sup> above the water table and 1.92 Mg/m<sup>3</sup> below the water table.

<sup>2</sup>Based on average values determined by Seed and Idriss (1971).

<sup>3</sup>Assuming peak horizontal ground surface acceleration is 0.13 g.



Fig. E.7 - Shear Wave Velocity and Soil Profiles for the Marina District School Site (Kayen et al., 1990).



Fig. E.8 - General Configuration of the Downhole Seismic Test Using the Pseudo-Interval Method to Calculate Shear Wave Velocity.

### **E.3 SUMMARY**

The case history data described in this chapter are limited to level and gently sloping sites with the following characteristics:

- (1) average critical layer depths less than 10 m;
- (2) uncemented soils of Holocene age;
- (3) ground water table depths between 0.5 m and 6 m; and
- (4) all  $V_S$  measurements from below the water table.

Of the 225 case histories, 57 are for soils with  $FC \le 5$  %, 98 for soils with FC = 6 % to 34 %, and 70 with  $FC \ge 35$  %. About 20 % of the case histories are for soils containing more than 10 % gravel. Nearly 50 % of the case histories are for earthquake magnitudes near 7.

Measurement Number (1)	Average Depth, m (2)	Measured Shear Wave Velocity <sup>1</sup> , <i>Vs</i> , m/s (3)	Total Overburden Stress <sup>2</sup> , kPa (4)	Effective Overburden Stress <sup>2</sup> , kPa (5)	Stress Reduction Coefficient <sup>3</sup> , $r_d$ (6)	Cyclic Stress Ratio <sup>4</sup> , <i>CSR</i> (7)	Overburden Stress- Corrected Shear Wave Velocity, $V_{S1}$ , m/s (8)
1	3.02	87	52.6	49.9	0.98	0.10	104
2	3.94	136	70.0	58.2	0.97	0.11	156
Average	3.5	112	61.3	54.1	0.98	0.11	130

# Table E.4 - Sample Calculations for the Marina District School Site and the 1989 Loma Prieta Earthquake.

<sup>1</sup>Based on pseudo-interval method.

<sup>2</sup>Assuming water table at 2.7 m; and material densities are 1.76 Mg/m<sup>3</sup> above the water table and 1.92 Mg/m<sup>3</sup> below the water table.

<sup>3</sup>Based on average values determined by Seed and Idriss (1971).

<sup>4</sup>Assuming peak horizontal ground surface acceleration is 0.15 g.

### **APPENDIX F**

## DEVELOPMENT OF LIQUEFACTION RESISTANCE CURVES FROM CASE HISTORY DATA

In the process of developing the liquefaction evaluation chart shown in Fig. 2.3 all case histories were initially plotted on the same chart. This aggregation was accomplished through an adjustment procedure; that is, the CSR values in each case history were adjusted to an earthquake with  $M_w = 7.5$  by dividing by Eq. (2.9) with n = -2.56. As done in penetration evaluation procedures, the sandy soil case histories were separated into three categories: (1) sands with average  $FC \le 5$  %; (2) sands with average FC = 6 % to 34 %; and (3) sands and silts with average  $FC \ge 35$  %. For consistency, the gravelly soil case histories also were divided into the same three categories based on fines content. However, no case histories exist in the database with gravel having  $FC \ge 35$  %. All data are plotted in Fig. F.1 along with the recommended  $CRR-V_{SI}$  curves. Development of these curves is discussed in this appendix.

The shape of the  $CRR-V_{SI}$  curves shown in Fig. F.1 is based on a modified relationship between shear wave velocity and cyclic stress ratio for constant average cyclic shear strain suggested by R. Dobry. The modified relationship is expressed as (Andrus and Stokoe, 1997):

$$CRR_{7.5} = \left\{ a \left( \frac{V_{S1}}{100} \right)^2 + b \left( \frac{1}{V_{S1}^* - V_{S1}} - \frac{1}{V_{S1}^*} \right) \right\}$$
(F.1)

where

 $CRR_{7.5} = CRR$  for magnitude 7.5 earthquakes,  $V_{S1}^* =$  the limiting upper value of  $V_{S1}$  for liquefaction occurrence, and a, b = curve fitting parameters.

As discussed in Section D.5, the first (quadratic) term of Eq. (F.1) is a form of Dobry's relationship given by Eq. (D.9). The second term is a hyperbola with a small value at low values of  $V_{SI}$ , and a very large value as  $V_{SI}$  approaches  $V_{S1}^*$ .



 Fig. F.1 - Curves Recommended for Calculation of CRR from Shear Wave Velocity Measurements Along with Case History Data Based on Lower Bound Values of MSF for the Range Recommended by the 1996 NCEER Workshop (Youd et al., 1997) and r<sub>d</sub> Developed by Seed and Idriss (1971).

### F.1 LIMITING UPPER $V_{S1}$ VALUE FOR LIQUEFACTION OCCURRENCE

As shown in Fig. F.1, CSR-values above about 0.35 are limited in the case history data. Thus, current estimates of  $V_{S1}^*$  rely, in part, on penetration-shear wave velocity correlations and, in part, on the data trend in Fig. F.1.

### F.1.1 Sandy Soils

In the SPT-based procedure, a corrected blow count,  $(N_I)_{60}$ , of 30 is assumed as the limiting upper value for liquefaction occurrence in sands with  $\leq 5$  % silt and clay (Seed et al., 1985; Youd et al., 1997). Table F.1 presents estimates of equivalent  $V_{SI}$  for corrected blow count of 30. The correlation by Ohta and Goto (1978) modified to a blow count with a theoretical free-fall energy of 60 % (Seed et al., 1985) suggested equivalent  $V_{SI}$  values of 207 m/s for Holocene sands, assuming that a depth of 10 m is equivalent to an effective overburden stress of 100 kPa. The stress-corrected crosshole measurements compiled by Sykora (1987b) for Holocene sands and non-plastic silty sands below the ground water table, with  $(N_I)_{60}$  between 25 and 35, exhibited an average  $V_{SI}$  value of 206 m/s and standard deviation of 41 m/s. Finally, the case history data in this study were used to investigate the  $V_{SI}$  and  $(N_I)_{60}$  relationship for well-documented sand layers with less than 10 % fines. These data are presented in Fig. F.2 along with the best-fit relationship that can be expressed as:

$$V_{SI} = B_1 [(N_1)_{60}]^{B_2}$$
(F.2)

where  $B_1 = 93.2 \pm 6.5$  and  $B_2 = 0.231 \pm 0.022$  for soils with fines content < 10 %, and with  $V_{SI}$  in m/s and  $(N_1)_{60}$  in blows/0.3 m. The plotted data exhibit a mean  $V_{SI}$  value of 204 m/s at a  $(N_1)_{60}$  value of 30 and residual standard deviation,  $S_{res}$ , of 12 m/s.

In the CPT-based procedure, a normalized cone tip resistance,  $q_{cIN}$ , of 160 is assumed as the limiting upper value for liquefaction occurrence in sands with  $\leq 5$  % silt and clay (Youd et al., 1997; Robertson and Wride, 1998). Figure F.3 presents average values of  $V_{SI}$  and  $q_{cIN}$  for soil layers with less than 10 % fines at several sites listed in Table E.1. Also shown in Fig. F.3 is the best-fit relationship for the plotted data, which can be expressed as:

$$V_{SI} = B_1 [q_{c1N}]^{B_2}$$
(F.3)

where  $B_1 = 88.2 \pm 15.5$  and  $B_2 = 0.154 \pm 0.037$  for soils with fines content < 10 %, and with  $V_{SI}$  in m/s and  $q_{cIN}$  is normalized tip resistance based on procedures by Robertson and Wride (1998). As noted in Table F.2, the plotted data exhibit a mean  $V_{SI}$  value of 193 m/s at a  $q_{cIN}$  value of 160 and residual standard deviation of 19 m/s.

Reference (1)	Relationship (2)	Equivalent V <sub>SI</sub> Estimate (m/s) (3)	Assumptions (4)
Ohta & Goto (1978); also given in report by Sykora (1987a, page 29)	$V_{S} = 69 (N_{j})^{0.173} z^{0.195} F_{I}F_{2}$ $N_{j} = \text{SPT blow count}$ measured in Japanese practice $z = \text{depth, m}$ $F_{1} = 1.00 \text{ for Holocene-}$ age soils $F_{2} = 1.085 \text{ for sands;}$ $1.189 \text{ for gravel}$ best-fit relationship for 289 sets of SPT and $V_{S}$ measurements from Japan	207 for Holocene sands 227 for Holocene gravels	<ol> <li>N<sub>j</sub> = 60/67 N<sub>60</sub></li> <li>N<sub>60</sub> = 30</li> <li>z = 10 m is equivalent to an overburden stress of 100 kPa</li> <li>All measurements are from below the ground water table</li> </ol>
Sykora (1987b, page 90); This Report	Correlation between $(N_I)_{60}$ and crosshole $V_S$ , normalized to effective overburden stress, measurements for Holocene sands and non-plastic silty sands below the ground water table at sites in U.S.A.; 16 sets of measurements (with known SPT equipment)	206 for Holocene sands and non- plastic silty sands below the water table standard deviation is 41 m/s	<ol> <li>Average for V<sub>SI</sub> values with (N<sub>1</sub>)<sub>60</sub> between 25 and 35</li> <li>σ'<sub>ν</sub> = 100 kPa</li> </ol>
Rollins et al. (1998a)	$V_S = 53 (N_{60})^{0.19} (\sigma_{\nu})^{018}$ best-fit relationship using equivalent $N_{60}$ -values from Becker Penetration Tests and $V_S$ measurements; 186 points from 7 Holocene gravel sites	232 for Holocene gravels most of data lie within ±25 % of relationship	<ol> <li>N<sub>60</sub> = 30</li> <li>σ'<sub>ν</sub> = 100 kPa</li> <li>All measurements are from below the ground water table</li> </ol>
This Report (see Fig. F.2)	$V_{SI} = B_1 [(N_1)_{60}]^{B_2}$ $B_1 = 93.2 \pm 6.5$ $B_2 = 0.231 \pm 0.022$ best-fit relationship for uncemented, Holocene-age sands with less than 10 % non- plastic fines; 25 sets of average SPT and V <sub>S</sub> measurements all from below the water table	204 for Holocene clean sands below the water table residual standard deviation is 12 m/s	<ol> <li>Average for V<sub>SI</sub> with (N<sub>1</sub>)<sub>60</sub> = 30</li> <li>σ'<sub>ν</sub> = 100 kPa</li> <li>Corrected blow count based on procedures given in Seed et al. (1985) and Robertson and Wride (1997; 1998)</li> </ol>

# Table F.1 - Estimates of Equivalent $V_{SI}$ for Holocene Sands and Gravels Below the Ground Water Table with Corrected SPT Blow Count of 30.



Fig. F.2 - Variations in  $V_{SI}$  with  $(N_I)_{60}$  for Uncemented, Holocene-age Sands with Less than 10 % Non-Plastic Fines.



Fig. F.3 - Variations in  $V_{SI}$  with  $q_{cIN}$  for Uncemented, Holocene-age Sands with Less than 10 % Non-Plastic Fines.

Reference (1)	Relationship (2)	Equivalent V <sub>SI</sub> Estimate (m/s) (3)	Assumptions (4)
This Report (see Fig. F.3)	$V_{SI} = B_I (q_{cW})^{B_2}$ $B_I = 88.2 \pm 15.5$ $B_2 = 0.154 \pm 0.037$ best-fit relationship for uncemented, Holocene-age sands with less than 10 % non- plastic fines; 23 sets of average SPT and V <sub>S</sub> measurements all from below the water table	193 for Holocene clean sands below the water table residual standard deviation is 19 m/s	<ol> <li>Average for V<sub>SI</sub> with q<sub>cIN</sub> = 160</li> <li>σ'<sub>ν</sub> = 100 kPa</li> <li>Normalized tip resistance based on procedures given in Robertson and Wride (1997; 1998)</li> </ol>

Table F.2 - Estimates of Equivalent  $V_{SI}$  for Holocene Sands Below the Ground Water Table with Normalized Cone Tip Resistance of 160.

From these estimates, a  $V_{SI}$  value of 210 m/s is assumed equivalent to an  $(N_I)_{60}$  value of 30 in clean sands ( $\leq 5$  % fines). A limiting upper  $V_{SI}$  value of 210 m/s for cyclic liquefaction occurrence at CSR = 0.6 is less than the general consensus value of 230 m/s suggested at the 1998 MCEER Workshop. As a result, Figs. F.2 and F.3 were added specifically to provide additional evidence to support the use of 210 m/s in clean sands.

For sandy and silty soils with FC  $\geq 35$  %, the SPT-based chart by Seed et al. (1985) indicates a limiting upper  $(N_I)_{60}$  value of about 21 for cyclic liquefaction occurrence. Table F.3 presents estimates of equivalent  $V_{SI}$  for blow count of 21. The correlation by Ohta and Goto (1978) suggested equivalent  $V_{SI}$  values of 195 m/s for Holocene sands. The stress-corrected crosshole compiled by Sykora (1987b) for Holocene sands and non-plastic silty sands below the ground water table, with  $(N_I)_{60}$  between 16 and 26, exhibited an average value of 199 m/s and standard deviation of 36 m/s. From these estimates, a  $V_{SI}$  value of 195 m/s is assumed equivalent to an  $(N_I)_{60}$  value of 21 in non-plastic soils with FC  $\geq 35$  %.

To permit the  $CRR-V_{SI}$  curves for magnitude 7.5 earthquakes shown in Fig. F.1 to have  $V_{SI}$  values between 195 m/s and 210 m/s at CRR near 0.6, values of  $V_{S1}^*$  are assumed to range linearly from 200 m/s to 215 m/s, respectively. The relationship between  $V_{S1}^*$  and fines content, FC, can be expressed by:

$V_{S1}^* = 215 \text{ m/s}$	for sands with $FC \leq 5$ %	(F.4a)
$V_{S1}^* = 215 - 0.5(FC-5) \text{ m/s}$	for sands with 5 % $< FC < 35$ %	(F.4b)
$V_{S1}^* = 200 \text{ m/s}$	for sands and silts with $FC \ge 35$ %	(F.4c)

To illustrate how well the recommended  $CRR-V_{SI}$  curves defined by Eqs. (F.1) and (F.4) fit the case history data, the data separated by soil type, are presented in Figs. F.4 through F.7. The recommended curves provide reasonable bounds for all the case history data above a CSR value of 0.35, supporting the use of the suggested  $V_{S1}^*$  values for sands and silts, as well as gravels. The use of these  $V_{S1}^*$  values for gravels is discussed below.

Reference (1)	Relationship (2)	Equivalent V <sub>S1</sub> Estimate (m/s) (3)	Assumptions (4)
Ohta & Goto (1978); also given in report by Sykora (1987a, page 29)	$V_{S} = 69 (N_{j})^{0.173} z^{0.195} F_{1}F_{2}$ $N_{j} = \text{SPT blow count}$ measured in Japanese practice $z = \text{depth, m}$ $F_{1} = 1.00 \text{ for Holocene-}$ age soils $F_{2} = 1.085 \text{ for sands;}$ 1.189 for gravelbest-fit relationship for 289 sets of SPT and $V_{S}$ measurements from Japan	195 for Holocene sands 214 for Holocene gravels	<ol> <li>N<sub>j</sub> = 60/67 N<sub>60</sub></li> <li>N<sub>60</sub> = 21</li> <li>z = 10 m is equivalent to         <ul> <li>an overburden stress of 100 kPa</li> <li>All measurements are from below the ground water table</li> </ul> </li> </ol>
Sykora (1987b, page 90); This Report	Correlation between $(N_i)_{60}$ and crosshole $V_s$ , normalized to effective overburden stress, measurements for Holocene sands and non-plastic silty sands below the water table at sites in U.S.A.; 31 sets of measurements (with known SPT equipment)	199 for Holocene sands and non- plastic silty sands below the water table standard deviation is 36 m/s	<ol> <li>Average for V<sub>Sl</sub> values with (N<sub>l</sub>)<sub>60</sub> between 16 and 26</li> <li>σ'<sub>ν</sub> = 100 kPa</li> </ol>
Rollins et al. (1998a)	$V_s = 53 (N_{60})^{0.19} (\sigma'_{\nu})^{018}$ best-fit relationship using equivalent $N_{60}$ -values from Becker Penetration Tests and $V_s$ measurements; 186 points from 7 Holocene gravel sites	217 for Holocene gravels most of data lie within ±25 % of relationship	<ol> <li>N<sub>60</sub> = 21</li> <li>σ'<sub>ν</sub> = 100 kPa</li> <li>All measurements are from below the ground water table</li> </ol>

Table F.3 - Estimates of Equivalent  $V_{SI}$  for Holocene Sands and Gravels Below the GroundWater Table with Corrected SPT Blow Count of 21.



Fig. F.4 - Curve Recommended for Calculation of CRR from  $V_{SI}$  Measurements in Sands and Silts with  $FC \ge 35$  % along with Case History Data.



Fig. F.5 - Curve Recommended for Calculation of CRR from  $V_{SI}$  Measurements in Sands with FC = 20 % along with Case History Data.



Fig. F.6 - Curve Recommended for Calculation of CRR from  $V_{SI}$  Measurements in Sands with  $FC \le 5$  % along with Case History Data.



Fig. F.7 - Curves Recommended for Calculation of CRR from  $V_{S1}$  Measurements in Gravels with  $FC \le 5$  % and FC = 20 % along with Case History Data.

### F.1.2 Gravelly Soils

Although the  $V_{S1}^*$  values given in Eq. (F.4) were determined for sandy soils, the results presented in Fig. F.7 indicate that these limits also represent reasonable limits for gravelly soils divided into the same categories based on fines content. This might be considered rather surprising, based on the penetration- $V_S$  correlations presented in the literature for gravelly soils. For instance, as noted in Table F.1, the correlation by Ohta and Goto (1978) suggested a  $V_{S1}$ value of 227 m/s for Holocene gravels at an equivalent  $(N_1)_{60}$  of 30. Similarly, the correlation by Rollins et al. (1998a) provided a best-fit value of 232 m/s for Holocene gravels. On the other hand, all the liquefaction case history data shown in Figs. F.4 through F.7 exhibit  $V_{S1}$  values of about 200 m/s or less, suggesting that 230 m/s may be inappropriately high.

To investigate further the value of  $V_{S1}^*$  in gravelly soils, laboratory studies involving  $V_S$  measurements in gravelly soils were reviewed. Kokusho et al. (1995b) clearly showed that the shear wave velocity of gravelly soils varies greatly and is highly dependent on the particle gradation. Weston (1996) showed similar results for coarse sands with gravels. In both cases, the results show that increasing the uniformity coefficient can significantly increase the shear wave velocity in medium dense to dense gravels. On the other hand, very loose gravelly soils, even well-graded gravels, can exhibit shear wave velocities similar to those of loose sands (Kokusho et al., 1995b). The case history data presented in Fig. F.7 supports the premise that gravelly soils that are loose enough to exhibit significant liquefaction effects (boils, ground cracks, etc.) have shear wave velocities similar to loose sands. Hence, the authors recommend the boundaries developed for sandy soils as preliminary boundaries for gravelly soils. However, additional work is clearly needed to understand the relationship between  $V_{SI}$  and liquefaction resistance of gravels.

## F.2 CURVE FITTING PARAMETERS a AND b

The curve fitting parameters a and b in Eq. (F.1) can be approximated from the case history data assuming the values of  $V_{S1}^*$  given in Eq. (F.4) and a *MSF* relationship. Three *MSF* relationships representing the range of proposed magnitude scaling factors (see Section 2.3.1) are considered below to establish the values of a and b in Eq. (F.1).

### F.2.1 Magnitude Scaling Factors Recommended by 1996 NCEER Workshop

**F.2.1.1 Lower Bound of Recommended Range**—Figure F.1 presents the case history data for magnitude 5.9 to 8.3 earthquakes adjusted using the lower bound for the range of magnitude scaling factors recommended by the 1996 NCEER workshop (Youd et al., 1997). The lower bound is defined by Eq. (2.9) with n = -2.56, as discussed in Section 2.3.1. Also

shown in Fig. F.1 are three recommended  $CRR-V_{SI}$  curves for earthquakes with magnitude near 7.5 and various fines content. The three curves were determined through an iterative process of varying the values of a and b until nearly all the liquefaction case histories were bound by the curves with the least amount of non-liquefaction case histories in the liquefaction region. The final values of a and b used to draw the curves were 0.022 and 2.8, respectively.

Of the 99 liquefaction case histories, only two liquefaction case histories incorrectly lie in the no liquefaction region. The two liquefaction case histories shown in Fig. F.1 that incorrectly lie in the no liquefaction region are two sites at Treasure Island (UM05 and UM09) where liquefaction was marginal during the 1989 Loma Prieta earthquake ( $M_w = 7$ ). The sites are located along the perimeter of Treasure Island. Mapped liquefaction effects generated by the 1989 earthquake near the UM05 site are ground cracks with 50 to 90 mm of horizontal displacement (R. D. Hryciw, personal communication to R. D. Andrus, 1998; Power et al., 1998). The nearest mapped sand boil is located 60 m away from the site. At the UM09 site, as much as 90 mm of vertical displacement was observed adjacent to a building located 60 m inland from the site. These displacements are small compared to the meters of displacement that are expected to occur during larger ground shaking. Thus, liquefaction was marginal at the UM05 and UM09 sites, and sloping ground may have been a factor. It is interesting to note that similar incorrect evaluations also are obtained when one uses the SPT and CPT data for these two sites. The SPT- and CPT-based evaluations for the UM05 site are discussed in Section F.3.1.

**F.2.1.2 Upper Bound of Recommended Range**—Figure F.8 presents the case history data for magnitude 5.9 to 8.3 earthquakes adjusted using the upper bound for the range of magnitude scaling factors recommended by the 1996 NCEER Workshop (Youd et al., 1997). The upper bound is defined by Eq. (2.9) with n = -3.3, as discussed in Section 2.3.1. Also shown in Fig. F.8 are the three  $CRR-V_{Sl}$  curves from Fig. F.1. Many case histories plot lower in Fig. F.8 than in Fig. F.1, because the  $M_w$  is less than 7.5 for most of the data. The downward shift in the liquefaction data points near the curves at CRR of about 0.08 is less than 0.01. This difference is not significant, and is within the accuracy of the plotted case history data.

### F.2.2 Revised Magnitude Scaling Factors Proposed by Idriss (1999)

Figure F.9 presents the case history data for magnitude 5.9 to 8.3 earthquakes adjusted using the revised magnitude scaling factors and stress reduction coefficients proposed by Idriss (1999, as discussed in Section 2.3.1. Also shown in Fig. F.9 are the three  $CRR-V_{SI}$  curves from Fig. F.1. Many of the case history data shown in Fig. F.9 plot higher than case history data in Fig. F.1, because the  $M_w$  is less than 7.5 for most of the data. The upward shift in the liquefaction data points near the curves at CRR of about 0.08 is less than 0.01. Again, this difference is not significant, and is within the accuracy of the plotted case history data.



Fig. F.8 - Curves Recommended for Calculation of *CRR* from Shear Wave Velocity Measurements Along with Case History Data Based on **Upper Bound** Values of *MSF* for the Range Recommended by the 1996 NCEER Workshop (Youd et al., 1997) and  $r_d$  Developed by Seed and Idriss (1971).



Fig. F.9 - Curves Recommended for Calculation of *CRR* from Shear Wave Velocity Measurements Along with Case History Data Based on Revised Values of *MSF* and  $r_d$  Proposed by Idriss (1999).

### F.2.3 Comparison of Magnitude Scaling Factors

The proposed relationships for MSF can be compared directly by combining them with the appropriate stress reduction coefficient into one factor. This factor is the product of  $r_d$  and the reciprocal of MSF. Figure F.10 presents values of  $r_d/MSF$  for the range recommended by the 1996 NCEER Workshop (Youd et al., 1997) and those proposed by Idriss (1999). As shown in the figure, there is not much difference between the two sets of  $r_d/MSF$  values for magnitude of 7.5 and depth less than 11 m. At magnitudes near 5.5 and shallow depths, the difference between  $r_d/MSF$  values proposed by Idriss (1999) and values recommended by the 1996 NCEER Workshop is as much as 50 %. Thus, at magnitudes less than about 7, the difference in using values of MSF and  $r_d$  proposed by Idriss (1999) and those adopted by the NCEER Workshop (Youd et al., 1997) is significant in the calculation of CSR.

For example, Fig. F.11 presents two liquefaction resistance curves for earthquakes with magnitude near 5.5 and clean soils ( $FC \le 5$  %). The upper curve was obtained by multiplying values of *CRR* defining the curve for  $FC \le 5$ % in Fig. F.1 by 2.2, the lower *MSF* recommended by the 1996 NCEER Workshop for magnitude 5.5 earthquakes (see Eq. (2.9) with n = -2.56). The lower curve was obtained by multiplying values of *CRR* defining the curve for  $FC \le 5$ % in Fig. F.1 by 1.68, the *MSF* proposed by Idriss (1999) for magnitude 5.5 earthquakes (see Eq. (2.10)). Also shown in Fig. F.11 are the available case history data for clean sands determined using average stress reduction coefficients proposed by Seed and Idriss (1971) and Idriss (1998). The two curves in Fig. F.11 exhibit differences in *CRR* of about 0.02 at  $V_{SI} = 100$  m/s and 0.1 at  $V_{SI} = 200$  m/s.

#### F.3 RECOMMENDED CRR-V<sub>SI</sub> CURVES

From the discussion presented above, the recommended  $CRR-V_{SI}$  curves are defined by Eqs. (F.1), (2.9) and (F.4) with a = 0.022, b = 2.8, and n = -2.56. The recommended curves for moment magnitudes ranging from 5.5 though 8 are presented in Figs. F.12 through F.17, respectively, along with the case history data. The value of -2.56 for n is recommended for determining magnitude scaling factors because it provides more conservative  $CRR-V_{SI}$  curves than -3.3, which is the n value defining the upper bound of the range of MSFs suggested by the 1996 NCEER Workshop (Youd et al., 1997) for magnitudes less than 7.5. Although the magnitude scaling factors defined by Eq. (2.9) with n = -2.56 provide less conservative  $CRR-V_{SI}$ curves than the factors proposed by Idriss (1999) for magnitudes less than 7.5, the factors determined of Ambraseys (1988), I. M. Idriss (personal communication to T. L. Youd, 1995), Arango (1996), Youd and Noble (1997), and Andrus and Stokoe (1997; as indicated by the very conservative  $CRR-V_{SI}$  curves shown in Figs. F.13 and F.14) supported their use.



Fig. F.10 - Variation of  $r_d/MSF$  with Depth for Various Magnitudes and Proposed Relationships.



Fig. F.11 - Comparison of Liquefaction Resistance Curves and Case History Data for Procedures Recommended by the 1996 NCEER Workshop (Youd et al., 1997) and the Revised Procedures Proposed by Idriss (1999) for Clean Sands and Earthquakes with Magnitude Near 5.5.



Fig. F.12 - Case History Data for Earthquakes with Magnitude Near 5.5 Based on Overburden Stress-Corrected Shear Wave Velocity and Cyclic Stress Ratio with Recommended Liquefaction Resistance Curves.



Fig. F.13 - Case History Data for Earthquakes with Magnitude Near 6 Based on Overburden Stress-Corrected Shear Wave Velocity and Cyclic Stress Ratio with Recommended Liquefaction Resistance Curves.



Fig. F.14 - Case History Data for Earthquakes with Magnitude Near 6.5 Based on Overburden Stress-Corrected Shear Wave Velocity and Cyclic Stress Ratio with Recommended Liquefaction Resistance Curves.



Fig. F.15 - Case History Data For Earthquakes with Magnitude Near 7 Based on Overburden Stress-Corrected Shear Wave Velocity and Cyclic Stress Ratio with Recommended Liquefaction Resistance Curves.



Fig. F.16 - Case History Data for Earthquakes with Magnitude Near 7.5 Based on Overburden Stress-Corrected Shear Wave Velocity and Cyclic Stress Ratio with Recommended Liquefaction Resistance Curves.



Fig. F.17 - Case History Data for Earthquakes with Magnitude Near 8 Based on Overburden Stress-Corrected Shear Wave Velocity and Cyclic Stress Ratio with Recommended Liquefaction Resistance Curves.

The recommended  $CRR-V_{SI}$  curves shown in Figs. F.12 through F.17 are dashed above CRR of about 0.35 to indicate that they are based on limited field performance data. The curves do not extend much below 100 m/s, because there are no field data to support extending them to the origin. It is important to note that these boundary curves are for extreme behavior, where boils and ground cracks occur.

### F.4 CORRELATIONS BETWEEN V<sub>s1</sub> AND PENETRATION RESISTANCE

One can obtain correlations between  $V_{SI}$  and corrected penetration resistance from the recommended  $CRR-V_{SI}$  relationships given in Fig. F.1 and 1996 NCEER Workshop (Youd et al., 1997) recommended SPT- and CPT-based relationships for magnitude 7.5 earthquakes by plotting values with equal CRR.

### F.4.1 Corrected SPT Blow Count

Figure F.18 presents the correlation of  $V_{SI}$  with  $(N_I)_{60}$  for clean soils ( $\leq 5$  % fines), based on the recommended  $CRR-V_{SI}$  and  $CRR-(N_I)_{60}$  relationships. Also shown are the field data and mean curve for sands with less than 10 % non-plastic fines from Fig. F.2. The correlation derived from the CRR relationships lies between the mean and the mean  $\pm 1S_{res}$  curves. The flatter slope below  $(N_I)_{60}$  of 6 exhibited by the CRR-based correlation can be explained by different assumed minimal values of CRR. The  $CRR-V_{SI}$  relationship for magnitude 7.5 earthquakes and  $FC \leq 5$  % shown in Fig. F.1 provides a CRR of 0.033 for  $V_{SI} = 100$  m/s, the lowest  $V_{SI}$  value shown in the figure. The 1996 NCEER Workshop recommended a CRR value of 0.05 for  $(N_I)_{60} = 0$ . The difference between minimal values of CRR is small, and is near the accuracy of both procedures.

The CRR-based correlation shown in Fig. F.18, along with the plotted field data, provide a simple method of comparing the  $V_{SI}$ - and  $(N_I)_{60}$ -based liquefaction evaluation procedures. Both procedures provide similar predictions of liquefaction potential, when the data point lies on the CRR-based curve. When the data point plots below the CRR-based curve, the  $V_{SI}$ -based liquefaction evaluation procedure provides the more conservative prediction. When the data point plots above the CRR-based curve, the SPT-based liquefaction evaluation procedure provides the more conservative prediction. Because most of the data points shown in Fig. F.18 plot below the CRR-based curve, the  $V_{SI}$ -based procedure provides an overall more conservative prediction of liquefaction resistance than does the SPT-based procedure for these sites.



Fig. F.18 - Relationships Between  $(N_1)_{60}$  and  $V_{SI}$  for Clean Sands Implied by the Recommended *CRR-V<sub>SI</sub>* Relationship and the 1996 NCEER Workshop Recommended *CRR-(N<sub>1</sub>)*<sub>60</sub> Relationship (Youd et al., 1997) with Field Data for Sands with Less than 10 % Fines. The data point for the Treasure Island UM05 site, which incorrectly lies in the region of no liquefaction shown in Fig. F.1, plots just below the *CRR*-based curve, as shown in Fig. F.18. Thus, this case history also incorrectly plots in the region of no liquefaction on the SPT-based liquefaction evaluation chart. Furthermore, the SPT-based procedure provides a slightly less conservative prediction of liquefaction resistance than the shear-wave-based procedure for this case history.

Although the CRR-based curve shown in Fig. F.18 generally trends parallel to the mean curve, there is a small hump between corrected blow counts of 8 and 26. This hump suggests that either the  $CRR-V_{SI}$  relationship is more conservative or the  $CRR-(N_1)_{60}$  relationship is less conservative in this range.

Similarly, as shown in Fig. F.19, a correlation between  $V_{SI}$  with  $(N_I)_{60}$  for soils with  $\geq 35$ % fines can be derived from the recommended  $CRR-V_{SI}$  and  $CRR-(N_I)_{60}$  relationships. Figure F.19 provides the basis for the method of estimating the cementation and aging correction factor,  $K_c$ , suggested in Section 2.3.4 (see Fig. 2.6).

### F.4.2 Normalized Cone Tip Resistance

Figure F.20 presents the correlation of  $V_{SI}$  with  $q_{cIN}$  for clean sands with median grain size,  $D_{50}$ , between 0.25 mm and 2.0 mm, based on the recommended  $CRR-V_{SI}$  and  $CRR-q_{cIN}$ relationships. Also shown are the field data and mean curve for clean sands with less than 10 % non-plastic fines from Fig. F.3. The correlation derived from the CRR relationships lies between the mean and the mean  $+1S_{res}$  curves for  $V_{SI} = 170$  m/s, indicating that the  $V_{SI}$ -based procedure provides an overall more conservative prediction of liquefaction resistance than does the CPTbased procedure for these sites. For  $V_{SI} < 170$ , the CRR-based correlation lies close to the mean curve, indicating that both procedure provide an overall similar prediction. The slope of the CRR-based correlation below  $q_{cIN}$  of 20 may be explained by the different assumed minimal values of CRR, as discussed in Section F.4.1.

The data point for the Treasure Island UM05 site, which incorrectly lies in the region of no liquefaction shown in Fig. F.1, plots on the *CRR*-based curve, as shown in Fig. F.20. Thus, the CPT-based procedure provides a similar incorrect prediction of no liquefaction for this site.

Similarly, as shown in Fig. F.21, a correlation between  $V_{SI}$  with  $q_{cIN}$  for soils with  $\geq 35$ % fines can be derived from the recommended  $CRR-V_{SI}$  and  $CRR-(N_I)_{60}$  relationships. Figure F.21 provides the basis for the method of estimating the cementation and aging correction factor,  $K_c$ , suggested in Section 2.3.4 (see Fig. 2.7).


Fig. F.19 - Relationships Between  $(N_1)_{60}$  and  $V_{SI}$  Implied by the Recommended CRR- $V_{SI}$ Relationship and the 1996 NCEER Workshop Recommended CRR- $(N_1)_{60}$ Relationship (Youd et al., 1997).



Fig. F.20 - Relationships Between  $q_{cIN}$  and  $V_{SI}$  for Clean Sands Implied by the Recommended *CRR-V<sub>SI</sub>* Relationship and the 1996 NCEER Workshop Recommended *CRR-q<sub>cIN</sub>* Relationship (Youd et al., 1997) with Field Data for Sands with Less than 10 % Fines.



Fig. F.21 - Relationships Between  $q_{cIN}$  and  $V_{SI}$  Implied by the Recommended  $CRR-V_S$ Relationship and the 1996 NCEER Workshop Recommended  $CRR-q_{cIN}$ Relationship (Youd et al., 1997).

## **F.5 SUMMARY**

The development of the recommended  $CRR-V_{SI}$  curves was outlined in this appendix. The recommended curves are based on a modified relationship between shear wave velocity and cyclic stress ratio for constant average cyclic shear strain suggested by R. Dobry. They are defined by Eqs. (F.1), (F.4) and (2.9) with a = 0.022, b = 2.8, and n = -2.56. The curve fitting parameters a and b are determined through an iterative process that involved varying their values until nearly all the liquefaction case histories were bound by the curves with the least amount of non-liquefaction case histories in the liquefaction region. Three *MSF* relationships are considered in determining the values of a and b. Equation (F.4), which provides a relationship between the limiting upper  $V_{SI}$  value and fines content, is based, in part, on the case history data and, in part, on penetration shear wave velocity correlations. From penetration- $V_S$ correlations, the recommended  $CRR-V_{SI}$  curves appear to be somewhat more conservative than the penetration-based curves recommended by the 1996 NCEER Workshop (Youd et al., 1997).

## **APPENDIX G**

### **PROBABILITY-BASED LIQUEFACTION EVALUATION**

This appendix presents three probability models for the case history data listed in Appendix H. The probability models are based on the work of Juang et al. (2001a; 2002), and are derived using logistic regression and Bayesian interpretation techniques. They are compared with the deterministic evaluation curve by Andrus et al. (1999) for clean soils ( $FC \le 5$  %) shown in Fig. 2.3. The probability models provide a means of objectively calibrating the deterministic liquefaction curve.

To develop the probability models, values of  $V_{SI}$  are adjusted to a clean soil equivalent. The procedure for adjusting  $V_{SI}$  values involves two steps. First, a *CRR* value is determined using Eq. (2.8) for each case history. Second, for each value of *CRR*, a clean soil equivalent  $V_{SI}$ value is determined using Eq. (2.8) with  $V_{SI}^* = 215$  m/s. Thus, this adjustment procedure, maintains the ratio of *CRR* to *CSR* (or factor of safety). The adjustment procedure can be expressed by:

$$V_{SI,CS} = K_{fc} V_{SI} \tag{G.1}$$

where

 $V_{SI,CS}$  = the equivalent clean soil value of  $V_{SI}$ , and  $K_{fc}$  = a fines content correction to adjust  $V_{SI}$  values to a clean soil equivalent.

Values of  $K_{fc}$  can be approximated using the following equation (Juang et al., 2001a; 2002):

$K_{fc} = 1,$	for $FC \leq 5$ %	(G.2a)
$K_{fc} = 1 + (FC - 5) f(V_{Sl}),$	for $FC = 6 \%$ to 34 %	(G.2b)
$K_{fc} = 1 + 30 f(V_{Sl}),$	for $FC \ge 35$ %	(G.2c)

where

$$f(V_{SI}) = 0.009 - 0.0109 \left(\frac{V_{SI}}{100}\right) + 0.0038 \left(\frac{V_{SI}}{100}\right)^2$$
(G.3)

Equations (G.1) through (G.3) provide an approximate mathematical description of the adjustment procedure. The adjusted case history data are plotted in Figs. G.1 and G.2 along with two probability models determined using logistic regression. The logistic regression-based probability models, as well as a Bayesian-based probability model, are discussed below.

## G.1 LOGISTIC REGRESSION MODELS

### **G.1.1 Logistic Regression Model 1**

The first logistic regression-based probability model, called Model 1, is similar in form to the model used by Liao et al. (1988) for analyzing SPT-based case histories. The probability equation for Model 1 is given by (Juang et al., 2001a; 2002):

$$\ln\left[\frac{P_L}{1-P_L}\right] = a_1 + a_2 V_{SI,CS} + a_3 \ln(CSR_{7.5})$$
(G.4)

where

 $P_L$  = the probability that liquefaction will occur,  $a_1, a_2, a_3$  = regression coefficients, and  $CSR_{7.5}$  = CSR adjusted to  $M_w = 7.5$ .

The mean values of  $a_1$ ,  $a_2$ , and  $a_3$  are 14.8967, -0.0611, and 2.6418, respectively. The standard deviations associated with the coefficients are 2.1637, 0.0098, and 0.4268, respectively. The Nagelkerke coefficient (equivalent to  $\mathbb{R}^2$ ) of this regression is 0.58. Probability curves for Model 1 are presented in Fig. G.1. From the figure, Model 1 appears to provide reasonable  $P_L$  curves within the limits of most of the data. However, the Model 1 curves may be inappropriately too conservative at high values of  $V_{SI,CS}$  (say > 200 m/s), since a corrected velocity of 210 m/s is considered equivalent to a corrected blow count of 30 in clean sands and liquefaction is generally assumed not possible above this value.



Fig. G.1 - Logistic Regression Model 1 and Case History Data Adjusted for Fines Content. (after Juang et al., 2001a)

#### G.1.2 Logistic Regression Model 2

To investigate the influence that the form of a regression equation might have on  $P_L$  curves, the analysis is repeated using a slightly different equation. The probability equation for this second logistic regression model, called Model 2, is defined by (as suggested by William Guthrie, NIST, to R. D. Andrus, June 1998):

$$\ln\left[\frac{P_L}{1-P_L}\right] = b_1 + b_2 V_{SI,CS} + b_3 \ln(CSR_{7.5}) + b_4 \left[\ln(CSR_{7.5})\right]^2$$
(G.5)

where

 $b_1, b_2, b_3, b_4$  = regression coefficients.

The mean values of  $b_1$ ,  $b_2$ ,  $b_3$  and  $b_4$  determined by Juang et al. (2001a; 2002) are 10.0155, -0.0643, -3.9534, and -1.8381, respectively. The standard deviations associated with the coefficients are 2.6102, 0.0107, 2.1738, and 0.6302, respectively. The Nagelkerke coefficient of this regression is 0.61. A coefficient of 0.61 is slightly greater than 0.58, suggesting a slightly stronger correlation for Model 2 than Model 1. Figure G.2 presents  $P_L$  curves defined by Eq. (G.5). The Model 2 curves exhibit steeper-slopes than Model 1 curves above a CSR value of about 0.1. They reach a maximum  $V_{SI,CS}$  value at CSR of about 0.33. Above CSR of 0.33, the curves trend to the left, decreasing in  $V_{SI,CS}$  with increasing CSR. Nevertheless, the results clearly show that  $P_L$  curves determined by logistic regression depend on the form of the regression equation. While Model 2 provides another possible probability model, one would expect  $P_L$  curves to slope towards higher values of  $V_{SI,CS}$  with increasing CSR rather than extend vertically, as suggested by the dashed lines in Fig. G.2.

#### **G.2 BAYESIAN MAPPING MODEL**

Juang et al. (1999) pioneered a Bayesian interpretation approach for mapping factor of safety,  $F_s$ , to  $P_L$ . In their approach, values of  $F_s$  are first determined for the liquefaction and non-liquefaction case histories using a deterministic evaluation curve. The  $V_s$ -based curve shown in Fig. 2.3 is the deterministic curve used in this case. Values of  $P_L$  are then estimated from the probability density functions of  $F_s$  for liquefaction and non-liquefaction case histories using Bayes' theorem. With the assumption of equal prior probability, the  $P_L$ - $F_s$  mapping function can be expressed as (Chen and Juang, 2000; Juang et al., 2000a):

$$P_{L} = \frac{f_{L}(F_{S})}{f_{L}(F_{S}) + f_{NL}(F_{S})}$$
(G.6)



Fig. G.2 - Logistic Regression Model 2 and Case History Data Adjusted for Fines Content. (after Juang et al., 2001a)

where

 $f_L(F_S)$  = the probability density function of the calculated  $F_S$  for the liquefaction case histories, and

 $f_{NL}(F_S)$  = the probability density function of the calculated  $F_S$  for the non-liquefaction case histories.

An analysis of the 225 case histories yields the probability density functions shown in Figs. G.3a and G.3b for the liquefaction and non-liquefaction cases, respectively. Applying Eq. (G.6), the predicted probability of liquefaction for each case history is obtained. Values of  $P_L$  and  $F_S$  for the case histories are plotted in Fig. G.4. The relationship formed by the  $F_S - P_L$  values can be approximated by (modified from Juang et al., 2001a):

$$P_L = \frac{1}{1 + \left(\frac{F_S}{0.73}\right)^{3.4}} \tag{G.7}$$

In Eq. (G.7), a  $F_s$  value of 1 corresponds to points on the deterministic curve. Thus, on average, the Andrus et al. (1999) curve for clean soils (see Fig. 2.3) is characterized with a  $P_L$ value of 26 % based on Eq. (G.7). The value of 26 % is slightly less than 30 % initially determined by Juang et al. (2001a). Subsequent analysis (Juang et al., 2002) revealed that a few of the calculated  $P_L$  values corresponding to low  $F_s$  values (see Fig. G.4) were unreasonably influencing the coefficients given in Eq. G.7. Thus, Eq. G.7 has been modified slightly from the preliminary equation proposed by Juang et al. (2001a).

Figure G.5 compares the  $F_{S}$ - $P_L$  relationship defined by Eq. (G.7) for the  $V_S$ -based recommended curve (Andrus et al., 1999) with the  $F_{S}$ - $P_L$  relationship developed by Juang et al. (2000) for the SPT-based recommended curve (Seed et al., 1985; Youd et al., 2001). There is remarkable agreement between the  $V_{S}$ - and SPT-based relationships. From Fig. G.5, the SPT-based recommended curve is characterized with an average  $P_L$  value of 31 %. These findings suggest that the  $V_S$ -based deterministic evaluation curves are somewhat more conservative than the SPT-based curves.

Equation (G.7) provides an important link between the probabilistic and deterministic methods. One can obtain a family of  $P_L$  curves for probability-based design by combining Eqs. (2.8), (2.12) and (G.7). The family of  $P_L$  curves for magnitude 7.5 earthquakes and soils with  $FC \leq 5$  % is presented in Fig. G.6. These curves, called the Bayesian Mapping Model, slope to the right with increasing CSR, which seems reasonable. They converge to a  $V_{SI}$  value of 215 m/s, the assumed value of  $V_{SI}^*$  for clean soils, at high values of CSR.



(a) Liquefaction Cases



(b) Non-Liquefaction Cases

Fig. G.3 – Probability Density Functions and Calculated Factor of Safety Distributions for the 225 Case Histories.



Fig. G.4 – Calculated Values of  $F_s$  and  $P_L$  for the 225 Case Histories with Best-Fit Relationship (Juang et al., 2002). Note that  $F_s$  is based on the Recommended Deterministic Evaluation Curves Shown in Fig. 2.3, and  $P_L$  is based on Eq. (G.6).



Fig. G.5 - Relationship Between  $P_L$  and  $F_S$  for the  $V_S$ -based Procedure (Juang et al., 2002) and SPT-Based Procedure (Juang et al., 2000a) Determined Using Bayes' Theorem.



Fig. G.6 - Bayesian Mapping Model and Case History Data Adjusted for Fines Content. (Juang et al., 2002)

#### G.3 COMPARISON OF PROBABILITY MODELS

Figure G.7 compares the logistic regression model curves for  $P_L = 26$  % with the Bayesian Mapping Model curve for  $P_L = 26$  %, which corresponds to the deterministic curve developed by Andrus et al. (1999) for soils with  $FC \le 5$  %. The Bayesian Mapping Model curve lies between the two logistic regression curves below a  $V_{Slcs}$  value of about 195 m/s, indicating close agreement between the three probability models. Above 195 m/s, the Bayesian Mapping Model curve closely follows the logistic regression Model 2 curve. Thus, the logistic regression models support the Bayesian Mapping Model in characterizing the deterministic curve proposed by Andrus et al. (1999) and Andrus and Stokoe (2000) as a 26 % probability of liquefaction curve.

The tendency for the  $P_L$  curves to converge to some limiting upper value reflects the tendency of dense soils to exhibit dilative behavior at large strains, causing negative pore-water pressures. It seems reasonable that the  $P_L$  curves should not continue to diverge with increasing  $V_S$ , or penetration resistance, but should converge somewhat to reflect the behavior of dense soils, as suggested by the curves shown in Fig. G.6. The wider spread exhibited in logistic regression-based  $P_L$  curves at high values of  $V_S$  and CSR is believe to be the result of an inherent property of these models, and not a real-world phenomenon. Thus, the Bayesian Mapping Model (Fig. G.6) is considered to be an improvement over the logistic regression models, and is suggested for engineering risk-based design.



Fig. G.7 - Comparison of the Probability-Based Logistic Regression and Bayesian Mapping Models for  $P_L = 26$  % with the Deterministic Curve Developed by Andrus et al. (1999) for Clean Soils. (modified from Juang et al., 2001a)

## APPENDIX H

# SUMMARY OF CASE HISTORY DATA

Table H.1 presents a summary of case history data described in Appendix E, and used in Appendix F to establish the recommended liquefaction resistance curves. This database is expanded and modified from the database presented by Andrus and Stokoe (1997). Most of the modifications are minor with the intent to have the data conform to the guidelines presented in this document. The major modifications are based on new information or correction of an error in calculations. Some case histories included in the earlier database by Andrus and Stokoe have been omitted due to one of the three following reasons: (1) The reported average downhole  $V_s$  measurement is for a depth interval much greater than the identified critical layer. (2) The critical layer is likely older than 10 000 years and contains carbonate. (3) The location of the critical layer or field behavior is uncertain. The case history data presented in Table H.1 are essentially the same as the data presented in the draft guidelines (Andrus et al., 1999), with only a few minor changes. References for the case history data are given in Table E.1.

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Salinas River South, S-R1	Xhole 7		-	6.1 0.3	32 6.0	50	3 sand & sifty sam	d (SP to SM)	4	•	I	~	8.0 14	1.2 12	2.6 0.	93 13	124	0.22	0.24
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Ninata Sita C1	SASW 7	ŝ	-	1.2 0.1	16 1.6	8 6.	5 sand		\$ S	~	~	~	4.5	2.8	0.7	97 11	136	0.15	0.15
Ninata Site C2	SASW 7	Ŀ,	-	1.2 0.1	16 1.2	4	8 sand		<5 <5	~	~	┛	4.0 7	3.4	8.5 0.	11/26	147	0.10	0.0
		$\left  \right $								f	-+-		_	+	+	+			
1975 HAICHENG, PRC EARTHOUAK	W	+	+	_	+	_				+	+		-	+	+	-		$\uparrow$	].
		-	+	C 14	0	u u	5 cand to ality cla		61	6	~	9	8.6 15	9.7 9	0.3 0.	92 14	151	0.13	0.12
Chemical Fibre		<u>.</u>	+-		000	i e	5 clavev elit to sil	v clav	83	~	~	2	6.8 12	4.8 7	3.6 0.	95 10:	111	0.13	0.12
Construction Building		<u>.</u>	+	C			O eity sand to cla	vav silt	90	~	~	4	4.4 8	1.7 4	3.7 0.	97 10	124	0.14	0.13
Fisheries & Shipbuilding		2	-	0.0			B contr cit to cla	vav sitt	42	6	~	3	4.8 8	9.8 5	0.3 0.	97 91	117	0.13	0.13
Glass Fiber	Unole /				2		o olavov elt to el	v clav	92	6	~	2	0.2 19	1.6 10	0.9 0.	90 14:	143	0.13	0.12
Middle School		<u>.</u>	-	0.1	2		C clayer all to all	t class	22	6	6	-	3.0	4.7 3	15.2 0.	98 12	158	0.12	0.11
Paper Mill	Dhole 7	<u> </u>		1.0 0.	121 2.		D CIEVEY SIII 10 SII	y ciay											
Test Tune					ŏ	sposit	Type			۶	ଥ								
1 831 1 YUT					lu	3		A = Alluviał		Œ	= Rec	ent (<	500 year	(8					
Xhole = crosshole seismic test					- 1			ar - Allindal Andal		I	Hol	- Puero	< 10 00 21 00	0 vear	(a)				
Dhole = downhole seismic test					Ì		, hydraulic	AF = AllUVIAI, IIUVIAI		:	5 - 4	2002	} ? /		5				

SCPT = Seismic Cone Penetration Test SASW = Spectral-Analysis-of-Surface-Waves test Susp. = suspension logger test

FD = fill, dumped FU = fill, uncompacted FI = fill, improved

VDF = volcanic debris flow

CSR 7.5 = CSR/(M<sub>w</sub>/7.5)<sup>-2.56</sup> ? = unknown

	┢	$\vdash$	$\left  \right $	-	-	$\left  \right $	General Character	tistics of Critical Laver-		F	╞		Verage	Value	ē.	tical La	Ver		
	-	-	-	-	۴							lo. of							
		S	uf. W	ster	Ľ		·		Average		2	falues		-	Eff.	-			
	Test	<b>-</b>	і 1	ble am	ax. La	/er Th	lck-	Sol	Fres	Deposit		٩	-	/ert.	Vert.				5
Site	Type A	₩, ≹	<u>ŏ</u> <u>~:</u>	pth av	ð.	Dth R	198	Type	Content	Type	Age A	veragel	Depth S	tress	Stress	> P	8 <8	5	7.5
	+-	- 0		6			-		%		+		Ε	Ş	¢	Ē	/m /s/	-	
			#		<u>  </u>						╫╼								
1979 IMPERIAL VALLEY EARTHOUA	Щ.	┝┤	┼╢	$\left  \right $	┝┥	$\left  \right $				+	┢								
Heher Road Channel Fill R1.R2	Xhola A	L.		9	50 0		7 silv and (SM)		00	Ac			2	6 C 8	0 47	- 00			0000
Hahar Road Channel Fill, S-R1	Xhola 8	2	┼╴	0	505		2.7 sity and (SM)		00	2 H	- a			83.0	0.04	1 00 0			0.20
Heber Road Point Bar. R1-R2	Xhole 6	5	-	1.8 0.	50 1.	0	2.4 sand w/silt (SP-SA	8	10	- P	: œ	-	4 6	80.3	45.4	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	200		0.00
Heber Road Point Bar, S-R1	Xhole 6	5	0	1.8 0.	50 1.	0	2.4 sand w/slit (SP-SN	8	10	<b>۲</b>	: œ	4	3.4	60.3	45.4 0	0.98 17	3 2 4	0	0.29
Kornbloom	SASW 6	s	0	2.4 0.	12 2	10	3.5 sendy silt (ML)		75	4	æ	ŝ	4.2	75.1	58.1 0	0.97 10	5 12	0.10	0.07
McKim	SASW 6	Ś		1.5 0.	51 1.	5	3.5 slity sand (SM)		20	¥	æ	4	3.0	54.1	39.7 0	0.98 12	8 16	0 0.43	0.30
Radio Tower	SASW 6	ŝ	-	2.0 0.	21 2.	7	3,4 silty sand to sand	y silt (SM to ML)	35	¥	£	5	4.4	79.2	55.8 0	0.97 9	0	4 0.1	0.13
Valt Canal	SASW 6	Ŋ	0	2.7 0.	12 2	~	2.8 sand w/ slit to slity	v sand (SP-SM to SM)	13	Å	œ	4	4.0	70.4	58.4 0	0.98 10		8 0.0	0.06
Wildlife	SASW 6	, N	-	1.5 0.	13 2.	v v	4.3 silty sand to sand	y silt (SM to ML)	27	*	œ	ŝ	4.7	86.7	55.3 0	11 78.0	4 13	3 0.13	0.09
Wildlife, 1	Xhole 6	ŝ	_	1.5 0.	13 .2.	2	4.3 silty sand to sand	y slit (SM to ML)	27	₽	œ	4	4.6	83.8	53.9 0	0.97 12	7 14	8 0.13	0.09
Wildlife, 2	Xhole 8	ŝ	0	1.5 0.	13 2.	5	4.3 slity sand to sand	y silt (SM to ML)	27	\$	œ	4	4.6	83.8	53.9 0	0.97 12	4 14	6 0.13	0.09
	+	+	+						-		+		-+	1	-+		-	_	
1990 MID-CHIBA, JAPAN EARTHOU	불	+	+		+	+					+	+			+		+		
	404	-		0		-	1.3 eithy cand		00	Æ	a	+-	8	11.4	65.3 0	0.96 15	5 17	0.01	0.04
Owi Island No. 1, Iayer 1	AIOUO						C.U SILV GOILO		20		: 3	†-	14 0 0	0 2 2	124 5 0	77 10			0 04
Owi Island No. 1, layer 2	chole 5	<b>D</b> .		1.3 0.	180	-	3.0 Birly Sand		00	c	╞	-	1.0	0.00	0.43		2	2	5
ACCENTION AND CALIFORNIA	EARTH	N IVK	+-	+	+	+					╞		-		-		+		
1901 MESIMOULOWING CHE				-	+	-					$\vdash$								
Heber Road Channel Fill, R1-R2	Xhole 5	6	0	1.8 0.	02 2.	0	2.7 slity sand (SM)		22	۳	œ	4	3.5	63.2	46.8 0	0.98 13	11	0.0	0.01
Heber Road Channel Fill, S-R1	Xhole 5	6	0	1.8 0.	02 2.	0	2.7 sity sand (SM)		22	4	<b>.</b>	4	3.5	63.2	46.8 0	0.98 13	9 I C	0.0	10.0
Hahar Road Point Bar. R1-R2	Xhole 5	8	0	1.8 0.	02 1	8	2.4 sand w/silt (SP-SA	9	9	*	œ	4	3.4	60.3	45.4 0	0.98 16	14 20	0.0	10.0
Haber Road Point Ber. S-R1	Xhole 5	6	0	1.8 0.	02 1	80	2.4 sand w/silt (SP-SN	6	-	*	<b>E</b> 1	4	4.6	00.3	40.4	1 98.0	20		10.0
Kombloom	SASW 5	6.	1	2.4 0.	36 2.	S.	3.5 sandy slit (ML)		75	÷	I I	0	2 0		1.00	00000	2 4 4		0.0
McKim	SASW 5	8	0	1.5 0.	06 1	ŝ	3.5 slity sand (SM)		20	¥ ÷	τļ	4 u	, , ,		23.7	02.00			0 10
Radio Tower	SASW 5	8	-	2.0 0.	20 2	N	3.4 silty sand to sand	V sitt (SM to ML)	00	ž		0		10.6	58.4	01 AA 10		0 2	0.12
Vall Canal	SASW 5	8	_	2.7 0.	30 2	N.	2.8 sand W/ silt to silt	V Sand (SP-SM 10 SM)	20	24			47	86.7	55.3 0	97 11	4 13	3 0.20	3 0.14
Wildlife	SASW	8		1.5	27 2	- - -	4.3 BILY SEND IO BEIND	y still JOM to ML/	50	2	: œ	4	4.6	83.8	53.9 0	0.97 12	7 14	8 0.26	3 0.14
Wildlife, 1	Xhole 5	8		1.5	2 12	<u>.</u>	a.3 slity sarid to sarid	v eith (SM to MI )	57	Ł	Œ	4	4.6	83.8	53.9 0	0.97 12	4 14	8 0.2(	3 0.14
Wildlife, 2	Xhole 5	8	┛	1.5 0.	27 2	- 0	4.3 SIRY Sand to same				+								
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		-	-	-												<b>.</b>			
Test Type					<b>ژ</b> ل ر	eposi	L Type			¥	ମ୍ମ		1						
Xhole - crosshole seismic test					4			A = Alluviai		œ	Ш Ц Ц Ц Ц	cent (<	500 ye	ars)					
Dhole = downhole seismic test					u. I		II, hydraulic	AF = Alluvial, fluvial WDE = volcorio dobrie fl	mot	I	£ I	locene	(< 10 (	900 ye	ars)				
SCPT = Seismic Cone Penetration	l Test						nediumo 'ii			č	1	Ċ		·2.5(	_				
SASW = Spectral-Analysis-of-Su	rface-V	Vaves	test		<b>بلن</b> د	¶ = ⊃	II, uncompacted			י נ	С. но Но		HV(M,	(0.)					
Susp. = suspension logger test					LL.		, improved			2		UMOL							

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		Н	$\vdash$				-Ge	neral Characteristics of Critical L	ayer			Av	erage V	alue (	or Critic	cal La	/er		
		-				Top			_		ž	<b>o</b>	_	-					
		S	uf. W	ater		of			Average		Ş	ues		-	H.	-			
	Test		lq. Te	ible a	nax.	Layer 7	hick-	Soll	Fines	Deposit		Ľ	Ver	t. V	94.				5
Site	Type N	ш ≩	#.2 De	pth .	Ng.	Depth	Less	Type	Content	Type A	90 Ave	rage De	pth Stre	SS St	ess ro	5 VS	Vs1	8	7.5
		크	7	-	 														
		의	Z	ε	8	E	E		%		_		<del>ک</del> د	- -	ጜ	È	s m/s		
1983 BORAH PEAK, IDAHO EARTH	QUAKE	$\left  \right $	┝╴┥	┟╴┦															
		-									_			$\left  \right $		-			
Andersen Bar, SA-1	MSWS	D.		8.0	0.29	8.0	2.4 S8	ndy gravel (GP-GW)	<5 <5	7	E		1.9 3	0.0	27.8 0.	99 10	5 14	0.25	0.20
Andersen Bar, X1-X2	Xhole 6		-	0.8	0.29	0.8	2.4 88	ndy gravel (GP-GW)	4 5 4	₽	æ	8	2.0 4(	9.6	28.7 0.	99 10	6 14:	0.26	0.21
Goddard Ranch, SA-2	SASW 8	<b>6</b>	┛	1.2	0.30	- 5	2.0 Sa	ndy gravel (GP)	л С	Å	I	2	2.4 45	9.7	37.3 0.	97 12	2 15	0.26	0.20
Goddard Ranch, SA-4	SASW 8	6	╾┤		0.30	1.2	2.0 Sa	ndy gravel (GP)	< 25	¥	I	2	2.4 45	9.8	37.4 0.	97 10	5 13(	0.25	0.20
Mackay Dam, Toe	SASW 6	8	0	2.3	0.23	2.3	2.7 sli	y sandy gravel (GW to GM)	8	~	I	2	2.8 57	2.2	52.3 0.	98 27	0 316	0.16	0.13
North Gravel Bar, Bar Site	SASW 8	0	•	0.	0.46	8.	1.2 SB	ndy gravel	~<5	+ ×	4	2	2.4 51	0.1	<b>36.0</b> 0.	99 20	6 26	0.41	0.33
North Gravel Bar, Terrace	SASW 8	8	•	3.0	0.46	3.0	1.3 sa	ndy gravel	~<5	A	4	2	3.7 7.5	5.2	<b>19.2</b> 0.	98 27	4 301	0.32	0.26
Pence Ranch, SA-1	SASW 6	0.		1.7	0.36	1.8	1.9 gr	welly sand to sandy gravel	<5	Æ	r	4	2.8 57	7.3	16.2 0.	98 10	3 125	0.26	0.23
Pence Ranch, SA-2	SASW 8	8	-	1.5	0.36	1.5	2.8 gri	welly sand to sandy gravel	<5	AF	I	9	3.1 60	.3	14.7 0.	98 9	411	0.30	0.25
Pence Ranch, SA-3	SASW 8	6.	1	1.4	9.36	1.4	1.8 gra	welly sand to sandy gravel	<5	AF	I	3	2.4 45	8.9	16.8 0.	98 10	2 131	0.25	0.23
Pence Ranch, SA-4	SASW 8	6.	+	1.8	0.36	1.8	2.8 gn	welly sand to sandy gravel	<5	AF	I	2	3.1 62	1.1	19.4 0.	98 10	9 13	0.26	0.23
Pence Ranch, SA-5	SASW 6	6.		1.5	9.36	1.5	1.9 gn	ivelity sand to sandy grave!	<5	AF	I	2	3.0 60	0.5	15.6 0.	98 12	3 15	0.25	0.24
Pence Ranch, SA-A	SASW 6	6.	-	2.0	9.36	2.0	1.7 gm	ively sand to sandy gravel	ŝ	₹	I	-	2.8 57	5.5	16.3 0.	98 13	4 16	0.26	0.23
Pence Ranch, SA-B	SASW 6	8		1.5	0.36	1.5	1.7 gra	welly sand to sandy gravel	<ul><li>5</li></ul>	¥	II	0	2.1 36	8.	12.9 0.	99 12	9 170	0.27	0.21
Pence Ranch, SA-C	SASW 6	6		1.5	0.36	1.5	1.9 an	welly sand to sandy gravel	<5	¥	I	~	2.1 36	3.4	12.4 0.	99 10	7 142	0.27	0.21
Pence Ranch. SA-D	SASW 6	0.	-	1.5	0.36	1.5	1.7 gr	welly sand to sandy gravel	<5	Å	Ŧ	2	2.1 35	4.0	13.8 0.	99 13	1 173	0.26	0.21
Parra Barch SA-F	SASW 8	9		1.7	3.36	1.7	1.5 an	welly sand to sandy gravel	<5	Å	I	2	2.3 45	3.3	18.3 0.	99 12	2 155	0.26	0.21
Parry Ranch XD-XF	Xhole 6	G		1.5	3.36	2.8	1.0 01	welly sand to sandy gravel	<5	¥	T	0	3.4 64	2	6.8 0.	98 14	6 176	0.32	0.28
							-							-		-	<u> </u>		
1985 CHIRA-IBARAGI, JAPAN EAR	THOUAKE	-	-	+	<u>+</u> -	f	 			-				$\left  \right $		$\left  \right $			
		-											-+	+	_				
Owi Island No. 1. laver 1	Dhole 6	0	0	1.3	0.05	4.5	3.3 sil	y sand	20	Ŧ		-	3.1 111	4	5.3 0.	96 15	5 17	0.0	0.03
Owi Island No. 1, layer 2	Dhole 6	0	0	1.3	0.05	13.0	<b>3.6 sil</b>	y sand	35	<	 	+ -	4.8 255	8.9	.4.5 0.	77 19	5 18	0.0	0.03
				+	+-		+		_		+			╀	╀	+	1	1	
10/26/85 TAIWAN EARTHQUAKE	EVENT		2	+		Ť	+				╞	<b> </b>		+	+				
10151010000	Yhola 5	e	- -	0.5	0.05	3.7	5.2 sil	y sand to sandy silt (SM-ML)	50	<	Ŧ	0	8.1 114	1.1	15.4 0.	96 13	7 16	0.0	0.03
LOUNTY LOST, LZ-LULU	Xhola 5			0.5	0.05	3.7	5.2 st	y sand to sandy silt (SM-ML)	50	<	T	0	5.1 114		15.4 0.	96 12	7 15	0.0	0.03
LORUNG LOOI, LZ-L/	Yhola 5		, , , ,	0.5	0.05	4.1	4.8 sil	y sand to sandy silt (SM-ML)	50	<	Ŧ	3	8.1 112	-	15.4 0.	96 15	8 19	0.0	0.03
LOWING LOCI, LO-LO	Vhola 5			0 5	0.05	0 6	5.9 sil	v sand to sendy sift (SM-ML)	50	×	H	4	5.3 99	8	0.4	96 14	2 17	0.0	0.03
Lotung LSS1, L8-L4		?	<u> </u>	?	2	3	2							+	+	+	_		
			-		╞╼┥				-	-	-			+	-	+-	-		
							+				╞	-		+		_			
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		-+-	+	╉		+	+			+-	-		-	╀	+-				
		┥	┥	-	1														
						1				Δu									

FU = fill, uncompacted FI = fill, improved FH = fill, hydraulic FD = fill, dumped Deposit Type F = 1 SASW = Spectral-Analysis-of-Surface-Waves test Susp. = suspension logger test SCPT = Seismic Cone Penetration Test <u>Test Type</u> Xhole = crosshole seismic test Dhole = downhole seismic test

AF = Altuvial, fluvial VDF = volcanic debris flow A = Altuvlaf

<del>ABB</del> R = Recent (< 500 years) H = Holocene (< 10 000 years)

CSR 7.5 = CSR/(M\_7.5)<sup>-2.58</sup>

? = unknown

					-			aneral Characteristics of Critical Lay	yer		Η	Ĩ	Average	e Value	for Crit	ical La	Yer			
						0 D					-	Vo. of								- 7
			Sul. M	later		of	-		Average		-	/ahues			Eff.	-		-		- 1
	Test		LIq.   T	able a	max. L	ayer T	hlck-	Soll	Fines	Deposit	-	Ē		Vert.	Vert.	_	_		B	_
Site	Type	M	EN.? D	epthe	BV9.	epth	<b>BS</b>	Type	Content	Type /	A 96 A	verage	Depth	Stress	Stress	> P	s S	8	7.5	
	-			+-		+	1				+		1	Ś	4	+				T
				╞╢	╡	E	≡¦∣		8		-		E	E	Ę	Ē	S TH			-17
11/7/85 TAIWAN EARTHQUAKE		LSS	6	+		+	+			-	===		-†		-+	-+	_			1
Lotuno 1887, 12-15/16	Xhola	5	c	20	80	27	0	the sand to south oilt (CM-MI)	0	-	╡	•								T
Lotuno LSST 12-17	Xhola	2	0	2	200		100	the same to same all (SM-ML)			<b>c</b> 3	<b>n</b> 0		1.411	40.4	.96.1	16	0.0	3 0.0	<u>-</u> I
Lotino I SST 18-13	Xhola			200				the same to same all (CM-ML)			<b>c</b>  3				40.4	.96.1	112	0.0	0.0	Ē
Internal CCT 1 p.1 A	aled Y							the send to sendy sin (Sim-ML)		<	C :			1.4.1	45.4	.96.1	99 19	0.0	3 0.0	-1
FUILIN FOOT , LO'L*		0.0	>	0.0	20.0	0.0	8.0	ILLA SELICO SELICO SILL (SM-ML)	20	<	I	4	5.3	<b>8</b> .8	40.4 0	- 66	12 17	0.0	3 0.0	=
1/16/86 TAIWAN EARTHOUAKE	(EVENT	LSS <sup>-</sup>	€				┝╌┼													11
																				1
Lotung LSST, L2-L5/L6	eloux	8.8	0	0.5	0.22	3.7	5.2	lity sand to sandy silt (SM-ML)	50	<	I	6	6.1	114.1	45.4 0	.96	17 16	7 0.3	4 0.2	<u>iù</u>
Lotung LSST, L2-L7	elorX	8.6	0	0.5	0.22	3.7	5.28	ity sand to sandy silt (SM-ML)	50	•	I	6	6.1	114.1	45.4 0	.96 1	7 15	5 0.3	4 0.2	N)
Lotung LSST, L8-L3	alor X Nole	<b>8</b> .6	0	0.5	0.22	4.4	4.8	ity sand to sandy silt (SM-ML)	50	<	I	9	6.1	114.1	45.4 0	.96 1	99	1 0.3	4 0.2	io.
Lotung LSST, L8-L4	alortX	8.8	•	0.5	0.22	3.0	5.9	ity sand to sandy silt (SM-ML)	50	<	I	4	5.3	99.8	40.4	.98 1	12	6 0.3	4 0.2	4
					+	1				T	╈						+		_	- 1
4/8/86 TAIWAN EARTHQUAKE	EVENT	SSI		-+-	Ť	$\dagger$	$\uparrow$				╈		1	T			-		-	
1	CICHY CICHY		4		200		, c .	the cond to conduct it (CM-MI)	2		1	•	4	1111	45.4.0	00.1	7 1 8	0	000	10
Lorung Loot, L2-L3/L0				2			10	lity cand to candy all (SM-MI)	205		;  <u>-</u>			114.1	45.4.0	96.1	7 15		009	2
L			<b>,</b>					the sand to sandy all (SM-MI )	50		: I	0		1 4 1	45.4.0	98.1	819		000	e e
LOIUNG L351, L0-L3		7 <b>4</b>	- -	2		- 0		the sand to sandy silt (SM-ML)	50	•	; <b>I</b>	• 4	5.3	99.8	40.4 0	96 1	2 17	0.0	6 0.0	i n
Colurig Looi, Lo-L+		5	>	2	5		*				+						-	-		1
5/20/BE TAIWAN EARTHOUAKE	(EVENT	LSS1	7													╞┤				11
	 													•			- :			1
Lotuno LSST. L2-L5/L6	alortX	6.6	0	0.5	0.18	3.7	5.2	lty sand to sandy slit (SM-ML)	50	<	I		9.1	114.1	45.4 0	.96.1	18	7 0.2	8 0.2	<u>o</u> ī
Lotuna LSST. L2-L7	Xhole	8.6	0	0.5	0.18	3.7	5.2	lity sand to sandy slit (SM-ML)	50	<		9	6	114.1	45.4 0	96 1	7 15	0.5	80.0	0
Lotuna LSST. L8-L3	<b>BlortX</b>	6.6	0	0.5	0.18	4.1	4.8	lity sand to sandy silt (SM-ML)	50	<	I I		- i 9	114.1	45.4 0	96.1			2.0	510
Lotung LSST, L8-L4	Xhole	6.6	0	0.5	0.18	3.0	5.9	lity sand to sandy silt (SM-ML)	20	<	T	4	<b>0</b> .3	89.8	40.4	1 98.	2	8 0.X	N.N.	21
		001	10	╎	+	+-	1			+										1.1
5/20/86 TAIWAN EAHIHUUAKE		3	5	+	╉	†-	+									-				T
Lating LST 12-1546	Xhole	6.2	0	0.5	0.04	3.7	5.2	ity sand to sandy silt (SM-ML)	50	<	I		<b>8</b> .1	114.1	45.4 0	.96 1:	1 18			1
Louing 5001 12-17	alotX	6.2	0	0.5	0.04	3.7	5.2	lity sand to sandy silt (SM-ML)	50	<	II:				40.4	000				
Lotino I SCT 18-13	Xhole	8.2	0	0.5	0.04	4.1	4.8	lity sand to sandy silt (SM-ML)	50	٩.	I.	, ,			+.04		0			
LOUGH	Xhole	6.2	0	0.5	0.04	3.0	5.9	lity sand to sandy silt (SM-ML)	50	<	≖	4	5.0	88.0	40.4	0.	2		2	
													1	T	╉		+	+	+	Т
											╞				-	+	+		+	Т
				+	+		T				╞					╞	-	-		
				-	-	1	1													
Test Tuna						Depo	sit Tyr	Đ		<i< td=""><td>믱</td><td></td><td></td><td></td><td></td><td></td><td></td><td></td><td></td><td></td></i<>	믱									
THE THE STATE STATE AND AND						*	=	A = Alluvial			۳ ۳	cent (<	500 y	ears)						
YU018 = CLOSSI JUE SEISI (193							-fill h.	drautic AF = Altuvial, fluvial	-	I	웃	locene	(< 10	000 Xe	IBIS)					
Dhole = downhole seismic test						Ē	í. Í		de flour											

CSR 7.5 = CSR/(M\_/7.5)<sup>-2.56</sup> ? = unknown

F = Tit FH = "fill, hydraulic FD = fill, dumped FU = fill, improved

SCPT = Seismic Cone Penetration Test SASW = Spectral-Analysis-of-Surface-Waves test Susp. = suspension logger test

							Ī	teneral Characteristics of Critical Layer	i				Averag	e Value	lo C	likcal L	ayer -	$\left  \right $	$\left  \cdot \right $	Π
						Top					-	No. of					+	-		٦
			Suf.	Water		of			Average			Values			Eff.			-	_	
	Test		LIQ.	Table	amax.	Layer	Thick-	Soll	Fines	Deposit		ء		Vert.	Vert.				0	Б
Site	Type	M	Eff. 7	Depth	avg.	Depth	ness	Type	Content	Type	Age /	Average	Depth	Stress	Stress	2	S S	2	۲ ۳	S
			7 = ≺								+			-+		-	_	_	-	Τ
			Z # 0	E	8	ε	ε		%		=		E	ŝ	£	-	N/8 m	/8	_	1
7/30/86 TAIWAN EARTHOUAKE (	EVENT	LSS1	[ 12)										F					-		[
																	$\left  \right $			T
Lotung LSST, L2-L5/L6	Xhole	6.2	0	0.5	0.18	3.7	5.2	silty sand to sandy silt (SM-ML)	50	۲	I	e	6.1	114.1	45.4	0.96 1	37 1	67 0.	28 0	17
Lotung LSST, L2-L7	Xhole	8.2	0	0.5	0.18	3.7	5.2	slity sand to sandy silt (SM-ML)	50	<	I	<b>0</b>	6.1	114.1	45.4	0.96 1	27 1	55 0.	28 0	1
Lotung LSST, L8-L3	Xhole	6.2	0	0.5	0.18	4.1	4.8	silty sand to sandy silt (SM-ML)	50	4	I	e	6.1	114.1	45.4	0.96 1	56 1	010	28 0	1
Lotung LSST, L8-L4	<b>Nole</b>	6.2	0	0.5	0.18	3.0	5.9	silty sand to sandy silt (SM-ML)	50	۷	I	4	5.3	99.8	40.4	0.96 1	42 1	79 0.	27 0	17
7/30/86 TAIWAN EARTHQUAKE	EVENT	LSS L	13)										-							
Lotung LSST, L2-L5/L6	Shole	6.2	0	0.5	0.05	3.7	5.2	silty sand to sandy silt (SM-ML)	50	۷	I	3	6.1	114.1	45.4	0.96 1	37 1	<b>87 0</b> .	08 0	05
Lotung LSST, L2-L7	BlohX	8.2	0	0.5	0.05	3.7	5.2	silty sand to sandy silt (SM-ML)	50	<	I	9	6.1	114.1	45.4	0.96 1	27 1	55 0.	080	05
Lotung LSST, L8-L3	Xhole	6.2	0	0.5	0.05	4.1	4.8	sifty sand to sandy silt (SM-ML)	50	•	I	3	6.1	114.1	45.4	0.96 1	56 1	910	08 0	05
Lotung LSST, L8-L4	Xhole	6.2	0	0.5	0.05	3.0	5.9	silty sand to sandy silt (SM-ML)	50	×	I	4	5.3	99.8	40.4	0.96 1	42 1	79 0.	080	05
																		-		
11/14/88 TAIWAN EARTHQUAKE			3T 16)								╡						-	+		
Continue LIST 12-1 6/1 B	Abola	7.6	6	0.5	0.14	3.7	5.2	silty sand to sandy sitt (SM-ML)	50	•	I	9	6.1	114.1	45.4	0.96 1	37 1	57 0.	220	23
I atima I CCT 1 9.1 7	Xhola	7.6	c	0	0.14	3.7	5.2	sity sand to sandy sit (SM-ML)	50	•	I	9	8.1	114.1	45.4	0.96 1	27 1:	55 0.	22 0.	23
Louing Loon	Xhola	7.6	0	0.5	0.14	4.1	4.8	sity sand to sandy sitt (SM-ML)	50	×	I	e	6.1	114.1	45.4	0.96 1	56 1	91 0.	22 0.	23
Loting 1557, 18-14	Xhole	7.8	0	0.5	0.14	3.0	5.9	silty sand to sandy slit (SM-ML)	50	۷	I	4	5.3	9.6	40.4	0.96 1	42 1	79 0.	210	22
																		_		Τ
1987 CHIBA-TOHO-OKI, JAPAN EA	RTHOL	AKE													-+	-+	-+	+	-+	Τ
												Ī								
Sunamachi, Tokyo Bay	Dhole	6.5	0	6.2	0.10	8.2	5.8	sand with silt to silty sand	15	Ŧ	Œ	8	0.0	166.8	138.8	0.92 1	20	0 	80	4
1987 ELMORE RANCH EARTHOUA	- U																┝╌┼╴	$\left  - \right $		
			-			4	C	the cord (StM)	00	AF	α	4	3.5	63.2	46.8	0.981	31 1:	59 0.	030	6
Heber Road Channel Fill, R1-R2	elor X	5.6	•		0.03					<b>P</b>	α	4	3.5	63.2	46.8	0.98 1	33 1(	310.	03 0.	5
Heber Road Channel Fill, S-R1	eloux	8.0 0	0		0.0		2.4	any sand town	07	¥.	Œ	4	3.4	60.3	45.4	0.98 1	64 2(	010	03 0	5
Heber Road Point Bar, R1-R2		20	5	-	200	•		and wellt (SP-SM)	10	Ł	æ	4	3.4	60.3	45.4	0.98 1	73 2	110.	03	5
Heber Road Point Bar, S-R1	Ahole	2.C	-	-		- 0	j	andvelt (MI)	75	¥	œ	S	4.2	75.1	58.1 (	0.97 1	05 1:	0	10	9
Kombloom	MON	2.0			*	2	2		20	AF.	Œ	4	3.0	54.1	39.7 (	0.98 1	26 1(	30 0.	05 0.	8
McKim	SASW	9.9 1			0.		0.0	aity send to sandy silt (SM to ML)	35	¥	£	2	4.4	79.2	55.8 (	0.97	90 1	40	<u>10</u>	05
Radio Tower	MON	201					20	and w/ elit to slitv and (SP-SM to SM)	13	Ł	œ	4	4.0	70.4	58.4 (	0.98 1	011	0 9	10.0	02
Vall Canal	MSN O	ם ה נ		- -	2.10	- u -	4.3	althy cand to sandy slit (SM to ML)	27	¥	æ	S	4.7	86.7	55.3 (	0.97 1	14	33.0	130.	6
Wildlife	Mon is	2.C	2	- -	2.5			alty cand to sandy silt (SM to ML)	27	₹	œ	4	4.6	83.8	53.9 (	0.97 1	27 1/	.0 81	130.	01
Wildlife, 1		0	>		0.10	2 U 1 C	2.4	althy sand to sandy sit (SM to ML)	27	₹	œ	4	4.6	83.8	53.9 (	0.97 1	24 1	0 18	13 0.	5
Wildlife, 2	XNOI8	9.0			0.10	2.3	2													
1						Den	oslt Tv	e		4	90									

<u>Test Type</u> Xhole = crosshole seismic test Dhole = downhole seismic test SCPT = Seismic Cone Penetration Test SASW = Spectral-Analysis-of-Surface-Waves test Susp. = suspension logger test

<u>Deposit Type</u> F = fill A = Alluvial FH = fill, hydraulic AF = Alluvial, fluvial FD = fill, dumped VDF = volcanic debris flow FU = fill, uncompacted F1 = fill, improved

<u>Age</u> R = Recent (< 500 years) H = Holocene (< 10 000 years)

o 7 5 = CSR//M /7.5)<sup>-2,56</sup>

CSR 7.5 = CSR/(M\_/7.5)<sup>-2.56</sup> ? = unknown

		Π	H	H	Н	General Chara	icteristics of Critical Layer-			H	Ī	Average	Value	for Crll	ical La	/ <del> </del>		
	_			+	Top					╡	No. of	+	-+			-		
		Sul	Water	1	0			Average		1	Values			Ħ.	_	-		
	Test	Ľď.	Table &	amax.	Layer 1	Thick-	Soll	Fines	Deposit		E	_	Vert.	Vert.	_	_		₿
Site	Type Mw	E#:2	Depth	BVG.	Depth	ness	Type	Content	Type	Age A	veraget	Depth S	tress 5	Stress	ž P	Vs1	Ð	7.5
		2 = 0	ε	B	ε	E		8		+		ε	ŝ	KPa B	È	8 m/8		
				T														
1967 SUPERSTITION HILLS, CALIFC	INIA EART	HOUAK	U							╞═┿		$\left  \right $						
Heher Bred Chennel Fill B1-B2	Yholo R E	G	•	010	0	10/ F					-+-							
Habar Road Channel Fill S-R1	Xhole 6.5			0 10	20	2.7 sith and /SM		22		I C	4	<u></u>	63.2	46.8	.98 13	1 159	0.15	0.1
Heber Road Point Bar, R1-R2	Xhole 6.5	0	- 1.	0.18	8.1	2.4 sand w/silt (SF	0.SM)	10	F R		4 4	0.0	201.12	45.8 0	91 98 13	191 5	0.15	0.1
Heber Road Point Bar, S-R1	Xhole 6.5	0	1.8	0.18	1.8	2.4 sand w/silt (SH	(WS-d	10	* 4	: a	•	4	6.09	45.4 0	01 08.17	3 2 1 1	0.10	
Kornbloom	SASW 6.5	0	2.4	0.21	2.5	3.5 sandy slit (ML		75	Å	æ	2	4.2	75.1	58.1 0	97 10	5 120	0	0
McKim	<b>SASW 6.5</b>	0	1.5	0.19	1.5	3.5 silty sand (SM	(	20	Ł	œ	4	3.0	54.1	39.7 0	.98 12	6 160	0.16	0.11
Radio Tower	SASW 6.5	0	2.0	0.20	2.7	3.4 slity sand to s	andy slit (SM to ML)	35	¥	£	S	4.4	79.2	55.8 0	.97 9	0104	0.18	0.12
Vall Canat	SASW 6.5	•	2.7	0.20	2.7	2.8 sand w/ silt to	slity sand (SP-SM to SM)	13	₽	π	4	4.0	70.4	58.4 0	.98 10	1 116	0.15	0.10
Wildlife	SASW 6.5	-	1.5	0.20	2.5	4.3 slity sand to st	andy slit (SM to ML)	27	¥	œ	s	4.7	86.7	55.3 0	.97 11	4 133	0.20	0.14
Wildlife, 1	Xhole 6.5	-	1.5	0.20	2.5	4.3 slity sand to su	andy sitt (SM to ML)	27	7	œ	4	4.6	83.8	53.9 0	.97 12	7 148	0.19	0.13
Wildlife, 2	Xhole 6.5	-	1.5	0.20	2.5	4.3 slity sand to si	andy silt (SM to ML)	27	Ŗ	œ	4	4.6	83.8	53.9 0	.97 12	4 146	0.19	0.13
				+	┥					+		+						
1989 LOMA PRIETA, CALIFORNIA E	ARTHOUAK	w		╺╍┾						+		-+	-+	-	-+	_		
		-							1	-	-+-	-						
Bay Farm Island, Dike	SASW 7.0	0	3.6	0.27	3.6	2.8 sand with fines	s (SP-SM)	0	I I	I I	-	2.2	91.9	0.17	.96 20	4 219	0.20	0.16
Bay Farm Island, Dike S-R1	Xhole 7.0	0	3.6	0.27	3.6	2.8 sand with fines	(SP-SM)	-		E I	- - -	<b>4</b> .9	87.1	75.2 0	.97 19	3 207	0.20	0.18
Bay Farm Island, Dike R1-R2	Xhole 7.0	•	3.6	0.27	3.6	2.8 sand with fines	s (SP-SM)	0	E i	II.	4	4.9	81.4	72.4 0	.97 20	5 222	0.19	0.16
Bay Farm Island, Loop	SASW 7.0	-	3.5	0.27	3.5	1.7 sand with fine:	5	<12 <	Ŧ	<b>-</b>		3.8	66.7	61.4 0	.98 12	5 142	0.18	0.15
Bay Farm Island, Loop S-R1	Xhole 7.0	-	3.5	0.27	3.5	1.7 sand with fine:		<12	Ŧ	I I		4.0	0	08.2	8 98.		0.10	2.0
Bay Farm Island, Loop R1-R2	Xhole 7.0	-	3.5	0.27	3.5	1.7 sand with fine:	8	<12	Ŧ	<b>m</b> :			75.6	68.2 0	11 86.	3124	0.19	0.0
Coyote Creek, S-R1	Xhole 7.0	0	2.4	0.18	3.5	2.5 sand & gravel		ŝ	¥;	r :		<b>7</b>	0.0P	01.8	01 18.			
Coyote Creek, R1-R2	Xhole 7.0	•	2.4	0.18	3.5	2.5 sand & gravel		ŝ	¥ 4	r 3	N 0	4 4	2.0	0.00	CI /R.	1 1 1		1910
Coyote Creek, R1-R3	Xhole 7.0	•	2.4	0.18	3.5	2.5 sand & gravel			ł	╘╵╛	4	1 4	2.0.2	90.0 610	07 18	a	0 15	13
Coyote Creek, R2-R3	Xhole 7.0	0	2.4	0.18	3.5	2.5 sand & gravel		<b>°</b>	٤٩		200	•	2.0	50 8 03	07 15	176	000	010
Harbor Office, UC-12	SOPT 7.0	-	1.9	0.25	3.0	1.6 slity sand			- 3		4	- 4	2 4	82.2	04 12	0 1 20	0.12	0.10
Marina District, No. 2	SASW 7.0	-	8.8	0.15	5.8	7.1 sand to siny su	MU (SP-5M)	<b>;</b> ;	Ēđ		•	8 4	117.0	82.2 0	94 10	5 113	0.12	0.10
Marina District, No. 3	SASW 7.0	-	2.9	0.15	5 0	7.1 Sand to Siry St	END (SP 10 SM)	y y	đ	- a		3.9	6.9.9	59.6 0	.98 12	0 137	0.11	0.09
Marina District, No. 4	SASW 7.0	- (		0.10	202	4 1 cand (SP)		\$2 \$2	Dune?	2	-	7.9	140.6	120.5 0	.93 22	0 211	0.10	0.09
Marina District, No. 5	D. MON	>			0 1	4 8 and (CD)		2	5	æ	2	3.5	62.3	54.1 0	.98 11	2 130	0.11	0.09
Marina District, school	Dhole 7.0		20	0.04		4.0 sand (SP)		<5	Ŧ	Œ	9	7.2	131.5	90.4 0	.94 14	8 152	0.21	0.18
Port of Oakland, POU/-1		-   •			2	4 D cand (SP)		<5	Ŧ	æ	2	6.4	115.7	82.9 0	.95 15	7 165	0.21	0.17
Port of Oakland, POO7-2	SASW /.U		200	0.24	5.0	4.0 sand (SP)		<5	₽	œ	e	7.2	31.8	90.7 0	.94 15	1 155	0.21	0.18
Pon or Uakiano, POU-2	201 - 100 201 - 100 201 - 100	•	0	0.24	5.0	4.0 sand (SP)		<5	Ŧ	œ	9	7.0	27.4	88.5 0	.95 14	7 152	0.21	0.18
Port of Oakland, POU/-2, 3-11	Vhole 7 0		3.0	0.24	5.0	4.0 sand (SP)		20 V	Ŧ	æ	9	7.01	27.4	88.5 0	.95 18	1 186	0.21	0.18
POR OF URKIRIO, PUOL-2, NI-112		-																
Test Tvoe					Depo	sit Type			</td <td>읭</td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td>	읭								
<u>vhola – crosshola saismic tast</u>					u u u		A = Alluvial		Œ	= 70	cent (<	500 ye	ars)					
nhole = downhole seismic test					FH	fili, hydraulic	AF = Alluvial, fluvial		I	ድ "	locene	(< 10(	300 ye	ars)				
SCPT - Selemic Cone Penetration	ו Test				F0 =	fill, dumped	VDF = volcanic debris i	flow										
	Mana-Way	ae tee	÷		FUa	fill, uncompacted			C	<b>SR 7.</b>	5 = CSI	RV(M_	7.5) <sup>-2.56</sup>					
SAGW & Specifal-Allarysis-UI-O	1100010				1	11 Immoved			~	= unkr	UMOL							
Susp. = suspension logger lest					ĩ	III, III provod			•	i								

SASW = Spectral-Analysis-of-Surface-Waves test Susp. = suspension logger test

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						Gene	sral Character	istics of Critical Layer	]	Π	$\vdash$	Ĩ	Average	Value	fo Cđ	ical Lay	er	
					Top							No. of				_		
		Suf.	Water	L	10				Average		-	Values	-+		Ef.	-		
	Test	Ľ	Table	amax.	Layer	Thick-		Soll	Fines	Deposit		<u>_</u>	-	/ert.	Vert.			ଞ
Site	Type Mw	E#. 2	Dept	Bvg.	Depth	Ness		Type	Content	Type	Age	veragel	Depth S	tress 5	Stress	۲ S	V81 C	SH 7.5
		- 2 = 0	ε	a	E	E			%		+		E	ç	KPa B	Ě	s/m (s	
Port of Oakland, PO07-3	SOPT 7.0	-	3.0	0.24	5.5	1.5 sand	1 with silt (SP-	SM)	10	Ŧ	æ	-	6.2 1	12.7	81.3 0	.95 17(	3 185 0	21 0.17
Port of Richmond, POR-2	SOPT 7.0	-	3.5	5 0.16	4.0	4.0 silt	(ML)		57	Ŧ	æ	4	5.8 1	02.9	80.9 0	96 14	5 152 0	.13 0.10
Port of Richmond, POR-2	SASW 7.0	-	3.5	5 0.16	4.0	4.0 silt	(ML)		57	Ŧ	α	2	5.3	93.6	76.4 0	.97 11	7 125 0	.12 0.10
Port of Richmond, POR-2, S-R1	Xhole 7.0	-	3.5	5 0.16	4.0	4.0 slit	(ML)		57	Ŧ	œ	2	6.1	109.4	84.1 0	.95 14:	3 150 0	.13 0.11
Port of Richmond, POR-2, R1-R2	Xhole 7.0	-	3.5	5 0.16	4.0	4.0 silt	(ML)		57	Ŧ	æ	2	6.1	09.4	84.0 0	.95 13	5 141 0	.13 0.11
Port of Richmond, POR-3	SOT 7.0	-	3.5	0.16	4	3.0 silt (	o slity sand (i	AL to SM)	>25	Ŧ	æ	<b>6</b>	5.2	93.4	76.3 0	.96 11(	0 118 0	.13 0.11
Port of Richmond, POR-4	SOFT 7.0		5. 5. 6.	0.16	8. F	3.2 slity	sand (ML to	(WS	>32	Ŧ	œ :	0	5.2	93.4	76.3 0	.96 121	3 137 0	.12 0.10
Salinas River North B1-B2 Salinas River North B1-B2	Xhole 7.0		o e	0.15	 ס מ	2.3 SBN	the slift (Mit to S	(M)	44	<	r I		200	2.11	0 8 06 1	90 1 7	153 0	11 0.05
Salinas River North. R1-R3	Xhole 7.0	0	6	0.15	9.1	2.3 sanc	th silt (ML to ;	(W)	44	<	I	0	9.9	77.2	139.8 0	.90 200	0 184 0	11 0.05
Salinas River North, R2-R3	Xhole 7.0	0	6.1	0.15	9.1	2.3 SBNC	ty silt (ML to ;	(W)	44	×	I	0	9.9	177.2	139.8 0	90 19	9 183 0	11 0.09
Salinas River South, S-R1	Xhole 7.0	0	8.1	0.15	6.6	5.3 sanc	1 & silty sand	SP to SM)	14	×	I	8	8.0	141.2 1	122.6 0	.93 13	1124 0	.10 0.05
Salinas River South, R1-R2	Xhole 7.0	0	6.1	0.15	6.6	5.3 sanc	d & silty sand	(SP to SM)	4	V	I	2	8.0	41.2 1	122.6 0	.93 14:	3 141 0	.10 0.09
Salinas River South, R1-R3	Xhole 7.0	0	6.1	0.15	6.6	5.3 sanc	d & silty sand	(SP to SM)	4	4	I	~	8.0	141.2 1	122.6 0	.93 15(	3 151 0	.10 0.09
Salinas River South, R2-R3	Xhole 7.0	0	8.1	0.15	<b>9</b> .9	5.3 sanc	d & slity sand	(SP to SM)	14	<	I	~	8.0	41.2	122.8 0	.93 161	3 159 0	.10 0.05
Sandholdt Road, UC-4 Layer 1	SOPT 7.0	-	-	3 0.25	2.1	0.6 sanc			~	~	I		2.5	44.3	37.4 0	6 66	116 0	.19 0.16
Sandholdt Road, UC-4 Layer 2	SOTI 7.0	•	1.8	3 0.25	5.9	4.1 sanc			4	~	I	4	8.0	48.0	87.2 0	94 20	9 216 0	26 0.22
Sandholdt Road, UC-6	SOPT 7.0	0		0.25	3.0	4.3 sanc			-	~	I	2	4.7	85.5	56.4 0	.97 17	199 0	.24 0.20
Santa Cruz, SC02	SOPT 7.0	-	0.6	3 0.42	1.3	3.2 slity	sand		-357	۷.	I:	÷	2.9	53.6	31.2 0	98 12:	2 161 0	45 0.38
Santa Cruz, SC03	SOPT 7.0	-	5.1	0.42	2.1	2.4 sanc	d to silty sand		-127	•	<b>-</b>  :	<b>1</b> 00	3.4	59.8	47.9 0	.98 14:	0 175 0	33 0.28
Santa Cruz, SC04	SOTI 7.0	0	-	3 0.42	1.8	2.1 silty	sand		-357	•	I:	~	8.2	1.10	41.3	91 86.	1 202 0	33 0.20
Santa Cruz, SC05	SOTI 7.0	0	2.8	0.42	3.0	1.8 silty	sand to sand	V silt	>357	•	<b>I</b> ]	<b>ه</b> ا	4 C		00.00	- 70 - 4		22 0 22
Santa Cruz, SC13	SOT 7.0			0.42	0.2		sand		2002	<	2 1	- - -	0.0	41.2	0 2 00	99 12	169 0	36 0.30
Santa Cruz, SC14		-		0.42	0.0	Alles C. I			;-	"	: I		3.5	62.0	45.4 0	98 11	3 142 0	21 0.18
State Beach, UC-15	0'/ LOS	- -	~ 0	0.20		7 1 2970			-	. ~	I	~	5.5	99.8	68.2 0	.96 16	2 179 0	.22 0.18
State Beach, UC-16	21-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-	- -	; -	0 13	2 4	2.5 altv	sand(SM)	-	24	Ŧ	Œ	-	6.9	27.1	73.3 0	.95 12!	9 142 0	.14 0.11
TI Fire Station, Hedpain		-		0.13	4.5	2.5 silty	sand(SM)		24	Ŧ	æ	-	5.8	07.0	63.7 0	.96 13:	3 150 0	.14 0.11
TI FIRE Station, Gibbs et al.	SASW 7.0		-	1 0.13	4.5	2.5 silty	sand(SM)		24	Ŧ	œ	-	5.3	98.4	59.6 0	96 13	149 0	14 0.11
TI Fire Station	SOPT 7.0	-	1.4	1 0.13	4.5	2.5 slity	sand(SM)		24	Ξi	<b></b>	- •	5.3	98.4	59.6	96 13	0 152 0	13 0.11
TI Fire Station, B1-B4	Xhole 7.0	-	-	1 0.13	4.5	2.5 slity	sand(SM)		24	E	τœ	τ <b>ο</b> α		013	60.9	96 12:	9 146 0	14 0.11
TI Fire Station, B2-B3	Xhole 7.0	-	-	0.13	4.5	2.5 8117	Sand(SM)		24	E	: œ	0	5.5 1	101.3	60.9 0	.96 13:	2 149 0	.14 0.11
TI Fire Station, B2-B4	Xhole 7.0			0.13	4 1 1	VIIIS C. 7	sand/SM)		24	Ŧ	Œ	6	5.5 1	101.3	60.9 0	.96 13	148 0	.14 0.11
Ti Fire Station, B4-B5	Xhole 7.0	- -		0.10		2.5 silv	sand(SM)		24	Ŧ	E	7	5.8	97.0	63.5 0	.96 13(	0 145 0	.14 0.11
TI Fire Station, Portable		-  -		0 14	5	5.2 SBNC	1 to clevey silt	v sand	13	Ŧ	œ	17	9.1	74.7	99.8 0	.92 17:	3 173 0	15 0.12
TI Perimeter, UM03		-		0.14	3.3	2.3 sanc	I to silty sand		2	₹	æ	8	4.6	83.7	62.5 0	97 15	169 0	12 0.10
TI Perimeter, UMUS	SOPT 7.0		-	0.14	2.0	1.7 SBNC	i to slity sand		5	Ŧ	æ	5	2.9	53.4	38.7 0	.98 12	9 157 0	.12  0.10
										•	90							
Test Type						osit Type				4		.,		1020				
Yhole - crosshole seismic test					Ĩ			<b>A = Alluvial</b>		-		cent (<	an ve					
Dhole = downhole seismic test					ΞH	= fill, hydre	aulic	AF = Alluvial, fluvial		***	ድ 	locene	(< 10 (	300 ye	ars)			
SCPT = Saismic Cone Penetration	t Test				F0 =	= fill, dump	pa	VDF = volcanic debris	flow					2 6.11	_			
SASW - Snectral-Analysis-of-Su	irface-Wat	ves te:	st		FU	= fill, uncol	mpacted			0	SR 7.	S = CS	R/(M_/	7.5)****				
Susp. = suspension logger test					" "	fill, improv	ved			~	= nuki	uwou						

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		H	H		H		General Characteristics o	f Critical Layer-			H		Verage	Value	for Crit	Cal La	<u>Y</u> er		
				_	2						-	lo. ol				_			
		-	Sut. V	Vater	o			-	Average		4	alues	_	_	Eff.	_	_		
	Test		Lig.	rable am	ex. Lay	er Thic	c- Soll		Fines 1	Deposit		<u>_</u>	<u>ر</u>	ert. V	'ert.				₿
Site	Type	₹	E#.7	Jepth av	g. Dep	11 705	Type	-	Content	Type /	<b>Age</b>	verageC	Depth St	ress S	tress	> P	8 8	8	7.5
			×:		_	-					+		-			-	_	_	
			N × O	ε	E	ε			8		+		E	æ	¢,	Ē	/m /s	_	
Ti Perimeter, UM09	80H	7.0	-	2.7 0.	14 2.	9.	B sand to sitty sand		14	Ŧ	æ	12	4.6	83.1	64.8 0	.97 14	3 16	0.1	0.09
Ti Approach to Pier, SA-5b	SASW	0.2	0	2.0 0.	14 2.	0 10.	B sand to sand with silt (SP	to SP-SM)	S	œ	æ	0	6.6 1	23.7	78.1 0	94 18	9 20	0	0.11
TI Approach to Pier, SA-6	SASW	0.7	0	2.0 0.	14 2.	0 10.	8 sand to sand with silt (SP	to SP-SM)	2	æ	æ	S	5.8 1	04.0	67.1 0	95 1 5	7 24	0	0 11
TI Approach to Pier, SA-7	SASW	7.0	•	2.0 0.	14 2.	0 10.	8 sand to sand with silt (SP	to SP-SM)	S	E	œ	0	8.8	23.7	78.1 0	98 1	1 20		0 11
Bay Bridge Toli Plaza, S-R1	Xhole	7.0	F	3.0 0.	24 5.	2.	5 sand to slity sand (SP-SM		6,	₽	œ	4	8.4 1	18.0	82.5 0	9.5 1 4	315	000	017
Bay Bridge Toll Plaza, R1-R2	Xhole.	7.0	-	3.0 0.	24 5.	2.	5 sand to sifty sand (SP-SM		6 I	₽	æ	4	6.4 1	16.0	82.5 0	95 14	4 15	0.2	0.17
Bay Bridge Toll Plaza, SFOBB-1	80H	7.0	Ŧ	3.0 0.	24 5.	2.	5 sand to sity sand (SP-SM)		6-	5	æ	2	6.5 1	17.8	83.4 0	95 14	8 15	0.2	0.18
Bay Bridge Toll Plaza, SFOBB-2	<u>В</u>	7.0	-	3.0 0.	24 6.	3.	0 sand to sity sand (SP-SM)		~13	₽	œ	4	7.5	98.9	92.4 0	.94 14	8 15	0.2	2 0.18
		1			4			-					_	_		_	_		
1993 KUSHIRO-OKI, JAPAN EART	HOUAK		ſ	+	+	-					-		-		-				
	ł		1						•	+	╡	•	-					•	
KUBNITO POR, NO. 2	insp.	200	- -				O SEND WILL SILL (SP-SM)						1	8. e	41.9	21 78.	8 2		0.60
Kushiro Port, No. D	0875	8.3	-	1.8 U.	41		Dubs 9		0	2	<u>r</u>	N	4.0	6.28	56.6 0	.97 13	2 16	6.0 1	0.46
		S	1445		╞	-				1	+		+	-		_	_		
1883 HOKKAILXO-NANSEI-OKI, JAI	ANEAL	2	JAR	╀	╇	+					+	1				+		_	
Bonelon Ucuies BH.4	Dhola	a	-		10	0	5 sandy graval		\$2 V	ġ	I	~	2.0	40.4	30.4 0	66	4	0.1	3 0.20
Hakadata Dart No. 1	Siso	8.3	•	1.2	15 2.		0 silty sand to sand with silt	(SM to SP-SM)	20	6	~	~	7.0 1	12.5	55.7 0	94 13	1 15	0.1	3 0.23
Hakodata Dort No 2	Susn.	8.3	0	1.4 0.	15 2.	80	2 sand with silt (SP-SM)		60	6	~	80	6.5 1	07.7	57.8 0	.95 14	3 16	0.1	7 0.22
Hakodata Port No 3	Suso.	8.3	-	1.4 0.	15 1.	=	5 slity sand (SM)		54	6	2	11	7.0 1	17.1	62.2 0	.93 12	4 14	 0	3 0.21
											-		-			-			
1994 NORTHRIDGE, CALIFORNIA (	SARTHO	AM	Ť		_							+	+	-	+		-	_	
	1000	1	•	0 7 0	2		D eithy sand		-10	<	6	e	5.6 1	00.6	78.4 0	97 16	0 17	0.4	0.30
Rory Lane, M-20	55	- n	- -		210		3 sitty sand		-10	<	~	2	5.4	97.0	76.7 0	.96 15	2 16	4.0	0.30
HOIY LANG, M-32	36	2		340	51 3	~	1 silty sand		~10	<	~	-	4.4	78.6	67.9 0	.97 12	9 14	0.3	7 0.28
Hory Larie, m.30	; ;	;	-								+				_	+	_	_	
1995 HYOGOKEN-NANBU, JAPAN	EARTHK	NAM									+				-	-	-		
		1							a	6	α	•	8.8	98.8	62.8 0	96 16	0 18	0.4	9 0,39
Hanshin Expressway 5, 3	Susp.	8.0		2.10.	20		E gravery sury sarra (SM)	(MIC OI	9	2	: œ	4	5.5 1	00.8	64.6 0	.96 11	2 12	3 0.5	3 0.45
Hanshin Expressway 5, 10		20.0			000	» ▼	1 elity candy gravel		-12	£	æ	4	6.5 1	15.7	94.2 0	.95 17	5 17	9.0	9 0.38
Hanshin Expressway 5, 14		200					A riravaliv sifty sand to 0rav	eliv sand	6	£	æ	5	8.8 1	51.6	96.7 0	.92 17	2 17	9.0	0.48
Hanshin Expressway 5, 25		200	- -				a eith sand to sandy gravel	with silt	12	6	~	8	7.2 1	23.1	77.8 0	.94 16	717	0.0	0.48
Hanshin Expressway 5, 29		2.0	- <		000		A elt eand graval		-10	~	I	2	6.6 1	28.5	83.4 0	.94 13	9 1 4	0.1	0.08
KNK		20.0			100		I prevely sand-sandy grave	i with silt	2	6	6	9	7.0 1	26.1	93.8 0	.95 17	5 17	3 0.5	0.41
Kobe-Nishinomiya EWY, 3			- -				Sintavally aand with silt		8	6	~	-	3.3	60.9	49.0 0	.98 20	0 23	0.4	2 0.34
Kobe-Nishinomiya EWY, 17		2 C	5		1000		I gravely send-sendy grave	I with silt		6	6	•	4.0	71.5	60.8 0	98 13	0 14	4.0	7 0.38
Kobe-Nishinomiya EWY, 23	SUSD	A.A		2.8 0.	031 4.	<b>1</b>													
Test Type					٥	posit.	<u>7</u> 00			2	<b>ല</b>								
Vhola – crosshola sajsmic test					ш	= =	A = Altu	viał		œ	# #	sent (<	500 yet	JLS)					
					ū	I = fill.	hydrautic $AF = AI$	liuvial. fluvial		I	유	ocene	(< 10 0	00 yea	(2				

CSR 7.5 = CSR/(M\_/7.5)<sup>-2.56</sup> 7 = unknown

A = Altuvial AF = Altuvial, fluvial VDF = volcanic debris flow

FH = fill, hydraulic FD = fill, dumped FU = fill, uncompacted F1 = fill, improved

F = 1

SCPT = Seismic Cone Penetration Test Dhole = downhole seismic test

SASW = Spectral-Analysis-of-Surface-Waves test Susp. = suspension logger test

							Ĭ	teneral Characteristics of Critical Layer			F		Average	Value	for Cri	tical L	ayer -	-	$\vdash$	
						Top						No. of								
			Suf. 1	Water		Jo			Average			Values			EH.					
	Test		Гd	Table a	emex.	Layer	Thick-	Soit	Fines	Deposit		5	_	/ert. V	/ert.				ซ	Æ
Site	Type	M	E#.7	Depth	avg.	Depth	ness	Туре	Content	Type	Age	Average	<b>Depth</b> S	tress S	tress	P	Vs V	81 G	F 1	S
			1 = Y																	
			N # 0	ε	0	ε	E		*				E	kPa	KPa A	E	n/s m	/8		
Kobe-Nishinomiya EWY, 28	Susp.	6.9	0	2.2	0.65	3.0	3.0	sand with gravel	~	~	~	-	6.01	10.7	74.0 0	0.96 1	80 1	73 0	60	48
Kobe Port, 7C	Dhole	<b>6</b> .9	-	4.4	0.55	7.5	15.0	sand to sifty sand	9	₽	œ	-	15.0 2	99.3 1	95.4 (	0.88	801	35 0	480	39
LPG Tank Yard, Kobe	Dhole	<b>8</b> .9	-	1.5	0.50	2.0	6.0	gravely sand with sit	9	æ	æ	-	4.0	79.3	54.6	98.0		0	48 0	8 6
Port Island, 1C	Dhole	6.9	-	3.0	0.50	4.0	12.2	gravelly sand to sity gravelly send	13	e	œ	-	8.6	71.4 1	16.3	0.88 1	701	640	42 0	34
Port Island, 2C	Dhole	6.9	-	1.5	0.50	5.0	10.5	sand with silt	10	£	æ	-	11.5	32.3 1	34.4	0.87 2	100	86.0	490	40
Port Island, Common Factory	Dhofe	6.9	0	3.2	0.50	3.2	1.8	gravely sand with silt (SP-SM)	9	E	œ	12	9.0	183.3 1	26.4	0.90 2	14 2	020	410	33
Port Island, Common Factory	Susp.	<b>8</b> .9	0	3.2	0.50	3.2	11.8	gravely sand with silt (SP-SM)	8	Ħ	œ	22	9.2	1 88.6 1	29.3	0.901	93 1	83 0	410	33
Port Island, Dhole Array '91	Dhola	8.9	-	2.4	0.50	2.4	14.6	sandy gravel with silt	10	£	œ	2	7.2	40.5	93.0	0.92 1	90 2	010	40 0.	32
Port Island, Dhole Array '95	Dhole	6.9	-	2.4	0.50	2.4	14.6	sandy gravel with slit	10	6	œ	2	7.8	52.8	99.4 (	1 80.0	65 1	65 0	45 0.	36
XOX	Dhole	6.9	•	7.0	0.45	7.0	5.0	silt, sand, gravel	-	8	8	-	8.5	51.6 1	36.9	0.93 1	90 1	76 0	30 0.	24

<u>Test Type</u> Xhole = crosshole seismic test Dhole = downhole seismic test SCPT = Seismic Cone Penetration Test SASW = Spectral-Analysis-of-Surface-Waves test Susp. = suspension iogger test

<u>Deposit Type</u> F = 111 FH = fill, hydraufic FD = fill, dumped FU = fill, uncompacted F1 = fill, improved

A = Altuvial AF = Altuvial, fluvial VDF = volcanic debris flow

CSR 7.5 = CSR/(M<sub>w</sub>/7.5)<sup>-2.58</sup> ? = unknown

<u>Age</u> R = Recent (< 500 years) H = Holocene (< 10 000 years)