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Strong-Motion Duration of Earthquakes

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STRONG-MOTION DURATION OF EARTHQUAKES

by Erik H. Vanmarcke and Shih-sheng P. Lai

July 1977

Sponsored by National Science Foundation Division of Advanced Environmental Research and Technology Grant ENV 76-19021

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Massachusetts Institute of Technology Department of Civil Engineering Constructed Facilities Division Cambridge, Massachusetts 02139

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ABSTRACT

A sample procedure is proposed to estimate the strong motion duration of earthquake ground motions. The proposed strong-motion duration is nearly proportional to the quantity $1_0/e_{max}^2$, in which a_{max} is the maximum ground acceleration and I_0 is the integral of the squared accelerations. The proportionality factor is weakly dependent on the predominant period of the ground motion. The procedure has been applied to 140 horizontal components of strong earthquake ground motions. The results are tubulated and summarized in the form of histograms. The dependence of strong-motion duration on earthquake-magnitude and epicentral distance is also examined.

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PREFACE

This is the tenth report prepared under the research project entitled "Evaluation of Seismic Safety of Buildings," supported by National Science Foundation Grant ATA 74-06935 and its continuation Grant ENV 76-19021.

The purpose of the supporting project is to evaluate the effectiveness of the total seismic design process, which consists of steps beginning with seismic risk analysis through dynamic analysis and the design of st uctural components. The project seeks to answer the question: "Given a set of procedures for these steps, what is the actual degree of protection against earthquake damage provided?" Alternative methods of analysis and design are being considered. Specifically, these alternatives are built around theree methods of dynamic analysis: (1) time-history analysis, (2) response spectrum modal analysis, and (3) random vibration analysis.

The formal reports produced thus far are:

- Arnold, Peter, Vanmarcke, Erik H., and Gazetas, George, "Frequency Concent of Ground Motions during the 1971 San Fernando Earchquake," M.I.T. Department of Civil Engineering Research Report R76-3, Order No. 526, January 1976.
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- Gasparini, Dario A., "On the Safety Provided by Alternate Seismic Design Methods," Supervised by Erik H. Vanmarcke and John M. Biggs, M.I.T. Department of Civil Engineering Research Report R77-22, Order No. 573, July 1977.

The project is supervised by Professors John M. Biggs and Erik H. Vanmarcke of the Civil Engineering Department. They have been assisted by Dr. Stavros Anagnostopoulos, a Research Associate in the Department. The research assistants on the project have been Peter Arnold, George Gazetas, Dario Gasparini, Robert Frank, William Luyties, Richard Haviland, Shih-sheng P. Lai, Ricardo Binder, and James Robinson.

STRONG-MOTION DURATION OF EARTHQUAKES by Erik H. Vanmarcke^a and Shih-sheng P. Lai^b

INTRODUCTION

The duration of strong shaking during an earthquake plays an important role in many seismic engineering problems. The response of very lightly damped linear systems and of yielding or strength-degrading nonlinear systems depends significantly on the duration of shaking. Duration is also a key parameter affecting the likelihood of occurrence of the phenomena of low-cycle fatigue or soil liquefaction during earthquakes.

No single quantitative measure of the strong-motion duration is in common: usage in earthquake engineering. Studies of the dependence of duration on magnitude (Housner, 1965) and on distance and magnitude (Esteva and Rosenblueth, 1964) are not based on quantitative definitions of duration. Two crude but simple measures of duration have been mentioned in the literature. The first defines duration ar the time interval between the first and last peaks equal to or greater than a given level, usually 0.05 g, on the accelerogram (Page et al., 1975). The second definition is based on cumulative energy obtained by adding squared accelerations: duration is the time interval required to accumulate a prescribed fraction of the total energy, for example, 95 percent (Husid et al., 1969) or 90 percent (Trifunac and Brady, 1975). Bolt (1973) suggested that both these definitions could also be applied to sinusoidal components of the earthquake motion occurring within different narrow frequency pands.

In this paper, a new method is proposed for estimating the strong-motion duration of earthquakes, and it is applied to the horizontal components of each of 70 strong-motion records published by C.I.T.

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BACKGROUND

In this section, we examine briefly two well known and closely related functions which describe the frequency content of earthquake ground acceleration. These are helpful in clarifying the relationship between various measures of the severity of earthquake motion and in introducing the proposed approach to obtaining the duration of strong shaking.

The Fourier amplitude spectrum (F.a.s.) of the ground acceleration a(t) is the absolute value of the Fourier transform of a(t).

$$f(\omega) = \begin{vmatrix} \infty \\ \int \\ -\infty \\ -\infty \\ a(t) e^{-i\omega t} dt \end{vmatrix} = \begin{vmatrix} t_0 \\ \int \\ 0 \\ a(t) e^{-i\omega t} dt \end{vmatrix}$$
(1)

in which $\omega = \text{frequency of vibration (rad/sec.)}$, $i = \sqrt{-1}$, and $t_o = \text{length of}$ the digitized accelerogram (in seconds). The <u>squared</u> Fourier amplitude spectrum $f^2(\omega)$, indicates how the total energy in the earthquake motion is distributed over the frequency axis. Its integral over all frequencies is directly related to the total motion "energy" I_o , or the so-called Arias Intensity (1970), as follows:

$$I_{0} = \int_{0}^{t_{0}} a^{2}(t) dt = \int_{-\infty}^{\infty} a^{2}(t) dt = \frac{1}{2\pi} \int_{-\infty}^{\infty} f^{2}(\omega) d\omega = \frac{1}{\pi} \int_{0}^{\infty} f^{2}(\omega) d\omega \qquad (2)$$

The equality in the center of Eq. 2 is Parseval's relation, while the equality on the right side results from the fact that since a(t) is real. F.a.s. is an even function of frequency (i.e., $f(-\omega) = f(\omega)$).

A closely related function is the estimated <u>spectral density function</u> (s.d.f.) $G(\omega)$ which indicates how the ground motion <u>power</u>, or energy per unit time, is distributed over all frequencies. We have (Bendat, 1958):

$$G(\omega) = \frac{1}{s_0} \frac{1}{\pi} f^2(\omega)$$
 (3)

in which \boldsymbol{s}_{0} is the yet to be determined strong-motion duration. It must be

stressed that the exact relationship between the two functions is rather delicate. $G(\omega)$ obtained by Eq. 3 is really an <u>estimate</u> of the true spectral density function. The latter is a statistical quantity describing the <u>expected</u> distribution of the earthquake power over frequency. Implied by Eq. 3 is an idealization of the earthquake as a segment of limited duration (s₀) of a random process with constant average frequency content.

A basic property of the s.d.f. is that the mean square acceleration σ_0^2 (which equals the variance, since the mean acceleration is zero) is obtained by integrating $G(\omega)$ over all frequencies.

$$\sigma_{0}^{2} = \int_{0}^{\infty} G(\omega) d\omega \qquad (4)$$

in which $\sigma_0 = r.m.s.$ (root-mean-square) strong-motion acceleration. Combining Eqs. 2, 3 and 4 yields the expected relationship

$$I_{0} = s_{0} \sigma_{0}^{2}$$
 (5)

In words, the total ground motion intensity I_0 is distributed uniformly, at constant average power σ_0^2 , over the strong-motion interval s_0 .

STRONG-MOTION DURATION

The Arias Intensity I_0 and the maximum amplitudes of acceleration (a_{max}) and velocity (v_{max}) are well-known measures of the severity of an earthquake motion a(t). Approximate procedures also exist for estimating the predominant frequency ω_0 of an accelerogram, e.g., counting the equivalent number of cycles in a(t), or more crudely, evaluating the ratio a_{max}/v_{max} . The question at hand is how to determine quantitatively the strong-motion duration s_0 and the corresponding strong-motion r.m.s. acceleration σ_0 . Eq. 5 suggests that large values of s_0 imply small r.m.s. values σ_0 . For example, when s_0 equals t_0 , we obtain $\sigma_0 = [I_0/t_0]^{\frac{1}{2}}$. For very small values of s_0 , the r.m.s. acceleration can become as large as a_{max} , which happens when $s_0 = I_0/a_{max}^2$. These extreme choices for s_0 are undesirable because they imply values of σ_0 which are not related in a consistent manner to the maximum acceleration a_{max} . The solid line in Fig. 1 shows the peak factor a_{max}/σ_0 as a function of s_0 for the N-5 Component of the El Centro. 1940 earthquake. Note that the ratio (a_{max}/σ_0) may be as high as 7.5 or as low as 1, depending on which value of s_0 is chosen. This observation is the key to the proposed method for evaluating strong-motion durations of earthquake records. The idea is that an approximate relationship must exist between σ_0 and a_{max} . The relationship is of course of a probabilistic nature. The theory of stationary Gaussian random functions provides a prediction of the most probable value of the peak factor a_{max}/σ_0 during a known time interval s_0 of steady strong shaking. Specifically, the value of a_{max}/σ_0 , which is exceeded <u>once on the average</u> during the interval s_0 (or which has a probability e^{-1} of not being exceeded during s_0), is approximately⁺:

$$(a_{max}/\sigma_0) = \sqrt{2 \ln (2 s_0/T_0)} \qquad s_0 \ge T_0 \qquad (6)$$

in which $T_0 = 2\pi/\omega_0$ = predominant period of the earthquake motion, and s_0/T_0 = number of cycles during the time interval s_0 . The lower limit on s_0 insures that the peak factor will not be too low⁺; Eq. 6 is plotted in Fig. 1 for three different values of T_0 : 0.2 sec., 0.3 sec. and 0.4 sec. Note the relative insensitivity of the peak factor (predicted by Eq. 6) to the choice of T_0

Assuming T₀ is known, Eqs. 5 and 6 can be viewed as a system of two equations and two unknowns, s₀ and σ_0 . The solution for s₀ is implicit in the following equation:

$$s_o = [2 \ln (2 s_o / T_o)] (I_o / a_{max}^2) \qquad s_o \ge T_o$$
 (7)

The ratio (I_0/a_{max}^3) is available for all strong-motion earthquake records.

Eq. 6 is derived on the basis of the assumption that the crossings of a specified, relatively high threshold occur as a Poisson arrival process. The formula is inappropriate for very low thresholds. The condition $a_{max} \ge 1.2 \sigma_0$ implies $s_0 \ge T_0$, which is the bound set in Eq. 6.

The solution to Eq. 7 is plotted in Fig. 2. Note that the duration s_0 is nearly a linear function of I_0/a_{max}^2 for a given value of T_0 . For the El Centro, 1940 N - S component, the ratio I_0/a_{max}^2 is 0.975, and T_0 is about 0.3 sec. From Fig. 2, it follows that s_0 is about 7.5 sec. The corresponding r.m.s. acceleration (obtained by using Eq. 5) is $\sigma_0 = 0.12$ g and the peak factor is $(a_{max}/\sigma_0) = 0.33$ g/0.12 g = 2.75. The quantities s_0 , σ_0 and a_{max} are shown in Fig. 3.

The main features of the strong-motion duration s_0 are that (i) the total motion energy I_0 is preserved, and (ii) a consistent relationship between a_{max} and the strong-motion r.m.s. acceleration σ_0 is guaranteed.

INTERPRETATION BASED ON MOVING AVERAGE POWER

The strong-motion duration may also be derived by using a procedure based on moving average power. The procedure yields essentially the same results as those just presented (Eq. 7), but it sheds a different light on the concept and the interpretation of strong-motion duration.

We start by examining the cumulative intensity function I(t) studied by Husid et al. (1969):

$$I(t) = \int_{0}^{t} a^{2}(u) du \qquad (8)$$

It is zero when t = 0 and reaches its maximum value $I_0 = I(t_0)$, the Arias Intensity, at the end of the record a(t). For some earthquake records, the function I(t) increases relatively slowly and smoothly, and has a "flat" average slope. In other cases, the strong shaking phase of the accelerogram is relatively short and intense, and the function I(t) will have a "steep" (average) slope during the time interval of strong shaking. The average slope of the function I(t) is the average moving power:

$$P_{\Delta t}(t_1) = \frac{1}{\Delta t} \int_{t_1}^{t_1 + \Delta t} a^2(t) dt \qquad 0 \le t \le t_0 - \Delta t \qquad (9)$$

in which Δt is the averaging interval. This function peaks where the motion is most intense. An example is shown in Fig. 4. Of course, the appearance of the average moving power will strongly depend on the choice of integration interval Δt . The larger Δt , the smoother the average moving power $P_{\Delta t}(t)$.

Let $\sigma_{\Delta t}^2$ denote the peak value of the moving average power. The corresponding "equivalent duration" $s_{\Delta t}$ is equal to the ratio of the Arias Intensity I_0 to the peak average power $\sigma_{\Delta t}^2$.

$$s_{\Delta t} = I_0 / \sigma_{\Delta t}^2$$
 (10)

As shown in Fig. 4, the effect is to distribute the total energy I_0 uniformly at constant power $\sigma_{\Delta t}^2$. Finally, a peak factor is obtained:

$$a_{max} / \sigma_{\Delta t} = a_{max} / [I_0 / s_{\Delta t}]^{\frac{1}{2}}$$
(11)

For each integration interval Δt , we can compute the r.m.s. acceleration $\sigma_{\Delta t}$, the equivalent duration $s_{\Delta t}$, and the peak factor $(a_{max}/\sigma_{\Delta t})$. For very small integration intervals Δt , the r.m.s. acceleration $\sigma_{\Delta t}$ approaches the maximum acceleration a_{max} , and the peak factor approaches one. At the other extreme, when Δt equals the record length t_o , we have $\sigma_{\Delta t} = \sqrt{T_o/t_o}$.

For each value of Δt , a theoretical prediction can also be made for the peak factor, again based on the probable relationship between the standard deviation and the maximum value during segments of stationary Gaussian motions:

$$(a_{max} / \sigma_{\Delta t}) = \sqrt{2 \ln 2(s_{\Delta t} / T_o)}$$
(12)

 T_{o} is the predominant period, which is initially treated as constant in the procedure. (The initial value of T_{o} is obtained by dividing the time t_{o} by half the number of zero-crossings in the accelerogram).

The theoretical peak factor, plotted against Δt will vary more slowly than the peak factor computed by Eq. 11. The desired peak factor is the ordinate of the point where the two curves intersect. The corresponding values of $\sigma_{\Delta t}$ and $s_{\Delta t}$ may then be obtained from Eqs. 10 and 11.

Actually, T_0 depends on the choice of the time interval within which zero-crossings are counted. For some accelerograms, the predominant frequency is strongly influenced by the motion at the end of the record. Therefore, it seemed worthwhile to implement an iterative computational procedure in which only zero-crossings within the interval $s_{\Delta t}$ (centered at the time the peak acceleration occurs) are counted.

NUMERICAL RESULTS

The procedure just outlined has been applied to two horizontal components from each of 70 strong-motion records published by the California Institute of Technology. The records were selected by McGuire (1977) so as to be reasonably representative of a broad range of earthquake magnitudes, source-to-site distances, and site conditions. He also classified the records into two categories: "rock" (11 records) and "soil" (59 records)⁺.

Table I lists the results for all 140 accelerograms which are labeled as in the C.I.T. computer files. In addition to the strong-motion durations and the corresponding peak factors, Table I lists the Richter magnitude (M), epicentral distance (R), M.M.I. Intensity, maximum acceleration Arias Intensity (I_0) , the predominant period (T_0) and the soil/rock designation for each record.

[†]Magnitudes and epicentral distances are reported by McGuire and Barnhard (1977), and soil designations are given by McGuire (1977).

Table II compares various durations for a subset of records. For each record, the time intervals based on the 5 to 95% energy fraction (as proposed by Trifunac and Brady, 1975), as well as those based on other energy fractions (10 to 90%; 20 to 80%; 25 to 75%; and 33 to 67%), and the proposed strong-motion duration, are listed side by side. Table II suggests strongly that the strong-motion durations cannot be related consistently to the time intervals corresponding by any specific fraction of total energy. By determining the energy fractions which correspond most closely to the strong-motion duration shorter than the ".33 + .57%" time interval and 4 records with a duration exceeding the ".05 + .95%" time interval.

Histograms of the peak factors, the predominant periods, and the strongmotion durations are shown in Figs. 5, 6, and 7, respectively. The peak factors have a mean of 2.67 and a standard deviation of 0.381; for the predominant periods, the mean is 0.34 sec and the standard deviation 0.18 sec; the mean strong-motion duration is 9.27 sec and its standard deviation is 8.76 sec.

For earthquake records with strong-motion durations below 6 seconds (when the slope of I(t) is relatively "steep"), peak factors range between 2.2 and 2.6. For records with strong-motion durations exceeding 6 seconds (when the slope of I(t) is relatively "flat"), peak factors range from 2.6 to 3.2.

Table I contains the information needed to conduct a number of ground motion parameter studies which involve strong-motion duration. In particular, consider the relationship shown in Fig. 8 between strong-motion duration and earthquake magnitude for different ranges of epicentral distance. For comparison, two well-known prediction equations for strong-motion durations are

also plotted in Fig. 6. The solid line (Fig. 8) is the linear relationship suggested by Housner (1965) for "the strong phase of ground shaking" (evaluated subjectively) against magnitude M:

$$s = 11 M - 53$$
 (13)

Esteva and Rosenblueth (1964) estimated the duration of an "equivalent ground motion with uniform intensity per unit time" as follows:

$$s = 0.02 e^{0.74 M} + 0.3 R$$
 (14)

where R is the focal distance in kilometers. The durations predicted by Eq. 14 are plotted (as dotted lines) against magnitude for two values of R in Fig. 8.

CONCLUSIONS

A new and simple procedure is presented for estimating the strong-motion duration \boldsymbol{s}_{n} and the r.m.s. strong-motion acceleration $\boldsymbol{\sigma}_{n}$ of earthquake ground motion records. The strong-motion duration is found to be nearly proportional to the quantity $I_{0}^{}/a_{max}^{2},$ in which $a_{max}^{}$ is the maximum ground acceleration and $I_{\rm c}$ is the Arias Intensity or the integral of the squared accelerations. The proportionality factor is equal to the square of the peak factor (the ratio of the maximum to the r.m.s. strong-motion acceleration); the latter averages about 2.65, and the proportionality factor about 7. A less important factor influencing the relationship between s and I_0/a_{max}^2 is the predominant period of the strong phase of the accelerogram. The main features of the proposed procedure to obtain the strong-motion duration s_{o} and the corresponding r.m.s. strong-motion acceleration $\sigma_{_{\rm O}}$ is that (i) it preserves the motion intensity J_0 , i.e., $I_0 = \sigma_0^2 s_0$, and (ii) it guarantees a consistent and predictable relationship between a_{max} and σ_0 , i.e., $(a_{max}/\sigma_0) = \sqrt{2 \ln 2(s_0/T_0)}$. This expression predicts the maximum acceleration which will occur once on the average during the strong-motion time interval s_n.

The proposed procedure could also be used to obtain equivalent durations for different frequency-filtered components of earthquake records as suggested by Bolt (1973). For example, one may wish to restrict attention to long-period ground motion components, or to components operating within a narrow band of frequencies in the neighborhood of the fundamental frequency of a structure of interest.

The proposed procedure has been applied to 140 horizontal components of strong earthquake ground motions. The results are tabulated and summarized in the form of histograms of strong-motion durations and peak factors. The dependence of strong-motion duration on earthquake magnitude and epicentral distance is also examined.

It is believed that the proposed definition of strong-motion duration fills a significant gap in earthquake engineering research. It will permit a more confident approach to a number of important seismic engineering problems in which strong-motion duration plays a key role. It is hoped that it will also be of service in the development of improved procedures for specifying design ground motions by permitting quantitative treatment of the correlation between duration and other ground motion parameters, as a function of earthquake source parameters and distance.

ACKNOWLEDGEMENTS

The authors gratefully acknowledge the assistance of C.A. Cornell, T. Shakal and H. Banon in acquiring the earthquake records used in this study. The information on duration (presented in Tables 2 and 3) based on different fractions of the Arias Intensity were obtained by C.A. Cornell and T. Shakal, and have not yet been published elsewhere.

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or recommendations expressed in this publication are those of the authors and do not necessarily reflect the views of the National Science Foundation.

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LIST OF FIGURE CAPTIONS

Figure 1	Predicted	and	Compu	ted Valu	es of	the	Peak	Factor	for	the	N-S
	Component	oft	the El	Centru,	1970	Eart	thqual	ke Recol	rd.		

- Figure 2 Relationship between the Proposed Strong-Motion Duration s_0 and the Parameter I_0/a_{max}^2 for any Accelerogram.
- Figure 3 Maximum Acceleration (a_{max}), Strong-Motion Duration (s_c), and R.M.S. Strong-Motion Acceleration for the N-S Component of the El Centro, 1940 Earthquake Record.
- Figure 4 Moving Average Power $P_{\Delta t}(t)$ whose Maximum Value is $\sigma_{\Delta t}^2$.
- Figure 5 Histogram of Peak Factors.
- Figure 6 Histogram of Predominant Periods.
- Figure 7 Histogram of Strong-Motion Durations.
- Figure 8 Strong-Motion Duration as a Function of Richter Magnitudes for Different Ranges of Epicentral Distance.

















Figure 8

TABLES

	EARTHQUAKE	COMP.	Σ	R (km)	IW	a _{max} (g)	$ \left[10^{4} \frac{\text{cm}^{2}}{\text{sec}^{3}}\right] $	t (sec)	PEAK FACTOR	s o DURATION	SITE s = soil r ≅ rock
Northwes Califur:	st Dia	S44W N46W	5.0	53	ß	0.104 0.112	0.5616 0.7047	0.18 0.24	2.76 2.65	4.11 4.12	s s
Kern Co	unty	SOOE S90M	7.7	127	7	0.047 0.053	0.4068 0.7132	0.44 0.52	2.92 3.00	16.00 23.59	s s
Kern Co	unty	N21E S69E	7.7	43	7	0.156	3.4020 3.6887	0.26 0.24	3.05 3.00	13.66 10.72	ທ່າ
Kern Co	ounty	N42E S48E	7.7	68	7	0.090	1.485 1.802	0.43 0.66	2.95 2.45	16.80 6.62	s s
Kern C	ounty	SOOM N90E	7.7	611	2	0.055 0.044	0.5191 0.6473	0.40 0.40	3.03 3.21	19.32 35.41	s s
Kern C	ounty	SOOM N90E	7.7	611	7	0.059	0.6343 0.6372	0.35 0.37	3.03 3.28	17.09 40.08	s s
Eureka		SOOM N90E	6.6	24	7	0.168 0.258	2.1154 4.4298	0.32 0.42	2.66 2.45	5.47 4.23	s s
Eureki		N44E N46W	6.6	40	7	0.159	3.4131 2.3542	0.39 0.37	2.86 2.42	11.53 3.52	s s
San Ji	ose	N31W N59E	5.8	10	7	0.102 0.108	0.4675 0.3323	0.25 0.19	2.52 2.43	2.97 1.78	s s
San Franc	isco	NO9M NB1E	5.3	15	9	0.043 0.046	0.1012 0.8769	0.21 0.27	2.70 2.40	4.07 2.42	in ni
San Franc	isco	N10E S80E	5.3	12	6	0.083	0.1751 0.3114	0.21 0.16	2.27 2.54	1.34 2.02	55

TABLE I

SITE s = soil ION r = rock	66 S S S	8 s	35 s 	55 s s	30 73 s	94 s	53 s 76 s	30 19 s	74 S 64 S	3 4 s	67 S 01 S
so DJRAT	5.7 5.7	0.4 	18.0	31.10	18.12	5.0	mvò	13.	'; 	1.	6.4
PEAK FACTOR	2.42	2.51 2.74	2.94	2.73	2.87	2.43	2.45	3.08	2.27	2.57 2.88	2.93 2.67
T ₀ (sec)	0.29 0.26	0.21	0.47 0.42	0.34	0.42	0.30	0.35	0.23	0.27	0 32 0.36	0.27 0.23
$\begin{bmatrix} I_0 \\ 10^4 \frac{\text{cm}^2}{\text{sec}^3} \end{bmatrix}$	0.3302 0.2122	0.0600 0.0360	0.8371 1.6319	1.5104 0.9603	0.1375 0.1818	1.4159	0.0617 0.0649	3.4234 3.8876	0.6262 0.4604	0.2566	0.3256 0.3296
a _{max} (g)	0.085	0.040 0.624	0,065	0.130	0.030 0.029	0.133	0.033	0.160	0.144	0.062	0.054 0.068
IMM	ور	وب	2	Q	<u>ں</u>	Q	ഹ	9	9	و	9
R (km)	14	24	21	64	96	53	38	64	55	104	43
Σ	5.3	5.3	5.6	6.5	6.5	6.3	5.4	6.5	5.5	6.6	5.4
COMP.	S09E S81W	N26E S64E	MLOS M68N	MOCS	M00S	SORM N82W	NOOE N90W	NOOS SOOM	N45E S45E	N45E S45E	N44E S46E
EARTHQUAKE	San Francisco	San Francisco	Hollister	Borrego Mountain	Borrego Mountain	Leng Beach	Scuthern California	Lower California	lst N.W. California	2nd N.W. California	Northern California
C.1.T. NO.	A016-1 A016-2	A017-1 A017-2	A018-1 A018-2	A019-1 A019-2	AG20-1 A020-2	2021-1 8021-2	8023-1 8023-2	B024~1 B024-2	8026-1 8026-2	8027-1 8027-2	8030-1 8030-2

TABLE I (Continued)

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SITE s = soil r = rock	N N	ν v	N N	s S	<u>5. 5</u>	See See	v) v)	۲ F	s s	2 2	ν v
s _o DURATION	3.64 4.12	1.66 0.95	2.37	12.60 9.55	1.39 0.65	8.27 8.07	4.77 5.58	7.01 5.65	9.93 22.88	2.71	4.77 6.14
PEAK FACTOR	2.47	2.26 1.80	2.59	3.12 3.09	2.27 1.65	2.92 2.92	2.58 2.60	2.77 2.65	2.78 3.07	2.43 2.85	2.86 2.94
To (sec)	0 34 0.26	0.26 0.38	0.16 0.18	0.19 0.16	0.21	0.23 0.23	0.34 0.38	0.30	0.41	0.28	0.16 0.16
$\left[10^{4} \frac{\text{cm}^{2}}{\text{sec}^{3}}\right]$	0.2434 0.2653	3.8963	1.9147 2.3601	0.3357 0.3855	1.9059 2.8140	0.0195 0.0123	0.0300 0.0314	0.1508 0.1737	7.9804 4.2129	4.3301 6.1760	0.0425 0.0451
a ^m ax (g)	0.065 0.068	0.355	0.237 0.275	0.053 0.064	0.269 0.347	0.014	0.021 0.020	0.041 0.046	0.255	0.135	0.027 0.025
IWW	و	9	9	9	9	Q	ъ	ц.			
(km) (km)	43	2	0	38	Q	77	5]	134	20	59	6 8
X	5.9	5.3	5.3	5.3	5.3	5.3	5.3	6.5	6.6	6.6	6.6
COMP.	N21E S69E	NO5W N85E	N50E N40W	N50E N40W	N65W S25W	N36H S54W	S11E N79E	N33E N57W	MOON	N21E N69W	S00W N90E
LARTHQUAKE	Wheeler Ridge	Parkfield	Parkfield	Parkfie!d	Parkfield	Parkfield	2nd Northern California	Borrego Mountain	San Fernando	San Fernando	San Fernando
C.1.1. NO.	8031-1 8031-2	8034-1 8034-2	B035-1 B035-2	B036-1 B036-2	B037-1 B037-2	B038-1 B038-2	B039-1 B039-2	B040-1 B040-2	C048-1 C048-2	D056-1 D056-2	E071-1 E071-2

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SITE s = snil r = rock	v v	s s	s sa	s s	ທ ທ	ŚŚ	v v	s s	s s	in vi	ט ט
s _o	8.09	16.42	17.63	13.90	8.99	13.28	1.37	48.70	6.86	2.03	2.79
Duration	11.09	7.51	20.24	17.72	16.26	11.37	9.95	35.85	6.10	2.80	3.33
PEAK	2.77	3.11	3.12	3.00	2.78	2.96	2.36	3.13	2.73	2.33	2.43
FACTOR	2.96	2.98	3.15		3.08	2.85	3.04	3.14	2.79	2.33	2.54
T _o	0.35	0.26	0.27	0.30	0.38	0.33	0.17	0.73	0.33	0.27	0.29
(sec)	0.27	0.18	0.28	0.39	0.29	0.40	0.20	0.51	0.25	0.37	0.27
$ \begin{bmatrix} I_0 \\ 10^4 \frac{cm^2}{sec^3} \end{bmatrix} $	1.1561 0.E030	2.0904 1.5509	0.6205 0.4517	0.1415 0.1523	0.0060	0.0889 0.1031	0.0233	0.0907 0.0799	0.1332 0.1005	2.0100 1.5176	0.6519 0.6519
a _{max} (g)	0.107 0.082	0.113 0.139	0.060	0.031 0.028	0.007	0.025 0.028	0.031	0.014	0.039 0.037	0.238	0.121 0.116
IWW			~	9	Q	4	Q	9	2	œ	9
R (km)	46	33	48	30	Ξ	150	22	147	80	16	50
Σ	6.6	6.6	6.5	5.6	5.5	6.3	2	6.3	5.75	5.9	6.4
COMP .	N83E	S60E	North	North	North	North	North	North	N45W	N45E	N45W
	S07W	S30W	East	East	East	East	East	East	S45W	S45E	S45W
EARTHQUAKE	San	San	Borrego	Imperial	Imperial	Lower	Imp. County	Gulf of	Humboldt	Santa	Northern
	Fernando	Fernando	Valley	Valley	Valley	California	Foreshock	California	Bay	Barbara	California
C. I. T.	F086-1	G114-1	T286-1	T287-1	T288-1	T289-1	T290-1	T293-1	U298-1	U299-1	U300-1
NO.	F086-2	G114-2	T286-2	T287-2	T288-2	T289-2	T290-2	T293-2	U298-2	U299-2	U300-2

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C.I.T. NO.	EARTHQUAKE	COMP.	Σ	R (km)	IWW	a _{max} (g)	$ \begin{bmatrix} I_{0} \\ 10^{4} \frac{\text{cm}^{2}}{\text{sec}^{3}} \end{bmatrix} $	T _o (sec)	PEAK FACTOR	s _o Duration	SITE s = soil r = rock
U301-1 U301-2	Northern California	W89W Sotu	5.3	21	7	0.197 0.122	1.0933 1.1072	0.39 0.38	1.66 2.57	0.78 5.08	ט ט
U305-1 U305-2	Central California	N89M S01M	5.3	27	5	0.053	0.2983 0.3082	0.32	2.83 2.91	8.70 10.82	s s
U207-1 U307-2	Central California	NB9W S01W	5.0	Q	ور	0.057 0.036	0.2963 0.1514	0.35	2.73 2.90	7.12 10.25	s s
U308-1 U308-2	Northern California	N46W S44W	5.7	29	Q	0.059 0.075	0.2011 0.2351	0.26	2.62 2.38	4.00 2.54	s s
U309-1 U309-2	Central Cal. Aftershock	M89M NLOS	5.6	21	~	0.172 0.076	1.0012 0.6454	0.37 0.43	1.98 2.71	1.43 8.46	s s
U311-1 U311-2	Parkfield	N21E Sége	5.3	131	4	0.008	0.0134 0.0234	1.00 0.76	2.73 2.67	21.04 13.49	s s
U312-1 U312-2	Ferndale	N46W S44W	5.8	32	Q	0.105 0.237	0.4672 0.6584	0.27 0.20	2.46 1.48	2.66 0.30	<i>ა</i> თ
U313-1 U313-2	Northern California	M68N SOTW	5.2	39	ъ	0.013	0.0357 0.0448	0.43	2.99 2.92	18.60 14.74	ທທ
V314-1 V314-2	Long Beach	N39E N51W	6.3	59	~	0.064	1.0251 1.4550	0.72 0.39	2.86 2.91	21.62 13.48	<u>۲</u> ۲
V315-1 V315-2	Long Beach	South	6.3	27	7-9	0.196 0.159	2.9709 2.5643	0.30	2.71 2.90	5.8] 8.88	v v
V316-1 V316-2	Torrance Gardena	North East	5.4	ى	و	0.040 C.055	0.1376 0.2399	0.35 1.54	2.69 1.17	6.47 1.52	s s

SITE s = soil r = rock	л v	<u>د د</u>	s S	s S S	<u>د</u> د	N N	<u>د</u> د	5- 5-	5 5	s sa	v v.
s _o DIJRATION	7.07 13.64	3.40 7.05	0.67 0.49	2.31 1.52	1.63 1.27	8.59 9.07	2.42 1.26	2.98 2.26	3.36 1.90	7.32 9.12	8.18 7.05
PEAK FACTOR	2.69 2.92	2.49 2.73	1.51 1.19	2.36 2.05	2.39 2.39	2.78 2.71	2. 4 3 2.30	2.80 2.66	2.74 2.53	2.85 2.93	2.91 2.92
T ₀ (sec)	0.38 0.38	0.31	0.43	0.29 0.38	0.19	0.37 0.46	0.25 0.18	0.12	0.16 0.16	0.25	0.24 0.20
$\begin{bmatrix} I_0 \\ 10^4 \frac{\text{cm}^2}{\text{sec}^3} \end{bmatrix}$	0.0215 0.0201	0.1532 0.1183	0.6101 0.1506	0.0872 0.0782	0.0465 0.0283	0.0235 0.0189	0.8279 0.9002	0.185 4 0.0950	0.1412 0.1466	0.1452 0.1328	0.0364 0.0287
a _{max} (g)	0.015	0.054 0.036	0.167 0.089	0.046 0.048	0.041	0.015	0.142 0.198	0.071	0.057 0.071	0.041	0.020 0.019
IWW	Q	9	Ŷ	ę	ŝ	9	9	5-6	2-6	Q	2
R (km)	9	17	Ŷ	61	18	151	13	16	22	18	56
Σ	5.4	Q	ŝ	S	4.5	6.25- 6.5	5.4	5.4	5.4	5.4	5.4
COMP.	S50E S40W	N36W S54W	South West	N79E S11E	Sou th East	South East	S65E S25W	S85E S05W	S54E S36W	South East	North East
EARTHQUAKE	Torrance Gardena	Southern California	Southern California	Northern California	Southern California	Northern Californi a	Lytle Creek	Lytle Creek	Lytle Creek	Lytle Creek	Lytle Creek
C.I.T. NO.	V317-1 V317-2	V319-1 V319-2	V329-1 V329-2	V330-1 V330-2	V331-1 V331-2	V332-1 V332-2	W334-1 W334-2	W335-1 W335-2	W336-1 W336-2	W339-1 W339-2	W342-1 W342-2
				-							

SITE s = scil r = rock	s s	აა	s s	S S
50	15.41	34.50	12.01	22.21
DURATION	6.57	38.51	13.95	18.62
PEAK	3.04	3.09	2. 6 3	2.87
FACTOR	2.74	3.09	2.80	2.80
T o	0.31	0.73	0.75	0.72
(sec)	0.31	0.65	0.55	0.73
$\begin{bmatrix} 1 \\ 0^{4} \frac{\text{cm}^{2}}{\text{sec}^{3}} \end{bmatrix}$	0.0762	0.0660	ρ.0095	0.0905
	0.0704	0.0560	0.0087	0.0813
a _{max}	0.022	0.013	0.008	0.019
(g)	0.629		0.007	0.019
INN	ى	ۍ.	9	ъ
(km)	144	174	205	214
Σ	6.5	6.5	6.5	6.5
COMP.	South	S04F	S82E	N83W
	East	S86W	S08W	S07W
EARTHQUAKE	Burrego	Borrego	Borrego	Borrego
	Mountain	Mountain	Mountain	Mountain
C.1.T.	Y370-1	Y371-1	Y373-1	Y379-1
NO.	Y370-2	Y371-2	Y373-2	Y379-2

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C.1.T.	DURAT	IONS CORRE	SPONDING T	O FRACTION	OF I _O	
No.	5 - 95%	10 - 90%	20 - 80%	25 - 75%	33 - 67%	DURATION S
B037-1	5.48	3.08	0.70	0.32	0.18	1.39
B035-1	13.72	8.68	4.56	3.76	2.10	2.37
A016-1	27.48	5.54	1.92	1.16	0.84	2.76
B021-2	18,20	11.44	6.78	5.44	3.92	2.91
A 0 08-1	13.00	6.08	2.66	2.04	1.42	5.47
G114-2	18.94	15.80	9.86	8.58	5.40	7.51
A004-2	28.82	20.10	10.26	8.16	4.72	10.72
T288-2	38.14	34.46	23.66	20.00	7.64	15.26
Y379-1	49.64	42.26	27.40	25,28	15.58	22.21
A003-2	26.84	20.04	13.38	10.00	7.32	23.59
A019-2	39.82	32.06	19.28	14.62	19.94	31.55
A007-2	32.08	27.14	18.32	14.12	7.54	40.08

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TABLE II

Comparison of Durations Corresponding to Various Fractions of Total Energy and the Proposed Strong-Motion Duration s_0 for 12 Strong-Motion Earthquake Records.