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The John A. Blume Earthquake Engineering Center

Department of Civil Engineering
Stanford University

11/6/78

**A BAYESIAN APPROACH TO SEISMIC
HAZARD MAPPING; DEVELOPMENT
OF STABLE DESIGN PARAMETERS**

by

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This research was partially
supported by

Organisme de Contrôle
Technique de la Construction
Algiers, Algeria

and by The John A. Blume
Earthquake Engineering Center
Department of Civil Engineering
Stanford University
Stanford, California



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**Report No. 28
March 1978**

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REPORT DOCUMENTATION PAGE	1. REPORT NO. NSF/RA-780321	2.	3. Recipient's Accession No. PB290374
4. Title and Subtitle Bayesian Approach to Seismic Hazard Mapping, Development of Stable Design Parameters		5. Report Date March 1978	
7. Author(s) C.P. Mortgat, H.C. Shah		6.	
9. Performing Organization Name and Address Stanford University Department of Civil Engineering John A. Blume Earthquake Engineering Center Stanford, California 94305		8. Performing Organization Rept. No. 28	
12. Sponsoring Organization Name and Address Applied Science and Research Applications (ASRA) National Science Foundation 1800 G Street, N.W. Washington, DC 20550		10. Project/Task/Work Unit No.	
15. Supplementary Notes		11. Contract(C) or Grant(G) No. (C) (G)	
16. Abstract (Limit: 200 words) The Bayesian procedure, developed in this research and applied as a case study to Nicaragua, has the advantage of combining all the available historical, seismological and geological information in a rational and consistent way. The mapping parameters considered for this study are the peak ground acceleration and the duration of the strong ground motion. In order to be consistent with the probabilistic approach of the model and to take into consideration the large scatter for attenuation data, probabilistic treatment of the attenuation relationships in the form of coefficient of variation is presented. Comparison is made between the iso-acceleration maps developed in this study and previous results. An in-depth numerical study of 97 accelerograms is also presented in this report. This numerical study was conducted to better understand the statistical behavior of input accelerograms and response time histories and to develop better design parameters for use by engineers. Various possible applications and uses of the parameters developed are suggested.		13. Type of Report & Period Covered	
17. Document Analysis a. Descriptors Bayes theorem Attenuation b. Identifiers/Open-Ended Terms Nicaragua c. COSATI Field/Group		Accelerometers Reaction time Seismic waves Mapping Earthquakes	
18. Availability Statement NTIS.	19. Security Class (This Report)	21. No. of Pages 241	
	20. Security Class (This Page)	22. Price A11-AD1	

CAPITAL SYSTEMS GROUP, INC.
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ACKNOWLEDGMENTS

The authors of this report wish to express their gratitude to Professor J. R. Benjamin, Professor J. M. Gere and Professor T. C. Zsutty for their comments and criticisms. Many hours of discussions held with Dr. Anne Kiremidjian were helpful in developing the Bayesian risk model.

The partial support provided by the Contrôle Technique de la Construction of Algeria and the John A. Blume Earthquake Engineering Center is gratefully acknowledged.



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CHAPTER I
INTRODUCTION

1.1 Introduction

The energy released by earthquakes propagates in the earth's crust as body and surface waves. The intensity and duration of shaking of structures located in the path of these waves depends upon the intensity and duration of the seismic ground motion along with the characteristics of the structure. Structural failures resulting in considerable damage and loss frequently occur because of these motions and inadequate seismic resistance of the structures.

Earthquake engineers and planners often use the words risk and hazard interchangeably in their work. Seismic risk is regarded by many to be synonymous with seismic hazard. There is some danger in this ambiguity since these two words for seismic phenomenon have different meanings. Seismic hazard is defined as "expected occurrence of future adverse seismic event (earthquake)". Seismic risk is defined as "expected consequences of future seismic event". Consequences may be life loss, economic loss, function loss and damage. Loosely, it can be said that a seismic hazard involves "nature's punch" while a seismic risk involves interaction between "nature's punch" and human activity.

The intensity and duration of future earthquake ground motions are random and can therefore be known only in the probabilistic sense of the likelihood of exceeding a given level during a given time period. (Rosenblueth and Esteva, 1966; Benjamin, 1968; Cornell, 1968*). If economic planning and engineering design criteria are to be formulated

* References are given at the end of the report.

on a rational basis, then it is necessary to have the best available estimates of these future ground motions. The best practical representation of earthquake loadings for a given geographical region is in the form of seismic hazard maps -- where the earthquake effect is shown in terms of the most useful engineering parameters for design. Presently there is a great need for improvements in risk mapping techniques and in the description of the related engineering parameters.

Therefore the present dissertation is divided in two parts. The first part concentrates on seismic hazard mapping which can be best defined as the exposure to seismic loading at a given location. This exposure is expressed in terms of an effect and the probability of its occurrence. The second part concentrates on a study of stable design parameters. Its general purpose is to provide a statistical and probabilistic view of the response of structures to earthquake excitation. The attention is focused on response parameters which have a direct engineering value.

Many attempts have been made to quantitatively describe the intensity or the severity of an earthquake (St. Amand, 1961; Howell, 1970; Blume, 1970, 1971; Trifunac and Brady, 1975). The intensity is usually expressed in terms of

- the amount of energy released at the hypocenter
- the effect on structures
- some design related parameters.

The widely used Richter magnitude (M) scale for rating the magnitude of an earthquake was proposed by Richter (1935). This development initiated the rating of earthquakes according to their magnitude at the source independent of the ground shaking at other locations.

More recently the moment of an earthquake was proposed as a measure of the energy released (Aki, 1966). The modified Mercalli and Rossi-Forel intensity scales are based on the effect on structures of local ground shaking. Structural engineers have tried to quantify earthquakes in terms of parameters that are more closely related to the loads induced by a structural behavior such as peak value of the measured record of the earthquake (acceleration, velocity or displacement), frequency content and duration of the ground motion.

For several years hazard maps have been a common way to present expected seismic severity for a region. The level of sophistication of these maps varies greatly as illustrated by the following examples.

Consider the uniform building code (UBC) seismic zone map (1970), as shown in Fig. 1.1. Note that the title of the map indicates "Seismic Risk Map of the United States". This is a misnomer. This map indicates, to some scale, the future seismic hazard in different parts of the country. This map cannot and does not take into account any consequences due to a future seismic event. Also, as pointed out in the note appearing on the map, this map does not take into account the frequency of occurrence of earthquakes. Thus, for example, if one region had an intensity VIII event only once during the last 400 years, it is placed in Zone 3. At the same time, a region with ten intensity VIII or greater events in the last 100 years is also placed in Zone 3. In spite of such shortcomings, engineers use and depend on such maps in the designs. There is no attention paid to the amount of uncertainty in the loading and response of systems they design. It is very difficult to visualize perfect information with the use of such maps.

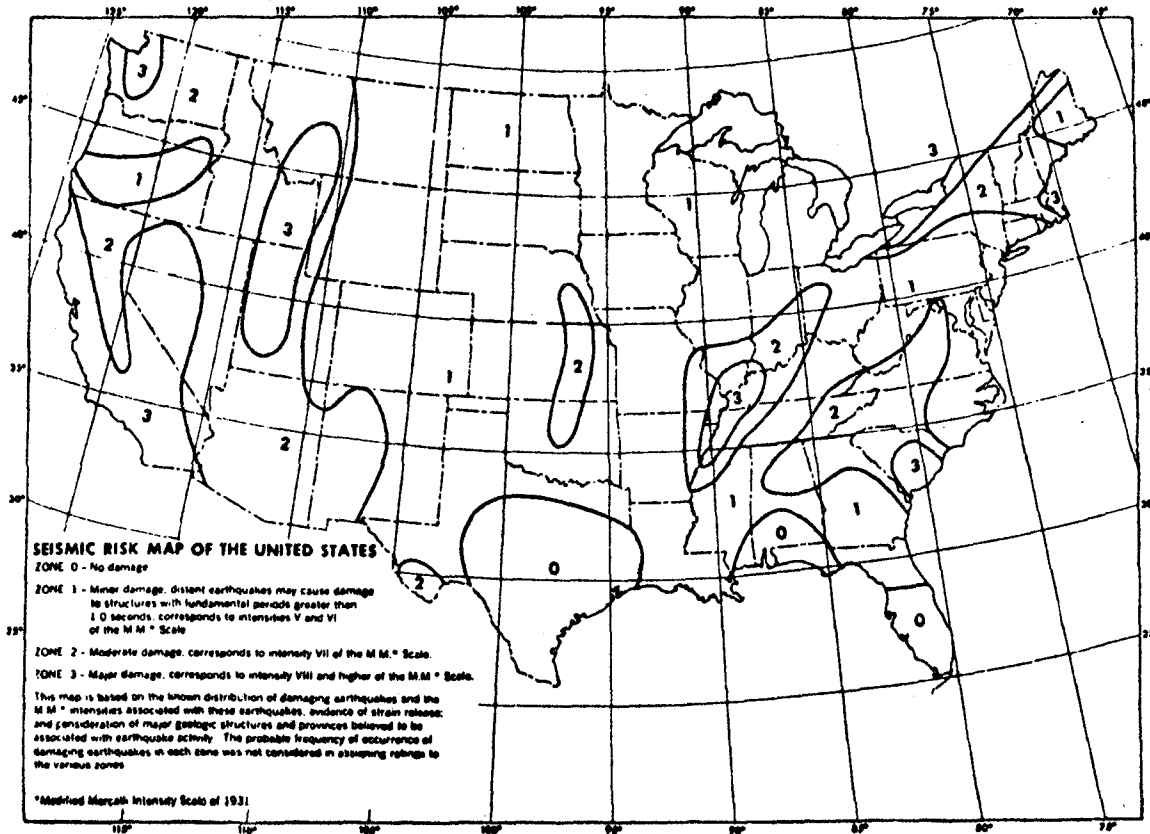


FIGURE 1.1. (UBC, 1970)

The evaluation of risk, using such maps, is not possible. Consider, for example, two locations in California, one near Sacramento and the other near Los Angeles. Also, consider that construction of similar facilities is planned at these two locations. Since the use and consequences of failure of these two facilities or structures located in two separate seismic regions is the same, the risk associated with the design should be similar. However, in using UBC zone maps, one would design both the facilities based on zone 3 factors. This means that either the facility near Los Angeles is underdesigned or the facility near Sacramento is overdesigned or both. For consistent seismic risk level for these two facilities, the site near Sacramento should have lower seismic design level compared to the site near Los Angeles. This brief argument points out the shortcomings of the current "risk maps".

Another recent development -- in the right direction -- is the work of Roger Greensfelder (1974) of the California Division of Mines and Geology. The revised August 1974 map of "Maximum Credible Rock Acceleration from Earthquakes in California" shows the peak rock acceleration levels in different parts of the state in the form of bedrock acceleration contours (Fig. 1.2). This is a hazard map. However, it has several major shortcomings. What is the return period for maximum credible rock acceleration? Is there a consistency in time? Are the frequencies the same for all levels for all regions? This information is very important for engineering design. Also, shall a designer use this map for warehouses, schools, hospitals? The map does not contain the information necessary for a consistent risk level based on use and consequences of failure of the structure. Thus, for a rational seismic

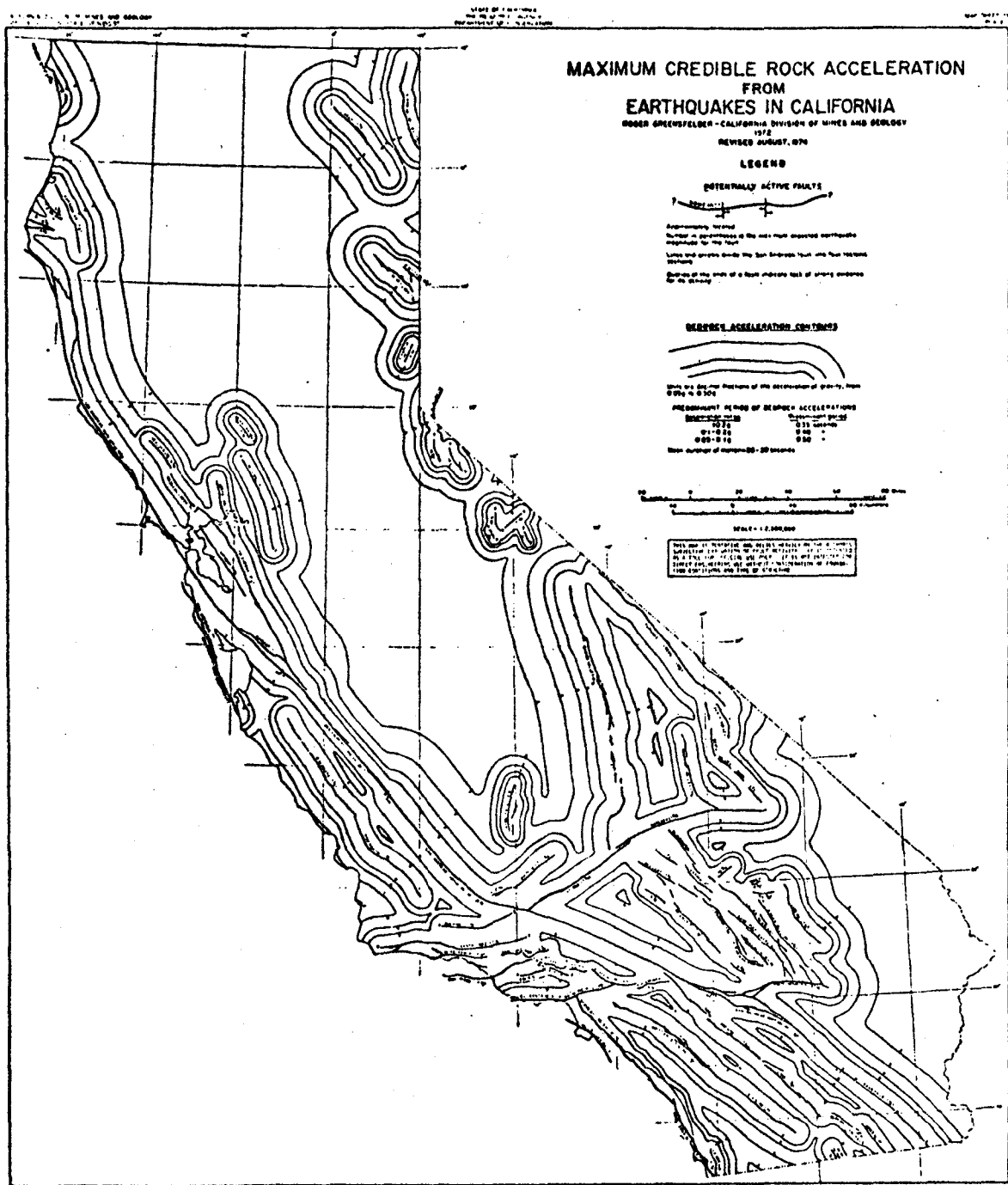


FIGURE 1.2. (Greensfelder, 1974)

risk analysis, a better methodology is needed.

Currently, there is considerable work done in the area of probabilistic estimation of seismic load parameters. In particular probabilistic forecasting in terms of iso-seismal and iso-acceleration maps has been studied. The usefulness of such maps is very much a function of the type and amount of information they contain. In order to best serve the needs of the structural engineering profession, any description of future earthquake ground motion should consist of information about the following:

1. Peak values (or other more descriptive parameters) of acceleration, velocity and displacement.
2. Frequency content.
3. Duration.

The following observations can be made:

- The peak values provide maximum amplitude but do not give any information about the other lower peaks contained in the record, whereas parameters such as mean and root mean square (RMS) values are based on the entire input record and therefore implicitly contain information about all the peaks and their distribution.
- The frequency content is generally represented by a response spectrum which provides the distribution of maximum response amplitudes at various frequencies (or periods).
- Duration is usually considered as the length of time over which "strong motion" is experienced.

None of the currently available procedures hazard mapping provide the complete information listed above. It is therefore the goal of the

first part of this report to increase the amount of information contained in hazard maps.

1.2 Current Procedures

1.2.1 Peak Amplitudes

Current procedures for estimating seismic hazard in terms of peak amplitudes are summarized in Fig. 1.3. In essence, they consist of the following steps:

Step 1. Identification of Seismic Sources.

Based on the geology and historic seismicity of the region, sources are identified as line sources (faults) or area sources. The largest earthquake associated with each source is established from the historic seismicity and geology (in terms of magnitude or intensity). Typical examples of this approach are given in the following references: Cornell and Vanmarcke (1969), Algermissen (1969, 1975), Shah, et al., (1975), Wiggins (1975), Der Kiureghian and Ang (1975), Liu and Fagel (1975), Kiremidjian (1975).

Step 2. Recurrence of Earthquakes

The recurrence of earthquakes of various magnitudes is based primarily on the historic seismicity. A straight line or a set of straight lines is fitted on the data using regression analysis. This method usually results in prediction uncertainties of large magnitudes where the data is scarce. Some variations have been proposed to allow for the lack of data; Esteva (1969); Wiggins (1975) who uses a Bayesian procedure at the level of the results once the analysis is complete. Vagliente (1973) developed a seismic Markov model.

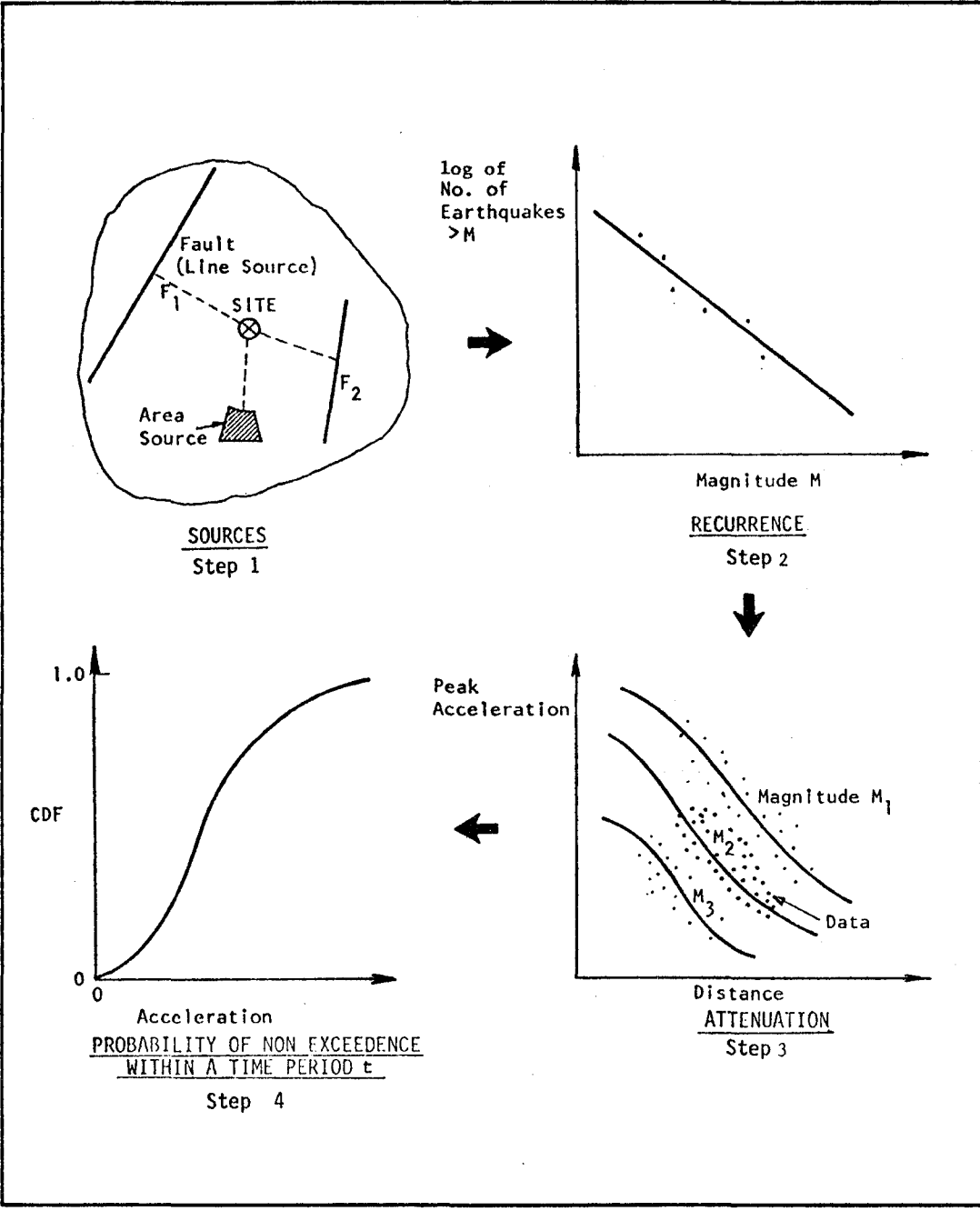


FIGURE 1.3. Current Approach to Hazard Mapping for Peak Values

Step 3. Selection of Attenuation Relationship

Using one of the numerous empirical attenuation relationships, the peak accelerations at a given site due to earthquakes of various sizes occurring at different source locations are estimated. The attenuation relationship is based on data of non-uniform quality since differences in recording techniques, local conditions, equipment reliability and human error are usually not taken into account. Most procedures utilize only the mean curves as determined from a regression analysis.

Step 4. Results

Utilizing the computations in Steps (1), (2) and (3) the probability that a certain acceleration will not be exceeded within a given time period t is determined. The results of the evaluation are presented in terms of iso-acceleration curves for selected levels of probabilities and time periods.

Fig. 1.4 shows a typical iso-acceleration map developed using the above procedure (Shah, et al., 1975). This map was developed for Nicaragua. The acceleration values of the iso-lines have a probability of 0.10 of being exceeded at least one time in fifty years. This is not a zoning map and values between the lines is obtained by interpolation. It should be pointed out that these maps by themselves do not help engineers in deciding the risk level they are taking or, for that matter, which map to use out of the many available for different time periods and different probabilities.

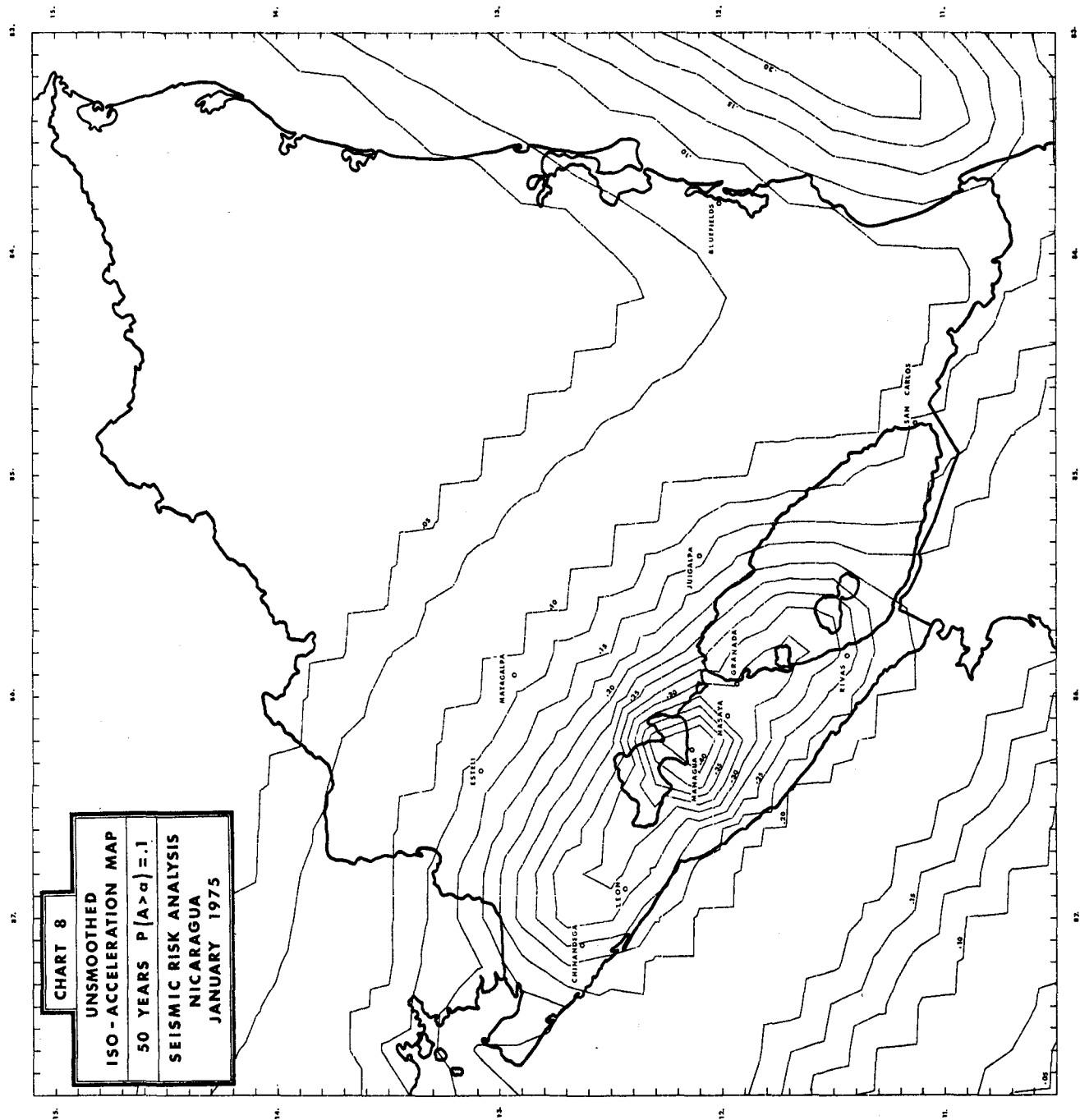


FIGURE 1.4. (Shah, et al. 1975)

1.2.2 Frequency Content

The two current approaches to mapping of frequency content are illustrated in Fig. 1.5 a and b respectively.

In the first approach, normalized response spectra are obtained for available accelerograms and a statistical analysis is made by either considering all spectra together (Clough, 1962; Blume, 1973) or by separating them by site conditions (Seed, et al., 1974). From the results (mean and coefficient of variation) of the statistical analysis response spectra for various probabilities of exceedence are established.

The second approach consists of the development of an attenuation (with distance) relationship for the peak response values for different periods or period bands. These attenuation relations are obtained by regression analysis on a number of response spectra. When combined with the recurrence relation at the sources, they can provide the probability distribution of peak responses for different periods or period bands at a given site (McGuire, 1975).

1.2.3 Duration

It is felt that this parameter is a most important measure of the damage producing capability of an earthquake and therefore it is to be incorporated in the maps developed herein. Some study of duration have been done (Trifunac and Brady, 1975; Dobry, et al., 1977) and some empirical relationships between magnitude, duration and distance have been obtained (Bolt, 1973). However, duration has never been included as a parameter in hazard map development.

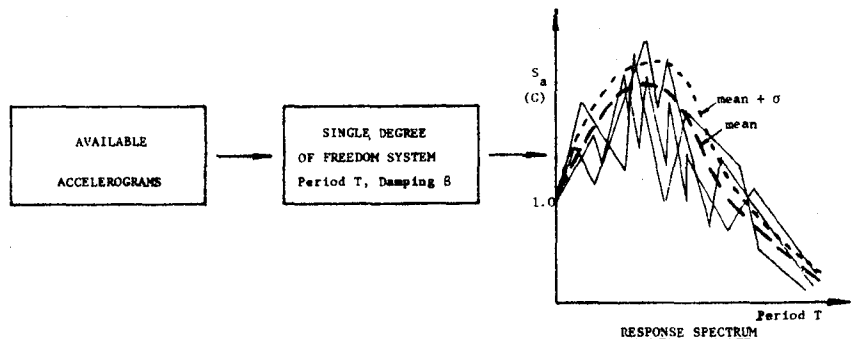


FIGURE 1.5(a). Statistics on Normalized Spectra

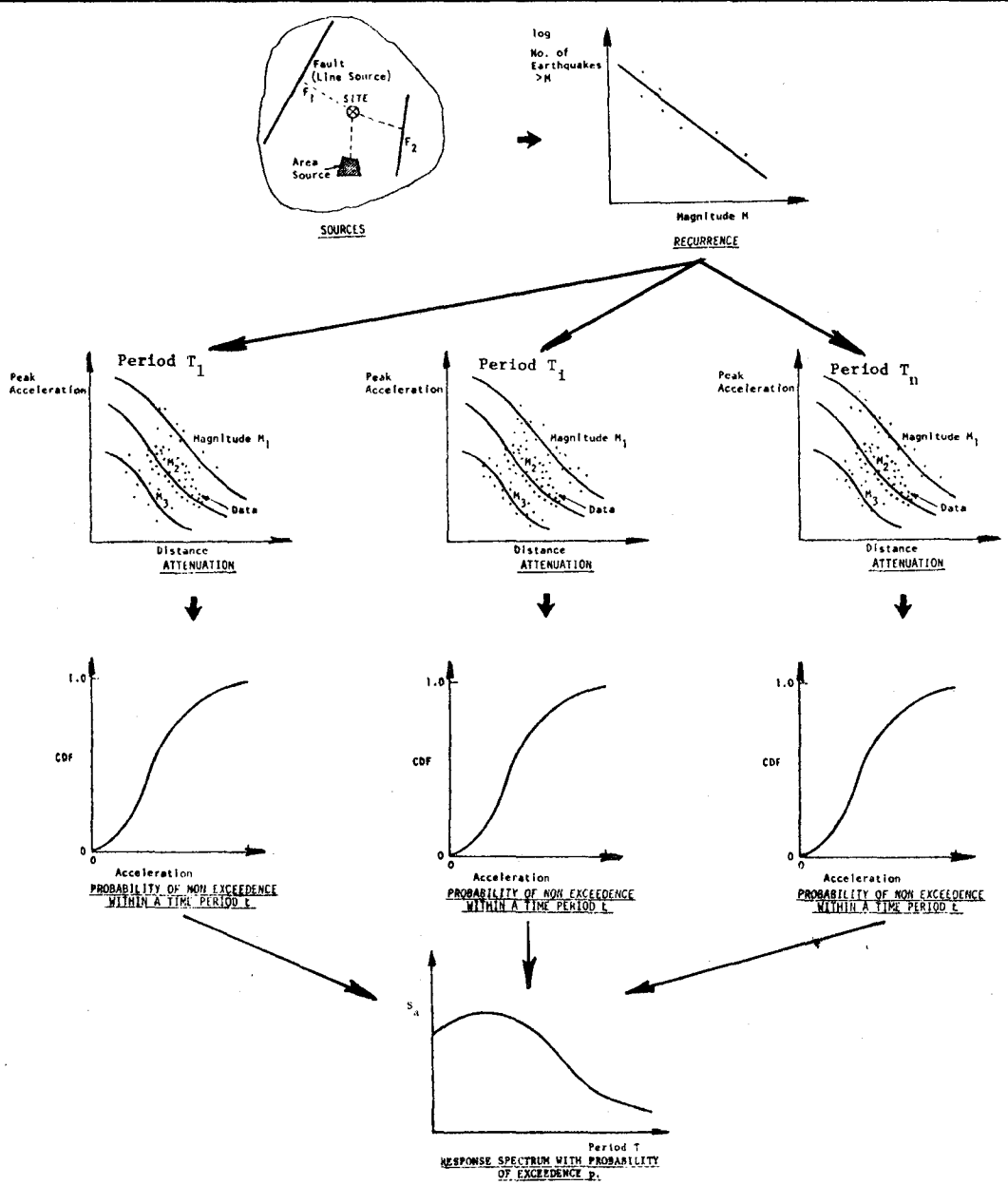


FIGURE 1.5(b). Attenuation of Spectral Values

FIGURE 1.5. Frequency Content Current Procedures

1.3 Limitations

Limitations exist in the descriptive parameters and in the procedures presently used. These limitations are discussed as follows. Several possible improvements are presented in the next section.

1.3.1 Peak Amplitudes

While it is well known that peak amplitudes are useful parameters in the description of seismic loading, they are only partial indicators of earthquake motion. They are often the chance result of a random transient phenomenon and therefore do not show good correlation with the remaining general behavior of the motion. Moreover, they show considerable scatter even for earthquake events considered as similar (same distance from causative fault, same magnitude event and same soil site conditions).

Better and more stable parameters based on sufficient (mean and RMS) statistics rather than extreme values are needed to define the amplitude content of the ground motion. Such parameters or sets of parameters can be obtained by considering earthquake records in their entirety.

1.3.2 Recurrence Relationships

Recurrence equations based on regional historic seismicity data may be inaccurate for individual faults and specific areas. It has been observed that the short historical data base cannot represent the true state of nature. Figure 1.6 shows a typical case where the lack of data poses a critical problem for the analyst. Reliance on historical frequency data alone could result in erroneous conclusions. The simple extrapolation of the fitted curve beyond the range of the data, as it is

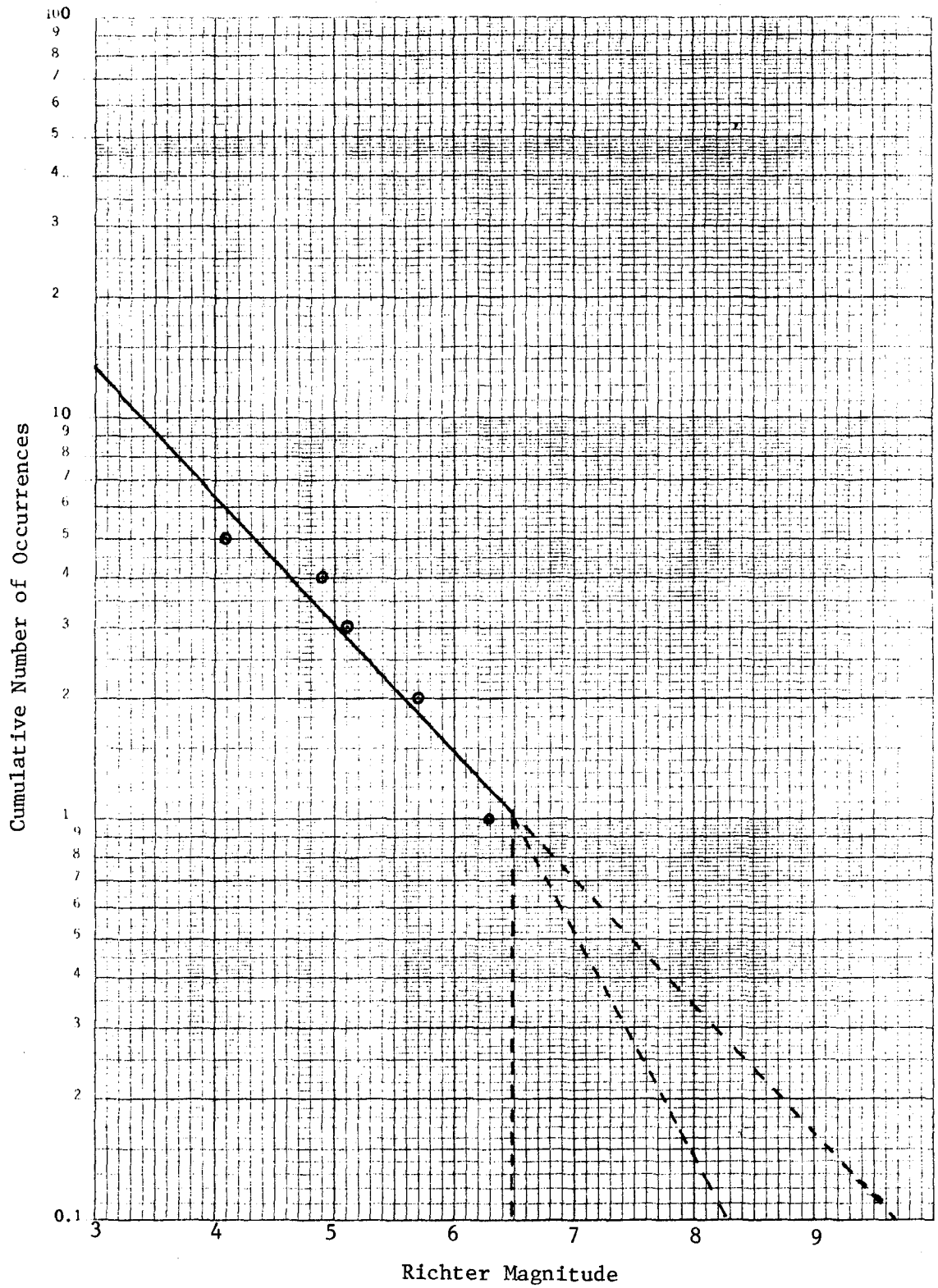


FIGURE 1.6. Typical Regression Graph

usually done in practice can be very dangerous since it is in this area that the magnitudes are governing any design decision.

The treatment of the data as discrete variables will help to solve this problem as presented in Chapter II. Inclusion of subjective geological and seismological opinion could increase the reliability and predictability of the source seismicity considerably. Bayesian statistics allowing combination of objective (or historical) data with subjective input in a rational and consistent way can greatly improve the accuracy of the predicted recurrence of a given magnitude.

1.3.3 Attenuation Relationships

Attenuation relationships for the variation of peak values with distance are significantly affected by the scatter of the peak values. Also a considerable amount of variation is due to non-uniform quality and amount of the basic data concerning past earthquakes. These conditions introduce large uncertainties in the prediction capabilities of the derived relation. While the probabilistic description of seismicity is well accepted and is best represented by hazard maps, it is often mentioned that the weakest link of the present description of seismicity resides in the uncertainties associated with the attenuation relationships. However, strangely enough, only the deterministic or mean value part of these relations is used and the possible random variation in the attenuation prediction is not considered at all. That is, a wide range of prediction uncertainty is completely ignored. This removes random uncertainty information where it is most needed and creates a break in the methodology. In order to be consistent with the probabilistic

approach, a probabilistic treatment of attenuation is needed and can be achieved by considering the probability distribution of the predicted peaks for a given distance.

1.3.4 Response Spectra

Response spectra derived from statistical averaging of spectral shapes or regression analyses for data over all period ranges do not consider the probability of occurrence of each sample of data. They do not recognize each earthquake type (fault type, distance from site, etc.) with their corresponding probability of occurrence. Therefore some bias is likely to occur (one earthquake type versus another) and the spectral statistics developed will provide an accurate estimate of probability of exceedence only for specific period ranges and may be significantly in error for other periods.

1.3.5 Duration

Duration is possibly the single most important factor in producing excessive damage. (H. M. Engle in Richter, 1958). However, it is very seldom mentioned explicitly as a ground motion parameter. The reason being that it is not a design parameter per se and consequently no direct use of this information has been made as yet. Treatment and use of this parameter would add valuable information to hazard mapping.

1.4 Scope of Current Work

The present work focuses on the elimination of some of the limitations presented above. Special attention is to be given to:

- The Bayesian modeling of the seismic sources.

- The probabilistic treatment of attenuation relationship both for acceleration and duration.
- The replacement of peak values by more stable and representative design parameters.

1.4.1 Seismic Hazard Mapping

The steps of the general model used for seismic mapping in this study are presented below.

Source Location

The location of the sources is determined by using recorded hypocentral position of past earthquakes for the period over which historical records are available. Geological and seismological information is introduced to supplement the data and to present a coherent picture of the seismicity of the region. The spatial distribution of hypocenters is then divided in or assigned to different sources so as to model the earthquake generating process.

Seismic Model

In the data presently available, the most commonly used measure of earthquake magnitude is the Richter magnitude (M). In the current procedures, the seismicity of a source is described by the mean rate of occurrence of events greater than a given magnitude (recurrence relationship). In the present model, this seismicity is obtained in two steps. First the occurrences are considered independently of the magnitudes, then the probability distribution of the magnitudes is introduced as explained below:

- Occurrences: With the assumption that earthquake occurrences form a Poisson process with mean rate of occurrence independent of magnitude, a probability distribution is obtained on the number of occurrences for a given period of time for a given source. The assumption of spatial and temporal independence is fairly well verified by data and is a common accepted practice in seismology. Moreover the amount of dependence due to the dual mechanism of stress accumulation and release has not been determined as yet.
- Magnitudes: Given that an event has occurred, a probability distribution of the Richter magnitudes (M) is determined from past data. The M are discretized at every $1/4$ unit as is commonly done in data recording. This representation has the advantage of getting away from the curve of line fitting. This method is specially valuable for regions where very little data is available. The probability corresponding to each magnitude can be used in a Bernoulli trial where one outcome will be an event of the magnitude considered (success) and the other an event of any other magnitude (failure). The following question can then be answered: "Given that n earthquakes will occur in future time t , what is the probability that there will be $0, 1, 2 \dots n$ events of any given magnitudes?" Combining these binomial conditional distributions with the Poisson distribution of occurrences, the distribution of the number of occurrences of each magnitude can be obtained.

Bayesian Statistics

Bayesian statistics are applied to the Poisson and binomial laws to supplement some of the incompleteness of the data. For example, with consideration of the fault length and the type of fault, geologists can determine the maximum magnitude earthquake that the source can generate. This information has to be taken into account even if no such earthquake had been recorded in the data. This can be done by assuming the mean rate of occurrences of the Poisson law to be a gamma probability distribution and the probability of success of the binomial law to be a beta probability distribution. This method has the advantage of including personal experience together with the data as well as updating the distribution as new data are made available.

Mapping Parameters

Two mapping parameters are used, namely the peak ground acceleration (PGA) and the duration of the ground motion. PGA is used since no other attenuation relationship is readily available in the literature. The methodology allows for the use of a more stable parameter such as RMS which would certainly improve the reliability of the model. More will be said about RMS and stable parameters in later sections.

Attenuation Relationships

The Esteva (1974) relationship is used to relate the PGA to M and hypocentral distance. Attenuation of duration is obtained using the relation developed by Bolt (1973).

Both relationships are treated probabilistically to take into consideration uncertainties in the attenuation decay. A first order probabilistic representation in the form of mean and coefficient of variation is introduced in this work.

Probabilistic Estimation of PGA and Duration

Combining the distributions that describe the seismicity at the sources with the two attenuation functions (for acceleration and duration), the probability of exceedence of any PGA or duration is to be obtained at the site.

Case Study

A case study of seismic mapping for the country of Nicaragua is presented.

1.4.2 Study of Stable Design Parameters

In order to develop a proper understanding of recorded time histories a detailed statistical analysis of relevant parameters associated with these time histories is necessary. Such an in-depth study is presented with the following goals in mind:

- To develop a stable parameter (or parameters) which can describe the behavior of the seismic recorded input or of response output of linear single degree of freedom systems.
- To relate the developed stable parameters to "design" parameters currently used by engineers; e. g., peak ground acceleration and response spectra values.
- To incorporate seismic hazard map development with the utilization of stable parameters and currently used design procedures.

With these goals in mind, a set of 97 earthquake time histories are studied. It should be pointed out that in order to study an earthquake time history in its entirety, one should investigate the statistical properties in the time domain and the frequency domain. The time domain analysis involves the study of the recorded motion in terms of the duration and other amplitude related statistical parameters. The most widely used frequency analysis involves the study of response time histories of linear one degree of freedom system (for a given damping and period) to the given earthquake record.

The current study investigates the following parameters of each of the 97 records considered.

(A) Input parameters obtained from recorded accelerograms

- Duration
- Mean acceleration
- Root mean square (RMS) of acceleration
- Probability distribution of peaks
- Parameter a_p defined as the acceleration which has probability p of being exceeded for a given record.

- The ratio $K_1 = a_p / \text{RMS}$

(B) Response parameters obtained from acceleration response history of a linear single degree of freedom system with damping ratios 5%, 10% and 20% and natural period range of .08 sec. to 5.0 sec.

- Duration
- Mean acceleration
- Peak response (corresponds to spectral value)

- RMS of acceleration
- Probability distribution of response peaks
- Parameter a_p defined as the response acceleration which has probability p of being exceeded for the given response.
- The ratio $K_1 = a_p / \text{RMS}$
- Cumulative potential energy per unit mass (ENGY) in response time history. (This will be further described in Chapter IV.)
- The ratio $K_2 = \text{RMS}^2 \cdot T^2 / (\text{ENGY} / \text{NBPK})$ in the above ratio, T is the period of the structure and NBPK is the number of peaks of the response time history.

In Chapter V and VI the stable behavior of the ratios K_1 and K_2 and their relevancy are described and discussed.

Development of an entirely new and fresh approach in earthquake engineering has its drawbacks. First, the practical implication of a new method of seismic hazard mapping for peak acceleration and duration cannot be assessed at this time. Second, the use and implication of new stable parameters describing the ground motion and the response together with the currently used parameters can only be achieved with time. At times, the reader of this work will feel that a "total story" involving the hazard mapping, the use of stable parameters and design procedure is not presented. It is not the intent of this report to present such a "story". It is hoped that the ideas developed in this work will be further studied and incorporated into a design methodology by future researchers and practicing engineers.

CHAPTER II

METHODOLOGY FOR SEISMIC HAZARD MAPPING

2.1 Introduction

This chapter concentrates on seismic hazard mapping which can be defined as the exposure to seismic loading at a given location. This exposure is expressed in terms of the severity of an effect and the probability of its occurrence.

The seismic hazard mapping model is divided in the following steps:

- Identification of Earthquake Sources

The locations of the sources are determined by using recorded hypocentral position of past earthquakes together with geological and seismological information.

- Development of an Earthquake Recurrence Model

The seismicity of a source is obtained in two steps. First the occurrence of events is considered independently of the magnitudes, then the probability distribution of the magnitudes is introduced.

- Choice of Seismic Ground Motion Parameters

Different mapping parameters are presented and discussed. Peak ground acceleration (PGA) and duration are chosen for this work together with corresponding attenuation relationships.

- Probabilistic Estimation of PGA and Duration

The probability of exceedence of any PGA and duration is obtained at the site considered by combining the probability distributions that describe the seismicity of the sources with the two attenuation functions.

Because of the limitations inherent in the current procedures of hazard mapping, special attention is given to:

- The Bayesian modeling of the earthquake recurrence phenomenon
- The use of two mapping parameters: peak acceleration and duration
- The probabilistic treatment of attenuation relationships

2.2 Identification of Seismic Sources

The locations of the sources are identified using the recorded hypocentral position of past earthquakes together with geological and seismological information. The spatial distribution of hypocenters is then divided in different sources as a function of their shape and seismicity. Two types of sources can be used to represent the seismicity of any region. They are the line and area sources.

Most of the earthquake epicenters around the world are located around major fault systems. Thus, the usual case of epicenters falling along a line gives rise to the so-called "line source". A number of straight lines may be necessary to model a fault to satisfy geometric considerations (curve shaped faults) or seismic considerations (variation in seismicity along a fault). The seismicity is assumed to be homogeneous over a segment.

In many parts of the world there are regions where the epicenters are not located along a line but are scattered all over a region (Gutenberg and Richter, 1954). This may be due to the existence of numerous faults criss-crossing the region or due to errors in estimation of epicentral locations. Thus, there are places where line sources may not fit the scatter of epicentral locations. For such cases, area

sources are used to determine probabilistic loading at the site. In the present model the area can either be a circle or a rectangle. The seismicity is assumed to be homogeneous over the area source.

2.3 Development of an Earthquake Recurrence Model (A Bayesian Approach)

2.3.1 Introduction

It has been observed that short or partial historical data bases cannot represent the true state of nature. Reliance on frequency data alone can result in erroneous conclusions. Inclusion of geological and seismological opinions can increase considerably the reliability and predictability of the seismicity. For this reason a Bayesian approach is adopted and used in this work.

In the data presently available, the most commonly used numerical measure of earthquake severity is the Richter magnitude. In this model, the magnitudes are discretized to every $1/4$ unit of Richter magnitude (M_i) as it is commonly done in data recording. The seismicity of the sources is described by the distribution of the number of occurrences of each magnitude. This representation permits the use of discrete models and has the advantage of getting away from data fitting which usually results in unacceptable uncertainties for large magnitudes where the data is scarce. Moreover, it leads to a direct and elegant use of Bayesian statistics.

The development is done in three steps:

- Assuming that earthquake occurrences form a Poisson process with mean rate of occurrence independent of magnitude, a distribution is obtained on the number of occurrences for the time period considered.

- Given that an event has occurred, a distribution on the magnitude of events is determined from past data. The process generating model can be assumed to be Bernoulli. The probability of success p_{M_i} corresponding to each trial is defined as the probability that the event that has occurred is of magnitude M_i . Thus the probability of failure $q_{M_i} = 1 - p_{M_i}$, at each trial is the probability that the event is not of magnitude M_i . The probability of having r events of magnitude M_i given that a total of n events have occurred can therefore be obtained using the binomial distribution.
- The distribution of the number of events of each magnitude independently of the number of trials is obtained by combining step one and two.

2.3.2 Occurrences

2.3.2.1 Poisson Model

Once the seismic sources have been located, it is assumed that earthquake occurrences on each source form a Poisson process with mean rate of occurrence independent of magnitudes. For earthquake events to follow the Poisson model, the following assumptions must be valid:

- (1) Earthquakes are spatially independent
- (2) Earthquakes are temporally independent
- (3) Probability that two seismic events will take place at the same time and at the same location approaches zero.

The first assumption implies that occurrence or non-occurrence of a seismic event at one site does not affect the occurrence or non-occurrence of another seismic event at some other site. The second

assumption implies that the seismic events do not have memory in time. The assumptions of spatial and temporal independence have been fairly well verified by data and are commonly accepted practices. The degree of dependence due to the dual mechanism of stress accumulation and release has not been determined with any amount of precision as yet and seems to fade away quite rapidly with time (Gardner and Knopolf, 1974). The third assumption implies that for a small time interval, Δt , more than one seismic event cannot occur on one source. This is a very realistic and good assumption which fits the physical phenomenon.

Hence, considering all the events of magnitude greater than an arbitrary lower bound, a distribution is obtained on the number of occurrences for a given period of time t . The lower bound is chosen such that earthquakes of magnitude smaller than the one specified have a negligible damage potential and can thus be disregarded. This is done for each seismic source.

In its most general form, the conditional Poisson law can be written as

$$p_N(n/\lambda) = \frac{e^{-\lambda t} (\lambda t)^n}{n!}, \quad t > 0; \quad n \text{ integer } \geq 0 \quad (2.1)$$

where $p_N(n/\lambda)$ = Probability of having n events in time period t , given λ .

n = Number of events.

λ = Mean rate of occurrence per unit of time.

$$\mu_N = \lambda$$

$$\sigma_N^2 = \lambda$$

2.3.2.2 Bayesian Probability

A Bayesian approach is used so that historical data and subjective information can be effectively combined and used in the analysis. If one assumes that the number of seismic events for a future time t follows a Poisson probability law, there is still uncertainty about the parameter λ , the mean rate of occurrence (Eq. 2.1). Therefore, λ is treated as a random variable. The probabilistic information on λ can be obtained through historical data or from the subjective knowledge of the analyst. The subjective probability distribution on λ is called the prior distribution. Using the historical information, one can obtain the sample likelihood function on λ . Combining the prior distribution and the sample likelihood function by means of Bayes' theorem, the posterior distribution on λ can be obtained (Benjamin and Cornell, 1970).

Let $f'_{\Lambda}(\lambda)$ be the prior probability distribution function on λ .

Let $L(\lambda)$ be the sample likelihood function on λ .

Then using Bayes' theorem the posterior distribution $f''_{\Lambda}(\lambda)$ is obtained as:

$$f''_{\Lambda}(\lambda) = N_1 L(\lambda) f'_{\Lambda}(\lambda) \quad (2.2)$$

where N_1 is a normalizing constant.

In the following sections, the concept of conjugate prior is used for analytical simplicity.

Prior Distribution on λ

For convenience the prior distribution for the random variable λ is chosen as the gamma distribution with parameters λ' and ν' . Since

the gamma distribution can fit a large variety of shapes this choice does not introduce any limitations in the model (Raiffa and Schlaifer, 1961). The parameters λ' and ν' are obtained directly from the subjective input and the prior can thus be written:

$$f'_{\Lambda}(\lambda) = \frac{\lambda'(\lambda')^{\nu'-1} e^{-\lambda'\lambda}}{\Gamma(\nu')} \quad \lambda \geq 0; \lambda' > 0; \nu' > 0 \quad (2.3)$$

where $\Gamma(\nu') = \int_0^{\infty} e^{-u} u^{\nu'-1} du$

$$\mu_{\Lambda} = \frac{\nu'}{\lambda'}$$

$$\sigma_{\Lambda}^2 = \frac{\nu'}{\lambda'^2}$$

Sample Likelihood Function

For any given source, the available data indicates that in the past T years, N earthquakes of Richter magnitude greater than the lower bound have occurred. This information is used in the construction of the sample likelihood function. Since the event-generating process is assumed to be a Poisson process, the sample likelihood function on λ is given by

$$L(\lambda/N, T) = \frac{e^{-\lambda T} (\lambda T)^N}{N!} \quad T > 0; N \geq 0 \quad (2.4)$$

where T = Time of the data base

N = Number of events greater than a fixed lower bound M observed during time period T.

Posterior Distribution on λ

Using equations 2.2 and 2.3 the posterior distribution on λ can be written as

$$f_{\Lambda}''(\lambda) = N_1 \cdot \frac{e^{-\lambda T} (\lambda T)^N}{N!} \cdot \frac{\lambda' (\lambda' \lambda)^{\nu' - 1} e^{-\lambda' \lambda}}{\Gamma(\nu')} \quad (2.5)$$

The value of the normalizing constant N_1 can be evaluated by noting that $f_{\Lambda}''(\lambda)$ is a probability distribution function. Thus,

$$\int_0^{\infty} f_{\Lambda}''(\lambda) d\lambda = 1.0 \quad (2.6)$$

Rearranging and substituting the values of N_1 in equation 2.5

$$f_{\Lambda}''(\lambda) = \frac{\lambda'' (\lambda'' \lambda)^{\nu'' - 1} e^{-\lambda'' \lambda}}{\Gamma(\nu'')} , \quad \lambda \geq 0; \lambda'' > 0; \nu'' > 0 \quad (2.7)$$

where $\lambda'' = \lambda' + T$

$$\nu'' = \nu' + N$$

$$\mu_{\Lambda} = \frac{\nu''}{\lambda''}$$

$$\sigma_{\Lambda}^2 = \frac{\nu''}{\lambda''^2}$$

Note that the posterior distribution of λ is also gamma type.

In equation 2.1, the conditional probability on the number of events n is based on λ . The unconditional or the marginal distribution on n can be obtained by using equation 2.7 together with equation 2.1 and integrating over all λ 's. Thus

$$\begin{aligned} p_N(n) &= \int_0^{\infty} p_N(n, \lambda) d\lambda \\ &= \int_0^{\infty} p_N(n/\lambda) f_{\Lambda}''(\lambda) d\lambda \\ &= \int_0^{\infty} \frac{e^{-\lambda t} (\lambda t)^n}{n!} \frac{\lambda'' (\lambda'' \lambda)^{\nu'' - 1} e^{-\lambda'' \lambda}}{\Gamma(\nu'')} d\lambda \end{aligned} \quad (2.8)$$

-continued-

$$= \int_0^{\infty} \frac{e^{-\lambda(t + \lambda'')} t^n \lambda''^{\nu''} \lambda^{n+\nu''-1}}{n! \Gamma(\nu'')} d\lambda \quad (2.8)$$

-continued-

which leads to

$$p_N(n) = \frac{\Gamma(n + \nu'')}{n! \Gamma(\nu'')} \cdot \frac{t^n \lambda''^{\nu''}}{(t + \lambda'')^{n+\nu''}} \quad (2.9)$$

for n integer ≥ 0

$$\nu'' > 0$$

$$\lambda'' > 0$$

$$t > 0$$

When ν'' has integer values, equation 2.9 can be rewritten

$$p_N(n) = \frac{(n + \nu'' - 1)!}{n! (\nu'' - 1)!} \frac{t^n \lambda''^{\nu''}}{(t + \lambda'')^{n+\nu''}} \quad (2.10)$$

for n integer > 0

$$\nu'' \text{ integer } > 0$$

$$\lambda'' > 0$$

$$t > 0$$

Equation 2.9 and 2.10 are called the marginal Bayesian distributions of n . These distributions, after taking into consideration the uncertainties on the mean rate of occurrence, give the probability of the number of events above a predetermined lower bound M , in time period t .

2.3.2.3 Weight of the Prior Distribution

The subjective information of the expert is expressed in the form of a prior distribution on λ . It is assumed that a gamma distribution

will fit the prior and the parameters λ' and ν' are obtained directly from the shape of this distribution. The physical meaning of λ' and ν' can be explained as follows: λ' can be interpreted as the equivalent time period over which the analyst bases his subjective input whereas ν' can be interpreted as the equivalent number of occurrences during this time period. Hence if $\lambda' = T$, both the prior and the data will have the same weight. If $\lambda' \gg T$, the weight of the prior overshadows the importance of the data. It implies that the expert has more confidence in his personal feeling than in the data. A complete analysis of the weight of the prior is given in Raiffa and Schlaifer (1961).

2.3.3 Richter Magnitude

2.3.3.1 Introduction

Up to this point, the seismicity of a source has been defined only by the distribution on the number of events that this source may generate in a given time period t . The next step consists in the representation of information on the severity of these events. Several parameters are suitable for this purpose such as energy release, intensity and Richter magnitude. The Richter magnitude is the parameter chosen herein since most of the available data is recorded in terms of this quantity. The Richter magnitudes are discretized every $1/4$ Richter unit (M_1) as commonly done in seismology. The number of different M_1 's expected to occur is thus finite. This allows for the use of a discrete model.

2.3.3.2 Bernoulli Model

A Bernoulli trial is used to model information on magnitudes. Given that an event has occurred, the probability that it is of any given Richter magnitude can be represented in terms of a Bernoulli trial. If the seismic event that has occurred is of the M under consideration, then the outcome of the Bernoulli trial is a success. Conversely, failure at each trial implies that the seismic event that has occurred is of M other than the one under consideration.

If p_{M_i} = probability of success at each trial corresponding to M_i .
and

$$q_{M_i} = 1 - p_{M_i}$$

= probability of failure at each trial,

then using the binomial law,

$$p_R(r_{M_i}/n, p_{M_i}) = C_n^{r_{M_i}} p_{M_i}^{r_{M_i}} (1 - p_{M_i})^{n-r_{M_i}} \quad (2.11)$$

for n integer > 0

$$r_{M_i} \text{ integer; } 0 \leq r_{M_i} \leq n$$

$$0 \leq p_{M_i} \leq 1$$

where $p_R(r_{M_i}/n, p_{M_i})$ is read as the probability that r_{M_i} events M_i will occur out of a total of n events given that the probability of occurrence of M_i is p_{M_i} at each trial,

$$\text{and } C_n^{r_{M_i}} = \frac{n!}{r_{M_i}!(n-r_{M_i})!}$$

A different probability p_{M_i} is obtained for each M_i considered in the model. A relation similar to equation 2.11 is thus obtained for each

of the Richter magnitudes. It is important to recognize that the probabilities p_{M_i} are mutually exclusive for the different magnitudes hence

$$\sum_{\text{all } M_i} p_{M_i} = 1.0 \quad (2.12)$$

2.3.3.3 Bayesian Probability.

Equation 2.11 represents the generating process for the number of events M_i . However, this information is conditional on the knowledge about p_{M_i} , the probability of success corresponding to M_i . To rationally incorporate the historical as well as subjective information on p_{M_i} , this parameter is treated as a random variable and a Bayesian formulation is used.

Prior Distribution on p_{M_i}

The conjugate prior distribution on p_{M_i} is assumed to be beta type. Since the normalized beta distribution is bounded between 0 and 1 and fits a large variety of shapes, this choice does not introduce any limitations in the model (Raiffa and Schlaifer, 1961). The beta form of this prior is thus given by:

$$f'_P(p_{M_i}) = \frac{1}{B'_{M_i}} p_{M_i}^{\xi'_{M_i} - 1} (1 - p_{M_i})^{\eta'_{M_i} - \xi'_{M_i} - 1} \quad (2.13)$$

for $0 \leq p_{M_i} \leq 1$

$$\xi'_{M_i} > 0$$

$$\eta'_{M_i} - \xi'_{M_i} > 0$$

where

$$B'_{M_1} = \frac{\Gamma(\xi'_{M_1})\Gamma(\eta'_{M_1} - \xi'_{M_1})}{\Gamma(\eta'_{M_1})}$$

$$\mu_{P_{M_1}} = \frac{\xi'_{M_1}}{\eta'_{M_1}}$$

$$\sigma_{P_{M_1}}^2 = \frac{(\xi'_{M_1})(\eta'_{M_1} - \xi'_{M_1})}{\eta'_{M_1}(\eta'_{M_1} + 1)}$$

where the parameters η'_{M_1} and ξ'_{M_1} are obtained from the subjective information.

A prior distribution of a similar form has to be assumed for each of the magnitudes considered. Here again, the prior distributions are not independent and the following condition must be satisfied:

$$\sum_{\text{all } M_1} \frac{\xi'_{M_1}}{\eta'_{M_1}} = 1.0 \quad (2.14)$$

Sample Likelihood

The usual form of the available data indicates that among the N earthquakes observed on a given source, R_{M_1} were M_1 . This information is used in the construction of the sample likelihood function. Noting that the generating process (eq. 2.11) is a binomial process, the sample likelihood function on p_{M_1} is given by

$$L(p_{M_1}/N, R_{M_1}) = p_{M_1}^{R_{M_1}} (1 - p_{M_1})^{N - R_{M_1}} \quad (2.15)$$

This operation must be repeated for each of the different magnitudes considered. Each of these sample likelihood functions must be combined with the corresponding prior distributions to obtain the posterior distributions.

Posterior Distribution on p_{M_1}

As described in Section 2.3.2.2, the posterior distribution on p_{M_1} is obtained from equations 2.13 and 2.15 and can be written as

$$f_P''(p_{M_1}) = N_1 \left[\begin{matrix} R_{M_1} & N-R_{M_1} \\ p_{M_1} & (1-p_{M_1}) \end{matrix} \right] \cdot \left[\begin{matrix} \xi_{M_1}'^{-1} & \eta_{M_1}'^{-\xi_{M_1}'^{-1}} \\ \frac{1}{B_{M_1}'} p_{M_1} & (1-p_{M_1}) \end{matrix} \right] \quad (2.16)$$

Applying the condition

$$\int_0^1 f_P''(p_{M_1}) dp_{M_1} = 1.0 \quad (2.17)$$

one obtains

$$f_P''(p_{M_1}) = \frac{1}{B_{M_1}''} \cdot p_{M_1}^{\xi_{M_1}''^{-1}} \cdot (1-p_{M_1})^{\eta_{M_1}''^{-\xi_{M_1}''^{-1}}} \quad (2.18)$$

$$\text{for } 0 \leq p_{M_1} \leq 1$$

$$\xi_{M_1}'' \geq 0$$

$$\eta_{M_1}'' - \xi_{M_1}'' \geq 0$$

$$\text{where } \xi_{M_1}'' = \xi_{M_1}' + R_{M_1}$$

$$\eta_{M_1}'' = \eta_{M_1}' + N$$

$$B_{M_1}'' = \frac{\Gamma(\xi_{M_1}'') \Gamma(\eta_{M_1}'' - \xi_{M_1}'')}{\Gamma(\eta_{M_1}'')}$$

Note that the posterior distribution on p_{M_1} is also beta type.

In equation 2.11, the conditional probability on the number of successes r_{M_1} is based on p_{M_1} and n . The condition on p_{M_1} can be

removed using equation 2.18 and integrating over all the values of

p_{M_i} as follows:

$$\begin{aligned}
 p_R(r_{M_i}/n) &= \int_0^1 p_R(r_{M_i}, p_{M_i}/n) dp_{M_i} \\
 &= \int_0^1 p_R(r_{M_i}/p_{M_i}, n) f_P''(p_{M_i}) dp_{M_i} \\
 &= \int_0^1 C_n^{r_{M_i}} p_{M_i}^{r_{M_i}-1} (1-p_{M_i})^{n-r_{M_i}} \frac{1}{B_{M_i}''} p_{M_i}^{\xi_{M_i}''-1} (1-p_{M_i})^{\eta_{M_i}''-\xi_{M_i}''-1} dp_{M_i} \\
 &= C_n^{r_{M_i}} \frac{1}{B_{M_i}''} \int_0^1 p_{M_i}^{r_{M_i}+\xi_{M_i}''-1} (1-p_{M_i})^{n+\eta_{M_i}''-(r_{M_i}+\xi_{M_i}'')-1} dp_{M_i} \quad (2.19)
 \end{aligned}$$

setting $r_{M_i} + \xi_{M_i}'' = \alpha_{M_i}$

$$n + \eta_{M_i}'' = \beta_{M_i}$$

one can write

$$p_R(r_{M_i}/n) = C_n^{r_{M_i}} \left[\frac{\Gamma(\eta_{M_i}'')}{\Gamma(\xi_{M_i}'') \Gamma(\eta_{M_i}'' - \xi_{M_i}'')} \cdot \frac{\Gamma(\alpha_{M_i}) \Gamma(\beta_{M_i} - \alpha_{M_i})}{\Gamma(\beta_{M_i})} \right] \quad (2.20)$$

for n integer > 0

r_{M_i} integer

$$0 \leq r_{M_i} \leq n.$$

The above expression is the distribution on the number of earthquakes of a fixed M_i given that n earthquakes have occurred. There is a similar distribution for each M_i considered.

2.3.3.4 Weight of the Prior Distribution

The parameters η'_{M_i} and ξ'_{M_i} are obtained directly from the shape of the prior distribution assumed to be of the beta type. The physical meaning of these parameters can be explained as follows: η'_{M_i} can be interpreted as the equivalent number of trials on which the expert bases his subjective input and ξ'_{M_i} as the equivalent number of successes. Hence if $\eta'_{M_i} = N$, both the prior and the data have the same weight. If $\eta'_{M_i} \ll N$, the weight of the data overshadows the importance of the prior. This implies that the expert has more confidence in the data than in his personal feelings.

2.3.4 Marginal Distribution on the Number of Magnitudes

In order to obtain the probability on the number of earthquakes of fixed magnitude regardless of the total number of occurrences, one removes the condition on the number of events n in equation 2.20 by taking the summation over all the events. The distribution on the number of events is given by equation 2.9. Hence

$$\begin{aligned}
 p_R(r_{M_i}) &= \sum_{n=0}^{\infty} p_R(r_{M_i}/n) p_N(n) \\
 &= \sum_{n=0}^{\infty} \left[C_n \frac{\Gamma(\eta''_{M_i})}{\Gamma(\xi''_{M_i}) \Gamma(\eta''_{M_i} - \xi''_{M_i})} \cdot \frac{\Gamma(r_{M_i} + \xi''_{M_i}) \Gamma(n + \eta''_{M_i} - r_{M_i} - \xi''_{M_i})}{\Gamma(n + \eta''_{M_i})} \right. \\
 &\quad \left. \frac{\Gamma(n + \nu'') t^{n \lambda''} \nu''}{n! \Gamma(\nu'') (t + \lambda'')^{n + \nu''}} \right] \quad (2.21)
 \end{aligned}$$

In the particular case where $r_{M_i} = 0$, the above expression becomes:

$$P(r_{M_i} = 0) = \sum_{n=0}^{\infty} \left[\frac{\Gamma(\eta_{M_i}''')}{\Gamma(\xi_{M_i}''')\Gamma(\eta_{M_i}'' - \xi_{M_i}''')} \cdot \frac{\Gamma(\xi_{M_i}''')\Gamma(n+\eta_{M_i}'' - \xi_{M_i}''')}{\Gamma(n + \eta_{M_i}''')} \cdot \frac{\Gamma(n + \nu'')t^{n\lambda''}\nu''}{n!\Gamma(\nu'')(t+\lambda'')^{n+\nu''}} \right] \quad (2.22)$$

In the particular case where ν'' , η_{M_i}'' , ξ_{M_i}'' are integer values, the gamma functions can be expressed as factorials and the two above expressions become

$$P_R(r_{M_i}) = \sum_{n=0}^{\infty} \left[\frac{1}{r_{M_i}! (n-r)!} \frac{(\eta_{M_i}'' - 1)!}{(\xi_{M_i}'' - 1)(\eta_{M_i}'' - \xi_{M_i}'' - 1)!} \frac{(r_{M_i} + \xi_{M_i}'' - 1)!(n + \eta_{M_i}'' - r_{M_i} - \xi_{M_i}'' - 1)!}{(n - \eta_{M_i}'' - 1)!} \cdot \frac{(n - \nu'' - 1)!}{(\nu'' - 1)!} \frac{t^{n\lambda''}\nu''}{(t + \lambda'')^{n+\nu''}} \right] \quad (2.23)$$

$$P(r_{M_i} = 0) = \sum_{n=0}^{\infty} \left[\frac{1}{n! \xi_{M_i}'' - 1)! (\eta_{M_i}'' - \xi_{M_i}'' - 1)!} \frac{(\eta_{M_i}'' - 1)!}{(\xi_{M_i}'' - 1)! (n - \eta_{M_i}'' - \xi_{M_i}'' - 1)!} \frac{(n - \nu'' - 1)!}{(n + \eta_{M_i}'' - 1)!} \frac{t^{n\lambda''}\nu''}{(t + \lambda'')^{n+\nu''}} \right] \quad (2.24)$$

These distributions describe totally the seismicity of the source considered in terms of the two parameters: magnitudes (M_i) and number of occurrences (n).

In most cases however, the objective involves the description of seismicity at a given site rather than of a source or a number of sources. The method commonly applied consists of the use of transfer functions to determine the effect of the combined sources on the site. The principle of this method has been described extensively in the

literature (Dalal, 1973; Shah, et al., 1975) and will not be presented in this study although it will be used with some modifications.

2.4 Choice of Seismic Ground Motion Parameters

2.4.1 Introduction

Seismic loadings at a given site cannot be directly related to earthquake magnitude measures, such as M , energy release, etc. At a given distance from the epicenter these loadings are dependent upon the ground motion characteristics and the structural characteristics. The ground motion characteristics can be broadly described by intensity, amplitude of motion, frequency content and duration (Gutenberg and Richter, 1956). The intensity is an overall measure of the severity of the ground shaking which may be used to normalize different ground motions for purposes of comparison. The recorded peak value of the acceleration, velocity or displacement is often used to describe the amplitude of ground motion. Other measures of amplitude such as RMS or percentile of exceedence are presented in Chapter IV of this dissertation. Frequency content is a measure of the relative predominance of various frequencies present in the motion. Duration is the length of time for which the significant ground motion lasts.

Since both intensity and frequency content are indicators of the ground motion severity, frequency content related intensity measures have been proposed (Housner).

The frequency content related intensity measures are more meaningful in the representation of the ground motion severity than the ground motion parameters, such as peak acceleration, velocities or

displacements. However, the later are more widely used. They are selected as indicators of the ground motion intensity because of the simplicity and convenience in their determinations and use. The peak ground acceleration is more widely used than the peak velocity and displacement.

The duration of motion is a parameter not often explicitly taken into consideration either in the intensity determination or in the specification of design criteria. Its importance has however been recognized (Richter, 1958; Bolt, 1973; Trifunac, 1976; Dolbry, et al., 1977).

Peak ground acceleration together with the duration of motion are adopted in this work as the intensity measure of the motion at a site. Various other parameters will be described in Chapter IV of this report. The peak ground acceleration parameter is adopted herein mainly to check as to how the present model compares with others using this parameter. In Chapter III results will be compared with the ones presented in Shah, et al., (1975).

2.4.2 Relation between Richter Magnitude and Ground Acceleration

Several empirical relationships between peak ground motion parameters and M have been proposed and employed in practice. All of these relationships have been derived by regression-type statistical analysis of the pertinent ground motion data.

Housner (1969) has proposed a procedure to determine peak ground accelerations incorporating magnitude, fault length, felt area and hypothetical elliptic shapes of felt areas. Wiggins (1964) has

derived a relationship involving peak ground motion parameters, earthquake energy, epicentral distance and site impedance which is defined as the product of the specific soil density and specific shear wave velocity. Esteva and Rosenblueth (1964) have derived and subsequently modified relationships involving hypocentral distance and M. Donovan (1974) has derived a relationship between the same parameters and has presented a summary of the different relationships used. Seed, et al., (1968, 1969) have derived average relationships for the characteristics of bedrock motion. Schnabel and Seed (1973) studied bedrock accelerations in the western United States.

It is worth noting that two approaches are generally proposed and used for specifying the ground motion parameters: one specifies the parameters for bedrock motion at the site and incorporates the site amplification explicitly, and the other specifies the parameters for the so-called firm site and adjusts them for a particular site if that site varies significantly from the firm site condition.

The definition of a firm site is rather qualitative and includes soil condition such as soft but unfissured rock, compacted granular soils, etc. There is however a definition of firm site based on the shear wave velocity. Many researchers use 1500 ft/sec or higher shear wave velocity to imply firm site condition.

The 1974 Esteva-Rosenblueth relationship has been adopted in this work only to allow comparison between the case study presented in Chapter III and Shah, et al., (1975) where this relationship was used. Some modifications have however been introduced.

Esteve and Rosenblueth have derived the following empirical relationship between M , hypocentral distance (R_h) and peak ground acceleration by regression analysis of ground motion data.

$$a = \frac{5000 \exp(.8M)}{(R_h + 40)^2} \quad (2.25)$$

where a = peak ground acceleration (cm/sec^2)

M = Richter magnitude

R_h = hypocentral distance (km)

The following observations regarding the Esteve-Rosenblueth relationship are pertinent:

- The seismic ground motion is highly complex in nature, and any empirical relationship, such as the one described above, can only be an approximation to the actual phenomenon. Therefore, careful judgment must be exercised to determine the applicability of this type of relationship and the associated uncertainties in its use for any given seismic region.
- This relationship is based on regression analysis of the pertinent ground motion data. Confidence intervals for the least square fit are quite wide because of a considerable data scatter. Therefore there is a high degree of uncertainty associated with the values as predicted by the relationship (Fig. 2.1).
- The relationship implies that sites located at an equal distance from the earthquake epicenter will record the same peak ground acceleration for a given earthquake. Thus the locus of equal

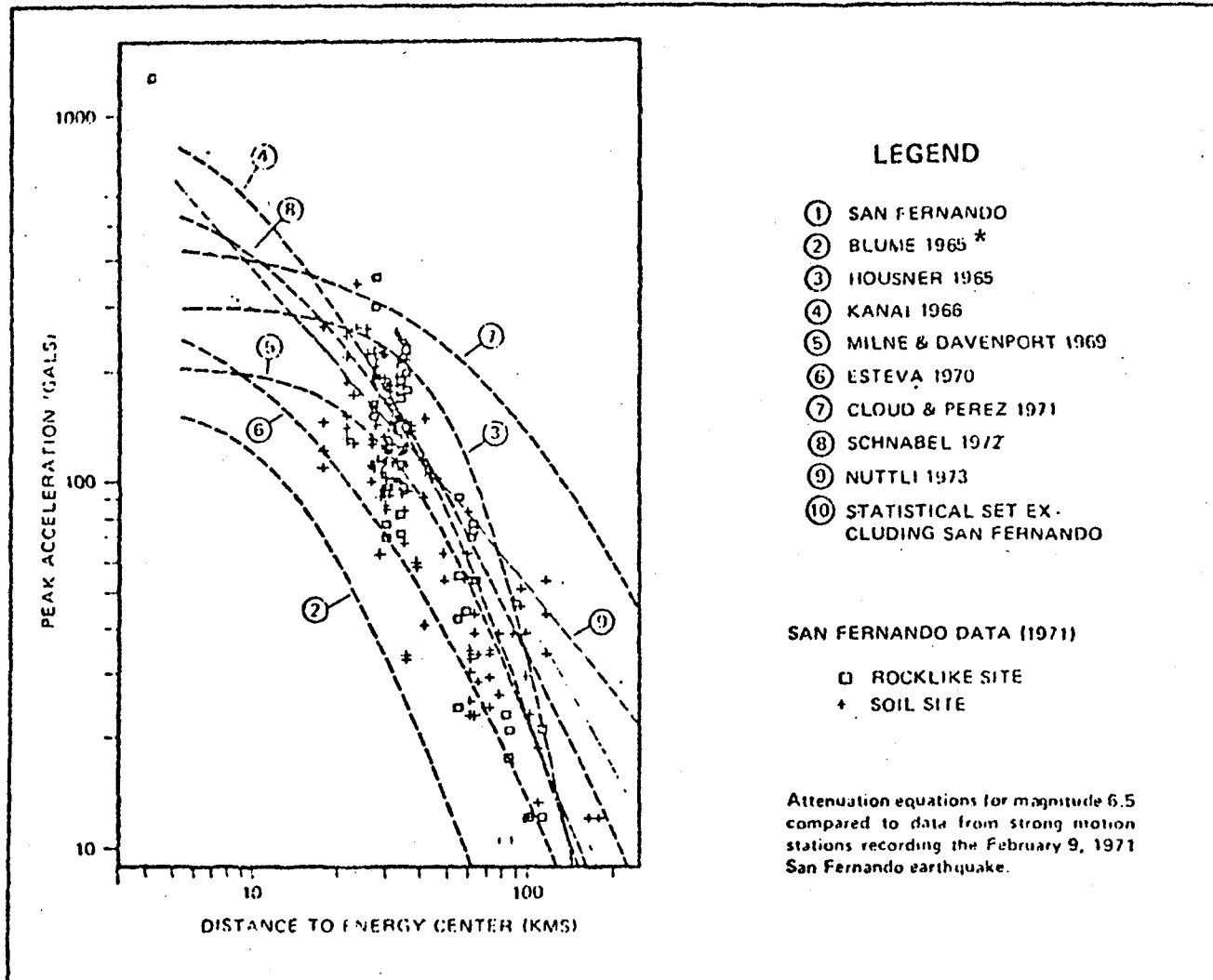


FIGURE 2.1 Attenuation Relationships.
(Donovan, 1974)

* This relation can provide results which are very close to mean data behavior if the soil characteristics for the region are recognized. The soil characteristics input was not used in the preparation of this figure.

peak ground acceleration as given by the relationship is a circle. In the case of large earthquakes which are usually associated with long fault breaks, the isoseismal contours are generally elongated and therefore almost elliptical. The Esteva-Rosenblueth relationship can be modified to incorporate the influence of extensive faulting, for example, by using the relationship between fault breakage and earthquake magnitude (Tocher, 1958). Also the isoseismal may not be circular because of the regional geological structure surrounding the site. Moreover such relationships can only be valid for larger epicentral distances. For "near source" sites use of such relationships by extrapolation gives erroneous results.

- Esteva and Rosenblueth have suggested that this relationship is applicable to firm sites only. Separate investigations should be made for sites with characteristics deviating considerably from those of the firm sites so as to appropriately incorporate the site influence (Wiggins, 1964).

The above observations regarding Esteva-Rosenblueth relationship are generally applicable for other relationships of this type and are given here to point out the limitations of such relationships.

In order to incorporate corrective measures for the first two observations mentioned above, the Esteva-Rosenblueth model is somewhat modified in the present report. Rather than using the relationship in a deterministic way as done previously, the value of an acceleration obtained for a given M and hypocentral distance is used as the mean value of a probability distribution.

The data (Fig. 2.1) does not suggest that any particular distribution is more suitable than the other. The main observation is that the scatter varies proportionally with the mean. For this reason a uniform distribution with a constant coefficient of variation is chosen.* The Esteva-Rosenblueth relationship can thus be plotted as shown in Fig. 2.2 and can be rewritten:

$$a_{\text{mean}} = \frac{5000 \exp(.8M)}{(R_h + 40)^2} \quad (2.26)$$

and for any M and R_h

$$a_{\text{max}} = a_{\text{mean}} + Q_1 a_{\text{mean}} = a_{\text{mean}} (1 + Q_1)$$

$$a_{\text{min}} = a_{\text{mean}} - Q_1 a_{\text{mean}} = a_{\text{mean}} (1 - Q_1)$$

with $0 \leq Q_1 \leq 1$.

The value of a_{max} or a_{min} can be specified by assigning the value of Q_1 .

In the present study, Q_1 is chosen to take values between 0 and 0.5.

In order to simplify the treatment of PGA in the model, this quantity is treated as a discrete variable as follows. Any particular value of PGA is rounded off to the nearest multiple of 0.02 G. Hence for any given M and R_h , the distribution of PGA can be written as:

$$P(a < a_{\text{min}}) = P(a < a_{\text{max}}) = 0 \quad (2.27)$$

$$P(a = A) = 1.0 / \left(\frac{a_{\text{max}} - a_{\text{min}}}{.02} + 1 \right) \quad (2.28)$$

for $a_{\text{min}} \leq A \leq a_{\text{max}}$.

where A, a, a_{min} and a_{max} have been rounded off to the nearest multiple of 0.02 G.

* In a more recent release of the model a log-normal distribution is used.

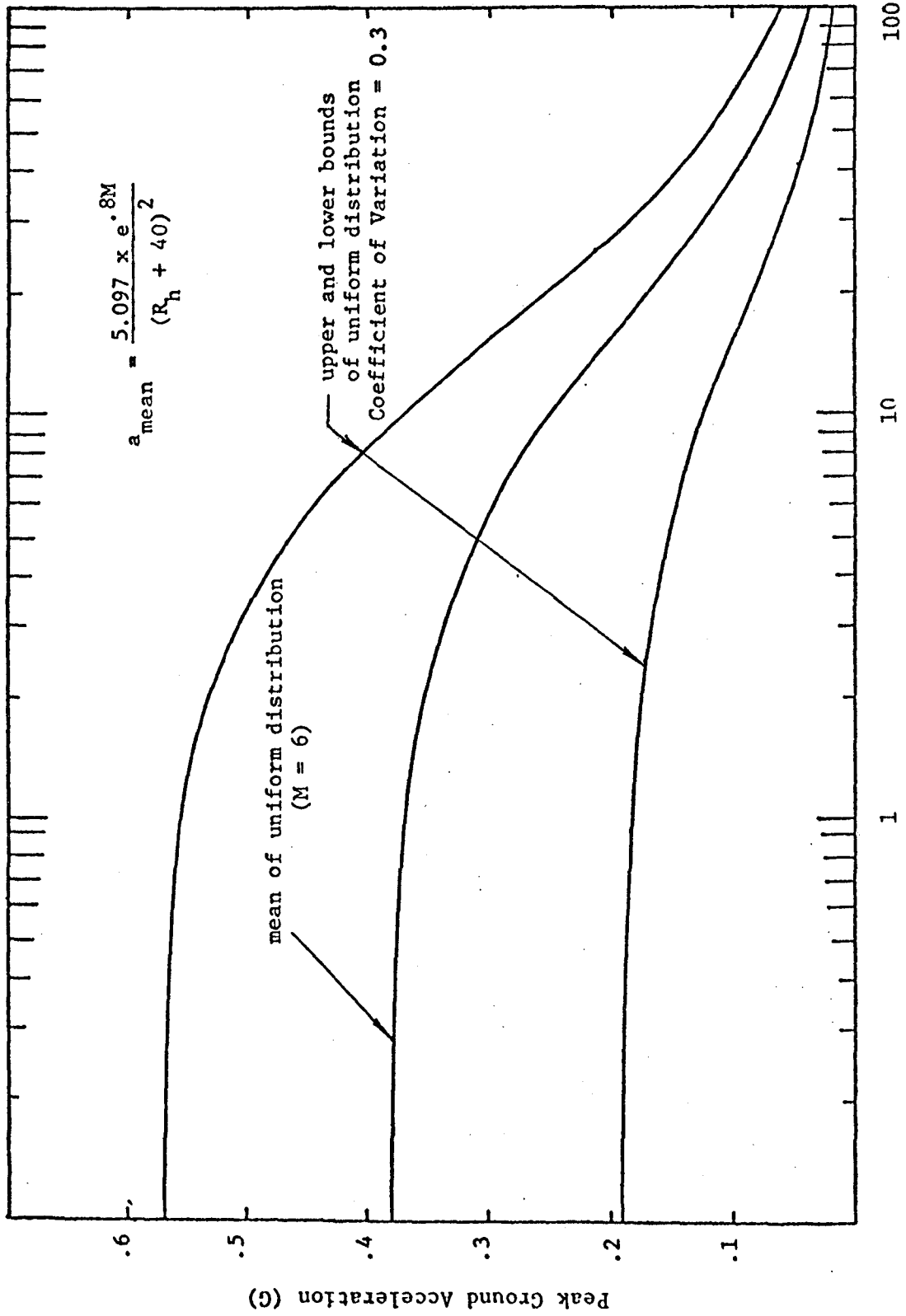


FIGURE 2.2. Esteva Attenuation Relationship with Uniform Distribution

2.4.3 Relation between Richter Magnitude and Duration

The duration parameter d of strong seismic shaking has not been investigated to any great extent. Its prediction from records is still rather rudimentary even though "duration is possibly the single most important factor in producing excessive damage" (H. M. Engle, in Richter, 1958). However several relationships have been presented.

Housner's (1965) relationship for "the strong phase of ground shaking" is essentially a linear law in terms of magnitude M

$$d = 11 M - 53 \quad M \geq 5 \quad (2.29)$$

where d = duration (sec)

M = Richter magnitude

Esteva-Rosenblueth (1964) define the duration s of an "equivalent" ground motion with uniform intensity per unit time as:

$$s = 0.02 \exp(0.74M) + 0.3r \quad (2.30)$$

where s = equivalent duration (sec)

M = Richter magnitude

r = distance from source (km)

For large M , this duration s is about half the duration d presented in equation 2.29.

The following quotation is applicable: "These formulae do express the key dependence of d and M which can be inferred at once from the rupture model of earthquakes. What the formulae lack is a stated

threshold of ground acceleration a to define "strong" and a treatment of frequency" (Bolt, 1973).

Two definitions of duration appear to be useful:

- (A) "Duration at particular frequency is the elapsed time between the first and last acceleration excursions greater than a given level (say 0.05 or 0.02 G). Bolt calls this interval the "bracketed duration". It is sometimes measured by cumulatively adding the squared accelerations and adopting the 95 percentile time interval (Husid, et al., 1969). However, particularly for earthquakes with a complex multiple source (Wyss and Brune, 1967), this definition often leads to a non-physical upper estimate. Trifunac (1976) and Dobry, et al., (1977) adopt this type of approach to obtain duration of earthquake records.
- (B) "Duration at a particular frequency is the total time for which acceleration at that frequency exceeds a given value". This interval called "uniform duration" by Bolt, may equal the corresponding "bracketed duration" or be much less. Uniform duration appears to have a greater mechanical significance with respect to actual structural response behavior.

The treatment of duration in this part of the dissertation is inspired by the paper "Duration of Strong Motion" by B. A. Bolt. Fifth World Conference on Earthquake Engineering, Rome, (1973).

Table 2.1 is reproduced from this paper. It gives the duration (sec) of earthquakes as a function of M and distance (km) from the source. Bolt asserts that observational data scatter indicates that

the chance of exceeding the tabulated values by 20% or more is about 0.1. This implies that the coefficient of variation is about 0.15.

The values given in Table 2.1 are plotted in Fig. 2.3.

The following expressions fit the data conservatively

$$d = d_{50} + .15(50 - R_h) \quad (2.31)$$

for $5 \leq R_h \leq 50$

$$d = d_{50} / (.00112(R_h - 50)^2 + 1) \quad (2.32)$$

for $R_h \geq 50$

where d = duration (sec)

d_{50} = duration (sec) at 50 km from the source for the M considered

R_h = hypocentral distance (km)

M distance	5.5	6.0	6.5	7.0	7.5	8.0	8.5
10	8	12	19	26	31	34	35
25	4	9	15	24	28	30	32
50	2	3	10	22	26	28	29
75	1	1	5	10	14	16	17
100	0	0	1	4	5	6	7
125	0	0	1	2	2	3	3
150	0	0	0	1	2	2	3
175	0	0	0	0	1	2	2
200	0	0	0	0	0	1	2

TABLE 2.1 Bracketed Duration (sec) ($\text{Acc} \geq 0.05g$;
 $\text{freq} \geq 2 \text{ Hz}$)
 distance = distance from source (km).

Reproduced from Bolt (1973).

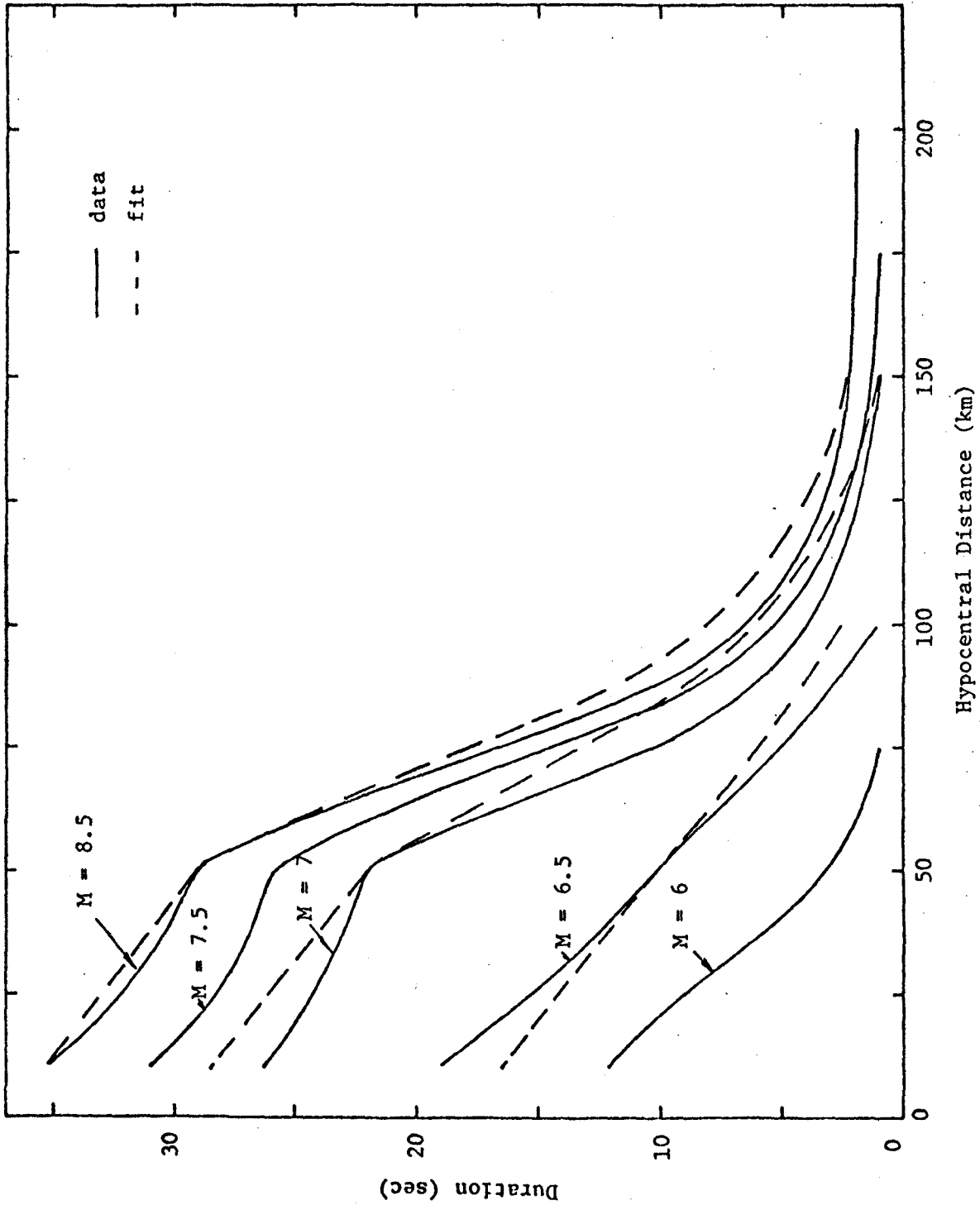


FIGURE 2.3. Bracketed Duration

The durations for different magnitudes at 50 km from the source (d_{50}) are given in Table 2.2 and are plotted in Figure 2.4. They are obtained directly from Table 2.1 using interpolation when the M is missing. They are the base values in the above expressions.

The observations listed in Section 2.4.2 regarding the Esteva-Rosenblueth relationship apply to this relation as well since this duration relationship has been obtained using the same method and the right hand side of the expression contains the same type of parameters. For this reason, the duration is also treated as a random variable with uniform distribution.* The mean is obtained from the above expressions (Eq. 2.31 and 2.32) and the coefficient of variation is kept constant. For any M and R_h , d_{mean} is given from the equations 2.31 and 2.32 where d_{50} is obtained from Table 2.2. The upper and lower bounds of the uniform distribution can be written

$$\begin{aligned}d_{\text{max}} &= d_{\text{mean}} + Q_2 d_{\text{mean}} = d_{\text{mean}} (1 + Q_2) \\d_{\text{min}} &= d_{\text{mean}} - Q_2 d_{\text{mean}} = d_{\text{mean}} (1 - Q_2) \\&\text{with } 0 \leq Q_2 \leq 1\end{aligned}$$

where the values of d_{max} and d_{min} can be specified by assigning the value of Q_2 .

In the present study Q_2 is chosen to take values between 0 and 0.5.

In order to simplify the treatment of duration in the model this quantity is treated as a discrete variable as follows. Any particular value of d is rounded off to the nearest multiple of 2 seconds. Hence for any given M and R_h the distribution of durations can be written as

* In a more recent release of the model a log-normal distribution is used.

Duration (sec)	M
1.5	4.0
1.5	4.25
1.5	4.50
1.5	4.75
2.0	5.0
2.0	5.25
2.0	5.50
2.5	5.75
3.0	6.0
5.5	6.25
10.0	6.50
17.0	6.75
22.0	7.0
24.25	7.25
26.0	7.50
27.25	7.75
28.0	8.0
28.5	8.25
29.0	8.50
29.5	8.75
29.5	9.0

TABLE 2.2 Bracketed Duration (sec) at 50 km from source
(Acc \geq 0.05 g; freq. \geq 2 Hz)

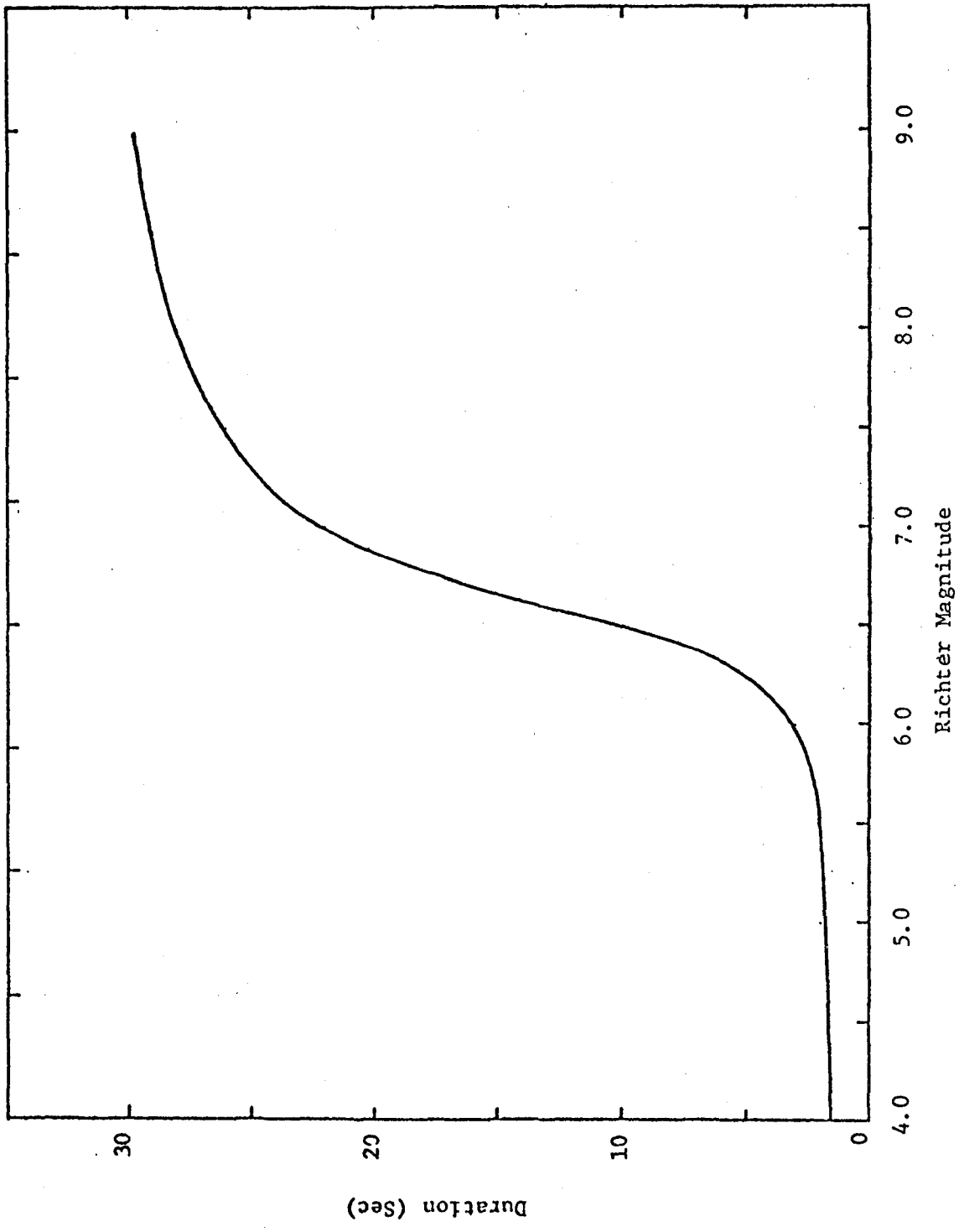


FIGURE 2.4. Bracketed Duration at 50 km from Hypocenter

$$P(d < d_{\min}) = P(d > d_{\max}) = 0 \quad (2.33)$$

$$P(d = D) = 1 / \left(\frac{d_{\max} - d_{\min}}{2} + 1 \right) \quad (2.34)$$

for $d_{\min} \leq D \leq d_{\max}$

where D , d , d_{\min} and d_{\max} have been rounded off to the nearest multiple 2 seconds.

2.5 Probabilistic Estimation of Peak Ground Acceleration and Duration

2.5.1 Introduction

Up to this point, the following steps have been discussed and developed:

- The seismic sources that can affect a site under consideration have been located. They can be line sources or area sources.
- The seismicity of each source (considered homogeneous over the source) has been determined. It is described in terms of the probability of occurrences of any number of events of any Richter Magnitude (M_1) for a given period of time. Bayesian concept is applied in this part of the analysis.
- Two attenuation equations have been chosen to convey the information from the sources to the site. The parameters chosen to describe the seismicity at the site are the peak ground acceleration and the duration of the shaking.

In general a site is surrounded by a number of sources. For design purposes, the probabilistic loading at the site is obtained by combining the effect of all the sources. The loading is expressed in terms of the probability of occurrence of at least one event of a given PGA and the probability of at least one event of a

given duration. Probability distribution functions or cumulative distribution functions for these two parameters are developed for each site for a given time period t . These two distributions totally describe the seismicity of the site in terms of the chosen parameters.

2.5.2 Contribution of a Segment

All the quantities (M , PGA, duration) used in the Section 2.4 are discretized to equal step increments. Hence, all the integral signs can be replaced by summations. Since the hypocentral distance is a parameter in both attenuation relationships the seismic source (usually a line or an area) must be divided into segments ΔL or ΔA to take into consideration the distance variation to the site from the source. The size of the segments is chosen small enough such that the approximation from a continuous to discrete computation is acceptable. The seismicity within a source remains the same from segment to segment. It is easily obtained from the data gathered for each source. The mean rate of occurrence of events (eq. 2.9) for the whole source becomes for a segment:

$$\lambda''' = (\lambda'' \Delta L) / L \quad (2.35)$$

where L = length or area of source

ΔL = length or area of segment.

The distribution on the number of events for each segment is obtained from equation 2.9 where λ'' is replaced by λ''' . The conditional distribution on magnitudes given M events (eq. 2.20) remains unchanged by the segmentation of the sources. The distribution of the

number of occurrences of each magnitude is given by equation 2.21. The probability of zero occurrence of each magnitude is given by equation 2.22. The same distribution applies for any segment of the source. The distribution of the number of occurrences of each M increment can be presented under a matrix form that describes the total seismicity of the segment (Table 2.3.).

Only a finite number of different magnitude events can occur on the segment (from the largest to the smallest magnitude considered). The number of occurrence associated with them. They are disregarded when this probability becomes negligible, say 10^{-8} . Hence the total number of events is finite and can easily be handled under the summation signs.

Since the distance from the segment to the site is known, all the parameters of the attenuation relationships (Eq. 2.26, 2.31, 2.32) are determined. Hence, for each M a uniform distribution on acceleration is obtained as follows:

$$a_{\text{mean}} = \frac{5000 e^{.8M}}{(R_h + 40)^2} \quad (2.36)$$

a_{min} and a_{max} are determined from the coefficient of variation and rounded off to the closest multiple of 0.02 G.

The contribution to acceleration greater or equal to a_i of all events M_j occurring on the same segment is written as:

$$P(A \geq a_i) = p \times P(M_j) + [1 - (1 - p)^2] P(2M_j) + \dots \\ + [1 - (1 - p)^n] P(nM_j) \quad (2.37)$$

		Richter Magnitude							
		4.0	4.25	4.50	4.75	--	--	M_i	--
Number of Occurrences	0								
	1								
	2								
	3							$P(3M_i)$	
	:								
	:								
	:								
	n							$P(nM_i)$	
	:								
	:								

TABLE 2.3. Seismicity of a Segment.

Probability of occurrence of any number of events
of given M_i .

where: $P(A \geq a_i)$ = probability of obtaining acceleration greater or equal to a_i at least once.

$P(kM_j)$ = probability of k occurrences of event M_j with $k = 1, 2 \dots n$

$p = P(A \geq a_i/M_j)$, probability of obtaining an acceleration greater or equal to a_i given an event M_j

Setting $q = 1 - p$, the above expression can be rewritten:

$$\begin{aligned} P(A \geq a_i) &= (1 - q) P(M_j) + (1 - q^2) P(2M_j) + \dots + (1 - q^n) P(nM_j) \\ &= P(M_j) + P(2M_j) + \dots + P(nM_j) - \\ &\quad q P(M_j) - q^2 P(2M_j) - \dots - q^n P(nM_j) \end{aligned} \quad (2.38)$$

$$P(A \geq a_i) = 1 - P(\text{no } M_j) - \sum_{k=1}^n q^k P(kM_j) \quad (2.39)$$

$$P(A \leq a_i) = P(\text{no } M_j) + \sum_{k=1}^n q^k P(kM_j) \quad (2.40)$$

with n chosen such that $q^k P(nM_j)$ can be neglected.

The above discussion assumes independence among events. Hence, the contribution of all possible events can be combined as follows:

$$P(A \geq a_i) \Big|_{\text{one segment}} = 1 - \prod_{\substack{\text{all} \\ M_j}} \left[1 - P(A \geq a_i) \Big|_{M_j} \right] \quad (2.41)$$

The whole range of magnitudes is covered starting with the largest one down to the smallest one that generates a noticeable effect at

the site ($M_j \geq M_{\min}$ as a function of distance). This eliminates the consideration of a large number of events.

Similarly for duration

$$d_{\text{mean}} = d_{50} + .15(50 - R_h) \quad R_h \leq 50 \text{ km} \quad (2.42a)$$

$$= \frac{d_{50}}{(.00112(R_h - 50)^2 + 1)} \quad R_h \geq 50 \text{ km} \quad (2.42b)$$

d_{\min} and d_{\max} are determined from the coefficient of variation and rounded off to the closest multiple of 2 seconds.

$$P(D \leq d_i) = P(\text{no } M_j) + \sum_{k=1}^r q^k P(kM_j) \quad (2.43)$$

with n chosen such that $q^n P(n M_j)$ can be neglected.

where: $P(D \geq d_i)$ = probability of obtaining
duration greater or equal
to d_i at least once.

$P(k M_j)$ = probability of k occurrences
of event M_j with $k = 1, 2, \dots, n$

$p = 1 - q = P(D \geq d_i / M_j)$, probability
of obtaining an acceleration
greater or equal to d_i
given an event M_j

and

$$P(D \geq d_i) | \text{one segment} = 1 - \prod_{\text{all } M_j} [1 - P(D \geq d_i)_{M_j}] \quad (2.44)$$

2.5.3 Contribution of One or Several Sources

As the events are assumed independent from segment to segment, the contribution of each segment of a source is combined as in equation 2.41 and 2.43.

$$P(A \geq a_i) \Big|_{\text{one source}} = 1 - \prod_{\text{all segments}} \left[1 - P(A \geq a_i) \Big|_{\text{one segment}} \right] \quad (2.44a)$$

$$P(D \geq d_i) \Big|_{\text{one source}} = 1 - \prod_{\text{all segments}} \left[1 - P(D \geq d_i) \Big|_{\text{one segment}} \right] \quad (2.44b)$$

The seismicity at the site in terms of probability of occurrence of each PGA value at least one time and each duration at least one time can be described by the two vectors in Table 2.4.

When several sources are considered, the same principle is applied for each source. Thus,

$$P(A \geq a_i) = 1 - \prod_{\text{all sources}} \left[1 - P(A \geq a_i) \Big|_{\text{one source}} \right] \quad (2.45a)$$

$$P(D \geq d_i) = 1 - \prod_{\text{all sources}} \left[1 - P(D \geq d_i) \Big|_{\text{one source}} \right] \quad (2.45b)$$

These expressions give the probability of occurrence at the site of at least one acceleration and at least one duration greater than a given level.

Once a cumulative distribution function is established for each node of a grid, a seismic exposure map can be prepared for any desired probability of nonexceedence. Smoothing procedures are used to obtain smooth contours at given intervals.

Such maps can be drawn for any of the parameters of interest. A detailed procedure for such a hazard mapping is given in Chapter III for the peak ground acceleration and duration. However, the methodology developed here is general enough that a map for spectral responses can be obtained if the attenuation information for such response parameters is known.

PGA (G)	Prob of at least one occurrence
0	
.02	
.04	
.06	
:	
:	
a_i	
:	
:	
a_n	

Duration (Sec)	Prob of at least one occurrence
0	
2	
4	
6	
:	
:	
d_i	
:	
:	
a_m	

TABLE 2.4. Probability of at least one occurrence of any
PGA a_i .

Probability of at least one occurrence of any
duration d_i .

CHAPTER III
APPLICATION TO NICARAGUA



3.1 Introduction

In this chapter seismic hazard maps are developed for Nicaragua using the model presented in Chapter II. Each step of the methodology is covered in the example:

- gathering of the available data
- location of the seismic sources
- Bayesian description of the seismicity
- use of probabilistic attenuation relationships
- development of iso-duration and iso-acceleration maps

This case study is presented to emphasize the following points:

- the simplicity of incorporating any subjective information with the historical data to obtain Bayesian estimation on fault seismicity.
- the versatility of the method in which various uncertainties can be incorporated in the attenuation relationships
- the similarity between the results obtained with this model and the ones presented in Shah, et al., (1975)*.

3.2 Seismic Data Base for Nicaragua

The earthquake data used in this study is the same as the data used in Shah, et al. (1975). Two main sources of information are considered. The NEIC-NOAA data file covering the period from January 1900 to August 1973 constitutes the primary source of information and is referred to hereafter as Source 1. The Catalogue of Nicaraguan Earthquakes, 1520-1973, by David J. Leeds (1973), is referred to as Source 2. The latter

* In the following pages, continuous reference is made to "A Study of Seismic Risk for Nicaragua, Part I", Shah, et al. (1975). It might be advantageous to the reader to make himself familiar with Chapters I through V of this publication.

is used to obtain:

- data about earthquakes not reported or incompletely reported in the NEIC-NOAA file (1900-1973);
- data about earthquakes associated with volcanic activity along the Cordillera de los Marrabios (1850-1973).

The data covering the period from 1520 to 1850 is not considered in the analysis. This information is mainly descriptive, of low reliability and difficulty evaluable in terms of presently used parameters. Moreover, it is biased in that sense that only few large events situated close to populated areas have been recorded. Hence, it cannot be used as such together with more homogeneous and better quality data. However its historical interest should not be neglected and could be used, for example, to help the analyst in formulating his subjective input. The time period of data gathering is thus 73 years for the entire country and 123 years for earthquakes associated with volcanic activity along the Cordillera de los Marrabios.

In spite of the complementarity of the two sources, a large number of events remain insufficiently documented to be used as such in the analysis. Rather than disregarding these events, the missing information is generated using a Monte Carlo process supplemented by judgment. It is felt that the total analysis benefits more than suffers from such an additional input.

The following remarks are valid for both sources:

- No critical study is made regarding the validity of the information and the reliability of the data.
- Whenever information as basic as epicentral location or magnitude is missing, the event is disregarded.

- Events with Richter magnitude smaller than 3.0 are not considered.

Source 1

When complete, the information contained in this source includes for a given event: time of occurrence, epicentral location (degrees), depth of hypocenter (km), and magnitude. The magnitude is in terms of one of the following:

- (1) CGS M_b average (Coastal Geodetic Survey body wave magnitude)
- (2) CGS M_s average (Coastal Geodetic Survey surface wave magnitude)
- (3) Richter magnitude M .

The acceleration attenuation relationships used in Chapter II are based on the Richter magnitude. Hence, when missing, this information is generated from M_b or M_s . It is known that for a given part of the world, the Richter magnitude and CGS M_b are linearly related such that

$$M = a + b M_b \quad (3.1)$$

In order to determine the coefficients a and b , a regression analysis is run for all the earthquakes of which M and M_b are known using the total data of Central America. The Richter magnitude is then obtained by substituting the value of M_b in equation 3.1.

Whenever data on depth of hypocenter are not available, a depth is assigned, as will be explained later in the chapter.

From Source 1, 419 events contain complete information; they are shown as a function of depth in Table 3.1.

Source 2

When complete, the information contained in this source includes for a given event: time of occurrence, epicenter location (degrees),

depth, Richter magnitude and sometimes a short description of the seismic event. The depth is either expressed in km or by a letter symbol N (0 - 60 km) or I (70 - 200 km). In the same way, the Richter magnitude is either expressed by its numerical value or by a letter symbol, as follows:

$$B - 7 \leq M \leq 7.7$$

$$C - 6 \leq M \leq 6.9$$

$$D - 5.3 \leq M \leq 5.9$$

$$E \quad M < 5.3$$

Through a simulation process, all the events taken from Source 2 are assigned a numerical Richter magnitude from letter magnitude.

Hence, an additional 196 events are obtained (including events from Source 1 with partial information), distributed as follows:

43 events associated with volcanic activity and with shallow hypocenters N (0 - 60 km).

40 events with shallow hypocenters N (0 - 60 km).

3 events with deep hypocenters I (70 - 200 km).

63 events with no data on depth.

47 events with numerical data on depth (km).

The 466 earthquakes with complete data (419 from Source 1 and 47 from Source 2), are plotted as a function of depth. Using those plots together with epicenter location, magnitude value, partial information on depth and judgment, the 149 remaining events are assigned appropriate depths. This led to a total data of 615 events ranging from

TABLE 3.1

Data from Source 1, Sorted According to Depth of Hypocenter
(419 Events)

Number of Earthquakes	Depth Range (km)
8	0 - 9
9	10 - 19
12	20 - 29
157	30 - 39
35	40 - 49
32	50 - 59
34	60 - 69
32	70 - 79
18	80 - 89
14	90 - 99
13	100 - 109
9	110 - 119
3	120 - 129
7	130 - 139
3	140 - 149
7	150 - 159
3	160 - 169
6	170 - 179
3	180 - 189
5	190 - 199
9	200 - 215

5 to 215 km in depth and from 3.0 to 7.7 in magnitude. Appendix A gives the listing of those earthquakes. In Figures 3.1 through 3.6 (Charts 2 through 7 from Shah, et al., 1975), they are plotted as a function of depth.

From these figures the general seismic pattern of Nicaragua can be divided into the following regions:

(1) The Benioff Zone dipping North East toward the Nicaraguan coast. This zone is marked by numerous earthquakes covering the whole range of magnitude (larger as depth increases) and extending several hundred kilometers into the earth's interior. The shallow earthquakes (~ 30 km) due to this source are from 30 to 100 km away from the coast. As the epicenters get closer to the coast, the hypocenters get deeper. Hence, under Managua the hypocenters of earthquakes situated on the Benioff Zone are very deep (100 - 200 km).

(2) In contrast, for the local seismic sources, such as the ones identified under Managua, the hypocenters are shallow (5 - 30 km). In magnitude scale, these sources do not generate major earthquakes such as those on the Benioff Zone. However, due to their shallowness and proximity to populated areas, they have caused extensive damage and loss of life in past history. The December 23, 1972 event was due to the local source under Managua.

(3) The line of volcanoes from Northwest to Southeast (Cordillera de los Marrabios) represents sources of future seismic activities. Volcanic eruptions are seldom by themselves sources of seismic activity, and in the past various earthquakes have been recorded preceding volcanic eruptions. For this reason earthquakes "associated" with volcanic activity were treated in the model as shallow tectonic earthquakes.

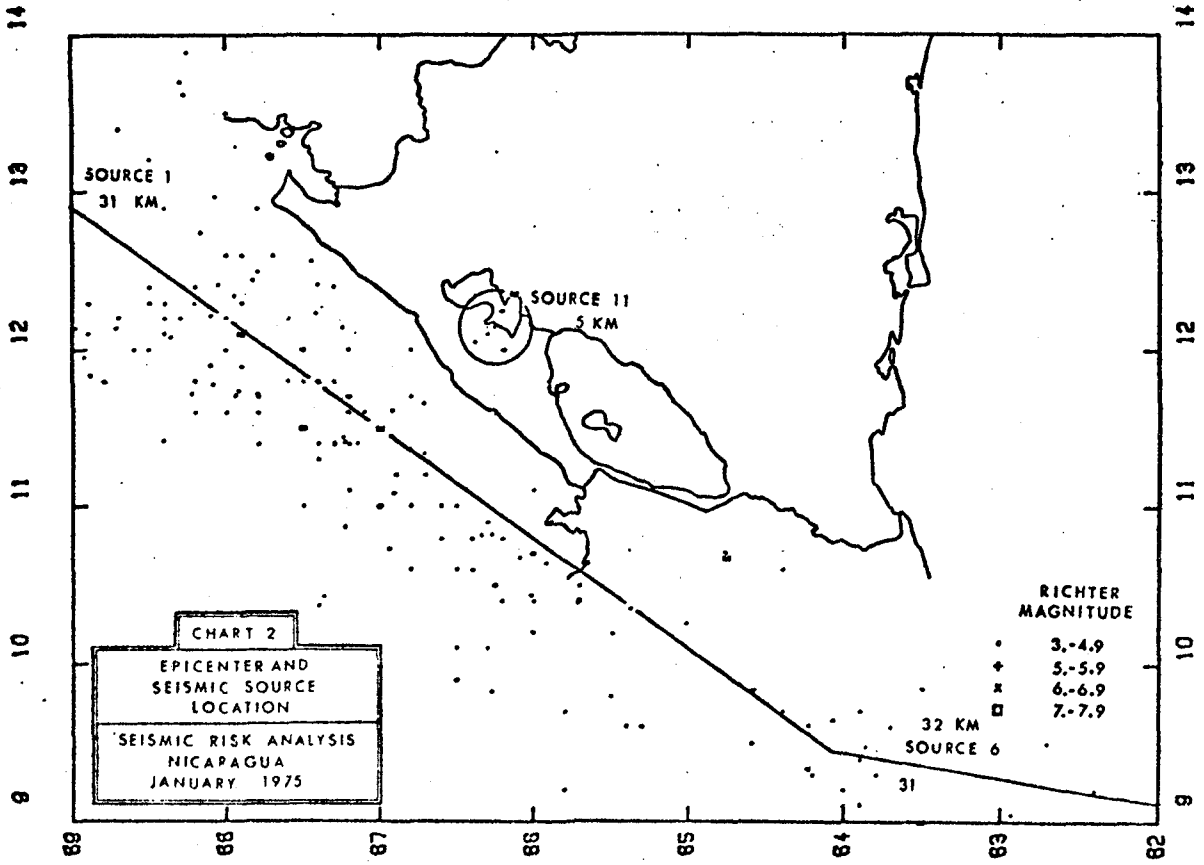


FIGURE 3.1

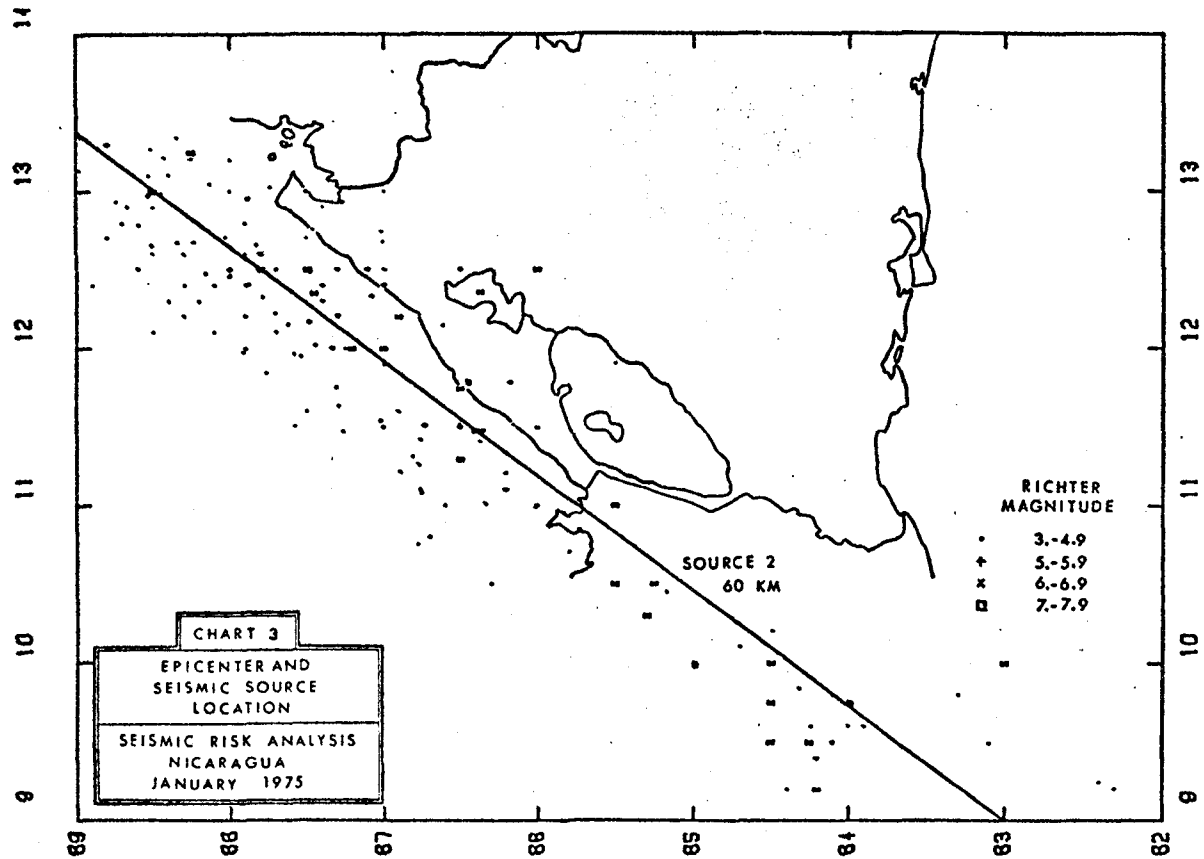


FIGURE 3.2

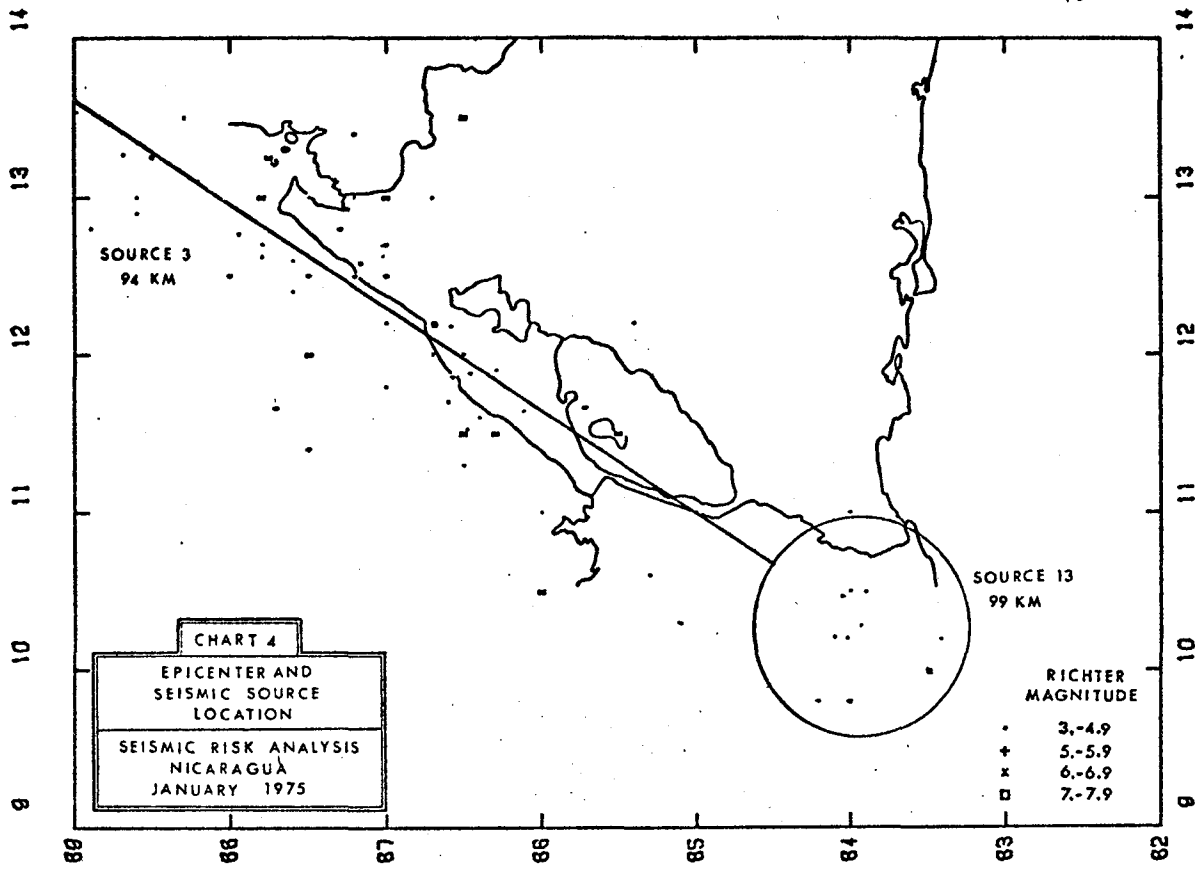


FIGURE 3.3

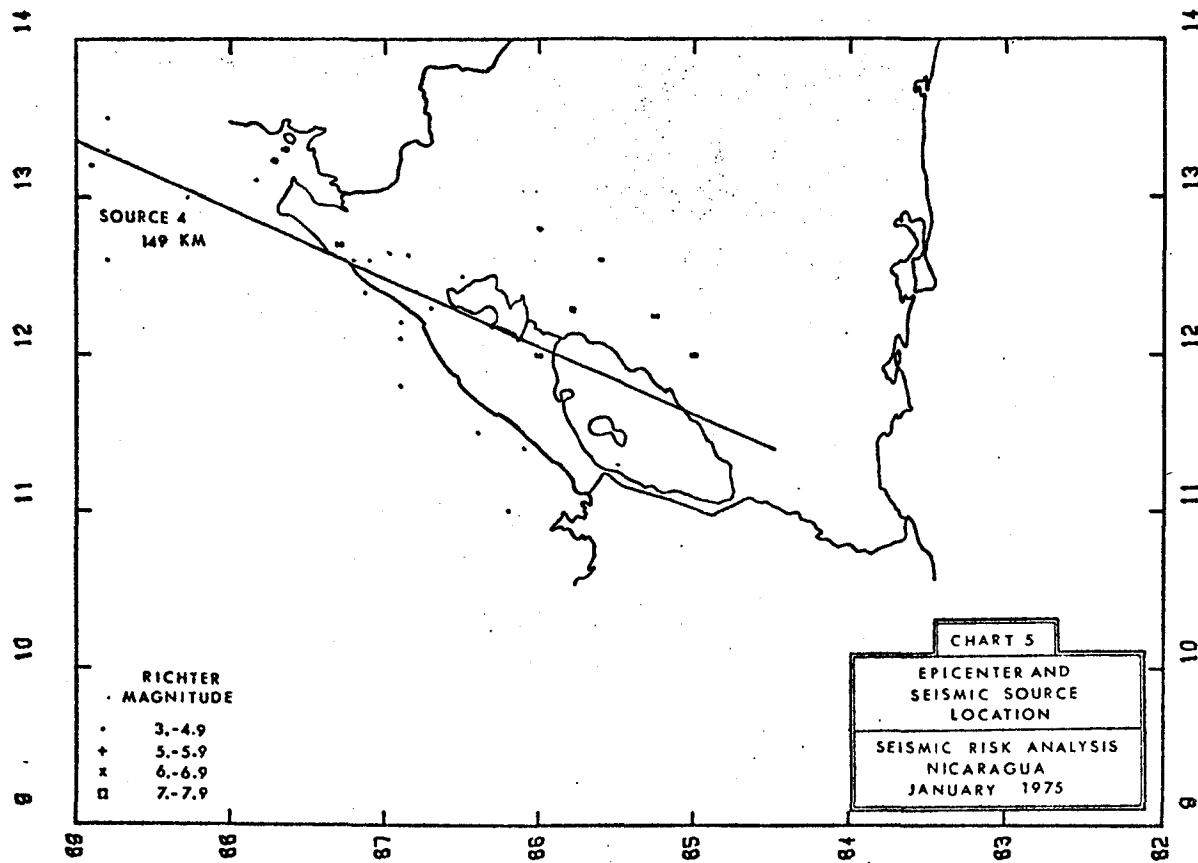


FIGURE 3.4

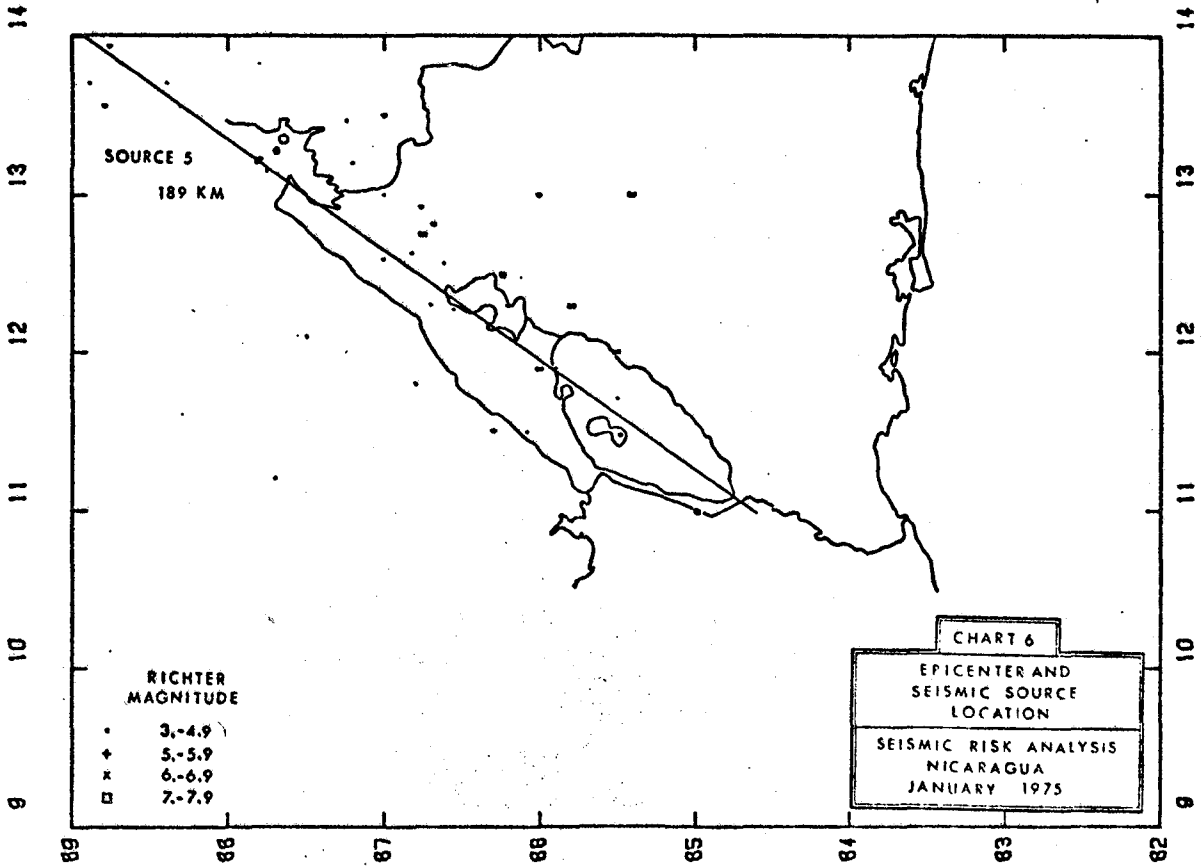


FIGURE 3.5

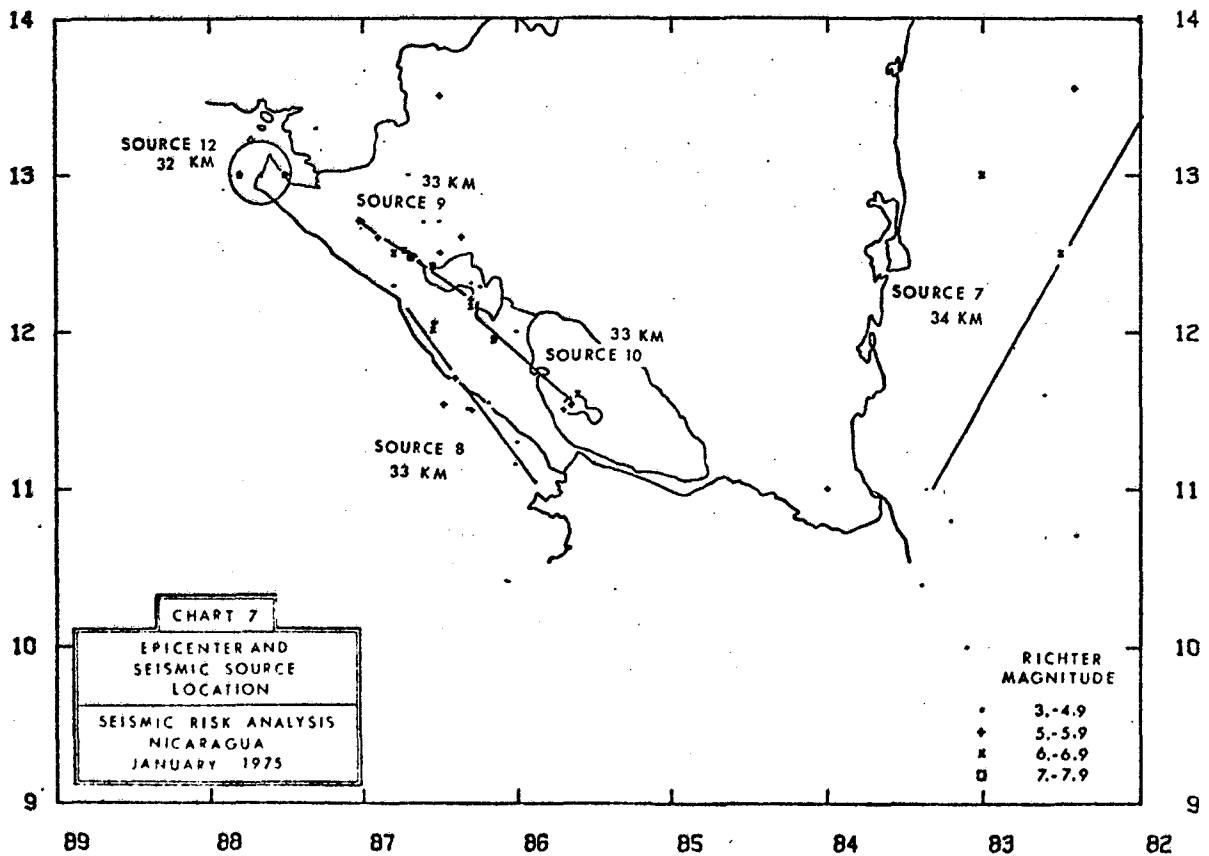


FIGURE 3.6

(4) Two shallow (~ 30 km) seismic regions, one more or less coinciding with the Pacific seashore between Lake Managua and the Costa Rica border, the other one in the Gulf of Fonseca.

(5) The Atlantic coast of low seismicity.

Source Location and Seismicity

Based on the above observations, the total number of events is divided into 13 seismic sources: Ten of these are line sources and three are area sources. Table 3.2 shows these 13 sources, the number of events and the depth range of each source.

Appendix A gives a listing of the earthquakes included in each source. Line sources are located by fitting a line through the data using regression analysis. For area sources, the centroid is obtained from the data and the radius taken as the distance from the centroid to the most distant epicenter in the source.

The depth of each source is computed as an average hypocentral depth of all the earthquakes included in the source. Earthquakes with no or limited depth information are not included in this averaging process. However, they are considered in determining the location and the seismicity of the sources. Figures 3.1 through 3.6 show the sources locations and depths.

Limitations

It can be said that there are limitations to the use of available data for the Nicaragua region. These limitations are given below.

TABLE 3.2

Seismic Sources for Nicaragua

Source	Number of Events	Names of Source	Depth (km.)
1 Line	159	Benioff	5 - 39
2 Line	186	Benioff & Costa Rica	40 - 79
3 Line	72	Benioff	80 - 109
4 Line	31	Benioff	110 - 159
5 Line	41	Benioff	160 - 215
6 Line	23	"Costa Rica"	5 - 39
7 Line	11	Atlantic	All Depths
8 Line	12	Pacific Coast Line	33
9 Line	57	Line of Volcanoes	33
10 Line	57	Line of Volcanoes	33
11 Area	5	Managua Area	5
12 Area	8	Gulf of Fonseca	33
13 Area	10	Costa Rica Area	80 - 109

1. 24% of the data contain incomplete information regarding the depth. This information was added from either judgment or by correlating the event with other data where the depth information was available.

2. 32% of the data have magnitude defined by a symbol. Numerical value of magnitude for these cases was obtained through simulation.

3. The reliability of the total data base was not evaluated.

(i) Some information was obtained from church and historical records.

(ii) Distribution of information over the country is biased. Populated areas have better records than sparsely populated areas. (No population → no records.)

(iii) Epicentral location could be in error due to lack of a good grid of recording system. It is hoped that the recording network presently installed by the Nicaraguan authorities, the U.S.G.S. and private organizations in Nicaragua will help in increasing the understanding of attenuation relationships and the accuracy of epicentral locations in the future. Such calibration may help in relocating the past events (Dewey, 1973).

It is felt that the work done by Dewey (1973) and others in calibrating the epicentral locations through the ESSO refinery record does not have sufficient experimental evidence as yet. Hence, no hypocenters are moved based on Dewey's work. (One exception is the 1931 earthquake-stadium fault.)

The above remarks point out the need for using the Bayesian model which can help to supplement the incompleteness of the data with subjective information.

3.3 Data Analysis and Subjective Input

As mentioned previously, the Richter magnitude is treated as a discrete variable. Its values are rounded off to the closest multiple of .25 on the Richter magnitude scale. These rounded off values are referred thereafter as M_i 's (4.0; 4.25; 4.5; 4.75; etc.). Events of M_i smaller than 4.0 are not considered in the model (although they are used in obtaining the subjective information on occurrences). Events of M_i smaller than 3.0 are totally disregarded.

The data is analyzed in two steps. In the first step information is obtained to determine the rate of occurrence of events independently of magnitude. This is used as an input to the Poisson-gamma model. In the second step information is gathered about the distribution of magnitudes of these events. For each M_i the probability of success given one trial is determined. A trial is defined as the occurrence of an earthquake. A success is defined as the earthquake being of the M_i considered while a failure is defined as the earthquake being of any other magnitude. This is used as the input to the Bernoulli-Beta model.

The analysis is based on two sources of information: the available data as presented in Appendix A and the subjective input introduced through Bayesian analysis. In order to make the comparison with Shah, et al. (1975), more meaningful, the subjective input is arbitrarily chosen as the input used for that study. One has to remember that one goal of the case study is to compare the behavior of the model developed in this dissertation with the other available models. This can only be done if the data used in comparing the two models is similar. The values of all the parameters used for the 13 sources are listed in Tables 3.3 through 3.14.

TABLE 3.3
Seismic Source Parameters

Source 1 (line) Benioff (5 - 39 km)				
Time Data Base (T): 73 years Number of Recorded Events (N): 82 v' from log-linear fit : 72 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $v'' = v' + N = 72 + 82 = 154$ $\eta''_{M_1} = \eta'_{M_1} + N = 72 + 82 = 154$				
Richter Magnitude (M_1)	Nb of Recorded Occurrences in M_1 bands (R_{M_1})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.7) (N_c)	Nb of Occurrences in M_1 bands (log-linear fit) (E'_{M_1})	$E'_{M_1} + R_{M_1}$ (E''_{M_1})
4.0	26	72.0	18.0	44.0
4.25	18	54.0	12.0	30.0
4.50	14	42.0	11.0	25.0
4.75	4	31.0	7.0	11.0
5.0	2	24.0	4.5	7.5
5.25	4	18.5	4.5	8.5
5.50	3	14.0	3.5	6.5
5.75	6	10.5	2.3	8.3
6.0	2	8.2	2.0	4.0
6.25	1	6.2	3.7	4.7
6.50	1	2.5	1.5	2.5
6.75	1	1.0	1.0	2.0
7.0				
7.25				
7.50				
7.75				

TABLE 3.4
Seismic Source Parameters

Source 2 (Line) Benioff & Costa Rica (40 - 79 km)				
Time Data Base (T) : 73 years Number of Recorded Events (N): 128 v' from log-linear fit : 125 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $v'' = v' + N = 125 + 128 = 254$ $\eta''_{M_1} = \eta'_{M_1} + N = 125 + 128 = 254$				
Richter Magnitude (M_1)	Nb of Recorded Occurrences in M_1 bands (R_{M_1})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.8) (N_c)	Nb of Occurrences in M_1 bands (log-linear fit) (E'_{M_1})	$E'_{M_1} + R_{M_1}$ (E''_{M_1})
4.0	26	125.0	20.0	46.0
4.24	9	105.0	19.0	28.0
4.50	22	86.0	15.0	37.0
4.75	6	71.0	12.0	18.0
5.0	10	59.0	11.0	21.0
5.25	7	48.0	7.0	14.0
5.50	10	41.0	7.0	17.0
5.75	5	34.0	5.5	10.5
6.0	13	28.5	5.0	18.0
6.25	3	23.5	4.0	7.0
6.50	3	19.5	3.5	6.5
6.75	7	16.0	2.5	9.5
7.0	3	13.5	5.9	8.9
7.25	4	7.6	6.6	10.6
7.50		1.0	1.0	1.0
7.75				

TABLE 3.5
Seismic Source Parameters

Source 3 (Line) Benioff (80 - 109 km)				
Time Data Base (T): 73 years Number of Recorded Events (N): 51 v' from log-linear fit: 50.5 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $v'' = v' + N = 50.5 + 51 = 101.5$ $n''_{M_i} = n'_{M_i} + N = 50.5 + 51 = 101.5$				
Richter Magnitude (M_i)	Nb of Recorded Occurrences in M_i bands (R_{M_i})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.9) (N_c)	Nb of Occurrences in M_i bands (log-linear fit) (ξ'_{M_i})	$\xi'_{M_i} + R_{M_i}$ (ξ''_{M_i})
4.0	8	50.5	5.5	13.5
4.25	2	45.0	4.5	6.5
4.50	3	41.5	4.0	7.0
4.75	4	36.5	3.5	7.5
5.0	3	33.0	3.0	6.0
5.25	2	30.0	3.0	5.0
5.50	5	27.0	3.0	8.0
5.75	7	24.0	2.5	9.5
6.0	8	21.5	2.0	10.0
6.25	3	19.5	2.0	5.0
6.50	1	17.5	1.5	2.5
6.75	2	16.0	10.4	12.4
7.0	2	5.6	4.2	6.2
7.25	1	1.4	1.05	2.05
7.50		0.35	0.25	0.25
7.75		0.1	0.1	0.10

TABLE 3.6
Seismic Source Parameters

Source 4 (Line) Benioff (110 - 159 km)				
Time Data Base (T): 73 years Number of Recorded Events (N): 16 v' from log-linear fit: 18 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $v'' = v' + N = 18 + 16 = 34$ $n''_{M_i} = n'_{M_i} + N = 18 + 16 = 34$				
Richter Magnitude (M_i)	Nb of Recorded Occurrences in M_i bands (R_{M_i})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.10) (N_c)	Nb of Occurrences in M_i bands (log-linear fit) (ξ'_{M_i})	$\xi'_{M_i} + R_{M_i}$ (ξ''_{M_i})
4.0	5	18.0	2.7	7.7
4.25	1	15.3	2.3	3.3
4.50	1	13.0	2.0	3.0
4.75	0	11.0	1.5	1.5
5.0	1	9.5	1.5	2.5
5.25	0	8.0	1.2	1.2
5.50	1	6.8	1.1	2.1
5.75	1	5.7	0.8	1.8
6.0	2	4.9	0.8	2.8
6.25	0	4.1	0.6	0.6
6.50	2	3.5	0.5	2.5
6.75	1	3.0	2.6	2.6
7.0	1	1.4	0.95	1.95
7.25		0.45	0.35	0.35
7.50		0.1	0.1	0.1
7.75				

TABLE 3.7
Seismic Source Parameters

Source 5 (Line) Benioff (160 - 215 km)				
Time Data Base (T) : 73 years Number of Recorded Events (N) : 31 ν' from log-linear fit : 31 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $\nu'' = \nu' + N = 31 + 31 = 62$ $\eta''_{M_i} = \eta'_{M_i} + N = 31 + 31 = 62$				
Richter Magnitude (M_i)	Nb of Recorded Occurrences in M_i bands (R_{M_i})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.11) (N_c)	Nb of Occurrences in M_i bands (log-linear fit) (E'_{M_i})	$E'_{M_i} + R_{M_i}$ (E''_{M_i})
4.0	5	31.0	5.0	10.0
4.25	2	26.0	4.0	6.0
4.50	3	22.0	4.0	7.0
4.75	2	18.0	3.0	5.0
5.0	5	15.0	2.5	7.5
5.25	3	12.5	2.0	5.0
5.50	1	10.5	1.7	2.7
5.75	2	8.8	1.5	3.5
6.0	2	7.3	1.3	3.3
6.25	1	6.0	.9	1.9
6.50	0	5.1	.9	.9
6.75	2	4.2	.6	2.6
7.0	0	3.6	.6	.6
7.25	2	3.0	.5	2.5
7.50	0	2.5	.4	.4
7.75	1	2.1	1.2	2.2
8.0	0	.9	.58	.58
8.25	0	.32	.29	.29
8.50	0	.03	.03	.03
8.75				

TABLE 3.8
Seismic Source Parameters

Source 6 (Line) Costa Rica (5 - 39 km)				
Time Data Base (T) : 73 years Number of Recorded Events (N) : 10 ν' from log-linear fit : 11 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $\nu'' = \nu' + N = 11 + 10 = 21$ $\eta''_{M_i} = \eta'_{M_i} + N = 11 + 10 = 21$				
Richter Magnitude (M_i)	Nb of Recorded Occurrences in M_i bands (R_{M_i})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.12) (N_c)	Nb of Occurrences in M_i bands (log-linear fit) (E'_{M_i})	$E'_{M_i} + R_{M_i}$ (E''_{M_i})
4.0	1	11.0	2.0	3.0
4.25	2	9.0	1.6	3.6
4.50	2	7.4	1.4	3.4
4.75	1	6.0	1.0	2.0
5.0	1	5.0	.9	1.9
5.25	0	4.1	.7	.7
5.50	1	3.4	.6	1.6
5.75	0	2.8	.5	.5
6.0	0	2.3	.4	.4
6.25	0	1.9	.35	.35
6.50	2	1.55	1.3	3.3
6.75		.25	.15	.15
7.0		.1	.1	.1
7.25				
7.50				
7.75				

TABLE 3.9
Seismic Source Parameters

Source 7 (Line) Atlantic (All depths)				
Time Data Base (T) : 73 years Number of Recorded Events (N): 7 v' from log-linear fit: 7.75 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $v'' = v' + N = 7.75 + 7 = 14.75$ $\eta''_{M_i} = \eta'_{M_i} + N = 7.75 + 7 = 14.75$				
Richter Magnitude (M_i)	Nb of Recorded Occurrences in M_i bands (R_{M_i})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.13) (N_c)	Nb of Occurrences in M_i bands (log-linear fit) (E'_{M_i})	$E'_{M_i} + R_{M_i}$ (E''_{M_i})
4.0	0	7.75	.6	.6
4.25	0	7.15	.6	.6
4.50	0	6.55	.5	.5
4.75	1	6.05	.5	1.5
5.0	2	5.55	.45	2.45
5.25	0	5.1	.4	.4
5.50	1	4.7	.4	1.4
5.75	0	4.3	.3	.3
6.0	1	4.0	.3	1.3
6.25	1	3.7	.4	1.4
6.50	0	3.3	1.9	1.9
6.75	0	1.4	.86	.86
7.0	1	.54	.32	1.32
7.25		.22	.12	.12
7.50		.1	.1	.1
7.75				

TABLE 3.10
Seismic Source Parameters

Source 8 (Line) Pacific Coast Line (33 km)				
Time Data Base (T): 73 years Number of Recorded Events (N): 10 v' from log-linear fit: 9.2 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $v'' = v' + N = 9.2 + 10 = 19.2$ $\eta''_{M_i} = \eta'_{M_i} + N = 9.2 + 10 = 19.2$				
Richter Magnitude (M_i)	Nb of Recorded Occurrences in M_i bands (R_{M_i})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.14) (N_c)	Nb of Occurrences in M_i bands (log-linear fit) (E'_{M_i})	$E'_{M_i} + R_{M_i}$ (E''_{M_i})
4.0	1	9.2	.8	1.8
4.25	1	8.4	.7	1.7
4.50	1	7.7	.7	1.7
4.75	1	7.0	.6	1.6
5.0	4	6.4	.6	4.6
5.25	0	5.8	.5	.5
5.50	1	5.3	.5	1.5
5.75	0	4.8	.4	.4
6.0	0	4.4	.4	.4
6.25	1	4.0	3.1	4.1
6.50		.9	.8	.8
6.75		.1	.1	.1
7.0				
7.25				
7.50				
7.75				

TABLE 3.11
Seismic Source Parameters

Sources 9 & 10 (Lines) Line of Volcanoes (33 km)				
Time Data Base (T): 123 years Number of Recorded Events (N): 54 ν' from log-linear fit: 54 $\lambda'' = \lambda' + T = 123 + 123 = 246$ $\nu'' = \nu' + N = 54 + 54 = 108$ (for both sources together); $\nu'' = 108/2 = 54$ for each source. $\eta''_{M_1} = \eta'_{M_1} + N = 54 + 54 = 108$				
Richter Magnitude (M_1)	Nb of Recorded Occurrences in M_1 bands (R_{M_1})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.15) (N_c)	Nb of Occurrences in M_1 bands (log-linear fit) (E'_{M_1})	$E'_{M_1} + R_{M_1}$ (E''_{M_1})
4.0	1	54.0	3.0	4.0
4.25	0	51.0	3.0	3.0
4.50	2	48.0	3.0	5.0
4.75	1	45.0	2.5	3.5
5.0	1	42.5	2.5	3.5
5.25	2	40.0	2.4	4.4
5.50	11	37.6	2.1	13.1
5.75	13	35.5	2.0	15.0
6.0	7	33.5	2.0	9.0
6.25	2	31.5	2.0	4.0
6.50	5	29.5	9.5	14.5
6.75	5	20.0	14.8	19.8
7.0	3	5.2	3.7	6.7
7.25	1	1.5	1.05	2.05
7.50		.45	.35	.35
7.75		.1	.1	.1

TABLE 3.12
Seismic Source Parameters

Source 11 (Area) Managua Area (5 km)				
Time Data Base (T): 73 years Number of Recorded Events (N): 5 ν' from log-linear fit: 6.5 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $\nu'' = \nu' + N = 6.5 + 5 = 11.5$ $\eta''_{M_1} = \eta'_{M_1} + N = 6.5 + 5 = 11.5$				
Richter Magnitude (M_1)	Nb of Recorded Occurrences in M_1 bands (R_{M_1})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.16) (N_c)	Nb of Occurrences in M_1 bands (log-linear fit) (E'_{M_1})	$E'_{M_1} + R_{M_1}$ (E''_{M_1})
4.0	1	6.5	1.1	2.1
4.25	0	5.4	.9	.9
4.50	0	4.5	.8	.8
4.75	1	3.7	.6	1.6
5.0	1	3.1	.5	1.5
5.25	0	2.6	.48	.48
5.50	1	2.12	.37	1.37
5.75	0	1.75	.27	.27
6.0	0	1.48	.27	.27
6.25	1	1.21	.41	1.41
6.50		.8	.8	.8
6.75				
7.0				
7.25				
7.50				
7.75				

TABLE 3.13
Seismic Source Parameters

Source 12. (Area) Gulf of Fonseca (33 km)				
Time Data Base (T) : 73 years Number of Recorded Events (N) : 7 ν' from log-linear fit : 7.7 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $\nu'' = \nu' + N = 7.7 + 7 = 14.7$ $\eta''_{M_i} = \eta'_{M_i} + N = 7.7 + 7 = 14.7$				
Richter Magnitude (M_i)	Nb of Recorded Occurrences in M_i bands (R_{M_i})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.17) (N_c)	Nb of Occurrences in M_i bands (log-linear fit) (ξ'_{M_i})	$\xi'_{M_i} + R_{M_i}$ (ξ''_{M_i})
4.0	0	7.70	.15	.15
4.25	0	7.55	.15	.15
4.50	0	7.4	.1	.1
4.75	0	7.3	.1	.1
5.0	0	7.2	.1	.1
5.25	0	7.1	.1	.1
5.50	4	7.0	.1	4.1
5.75	1	6.9	.1	1.1
6.0	1	6.8	4.8	5.8
6.25	1	2.0	1.9	2.9
6.50		.1	.1	.1
6.75				
7.0				
7.25				
7.50				
7.75				

TABLE 3.14
Seismic Source Parameters

Source 13 (Area) Costa Rica Area (80 - 109 km)				
Time Data Base (T) : 73 Years Number of Recorded Events (N) : 4 ν' from log-linear fit: 5.4 $\lambda'' = \lambda' + T = 73 + 73 = 146$ $\nu'' = \nu' + N = 5.4 + 4 = 9.4$ $\eta''_{M_i} = \eta'_{M_i} + N = 5.4 + 4 = 9.4$				
Richter Magnitude (M_i)	Nb of Recorded Occurrences in M_i bands (R_{M_i})	Cumulative Nb of Occurrences (log-linear fit Fig. 3.18) (N_c)	Nb of Occurrences in M_i bands (log-linear fit) (ξ'_{M_i})	$\xi'_{M_i} + R_{M_i}$ (ξ''_{M_i})
4.0	1	5.4	.8	1.8
4.25	0	4.6	.6	.6
4.50	1	4.0	.5	1.5
4.75	0	3.5	.5	.5
5.0	0	3.0	.4	.4
5.25	0	2.6	.4	.4
5.50	1	2.2	.3	1.3
5.75	0	1.9	.25	.25
6.0	0	1.65	.25	.25
6.25	0	1.4	.2	.2
6.50	0	1.2	.15	.15
6.75	0	1.05	.15	.15
7.0	1	.9	.5	1.5
7.25		.4	.3	.3
7.50		.1	.1	.1
7.75				

Poisson Model

The generating process for the number of occurrences is the Poisson model with uncertain mean rate of occurrence λ (eq. 2.1). The parameter λ is treated as a random variable and Bayesian statistics is applied to it.

The sample likelihood function on λ (eq. 2.4) is derived from the generating Poisson model. The parameters T and N of the sample likelihood function are determined from the available data. T represents the time base for which the data is available: 73 years for all sources except for the two volcano lines for which it is 123 years. N represents the total number of occurrences observed on the source considered during this time period.

The gamma prior distribution on λ (eq. 2.3) is based on the subjective input of the analyst. The numerical values of the parameters λ' and ν' are obtained from this subjective input. For this study, it is assumed that the values of λ' and ν' are respectively equal to T and N of the corresponding source. The implication of this assumption is that the subjective information of the expert is similar to the available data. In other words, the analyst has as much confidence in his subjective input as he has in the data.

Based on the values of λ' , ν' , T and N, the parameters λ'' and ν'' of the posterior distribution on λ (eq. 2.7) can be calculated for each source. It should be pointed out that in the absence of any subjective information (diffuse prior), the analysis can be carried out with objective data alone and in the absence of any objective data, the analysis can be

carried out with subjective information alone. Knowledge of λ'' and ν'' completely defines -- in a posterior sense -- the probability function for the mean rate of occurrence λ for the source considered during a future time t .

Convolving the conditional Poisson generating process for the number of occurrences with the posterior distribution on λ , the marginal distribution for the number of occurrences (eq. 2.9) is derived for each source considered. This distribution does not give any information on the magnitude of the occurrences. The next step is to obtain the posterior conditional distribution on magnitudes.

Bernoulli Trials

The generating process for the number r_{M_i} of events of any specific M_i given that a total of n events have occurred is the binomial model. However, the probability of success p_{M_i} for each trial has been assumed to be uncertain. The parameter p_{M_i} is treated as a random variable and Bayesian statistics is applied to it.

The sample likelihood function on p_{M_i} (eq. 2.15) is derived from the generating binomial process. From the available data, the parameters N and R_{M_i} of the sample likelihood function can be determined. N represents the total number of events recorded on the source under consideration and R_{M_i} represents the number of earthquakes of magnitude M_i (successes) recorded on the same source. R_{M_i} must be determined for each source and each M_i .

Using the conjugate beta prior (eq. 2.13) for the distribution on p_{M_i} , the parameters η'_{M_i} and ξ'_{M_i} are determined from the analyst's

subjective input. For this case study, it is assumed that the analytical recurrence relationship fitted to the data for each source constitutes the subjective input. For each individual source, the analytical relationship describing the recurrence of various M_i events is given by the following log-linear relationship

$$\ln N(M) = \alpha + \beta M$$

where $N(M)$ = Number of events above magnitude M

M = Richter magnitude

α and β are regression constants.

Figures 3.7 through 3.18 (Figures 3.2 through 3.13 from Shah, et al., 1975) show these recurrence relationships for all the seismic sources considered. Table 3.15 gives the summary of α and β values for each source and the cutoff point corresponding to $\ln N(M) = 0.1$. It can be seen from Table 3.15 as well as from Figures 3.7 through 3.18 that most of the analytical recurrence relationships fitted to the data are bi-linear.

The prior η'_{M_i} represents the subjective knowledge about the number of events for a source above the fixed lower bound ($M_i = 4.0$). As an example, consider the source 1. From Figure 3.7 the η'_{M_i} corresponding to this source is 72. ξ'_{M_i} represents the number of earthquakes of magnitude M_i . Again from Figure 3.7, for $M = 3.875$, $N_c = 72$ and for $M = 4.125$, $N_c = 54$ thus, for $M_i = 4.0$, ξ'_{M_i} is equal to $72 - 54 = 18$. Because of the definition of the prior, within each source, η'_{M_i} is constant for

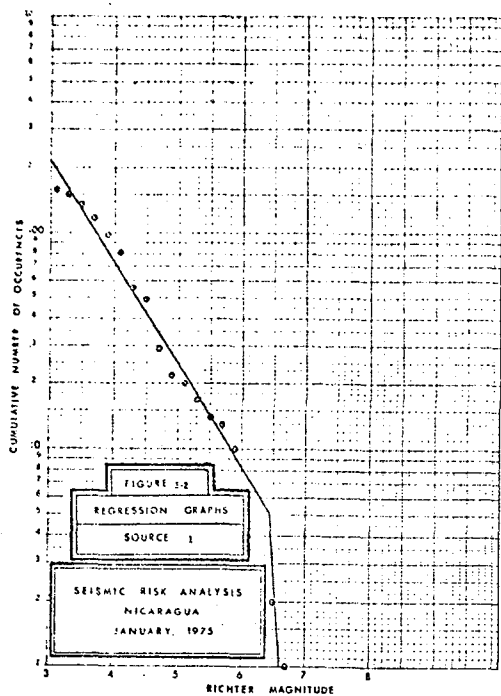


FIGURE 3.7

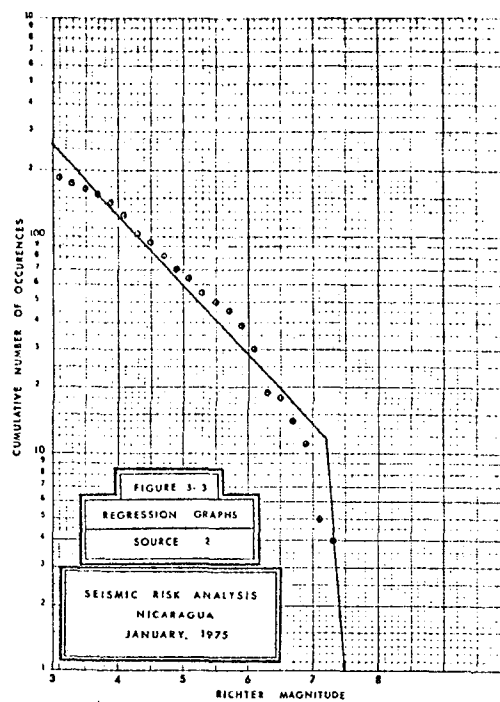


FIGURE 3.8

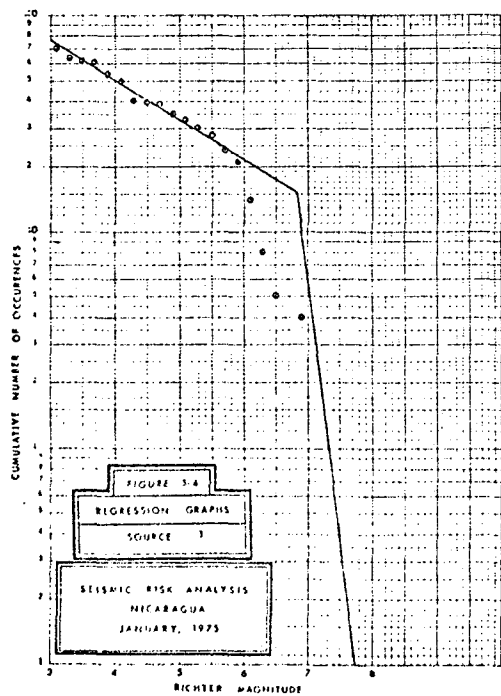


FIGURE 3.9

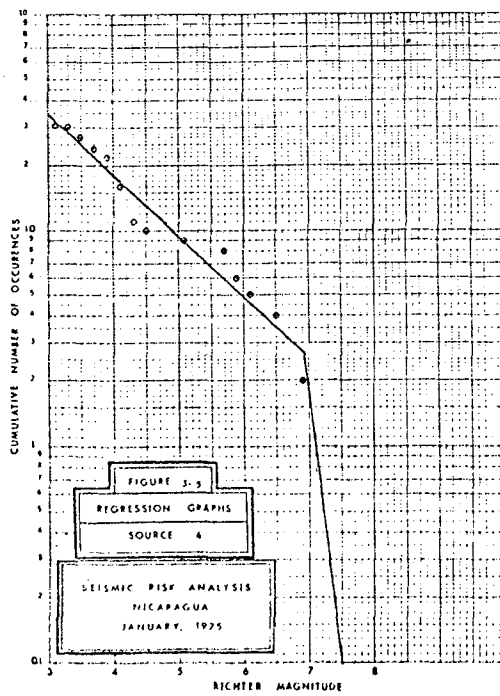


FIGURE 3.10

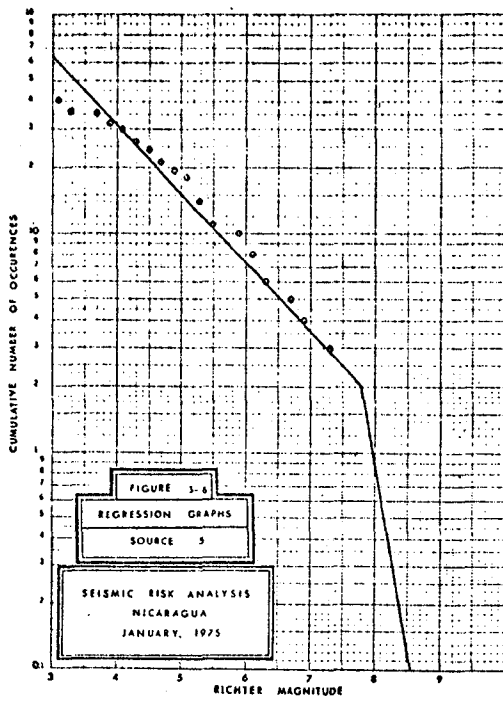


FIGURE 3.11

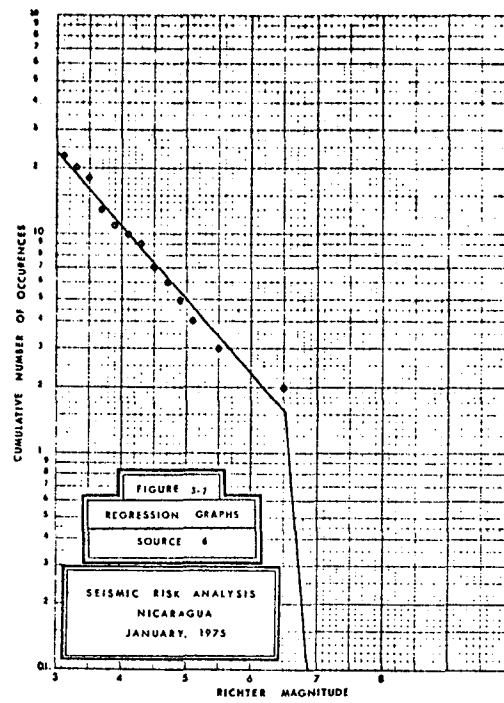


FIGURE 3.12

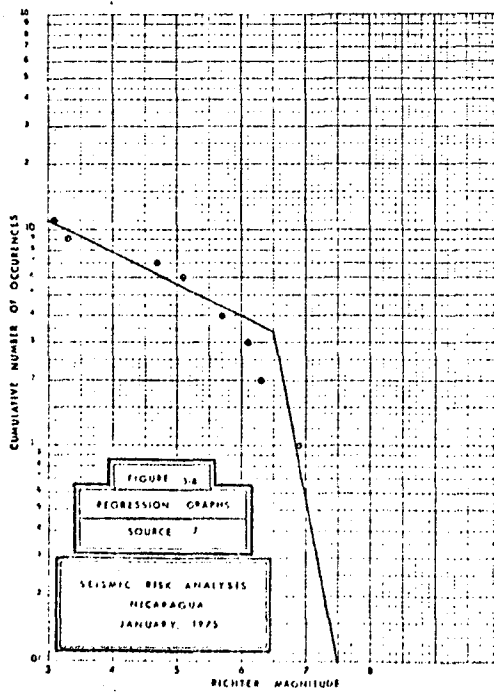


FIGURE 3.13

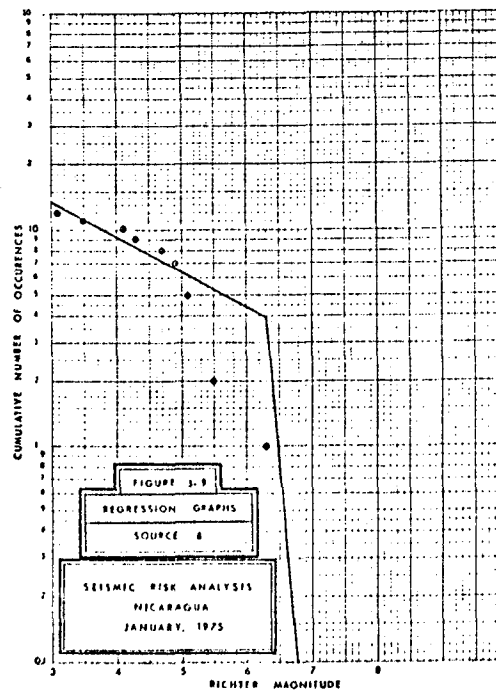


FIGURE 3.14

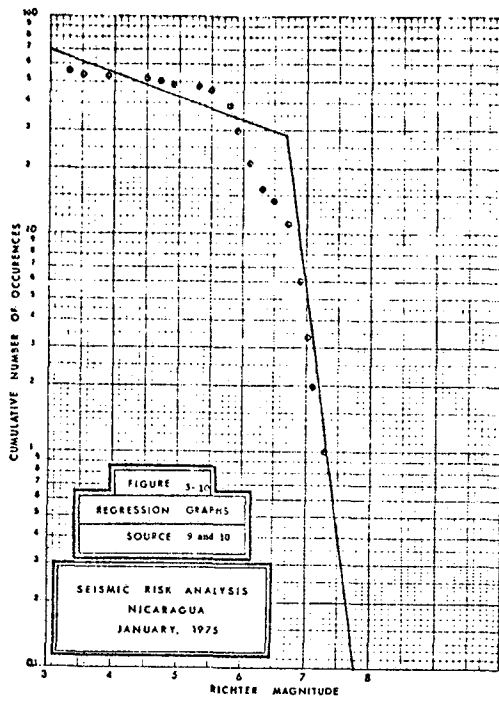


FIGURE 3.15

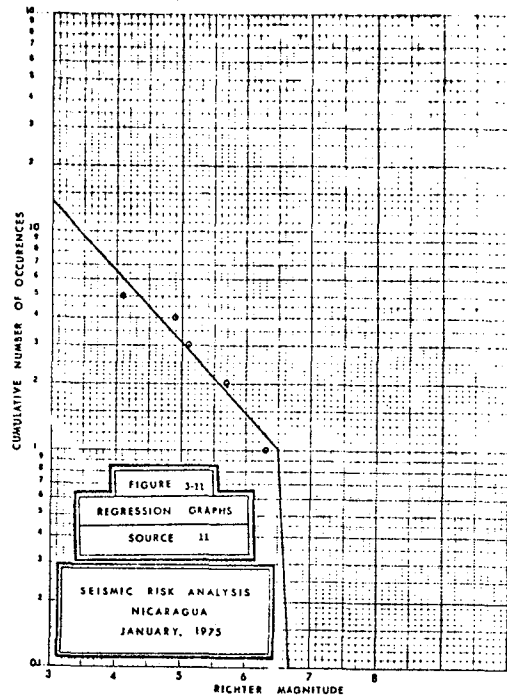


FIGURE 3.16

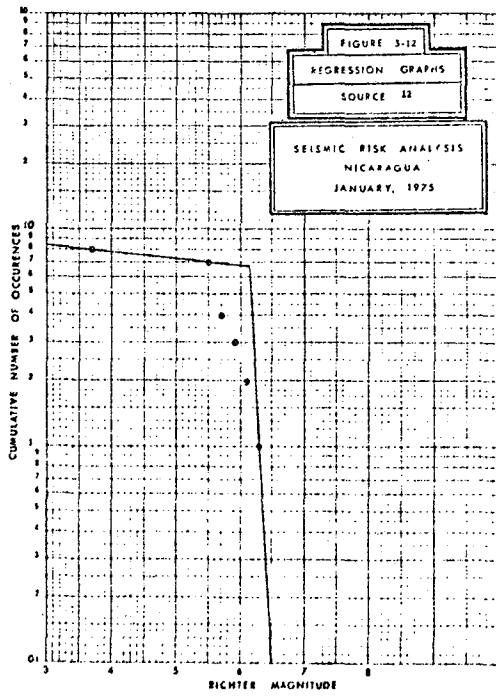


FIGURE 3.17

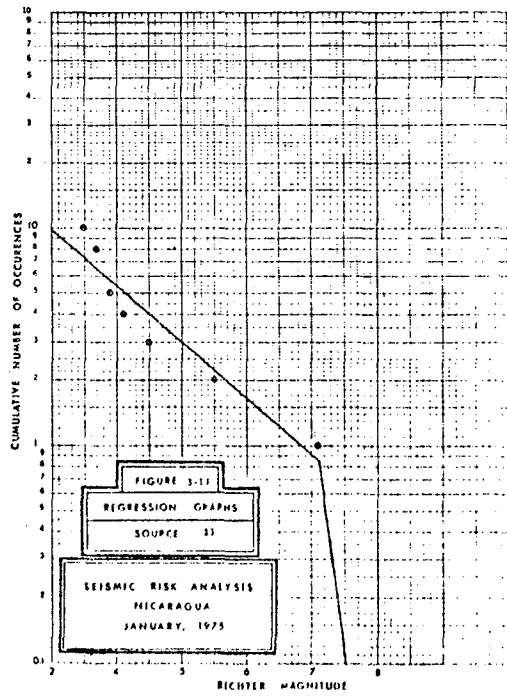


FIGURE 3.18

all M_i 's. If the prior had been input differently such as in the form of a distribution for each M_i , different η_{M_i}' could have been obtained. The condition

$$\sum_{\text{all } M_i} \frac{\xi_{M_i}'}{\eta_{M_i}'} = 1.0 \quad (2.14 \text{ repeated})$$

obviously satisfied in this case, should then have been checked and satisfied by normalizing these ratios.

Having determined the parameters of the sample likelihood function as well as the ones of the prior distribution, the posterior parameters η_{M_i}'' and ξ_{M_i}'' (eq. 2.18) can be obtained by using the concept of conjugate distribution. The knowledge of η_{M_i}'' and ξ_{M_i}'' completely defines -- in the posterior sense -- the probability distribution of the probability of success p_{M_i} of magnitude M_i on the source considered.

The marginal distribution on the number of successes M_i 's (eq. 2.20) is obtained by convolving the posterior distribution on p_{M_i} and the conditional generating process of r_{M_i} . However, this marginal distribution is still conditional on the number of events n .

Combining the distribution of r_{M_i} for given n (eq. 2.20), with the distribution on n (eq. 2.9) gives the marginal Bayesian distribution on r_{M_i} (eq. 2.21). This information completes the description of seismicity for a given source.

To obtain the probabilistic information on the peak ground acceleration and duration at the site, the above information on the

Table 3.15
Parameters of Recurrence Relationships

Source	α_1	β_1	α_2	β_2	Cutoff
1.	8.66	-1.09	30.06	-4.55	6.8
2.	7.78	-0.74	69.08	-9.21	7.8
3.	5.59	-0.42	42.02	-5.75	7.7
4.	5.49	-0.65	32.34	-4.55	7.5
5.	6.31	-0.72	42.16	-5.27	8.5
6.	5.47	-0.77	51.55	-7.82	6.9
7.	3.40	-0.33	24.15	-3.53	7.5
8.	3.73	-0.37	47.68	-7.57	6.8
9.	4.99	-0.24	39.65	-5.43	7.8
10.	4.99	-0.24	39.65	-5.43	7.8
11.	4.81	-0.74	80.79	-12.40	6.7
12.	2.35	-0.07	82.15	-13.04	6.5
13.	4.05	-0.59	39.27	-5.54	7.5

seismicity of various sources must be combined with the attenuation relationships presented in Chapter II.

As mentioned in that chapter, a uniform distribution with a constant coefficient of variation is used to describe the scatter of the attenuation data with respect to the mean.

Figure 3.19 shows the cumulative distribution function for the peak ground acceleration for Managua. The future time period considered is 50 years. The assumed coefficients of variation on the attenuation relationship are 0.0, 0.3 and 0.5. The dotted line indicates the results from a previous study (Shah, et al., 1975).

Figure 3.20 shows the cumulative distribution function for the duration for Managua. The time period considered is 50 years. The assumed coefficients of variation on the attenuation relationship are 0.0, 0.3 and 0.5. No similar results are available in the literature for comparison.

Comparing the cumulative distribution function from this study with the one obtained in Shah, et al., (1975)(Figure 3.19), the following observations can be made.

- With a zero coefficient of variation, the Bayesian approach leads to smaller accelerations. This can be explained by the following: the combination of the Poisson-gamma and binomial-beta models produces smaller probabilities than the Poisson model only. This is noticed in simple test runs in which the data is as similar as the difference in input format permits. Moreover, the data used in the Bayesian approach is less conservative than the one used in the

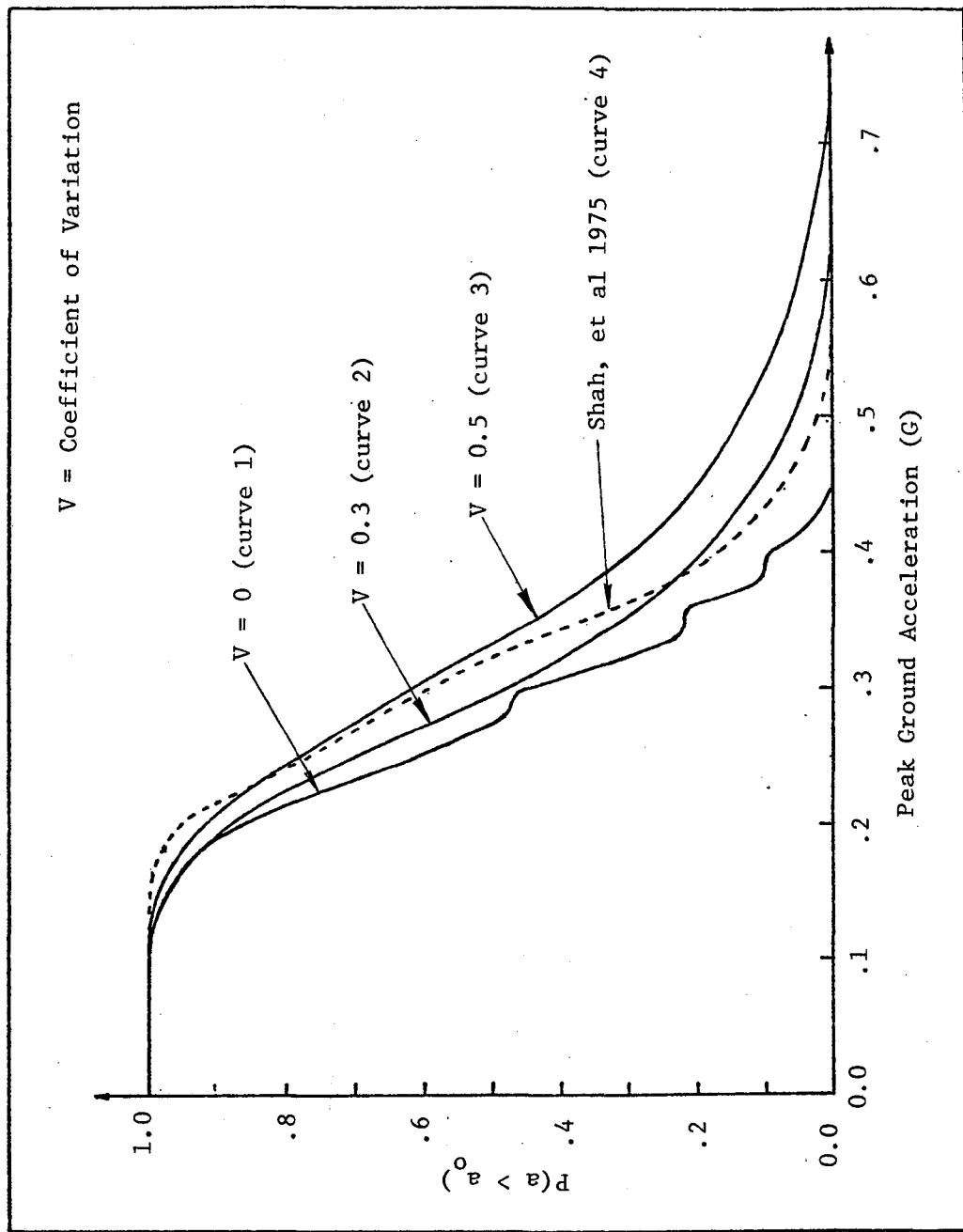
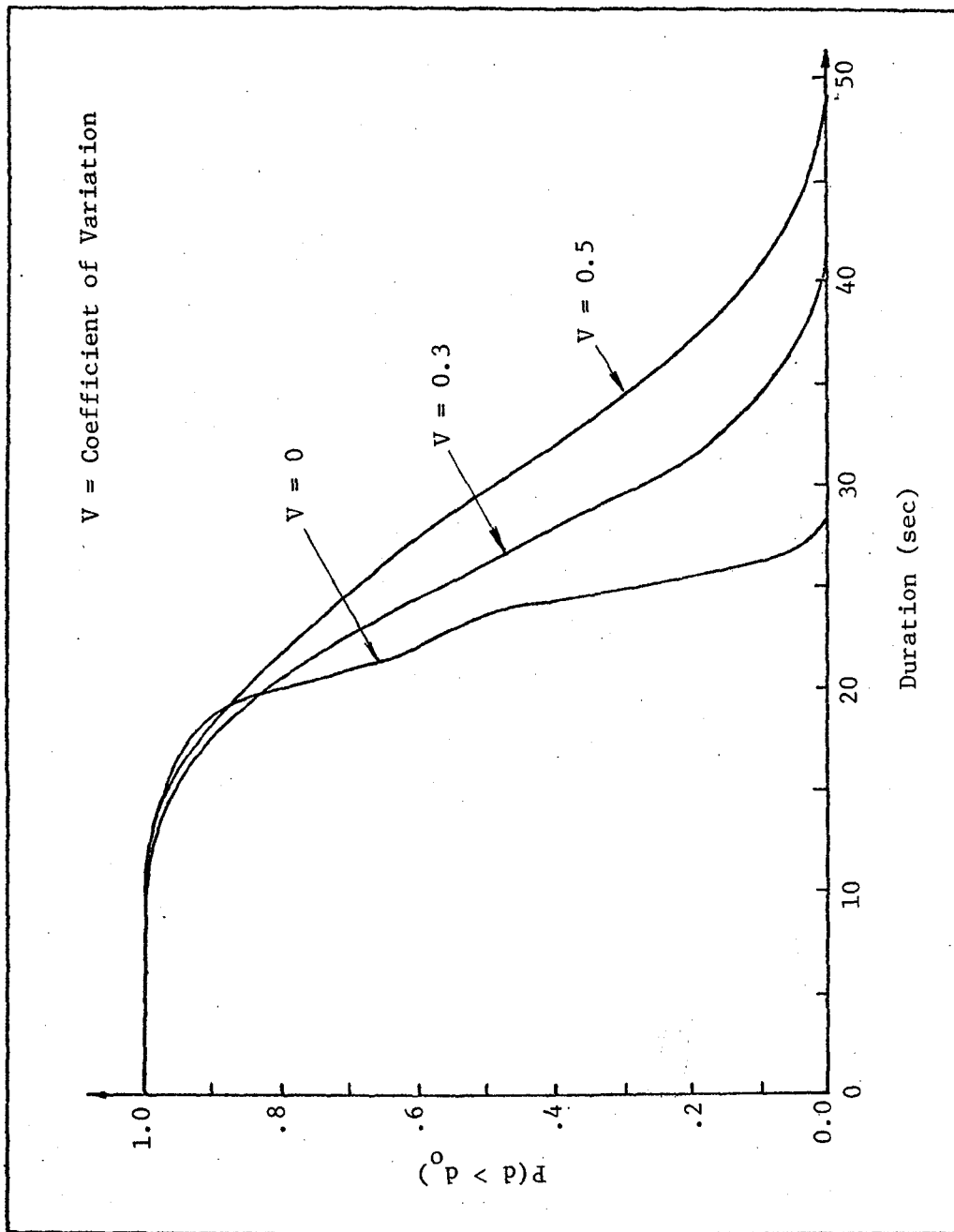


FIGURE 3.19. Probability of Peak Ground Acceleration $P(a > a_0)$

FIGURE 3.20. Probability of Duration $P(d > d_0)$

Poisson model. The log-linear recurrence relationships are the sole data to the Poisson model. In the present model the log-linear recurrences are combined with the available data through Bayesian statistics. This introduces a decrease in seismicity as explained below. For large M 's where insufficient information is available to obtain a good fit, it is common to adopt a conservative approach. Hence for those magnitudes, that are governing the results, the log-linear recurrence relationship often implies a seismicity greater than the one shown by the data itself. (Figures 3.7 through 3.18). Therefore the combination of the data and the log-linear fit results in seismicity smaller than the one described by the recurrence relationships only.

- Curve 1 is not as smooth as curve 4. Both models are very sensitive to the value of the upper cutoff on each source. In the present model it is input as an abrupt cutoff, whereas in the Poisson model it is input as a sharp but continuous decrease which for small probabilities extends beyond the cutoff value. Therefore the switch of the governing influence from one source to another as the PGA increases does not appear as sharply. This produces a smoother CDF. Moreover, the treatment of the acceleration as a discrete variable (curve 1) must of necessity introduce some round off approximations.
- Increasing the coefficient of variation of the attenuation relationship increases the probability of exceeding a given level of peak ground acceleration. This is expected. A coefficient of variation

of zero implies that the attenuation relationship has no scatter. As the coefficient of variation increases, the uncertainty associated with the attenuation relationship increases. This results in larger accelerations for the same M and distance.

- The probability of exceeding a given PGA approaches the results presented in Shah, et al. (1975) (Curve 4) as the coefficient of variation increases from 0.0 to 0.3. The versatility of the present model and its superiority over the currently available models is obvious.
- Similar behavior regarding duration is observed in Figure 3.20 as the coefficient of variation in duration attenuation increases, the probability of exceedence for a given duration also increases. No comparison can be made about this behavior since similar results are not available in the literature.

Figures 3.21 and 3.22 emphasize the behavior of the PGA and duration as a function of the coefficient of variation in PGA and duration attenuations: larger uncertainty in the attenuation relationships results in higher values of PGA or duration.

In developing some iso-acceleration maps for Nicaragua, the following cases are considered.

1. Economic life of 30 years. 10% chance of exceedence.
Coefficients of variation of 0.0 and 0.3.
2. Economic life of 50 years. 10% chance of exceedence.
Coefficients of variation of 0.0 and 0.3.

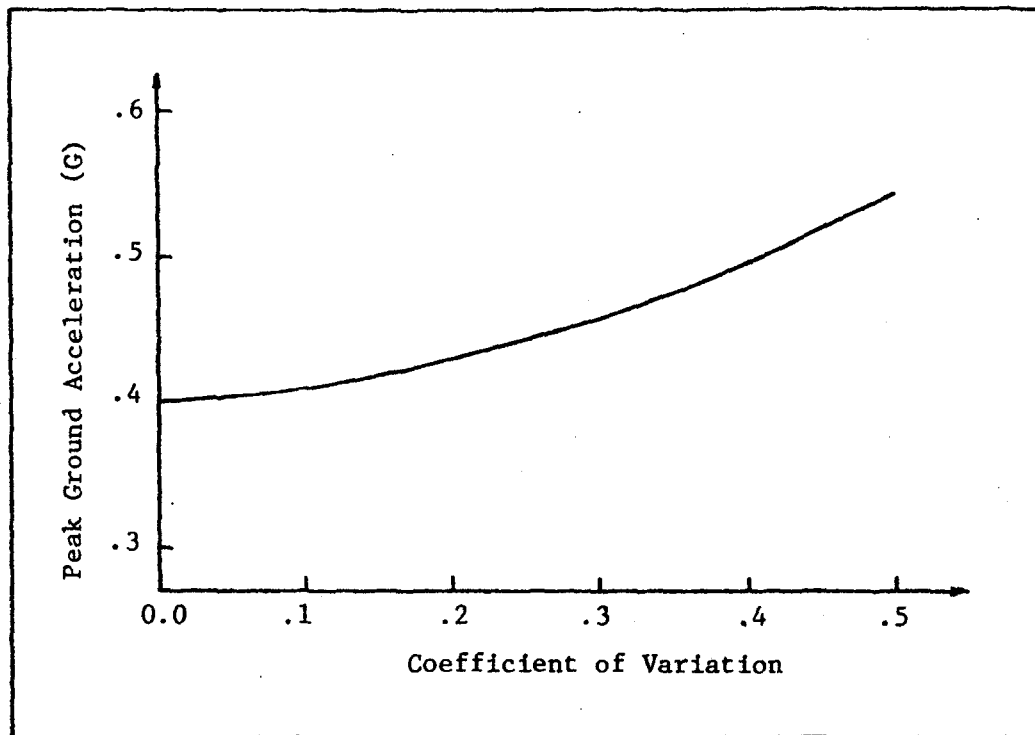


FIGURE 3.21. Peak Ground Acceleration $P(a > a_0) = 0.10$

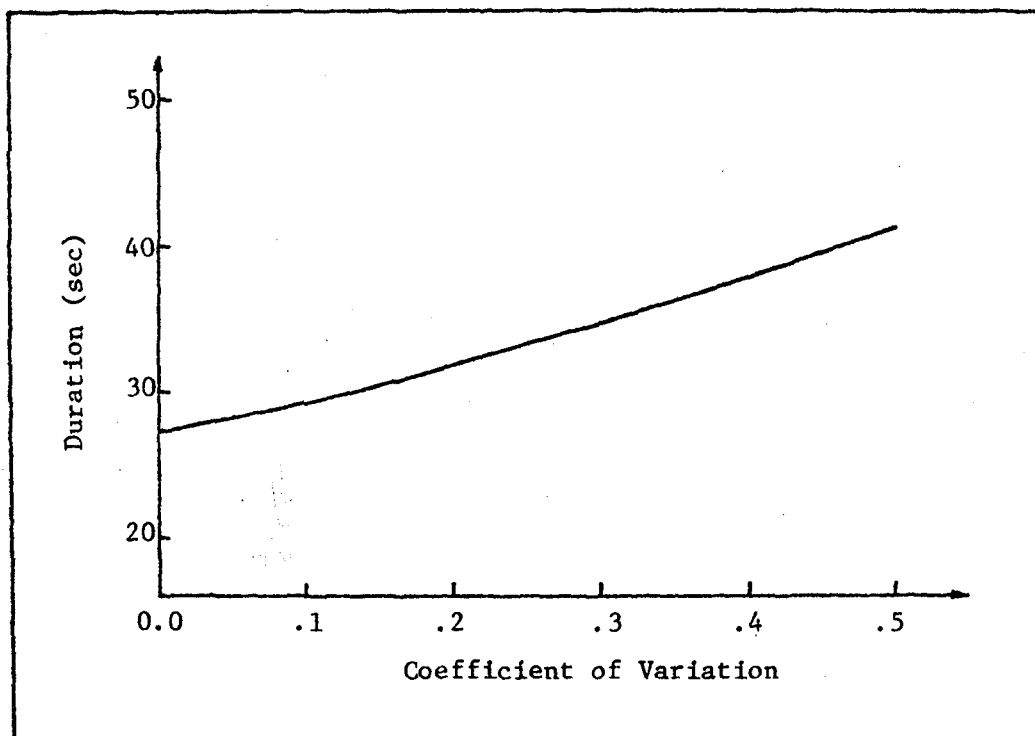


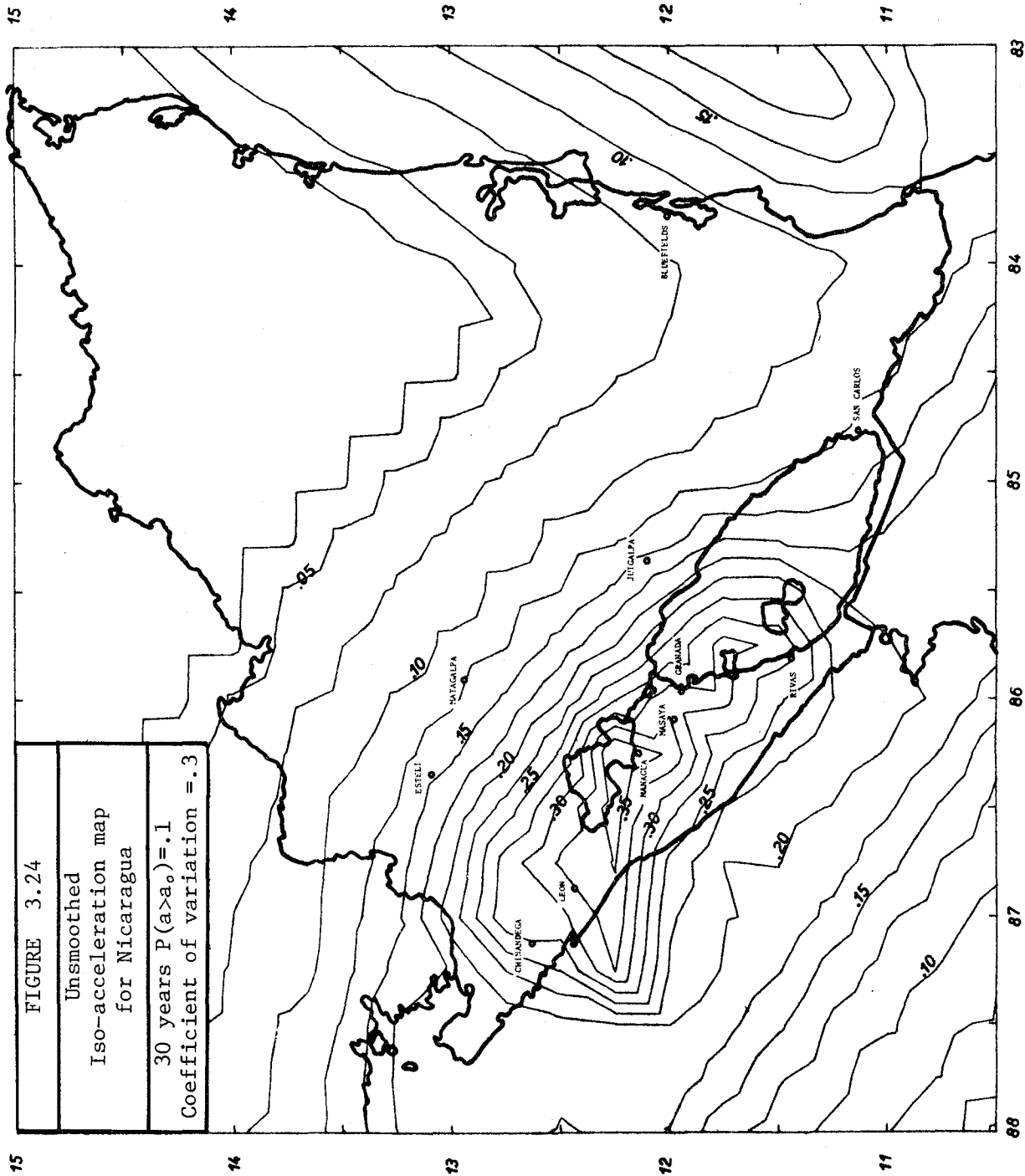
FIGURE 3.22. Duration $P(d > d_0) = 0.10$

Thus four iso-acceleration maps (Fig. 3.23 through 3.26) and four iso-duration maps (Fig. 3.27 through 3.30) are developed. The maps are drawn for return periods corresponding to 475 years. Comparing the iso-acceleration map of this study having a coefficient of variation of 0.3 and 475 years return period with the map developed in Shah, et al., (1975). The following comments can be made.

- For the same data base, the shapes of the iso-acceleration lines are very similar.
- The peak ground accelerations obtained from this study (for a coefficient of variation of 0.3) are slightly smaller.
- The iso-duration maps developed here are unique and no comparable results are available in the literature. However the values presented here look reasonable and in good agreement with past events.
- For a given region and a given return period the peak increases with the increase in coefficient of variation of the attenuation relationship.

3.4 Conclusions

In this chapter, a method of using the available data and the subjective information to obtain seismic hazard for a region is presented. It is shown that any geological information, no matter how subjective, can be incorporated in describing the seismicity of a fault or a source. Once such an information is incorporated in the hazard mapping methodology, any objection to the use of historical data alone can be removed. The purpose of presenting the Nicaragua seismic hazard map is not only to explain the Bayesian methodology but



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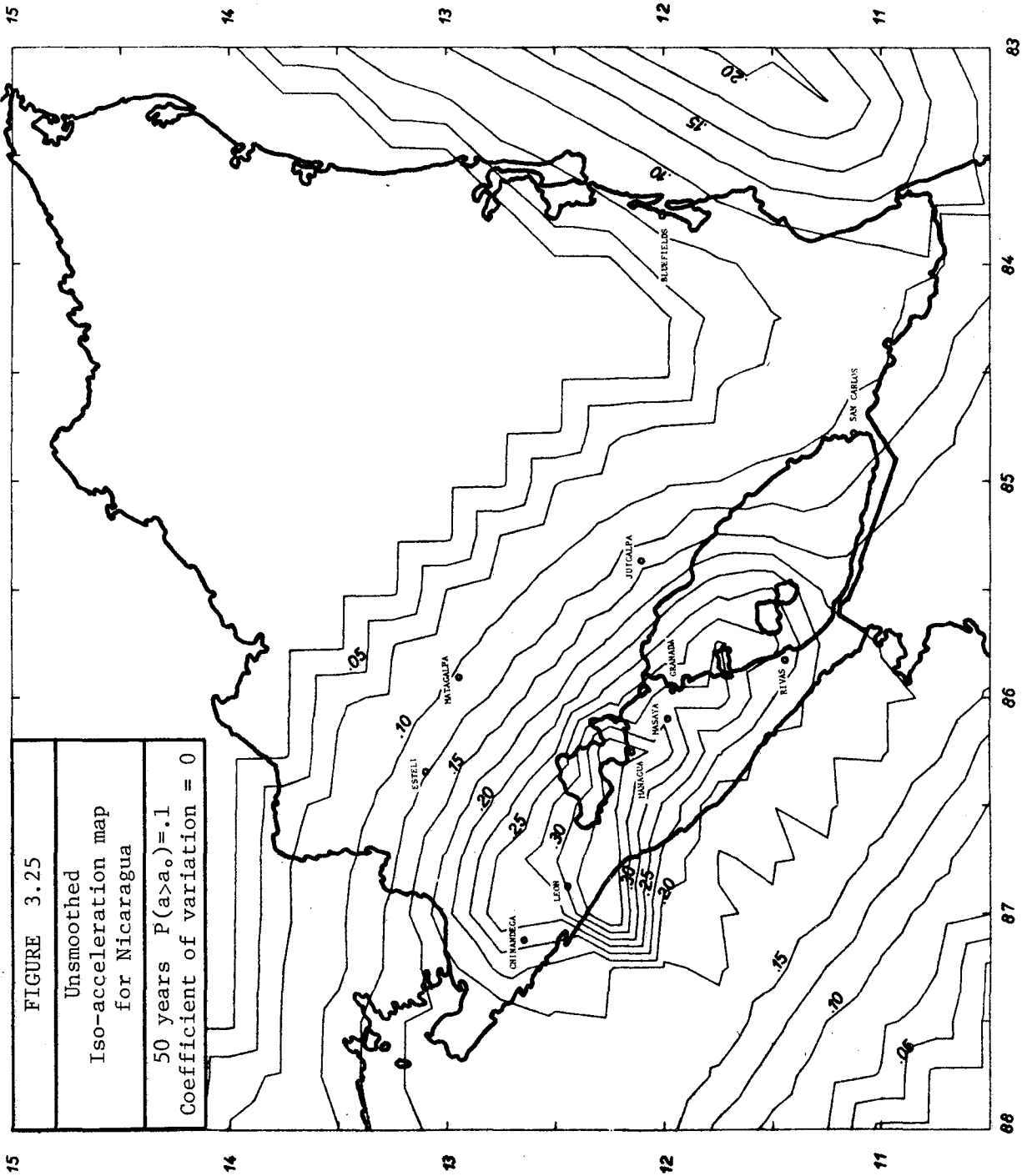
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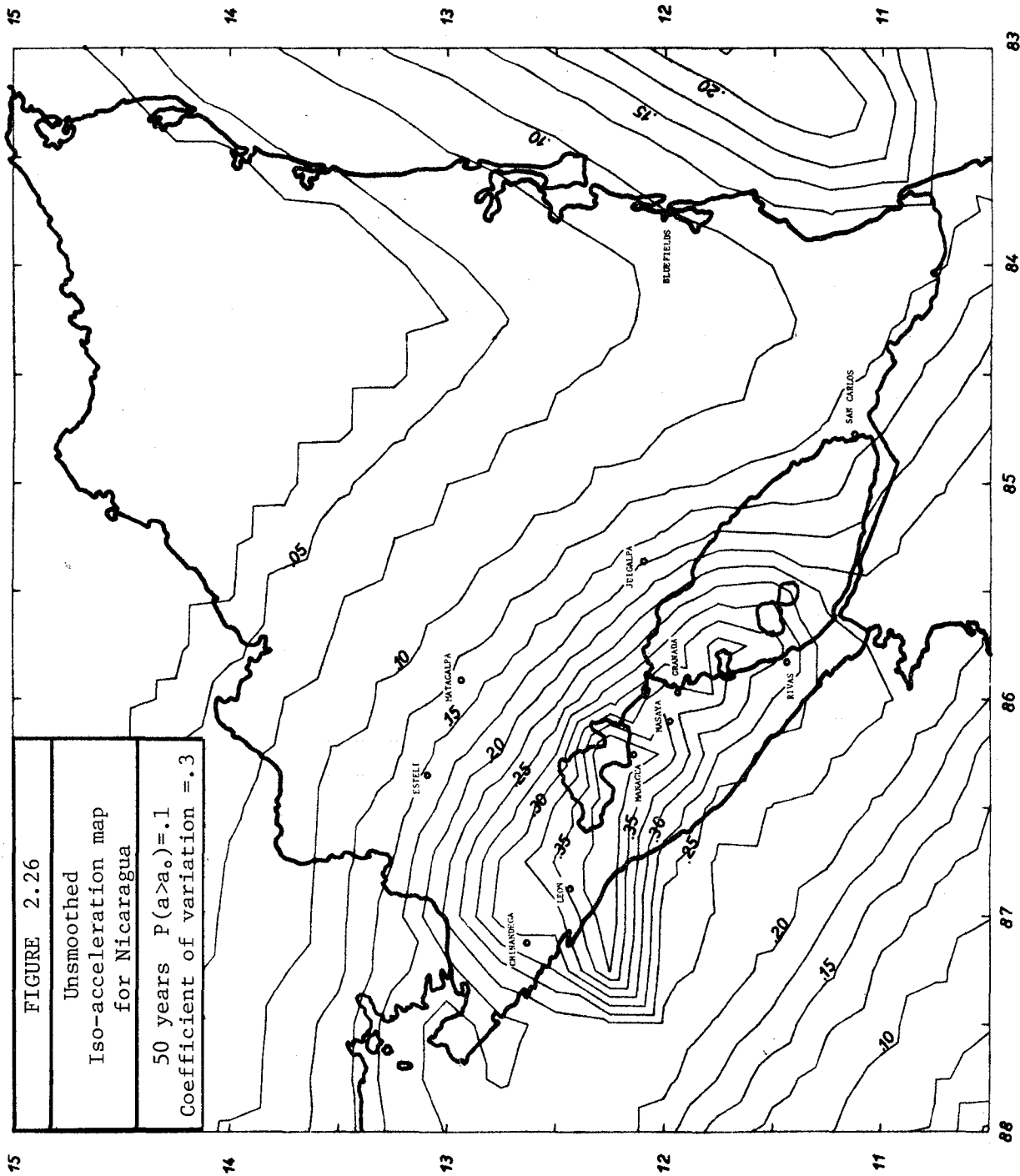
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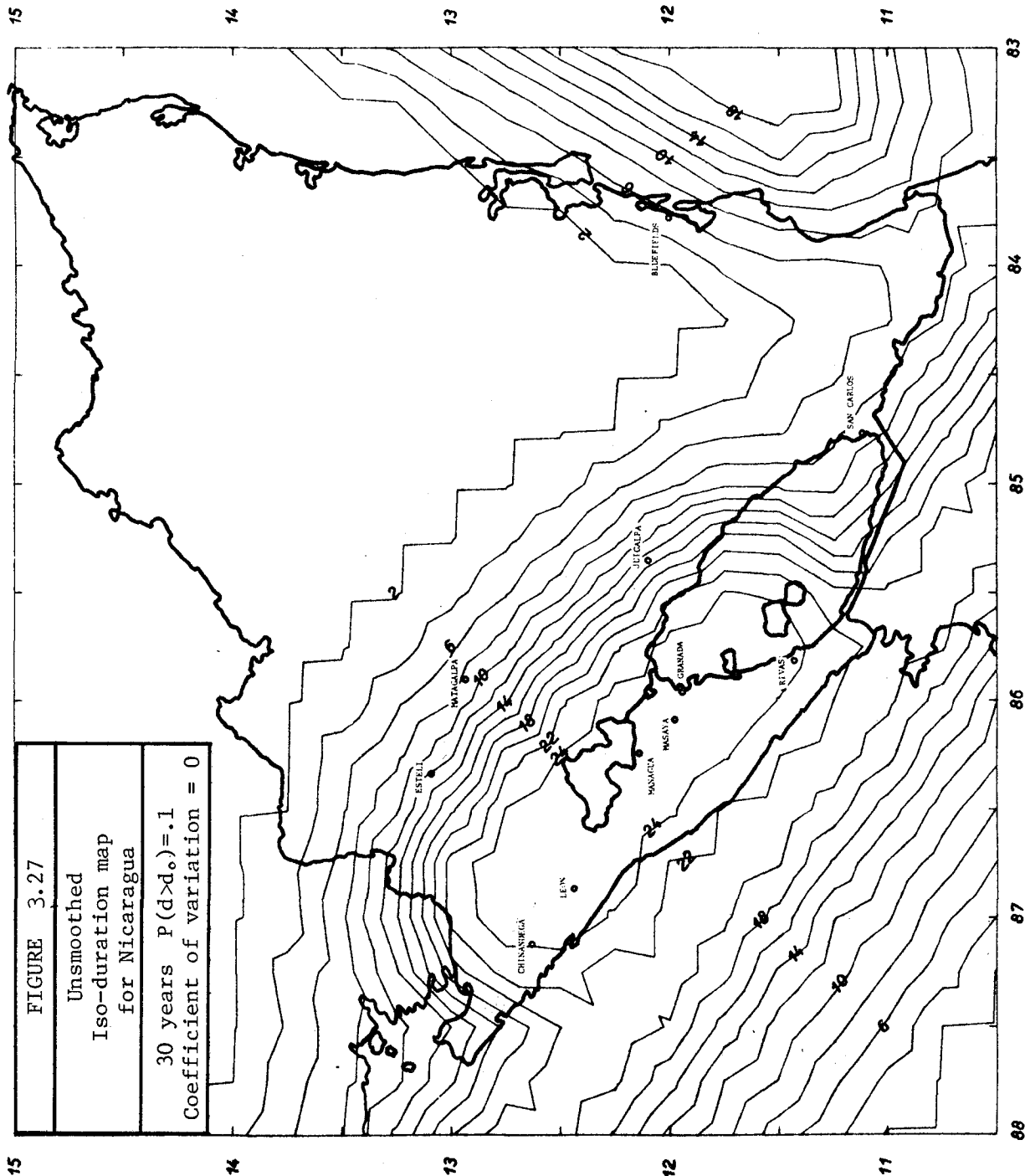
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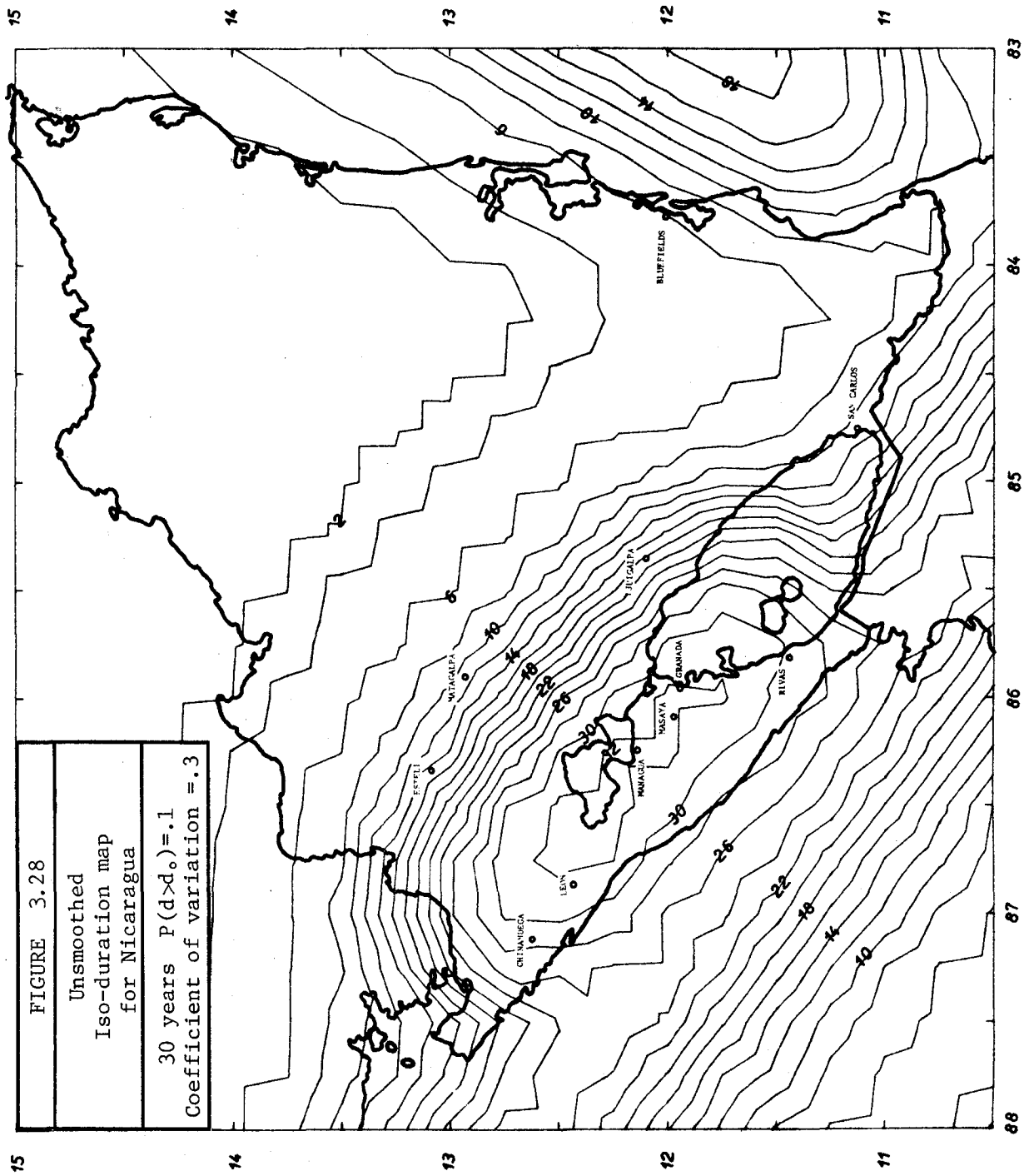
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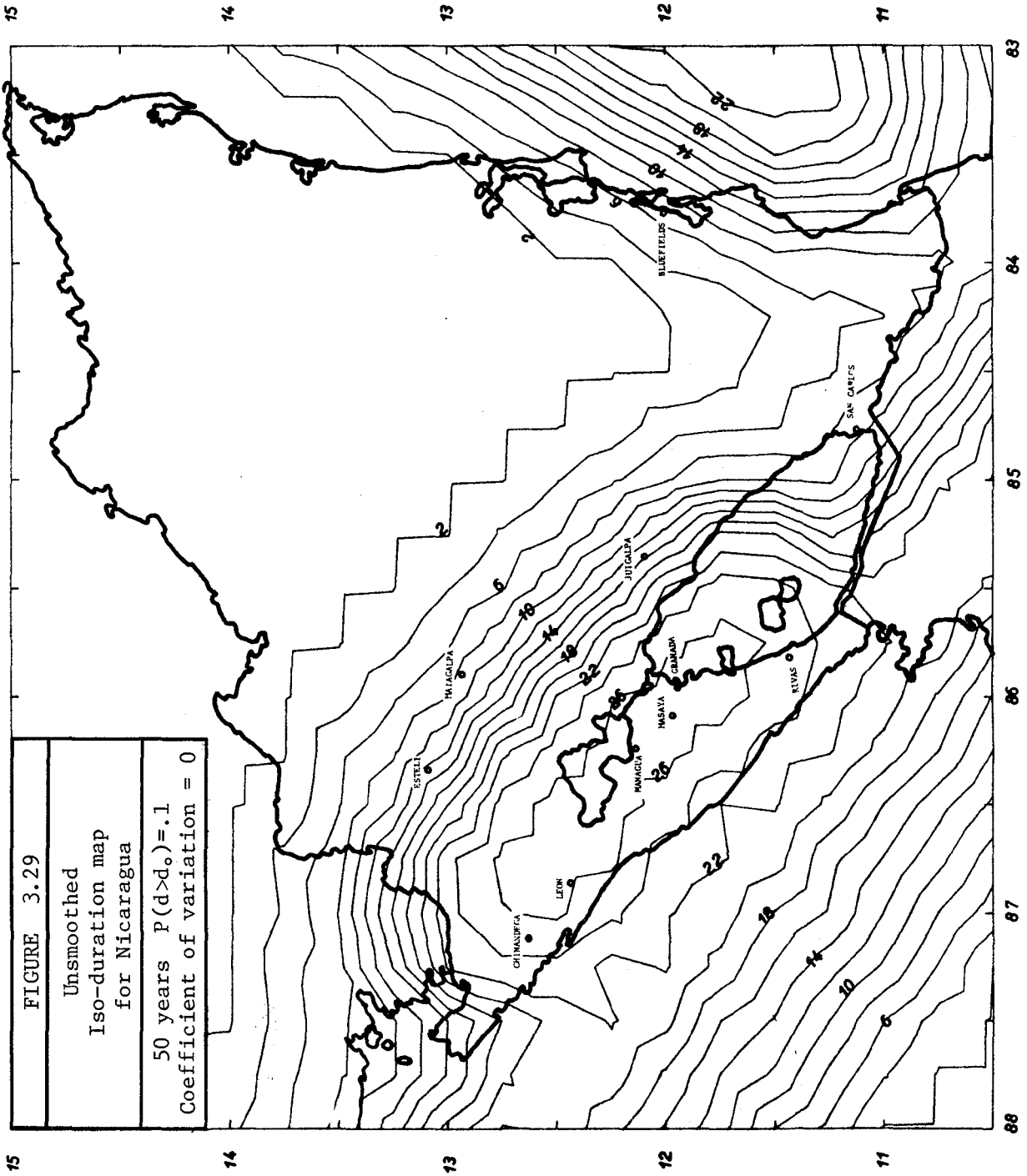
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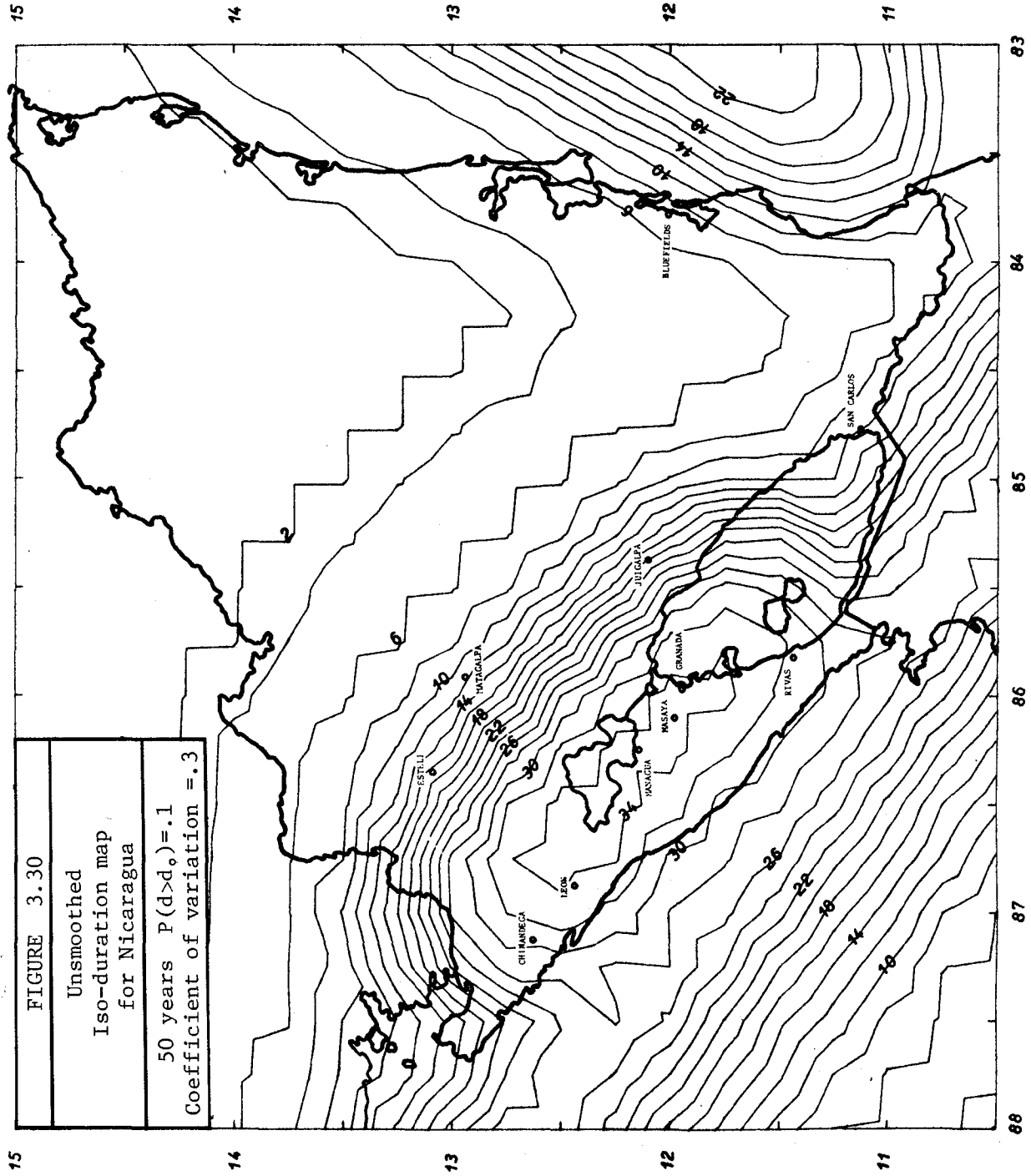
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also to emphasize the following points;

- The simplicity of incorporating any subjective information with the historical data to obtain Bayesian estimation on fault seismicity.
- The versatility of the method in which various uncertainties can be incorporated in obtaining the final probabilistic hazard information. In the example case, uncertainty in the attenuation equations was included through the coefficient of variation approach.
- The advantage of explicit probabilistic information on duration. This should be of great interest to engineers who wish to estimate the damage potential of a given event and also to researchers who are trying to develop seismic intensity parameters based on input energy, duration, amplitude, etc.
- The efficiency of the algorithm developed for mapping. Total information (CDF) for acceleration and duration was obtained for the 399 nodes of the grid. On a computer IBM 370/168, the execution time varied between 20 to 35 CPU seconds as a function of the value of the coefficient of variation in the attenuation relationships. (0.0 or 0.3).

It is hoped that the Bayesian methodology presented here will be applied to regions where there is very little historical data and where some subjective information on the local seismicity is available.



CHAPTER IV
ANALYSIS OF EARTHQUAKE RECORDS

4.1 Introduction

In Chapter I, Section 1.4.2 it is pointed out that an in depth study of earthquake records (accelerograms) involves time and frequency domain analyses. In the current analysis and design procedures, the time domain information is usually in the form of peak values (such as peak acceleration, velocity and displacement) and duration. The frequency domain information most widely used by engineers is in the form of response spectra (Newmark and Valestos, 1964). It is felt that such information is necessary but not sufficient and global for understanding earthquake ground motion. To develop a better understanding of currently available and digitized accelerograms, 97 records are analyzed in this study. As mentioned in Chapter I, the analysis involves the following parameters:

(A) Input parameters obtained from recorded accelerograms

- Duration (defined later in Section 4.2)
- Mean acceleration
- Root mean square (RMS) of acceleration
- Parameter a_p defined as the acceleration which has probability p of being exceeded for a given record.

bility p of being exceeded for a given record.

- The ratio $K_1 = a_p / \text{RMS}$

(B) Response parameters obtained from acceleration response histories of a linear single degree of freedom system with damping ratios of 5%, 10% and 20% and natural period range of 0.08 sec to 5.0 sec.

- Duration (defined later in Section 4.4)

- Mean acceleration
- Peak response (corresponds to spectral value)
- RMS of acceleration
- Probability distribution of response peaks
- Parameter a_p defined as the response acceleration which has probability p of being exceeded for the given response.

- The ratio $K_1 = a_p/\text{RMS}$
- Cumulative potential energy per unit mass (ENGY) in the response time history.

- The ratio $K_2 = \text{RMS}^2 \cdot T^2 / (\text{ENGY}/\text{NBPK})$, where T is the period of the structure and NBPK is the number of peaks of the response time history.

4.2 Description of the Accelerograms Used

It is assumed herein that the digitized and corrected records of acceleration represent a true and sufficient input. These records are obtained on magnetic tape from the Earthquake Engineering Research Laboratory, California Institute of Technology. They correspond to the Volume II Report Series (Hudson, et al., 1971-1974). They are digitized at equal time increments of 0.02 seconds.

A set of 49 earthquakes is considered in this study (Table 4.1). No scaling of the amplitude is used. However, since duration is recognized to be of importance, some convention is necessary to define the duration of an earthquake record. Surprisingly, no standard (and acceptable) definition of duration exists in the literature. Whatever few references that are available define the duration qualitatively

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#	RECORD NAME	CUTOFF (G)	DURATION (SEC)
A006	HOLLYWOOD 1952 BAS. SOUTH	.000	42.460
		.020	36.420
		.040	19.100
A007	HOLLYWOOD 1952 BAS. EAST	.060	.000
		.000	42.520
		.020	39.640
A008	HOLLYWOOD 1952 PE. SOUTH	.040	13.060
		.060	.000
		.000	76.660
A009	HOLLYWOOD 1952 PE. EAST	.020	49.560
		.040	16.720
		.060	.000
A010	EUREKA 1954 N11W	.000	70.000
		.020	26.000
		.040	8.860
A011	EUREKA 1954 N79E	.060	6.720
		.000	79.560
		.020	29.920
A012	FERNDALE 1954 N46W	.040	11.040
		.060	9.260
		.000	42.540
A013	FERNDALE 1954 N46W	.020	39.140
		.040	18.520
		.060	17.540
A014	SAN JOSE 1955 N31W	.000	42.420
		.020	27.520
		.040	24.180
A015	SAN JOSE 1955 N59E	.060	14.760
		.000	49.600
		.020	9.760
A016	SAN JOSE 1955 N59E	.040	1.820
		.060	1.360
		.000	51.760
A017	SAN JOSE 1955 N59E	.020	9.020
		.040	1.400
		.060	1.400

#	RECORD NAME	CUTOFF (G)	DURATION (SEC)
A001	EL CENTRO 1940 SOUTH	.000	54.780
		.020	51.140
		.040	50.520
A002	EL CENTRO 1940 WEST	.060	26.540
		.000	53.500
		.020	50.620
A003	FERNDALE 1951 S44W	.040	29.660
		.060	26.760
		.000	55.920
A004	FERNDALE 1951 S44W	.020	14.760
		.040	7.780
		.060	4.540
A005	PASADENA 1952 SOUTH	.000	55.920
		.020	13.440
		.040	9.580
A006	PASADENA 1952 WEST	.060	4.740
		.000	77.300
		.020	35.380
A007	PASADENA 1952 WEST	.040	17.560
		.060	.000
		.000	77.400
A008	TFT 1952 N21E	.020	37.500
		.040	19.160
		.060	.000
A009	TFT 1952 S69E	.000	54.400
		.020	44.920
		.040	29.900
A010	SANTA BARBARA 1952 N42E	.060	18.600
		.000	54.420
		.020	45.700
A011	SANTA BARBARA 1952 S48E	.040	18.840
		.060	17.480
		.000	75.520
A012	SANTA BARBARA 1952 S48E	.020	59.740
		.040	21.600
		.060	17.440
A013	SANTA BARBARA 1952 S48E	.000	75.500
		.020	43.560
		.040	20.680
A014	SANTA BARBARA 1952 S48E	.060	9.880

TABLE 4.1. Earthquake Durations for Different Cutoff Levels

#	RECORD NAME	CUTOFF (G)	DURATION (SEC)
A011	EL CENTRO 1956 SOUTH	.000	90.060
		.020	32.400
		.040	.000
A012	EL CENTRO 1956 WEST	.060	.000
		.000	90.000
		.020	24.860
A013	EL CENTRO AFTERSHOCK 1956 SOUTH	.040	10.600
		.060	.000
		.000	88.140
A014	EL CENTRO AFTERSHOCK 1956 WEST	.020	.000
		.040	.000
		.060	.000
A015	SOUTHERN PACIFIC 1957 N45E	.000	68.120
		.020	.000
		.040	.000
A016	SOUTHERN PACIFIC 1957 N45W	.060	.000
		.000	36.980
		.020	8.060
A017	ALEXANDER BLDG. 1957 N09W	.040	1.900
		.060	.000
		.000	39.060
A018	ALEXANDER BLDG. 1957 N81E	.020	5.120
		.040	2.020
		.060	.000
A019	GOLDEN GATE 1957 N10E	.000	59.720
		.020	3.660
		.040	2.780
A020	GOLDEN GATE 1957 S80E	.060	.000
		.000	39.620
		.020	29.360

#	RECORD NAME	CUTOFF (G)	DURATION (SEC)
A016	S F STATE BLDG 1957 S09E	.000	40.800
		.020	29.040
		.040	5.740
A017	S F STATE BLDG 1957 S81W	.060	2.820
		.000	40.740
		.020	26.960
A018	OAKLAND 1957 N26E	.040	2.660
		.060	.000
		.000	40.300
A019	OAKLAND 1957 S64E	.020	2.020
		.040	.000
		.060	.000
A020	HOLLISTER 1961 S01W	.000	40.500
		.020	20.160
		.040	12.030
A021	HOLLISTER 1961 N89W	.060	1.180
		.000	40.520
		.020	19.200
A022	BORRECO 1968 SOUTH	.040	12.900
		.060	10.120
		.000	87.440
A023	BORRECO 1968 WEST	.020	35.640
		.040	21.280
		.060	9.420
A024	SAN DIEGO 1968 SOUTH	.000	87.230
		.020	43.500
		.040	20.940
A025	SAN DIEGO 1968 EAST	.060	.000
		.000	79.220
		.020	13.440
A026	SAN DIEGO 1968 EAST	.040	.000
		.060	.000
		.000	79.260
A027	SAN DIEGO 1968 EAST	.020	11.340
		.040	.000
		.060	.000

TABLE 4.1. Earthquake Durations for Different Cutoff Levels (Cont.)

#	RECORD NAME	CUTOFF (G)	DURATION (SEC)
B026	FERNDALE 1938 S45W	.000	98.800
		.020	14.160
		.040	12.940
		.060	4.740
B027	FERNDALE 1938 N45W	.000	98.460
		.020	25.580
		.040	11.700
		.060	2.720
B028	WASHINGTON 1949 S02W	.000	89.780
		.020	22.140
		.040	.660
		.060	.000
B029	WASHINGTON 1949 N88W	.000	89.680
		.020	21.920
		.040	10.920
		.060	5.360
B030	OLYMPIA 1949 S04E	.000	74.880
		.020	6.180
		.040	.000
		.060	.000
B031	OLYMPIA 1949 S86W	.000	74.840
		.020	.420
		.040	.000
		.060	.000
B032	HELENA 1935 NORTH	.000	90.320
		.020	34.940
		.040	24.060
		.060	15.120
B033	HELENA 1935 EAST	.000	90.260
		.020	35.160
		.040	20.220
		.060	17.880
B034	FERNDALE 1941 S45W	.000	50.940
		.020	4.960
		.040	3.500
		.060	3.280
B035	FERNDALE 1941 N45W	.000	51.040
		.020	4.920
		.040	3.740
		.060	3.560

TABLE 4.1. Earthquake Durations for Different Cutoff Levels (Cont.)

#	RECORD NAME	CUTOFF (G)	DURATION (SEC)
B031	TAFT 1954 N21E	.000	65.420
		.020	13.200
		.040	6.280
		.060	5.840
		.000	65.380
		.020	15.100
B032	OLYMPIA 1965 S08E	.040	7.420
		.060	6.540
		.000	61.880
		.020	35.140
B033	PARKFIELD A.2 1966 N65E	.040	15.280
		.060	13.080
		.000	81.980
		.020	29.580
B034	PARKFIELD A.5 1966 N05W	.040	23.260
		.060	12.300
		.000	43.680
		.020	26.560
B035	PARKFIELD A.8 1966 N05E	.040	14.500
		.060	13.720
		.000	43.960
		.020	19.680
B036	PARKFIELD A.12 1966 N50E	.040	13.900
		.060	12.800
		.000	45.980
		.020	20.940
B037	PARKFIELD A.12 1966 N40W	.040	15.460
		.060	12.200
		.000	26.220
		.020	20.540
B038	SAN LUIS OBISPO 1966 N36W	.040	13.440
		.060	9.620
		.000	26.180
		.020	21.000
B039	EUREKA 1967 S11E	.040	10.820
		.060	7.260
		.000	45.020
		.020	24.620
B040	BORREGO 1968 N33E	.040	15.980
		.060	6.000
		.000	45.020
		.020	22.460
B041	BORREGO 1968 N57W	.040	16.000
		.060	6.000
		.000	45.020
		.020	22.460

TABLE 4.1. Earthquake Durations for Different Cutoff Levels (Cont.)

#	RECORD NAME	CUTOFF (G)	DURATION (SEC)
D058	HOLLYWOOD ST PE. 1971 SOUTH	.000	79.500
		.020	31.800
		.040	12.860
	HOLLYWOOD ST PE. 1971 EAST	.060	7.620
		.000	79.500
		.020	31.780
G106	CALTECH LAB 1971 SOUTH	.040	11.580
		.060	9.560
		.000	99.020
	CALTECH LAB 1971 WEST	.020	12.960
		.040	6.860
		.060	7.500
J141	LAKE HUGHES A1 1971 N21E	.000	99.020
		.020	16.380
		.040	9.420
	LAKE HUGHES A1 1971 S69E	.060	8.200
		.000	50.220
		.020	21.260
	LAKE HUGHES A4 1971 S69E	.040	15.360
		.060	5.020
		.000	66.280
J142	LAKE HUGHES A4 1971 S69E	.020	20.980
		.040	16.460
		.060	6.840
	LAKE HUGHES A4 1971 S21W	.000	37.040
		.020	16.680
		.040	6.500
J143	LAKE HUGHES A9 1971 N21E	.060	6.040
		.000	35.000
		.020	12.680
	LAKE HUGHES A9 1971 N69W	.040	4.920
		.060	3.580
		.000	35.040
	LAKE HUGHES A9 1971 N69W	.020	12.860
		.040	3.600
		.060	3.600

#	RECORD NAME	CUTOFF (G)	DURATION (SEC)
C051	PACOIMA DAX 1971 S16E	.000	41.840
		.020	40.420
		.040	34.200
	PACOIMA DAX 1971 S74W	.060	35.740
		.000	41.740
		.020	34.800
C055	ORION BLDG 1971 NORTH	.040	33.800
		.060	33.800
		.000	59.520
	ORION BLDG 1971 WEST	.020	36.660
		.040	20.780
		.060	19.600
C051	EAST FIRST ST 1971 N36E	.000	59.620
		.020	36.860
		.040	22.460
	EAST FIRST ST 1971 N54W	.060	20.800
		.000	52.560
		.020	50.040
C054	FIGUEROA ST 1971 N52W	.040	10.640
		.060	8.540
		.000	52.320
	FIGUEROA ST 1971 S38W	.020	16.700
		.040	8.040
		.060	6.820
C054	FIGUEROA ST 1971 N52W	.000	57.300
		.020	16.960
		.040	8.980
	FIGUEROA ST 1971 S38W	.060	7.780
		.000	57.320
		.020	19.820
D057	HOLLYWOOD ST BAS. 1971 SOUTH	.040	12.700
		.060	6.540
		.000	62.160
	HOLLYWOOD ST BAS. 1971 EAST	.020	18.220
		.040	12.920
		.060	10.180
	HOLLYWOOD ST BAS. 1971 EAST	.000	62.140
		.020	30.900
		.040	14.820
	HOLLYWOOD ST BAS. 1971 EAST	.060	9.560

TABLE 4.1. Earthquake Durations for Different Cutoff Levels (Cont.)

#	RECORD NAME	CUTOFF (G)	DURATION (SEC)
J144	LAKE HUGHES A12 1971 N21E	.000	36.640
		.020	22.620
		.040	15.000
		.060	12.120
	LAKE HUGHES A12 1971 N69W	.000	36.760
		.020	19.060
		.040	14.580
		.060	10.500
L166	LANKERSHIM 1971 NORTH	.000	65.360
		.020	19.260
		.040	10.500
		.060	7.540
	LANKERSHIM 1971 WEST	.000	65.200
		.020	19.020
		.040	8.840
		.060	7.460

TABLE 4.1. Earthquake Durations for Different Cutoff Levels
(Cont.)

(i.e., the time during which the damage is likely to occur). To develop a simple and usable definition of record duration, a sensitivity study of the effects of various levels of acceleration cutoffs on the duration is conducted. Table 4.1 gives the results of this study. After inspection of these results, it was decided to use 0.02G level of acceleration as the cutoff acceleration. In other words, the record is truncated at the time when the acceleration becomes and remains smaller than 0.02G. This is a rather arbitrary definition; however, for this study, it seems to be a reasonable definition.

Because of the above definition of duration, three earthquake records are totally eliminated, their peak ground accelerations being smaller than 0.02G. The total number of records analyzed is thus 97 noting that the Parkfield A.2 1966 earthquake has only one component recorded (record B033 in Hudson, et al., 1972). These records are chosen to provide a large spectrum of earthquakes recorded on different soil conditions and at various distances from the epicenters.

Only horizontal accelerations have been considered in the study. The parameters mean, RMS and a_p have been computed using both the peak values and the values corresponding to the equal time increments (the input and responses being digitized at equal time increments). The parameters values (mean, RMS and a_p) obtained from both methods are proportional. Hence they are related by a constant and the conclusions drawn for one set of parameters can be applied to the other. In the following sections only the results based on the incremental values are presented.

4.3 Statistics of the Input Parameters

The first step in the analysis of past earthquake data is to evaluate the various input and statistical parameters associated with the recorded and digitized accelerograms. Table 4.2 shows the results of this analysis. Each of the parameters analyzed is described below:

- NBPK (Column 4)

This parameter represents the number of peaks of the earthquake records. Strictly, NBPK represents the number of zero crossings of the given accelerogram. This information is useful in calculating the RMS of the record.

- RMS (Column 5)

This parameter is a statistical summary of all the incremental values recorded for a given accelerogram. In terms of the frequency domain, the RMS of the input is the ordinate of RMS vs. period graph at zero period. This information is extensively used in developing a stable parameter described in later sections. The RMS for all the records is calculated using the following expression.

$$\text{RMS} = \sqrt{\frac{\sum h_i^2}{n-1}} \quad (4.1)$$

where h_i = individual acceleration amplitude

n = total number of time increments.

This parameter is a better description of the accelerogram than the peak value representation currently used. Also, being a statistical summary, its variation is less than the variation of the peaks.

#	RECORDER NAME	PGA (G)	N3PK	RMS (G)	LAMBDA	P(A>A*)			K1	
						A* P=5%	A* P=10%	P=5%	P=10%	
A001	EL CENTRO 1940 SOUTH	.348	315	.048	35.43	.085	.065	1.760	1.353	
	EL CENTRO 1940 WEST	.214	296	.041	58.07	.079	.060	1.939	1.491	
A002	FERNDALE 1951 S44W	.104	106	.019	73.14	.041	.031	2.144	1.648	
	FERNDALE 1951 N46W	.112	98	.022	63.96	.047	.036	2.092	1.608	
A003	PASADENA 1952 SOUTH	.047	178	.010	143.68	.021	.016	2.079	1.598	
	PASADENA 1952 WEST	.053	157	.014	106.04	.028	.022	2.090	1.606	
A004	TAFT 1952 N21E	.156	298	.028	53.15	.056	.043	2.017	1.551	
	TAFT 1952 S69E	.179	300	.029	51.90	.058	.044	1.995	1.534	
A005	SANTA BARBARA 1952 N42E	.090	170	.019	76.74	.039	.030	2.014	1.548	
	SANTA BARBARA 1952 S48E	.131	177	.020	74.00	.040	.031	1.990	1.530	
A006	HOLLYWOOD 1952 BAS. SOUTH	.055	215	.012	110.50	.027	.021	2.201	1.692	
	HOLLYWOOD 1952 BAS. EAST	.044	201	.012	111.12	.027	.021	2.202	1.693	
A007	HOLLYWOOD 1952 PE. SOUTH	.059	245	.013	108.35	.028	.021	2.210	1.698	
	HOLLYWOOD 1952 PE. EAST	.042	230	.012	110.91	.027	.021	2.208	1.697	
A008	EUREKA 1954 N11W	.168	103	.031	54.73	.055	.042	1.761	1.353	
	EUREKA 1954 N79E	.258	104	.039	48.86	.061	.047	1.556	1.196	
A009	FERNDALE 1954 N44E	.159	100	.034	44.71	.067	.052	1.965	1.510	
	FERNDALE 1954 N46W	.201	110	.029	55.59	.054	.041	1.846	1.419	
A010	SAN JOSE 1955 N31W	.102	65	.021	71.65	.042	.032	1.946	1.495	
	SAN JOSE 1955 N59E	.108	73	.019	82.71	.036	.028	1.934	1.487	
A011	EL CENTRO 1956 SOUTH	.033	147	.010	133.57	.022	.017	2.315	1.780	
	EL CENTRO 1956 WEST	.051	100	.015	84.83	.035	.027	2.379	1.829	
A013	SOUTHERN PACIFIC 1957 N45E	.047	55	.011	117.65	.025	.020	2.512	1.777	
	SOUTHERN PACIFIC 1957 N45W	.046	34	.017	72.52	.041	.032	2.403	1.847	

TABLE 4.2. Listing of Input Parameters

#	MECCRU NAME	PGA (G)	MBPK	RMS (G)	LAMBDA	P(A>A*)			K1	
						A* P=5%	A* P=10%	A* P=10%	P=5%	P=10%
A014	ALEXANDER BLDG 1957 N09W	.043	56	.014	97.68	.031	.024	.024	2.197	1.688
	ALEXANDER BLDG 1957 N81E	.046	236	.005	341.02	.009	.007	.007	1.618	1.243
A015	GOLDEN GATE 1957 N10E	.083	45	.023	63.95	.047	.036	.036	2.018	1.551
	GOLDEN GATE 1957 S80E	.105	61	.027	57.64	.052	.040	.040	1.909	1.487
A016	S F STATE BLDG 1957 S09E	.065	192	.011	192.93	.016	.012	.012	1.445	1.110
	S F STATE BLDG 1957 S81W	.056	206	.009	229.04	.013	.010	.010	1.518	1.167
A017	OAKLAND 1957 N26E	.040	20	.014	93.25	.032	.025	.025	2.292	1.761
	OAKLAND 1957 S64E	.024	22	.009	149.88	.020	.015	.015	2.248	1.728
A018	HOLLISTER 1961 S01W	.065	62	.020	63.87	.047	.036	.036	2.319	1.782
	HOLLISTER 1961 N89W	.179	76	.029	48.73	.061	.047	.047	2.098	1.613
A019	BORREGO 1968 SOUTH	.130	187	.020	73.70	.041	.031	.031	2.065	1.587
	BORREGO 1968 WEST	.057	214	.014	93.39	.032	.025	.025	2.263	1.740
A020	SAN DIEGO 1968 SOUTH	.030	75	.007	214.71	.014	.011	.011	1.972	1.515
	SAN DIEGO 1968 EAST	.029	57	.007	214.13	.014	.011	.011	1.991	1.530
B021	LONG BEACH 1933 N08E	.133	69	.031	43.28	.069	.053	.053	2.239	1.721
	LONG BEACH 1933 S82W	.154	122	.022	67.65	.044	.034	.034	2.004	1.540
B022	HOLLYWOOD PE. 1933 SOUTH	.044	74	.013	93.14	.032	.025	.025	2.515	1.933
	HOLLYWOOD PE. 1933 WEST	.087	83	.026	50.21	.060	.046	.046	2.326	1.788
B023	HOLLYWOOD BAS. 1933 SOUTH	.033	47	.007	184.80	.016	.012	.012	2.213	1.701
	HOLLYWOOD BAS. 1933 EAST	.027	2	.015	61.56	.037	.028	.028	2.450	1.884
B024	EL CENTRO 1934 SOUTH	.160	223	.032	46.80	.064	.049	.049	2.026	1.557
	EL CENTRO 1934 WEST	.183	221	.054	44.96	.067	.051	.051	1.986	1.526
B025	HELENA 1935 NORTH	.146	62	.030	51.24	.058	.045	.045	1.918	1.474
	HELENA 1935 EAST	.145	67	.038	41.57	.072	.055	.055	1.874	1.440

TABLE 4.2. Listing of Input Parameters (Cont.)

#	RECORD NAME	PGA (G)	WRPK	RMS (G)	LALBDA	P(A>A*)		K1	
						A* P=5%	A* P=10%	P=5%	P=10%
B026	FERNDAL 1938 S45W	.144	89	.022	65.90	.045	.035	2.032	1.562
	FERNDAL 1938 N45W	.089	72	.022	64.96	.046	.035	2.118	1.628
B027	FERNDAL 1941 S45W	.062	54	.016	82.66	.036	.028	2.311	1.776
	FERNDAL 1941 N45W	.039	43	.015	83.52	.036	.028	2.460	1.891
B028	WASHINGTON 1949 S02W	.068	254	.015	92.65	.032	.025	2.141	1.646
	WASHINGTON 1949 N88W	.067	251	.013	107.21	.028	.021	2.152	1.654
B029	OLYMPIA 1949 S04E	.165	351	.032	49.57	.060	.046	1.906	1.465
	OLYMPIA 1949 S86W	.280	368	.038	42.07	.071	.055	1.666	1.434
B030	FERNDAL 1952 S44W	.054	81	.017	90.06	.033	.026	2.004	1.540
	FERNDAL 1952 N46W	.076	109	.014	108.17	.028	.021	2.005	1.541
B031	TAFT 1954 N21E	.065	115	.013	125.86	.024	.018	1.891	1.454
	TAFT 1954 S69E	.068	124	.013	107.58	.028	.021	2.083	1.601
B032	OLYMPIA 1965 S04E	.137	534	.023	68.05	.044	.034	1.892	1.454
	OLYMPIA 1965 S86W	.198	297	.031	48.98	.061	.047	1.977	1.519
B033	PARKFIELD A.2 1966 N65E	.489	121	.066	30.99	.097	.074	1.456	1.119
B034	PARKFIELD A.5 1966 N05W	.355	186	.045	40.99	.073	.056	1.634	1.256
	PARKFIELD A.5 1966 N85E	.434	157	.051	36.41	.082	.063	1.614	1.241
B035	PARKFIELD A.8 1966 N05E	.237	141	.031	48.76	.061	.047	1.990	1.529
	PARKFIELD A.8 1966 N40W	.275	162	.034	48.65	.062	.047	1.810	1.391
B036	PARKFIELD A.12 1966 N50E	.053	175	.012	108.91	.028	.021	2.277	1.750
	PARKFIELD A.12 1966 N40W	.064	188	.013	101.75	.029	.023	2.311	1.776
B037	TEMBLOR 1966 N65W	.269	84	.042	43.06	.070	.053	1.675	1.287
	TEMBLOR 1966 S25W	.347	80	.050	36.80	.081	.063	1.617	1.243
B040	BORRECO 1968 N33E	.041	192	.008	183.68	.016	.013	2.102	1.616
	BORRECO 1968 N57W	.046	141	.009	152.50	.020	.015	2.137	1.642

TABLE 4.2. Listing of Input Parameters (Cont.)

#	RECORD NAME	PGA (G)	NBPk	RMS (G)	LAMSDA	P(A>A*)			K1	
						A*	P=5%	A*	P=10%	P=5%
C041	PACOIMA DAM 1971 S16E	1.170	412	.116	19.35	.155	.119	1.332	1.024	
	PACOIMA DAM 1971 S74W	1.075	380	.120	17.56	.171	.131	1.422	1.093	
C048	ORION BLDG 1971 NORTH	.255	175	.046	34.22	.086	.067	1.900	1.460	
	ORION BLDG 1971 WEST	.134	158	.033	43.27	.069	.053	2.078	1.597	
C051	EAST FIRST ST 1971 N36E	.100	207	.021	73.34	.041	.031	1.921	1.476	
	EAST FIRST ST 1971 N	.125	129	.028	49.42	.061	.047	2.164	1.663	
C054	FIGUEROA ST 1971 N52W	.150	125	.028	55.14	.054	.042	1.911	1.469	
	FIGUEROA ST 1971 S38W	.119	150	.026	54.72	.055	.042	2.136	1.642	
D057	HOLLYWOOD ST BAS. 1971 SOUTH	.106	120	.028	50.10	.060	.046	2.158	1.659	
	HOLLYWOOD ST BAS. 1971 EAST	.151	137	.028	57.54	.052	.040	1.859	1.429	
D058	HOLLYWOOD ST PE. 1971 SOUTH	.171	283	.029	58.81	.051	.039	1.775	1.364	
	HOLLYWOOD ST PE. 1971 EAST	.211	225	.036	48.15	.062	.048	1.727	1.328	
G106	CALTECH LAB 1971 SOUTH	.089	143	.023	58.93	.051	.039	2.214	1.702	
	CALTECH LAB 1971 WEST	.192	177	.056	44.05	.068	.052	1.865	1.433	
J141	LAKE HUGHES A1 1971 N21E	.148	117	.028	54.23	.055	.042	1.998	1.535	
	LAKE HUGHES A1 1971 S69E	.111	142	.024	55.93	.054	.041	2.239	1.721	
J142	LAKE HUGHES A4 1971 S69E	.171	305	.027	57.18	.052	.040	1.917	1.474	
	LAKE HUGHES A4 1971 S21W	.146	248	.027	55.32	.054	.042	2.012	1.546	
J143	LAKE HUGHES A9 1971 N21E	.122	176	.024	64.48	.046	.036	1.973	1.517	
	LAKE HUGHES A9 1971 N69W	.112	176	.021	73.84	.041	.031	1.928	1.482	
J144	LAKE HUGHES A12 1971 N21E	.353	260	.050	34.25	.087	.067	1.738	1.336	
	LAKE HUGHES A12 1971 N69W	.283	218	.051	32.10	.093	.072	1.846	1.419	
L166	LANKERSHIM 1971 NORTH	.167	157	.025	59.30	.051	.039	2.006	1.544	
	LANKERSHIM 1971 WEST	.150	152	.031	49.31	.061	.047	1.940	1.491	

TABLE 4.2. Listing of Input Parameters (Cont.)

- Probability distribution function of amplitude values at equal time increments.

To understand the variability and distribution of the accelerogram amplitudes, various probability functions are tried considering the amplitude values as a random variable. In particular two distributions seem to fit the data quite well. They are the gamma and the exponential distributions. Figure 4.1 shows two typical CDF fits of gamma and exponential distributions to accelerograms. After trying these two distributions for all the accelerograms considered, it is decided to use the exponential form because it is simple to use and fits the data reasonably well. The probability distribution function for exponential shape (Fig. 4.2) is given by

$$f_A(a) = \lambda e^{-\lambda a} \quad (4.2)$$

where A = random variable defining the acceleration amplitude

a = the value A takes

λ = constant of the exponential distribution

It should be pointed out that once the distribution of the amplitudes is known, any probability statement regarding the acceleration can be made. As an example:

$$\begin{aligned} P(\text{Peak amplitude } A \geq a) &= \int_a^{\infty} \lambda e^{-\lambda A} dA \\ &= e^{-\lambda a} \end{aligned} \quad (4.3)$$

This information is extensively used in determining a_p defined below and in obtaining the stable parameter K_1

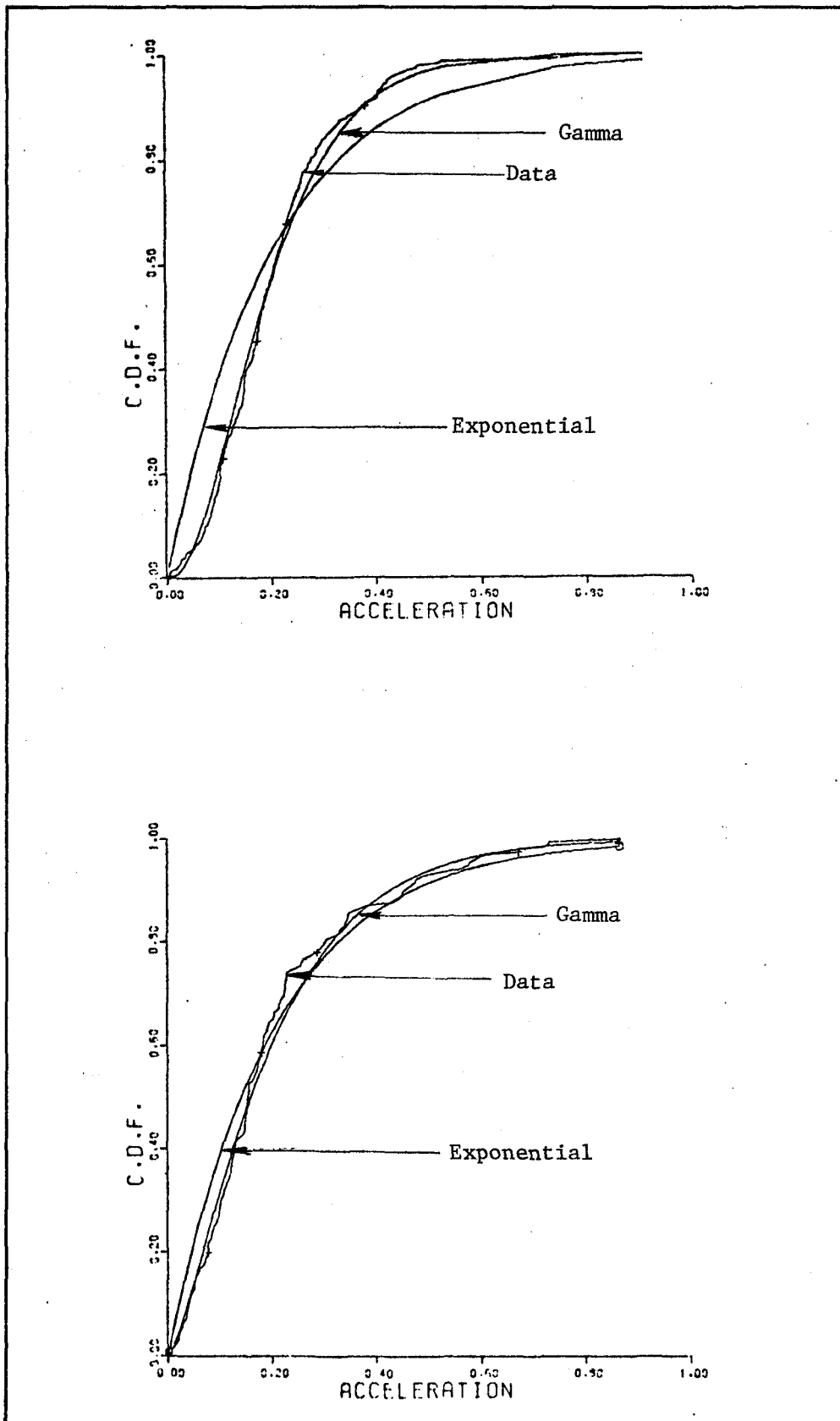


FIGURE 4.1. Two Typical Gamma and Exponential Fits of Earthquake Record Amplitudes.

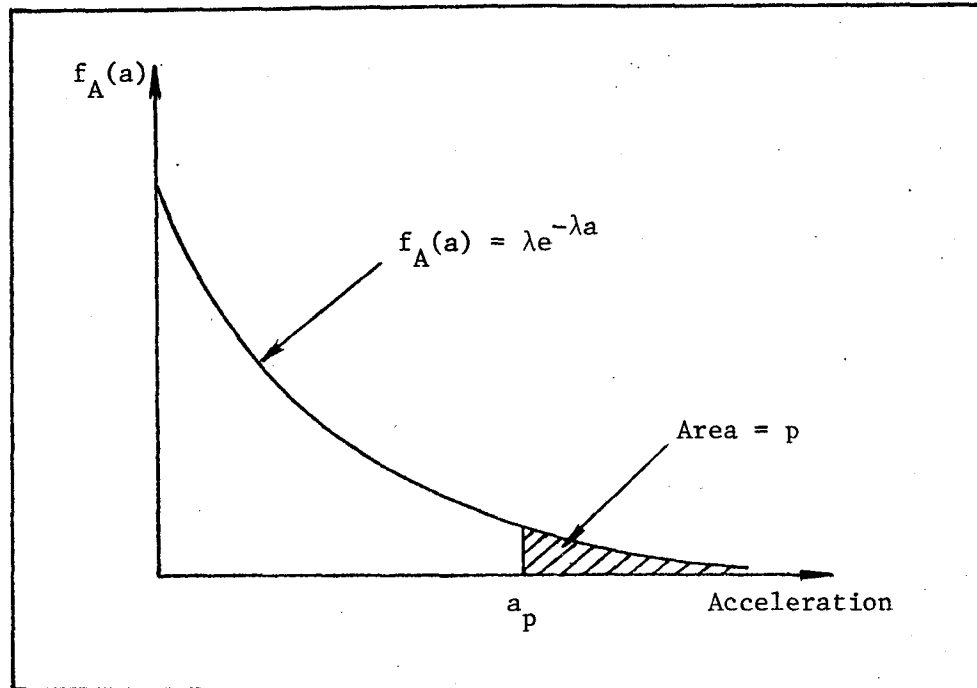


FIGURE 4.2. Exponential Distribution

- Mean acceleration (Column 6: λ)
This parameter is the average amplitude of the acceleration throughout the record. It is equal to $1/\lambda$ where λ is the parameter of the exponential distribution.
- a_p (Column 7: $p=5\%$; Column 8: $p=10\%$)
This variable represents the acceleration which has p probability of exceedence for a given accelerogram (Fig. 4.2).
- Ratio $K_1 = a_p/\text{RMS}$ (Column 9: $p=5\%$; Column 10: $p=10\%$)
As will be described in Section 4.4, this ratio is truly stable for all accelerograms with the exception of few records.

The reasons for developing the parameters presented in Table 4.2 are:

- To numerically summarize the relevant parameters describing the recorded seismic event.
- To look at the behavior of the summary parameters for trends and similarities
- To obtain the values of response parameters in the frequency domain at zero period
- To obtain a stable input parameter which could be used for a better description of the recorded event.

It should be noted that even though the values of a_p and RMS varies substantially from one accelerogram to another, the value of the ratio $K_1 = a_p/\text{RMS}$ remains essentially constant.

4.4 Statistics of the Response Parameters

As mentioned in Chapter I, the proper understanding of a given earthquake time history requires a time and frequency domain analysis. In Section 4.3, all the relevant time domain parameters as well as the frequency domain parameters for zero period are presented. In this section, the response parameters corresponding to the 97 accelerograms are evaluated and presented.

The response parameter most widely used by engineers is the response spectrum value . This value is a good representation of a single degree of freedom system to a given accelerogram. However, it is not sufficient information. For this reason, a complete analysis of the response parameters is presented in this section. Figure 4.3 shows

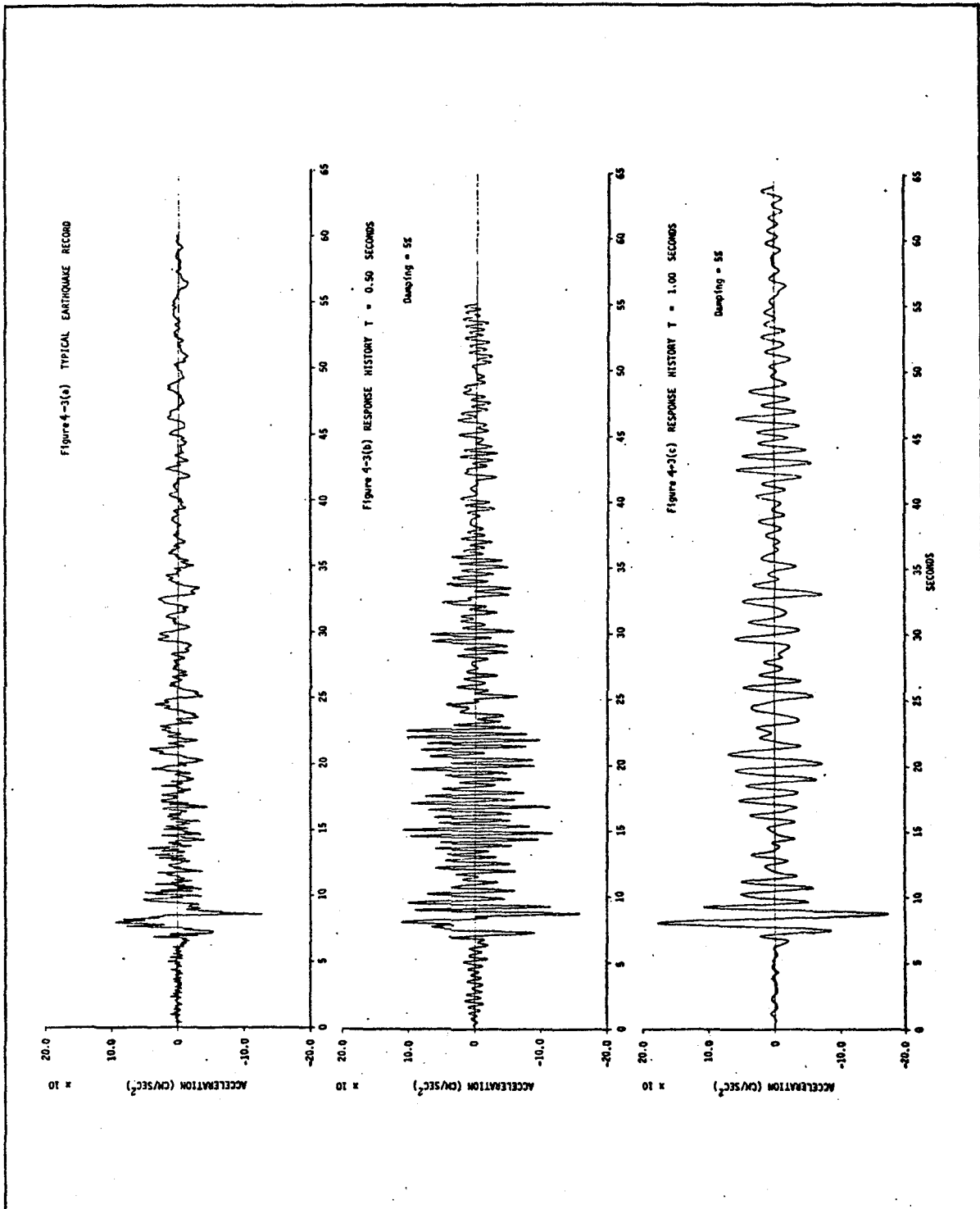


FIGURE 4.3. Typical Earthquake Record and Response Time Histories

a typical earthquake accelerogram and the responses of a one degree of freedom system (for given periods and damping). The current procedure is to look at the largest response peak, corresponding to a given damping ratio (say 5%) and period (say 0.50 sec and 1.0 sec) of the vibratory system. The highest response peaks then represent the response spectrum values corresponding to the period and damping considered.

In this study, the following parameters are evaluated for each response time history (acceleration only)

- Number of peaks (NBPK)
- Duration of the response
- Spectral acceleration response
- RMS of acceleration response
- Mean acceleration response
- Probability distribution function of response accelerations
- Parameters a_p defined as the response acceleration of a given one degree of freedom system subjected to a given accelerogram and which has probability p of being exceeded.
- Cumulative potential energy per unit mass (ENGY) in the response time acceleration history
- The ratio $K_1 = a_p/\text{RMS}$ for each response history
- The ratio $K_2 = \text{RMS}^2 \cdot T^2 / (\text{ENGY}/\text{NBPK})$ for each response history

For each of the 97 accelerograms, 60 single degree of freedom systems varying in periods from 0.08 sec to 5.0 sec are considered. The damping ratio of 5% is kept constant for all the cases. (In fewer

instances dampings of 10% and 20% are also considered.) Thus, for each accelerogram, 60 response time histories are analyzed to obtain the response parameters mentioned above.

The first decision to make in analyzing the responses concerns the duration of the response. Surprisingly, no work is available in the literature defining the duration of the response record. Unless a convention (or standard) is used to define the duration, it is not possible to evaluate parameters such as RMS, NBPK and ENGY. Hence, it is somewhat arbitrarily decided to terminate the response when the amplitude of the response acceleration peak reaches 10% of the highest response peak and does not exceed that value thereafter. This convention fixes the response duration. Another way of defining this parameter would be to consider the response duration as a function of input duration, the period of the oscillator under consideration and the damping of the system. More research is needed to rationally evaluate this parameter.

The response parameters analyzed are listed for all earthquakes in Appendix B. They are described below.

- NBPK (Column 2)

This parameter represents the number of zero crossings of the given response history. It is useful in computing the RMS and the factors K_1 and K_2 described later in this section.

- Duration (Column 3)

This parameter represents the duration of the response in seconds. It is the time for which the response acceleration amplitude (for a given oscillator) remains greater than 10% of the highest response acceleration.

- Spectral response acceleration (Column 4)

This parameter is the amplitude of the largest peak of the response. It corresponds to the widely used acceleration response spectrum.

- RMS (Column 5)

This parameter is a statistical summary of the acceleration amplitudes at equal time increments (Hald, 1952). For each period it is computed as

$$\text{RMS} = \sqrt{\frac{\sum h_i^2}{n-1}} \quad (4.4)$$

where h_i = acceleration amplitude at equal time increments

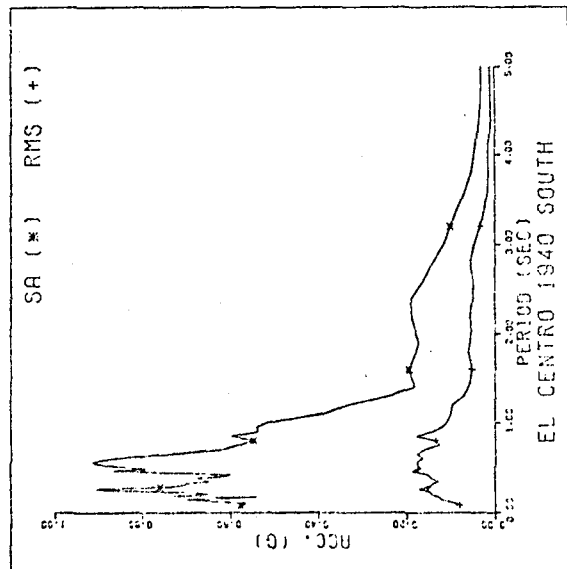
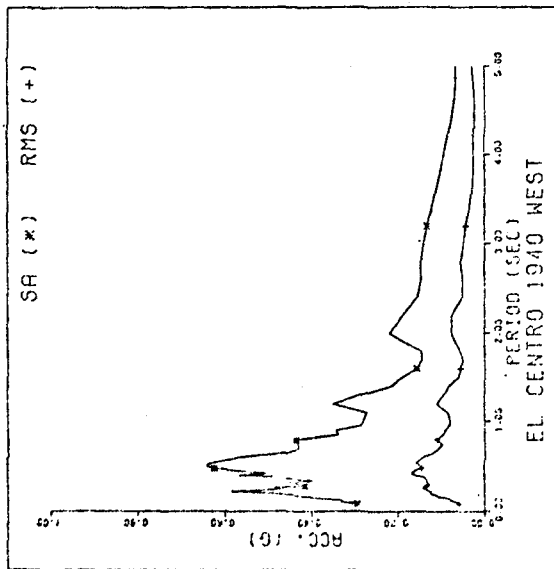
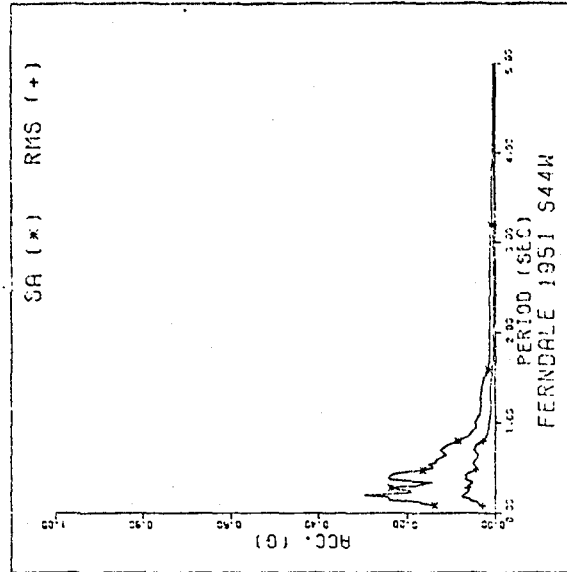
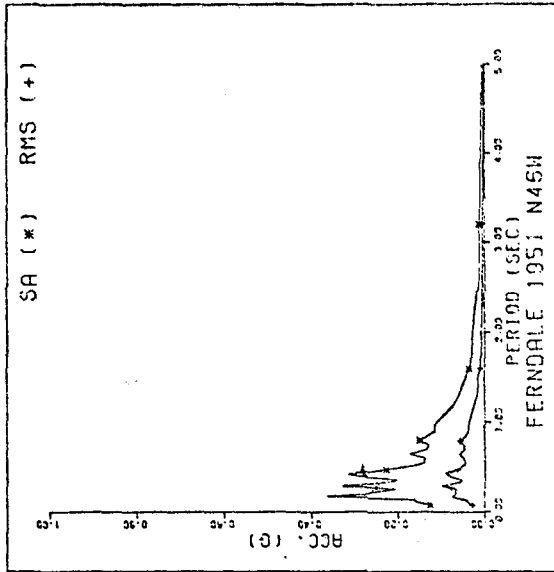
n = total number of increments

This parameter gives a better statistical description of the response than the peak values S_a . Being a statistical summary its variation is less than the variation of the peaks particularly for short period ($T < .2$ sec) which are more susceptible to phase shift as well as subtraction and addition of waves. Figure 4.4 gives to the same scale the corresponding S_a and RMS spectrum for all the records. The RMS is used in the determination of factors K_1 and K_2 .

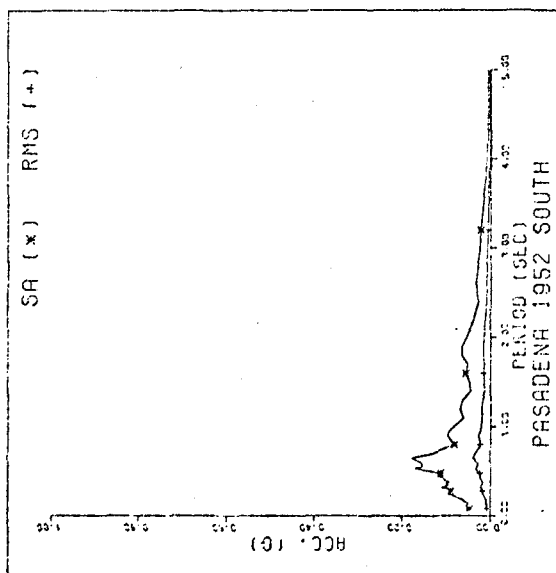
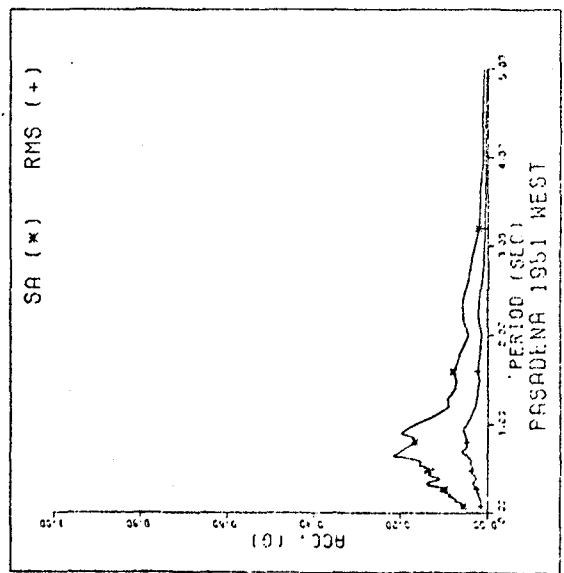
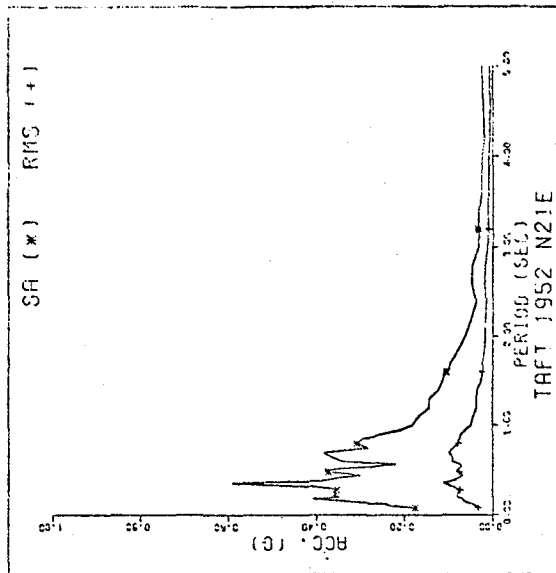
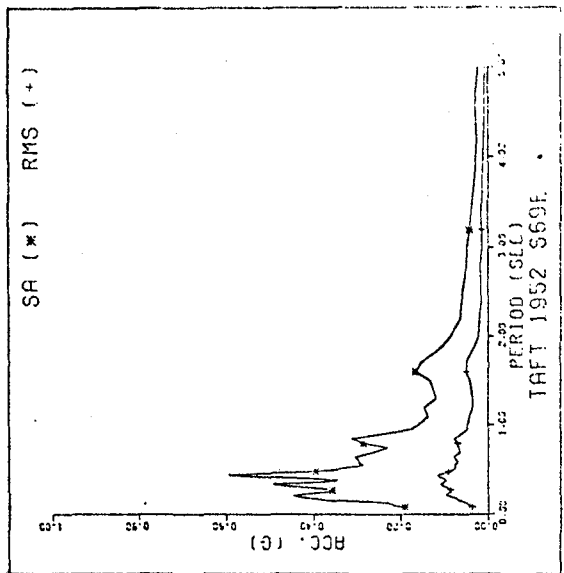
- Probability distribution function of acceleration amplitudes.

To understand the variability and distribution of the response acceleration, various probability functions are tried considering the amplitude values as a random variable. Here again, the gamma distribution shows a somewhat better fit than the exponential distribution. However for simplicity

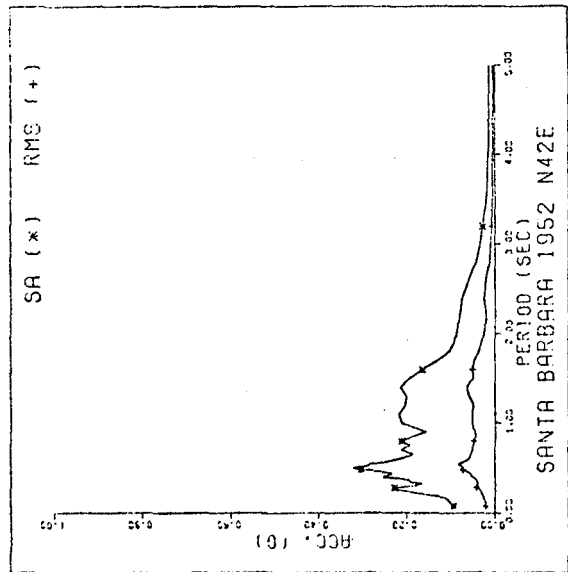
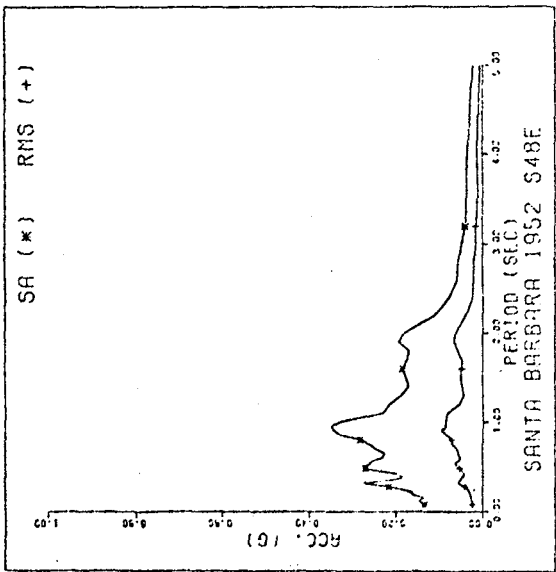
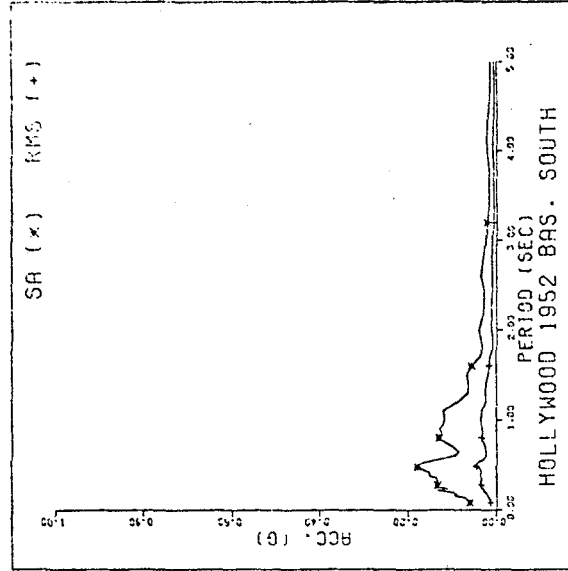
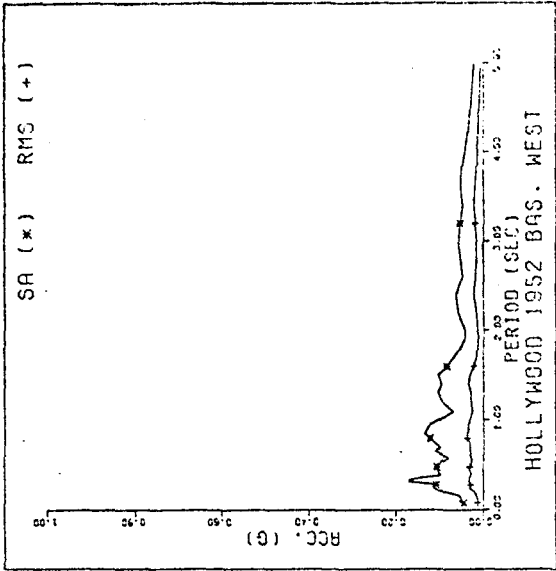
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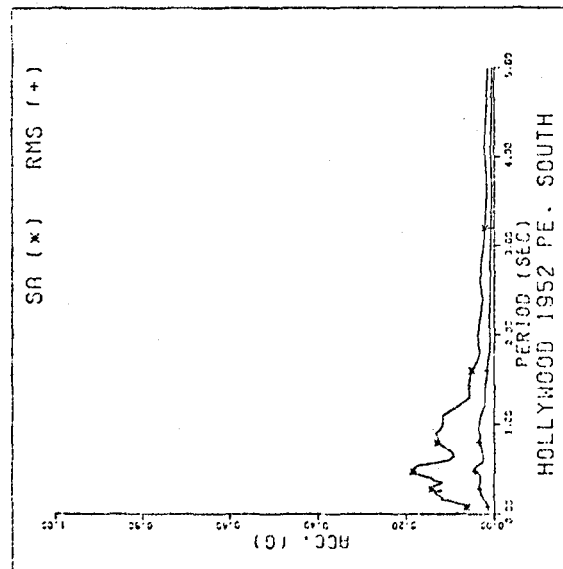
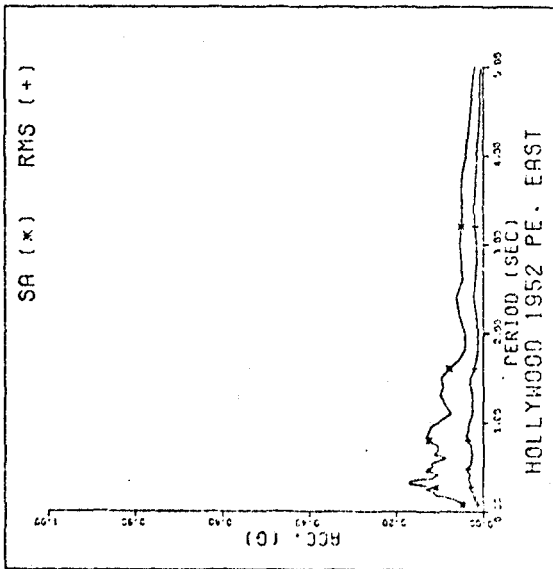
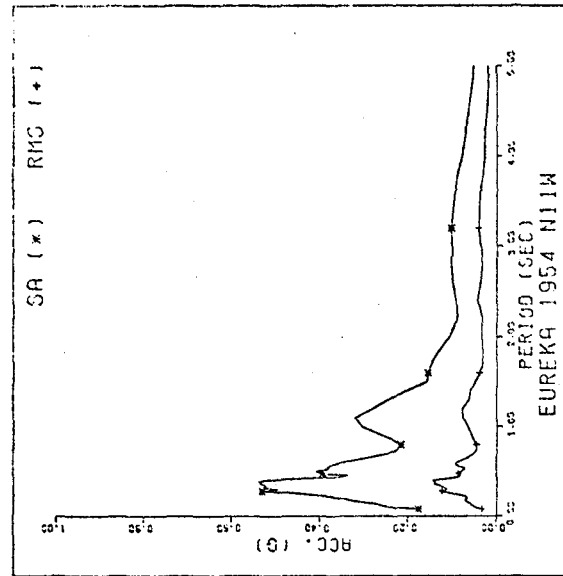
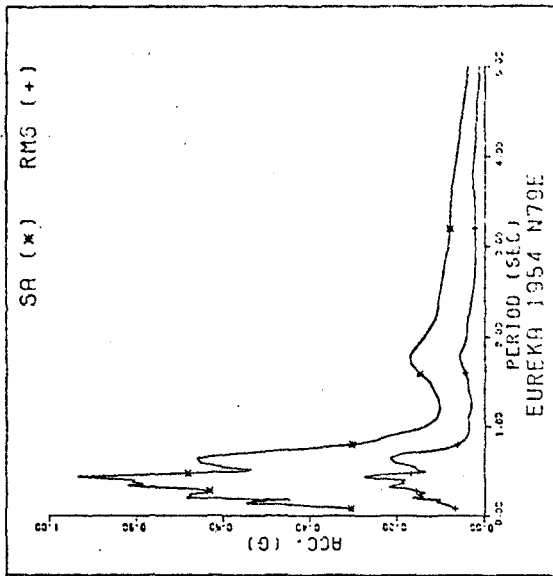
(A001) (A002) FIGURE 4.4. Acceleration and RMS Response Spectra



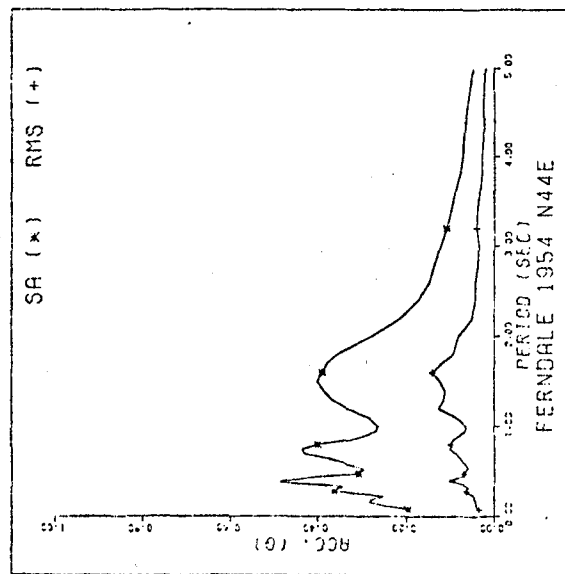
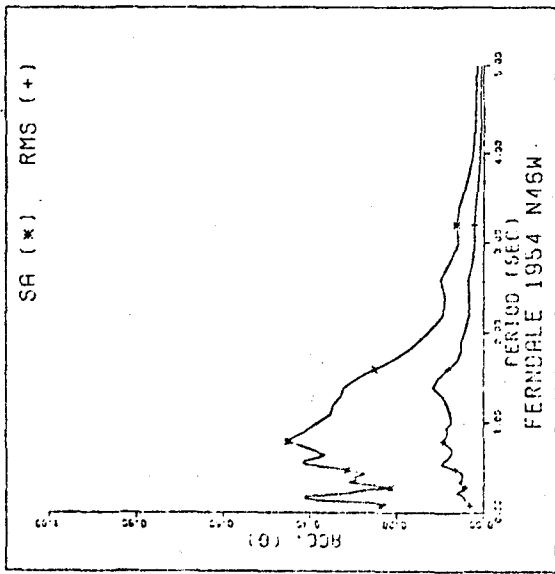
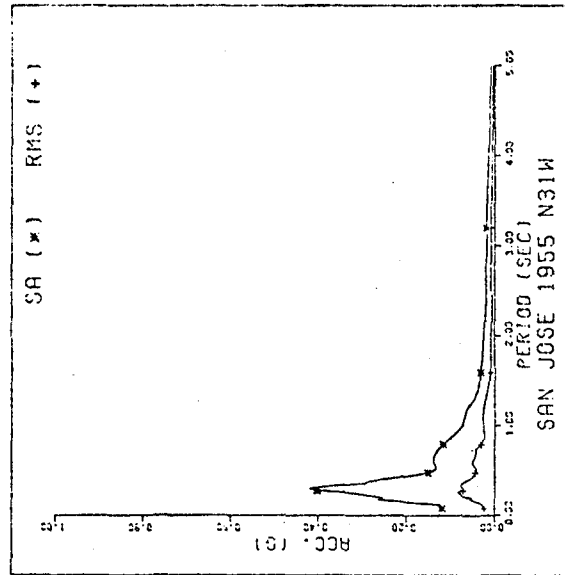
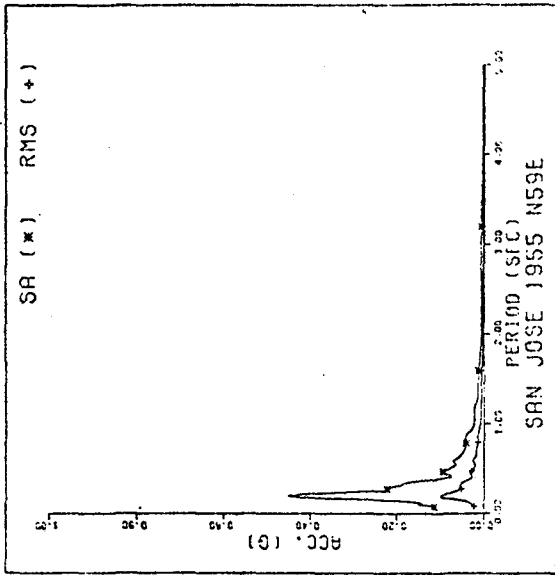
(A003) (A004) FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



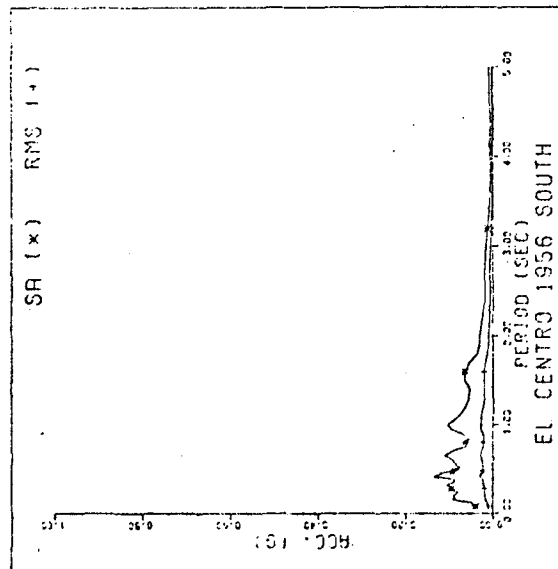
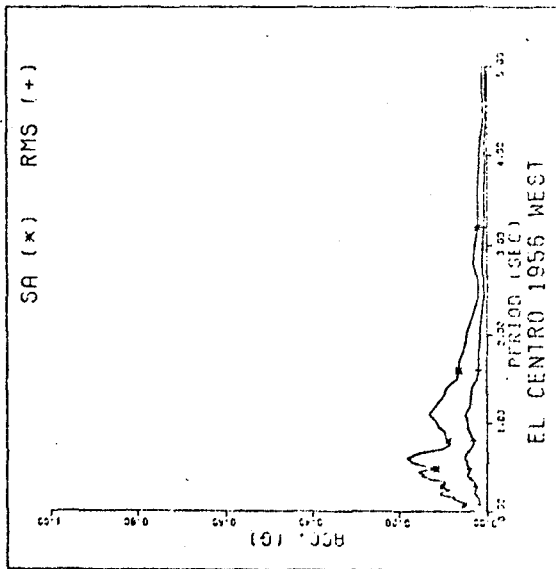
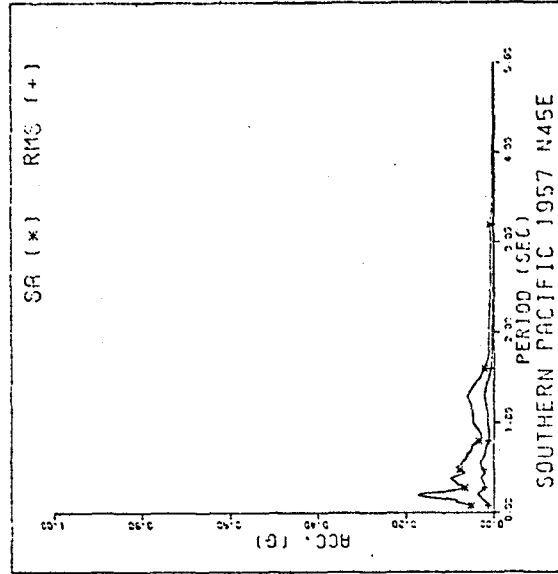
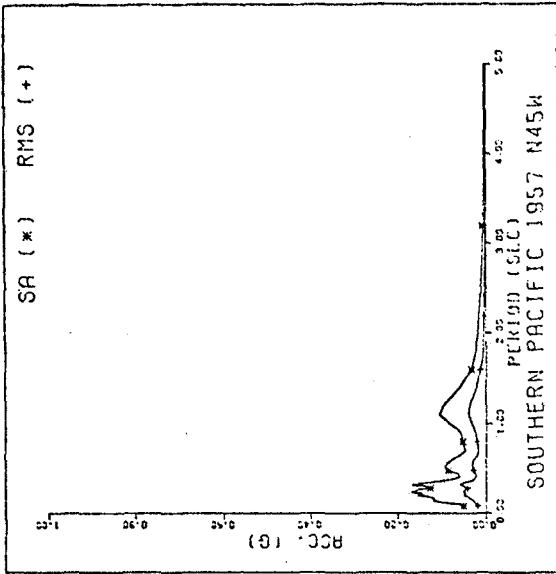
(A005) (A006)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



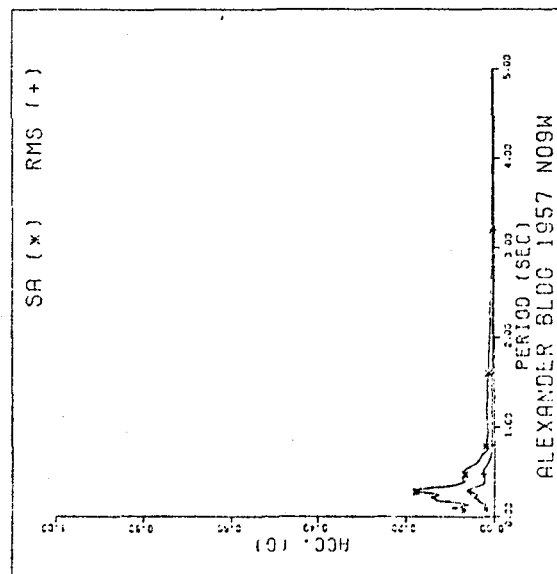
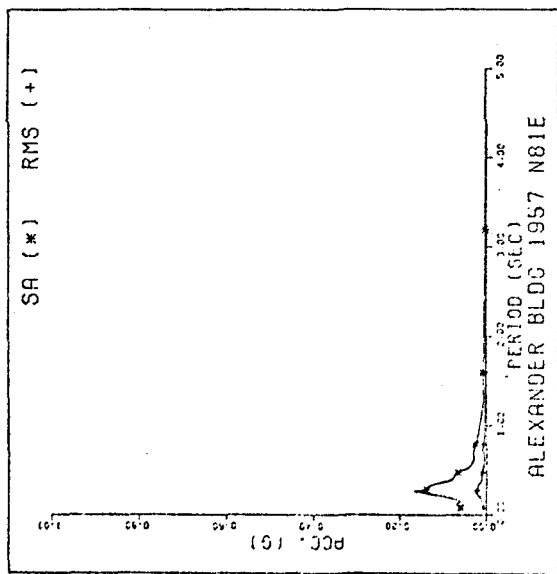
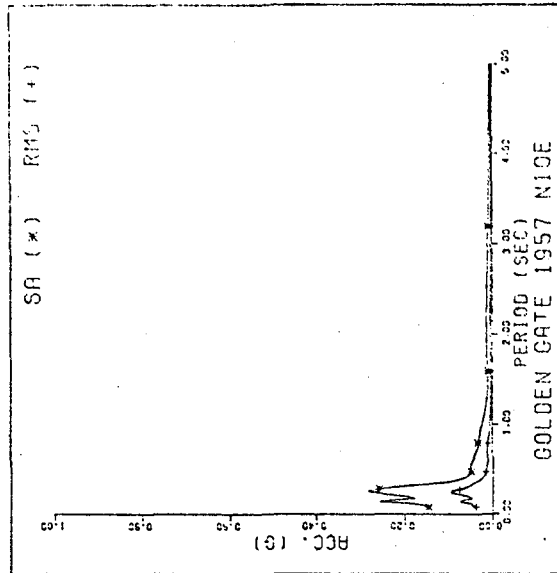
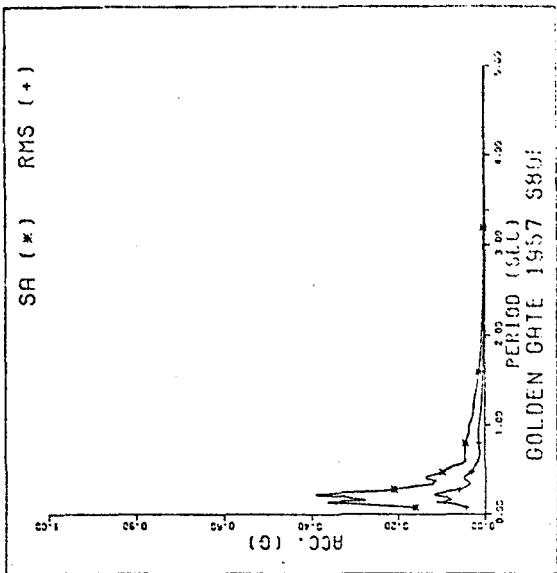
(A007) (A008) FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



(A009) (A010)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)

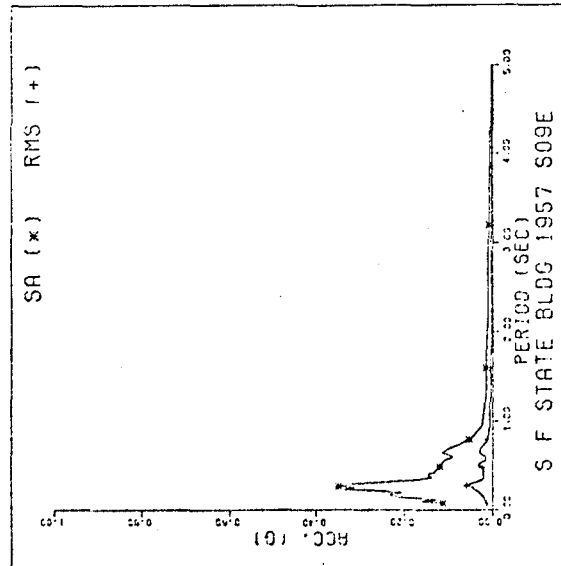
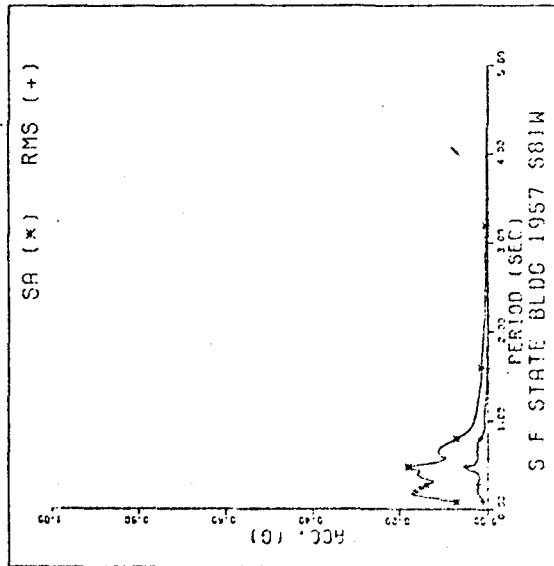
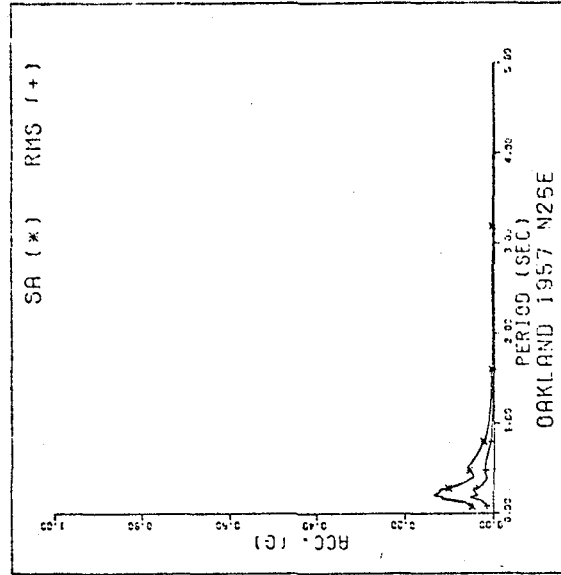
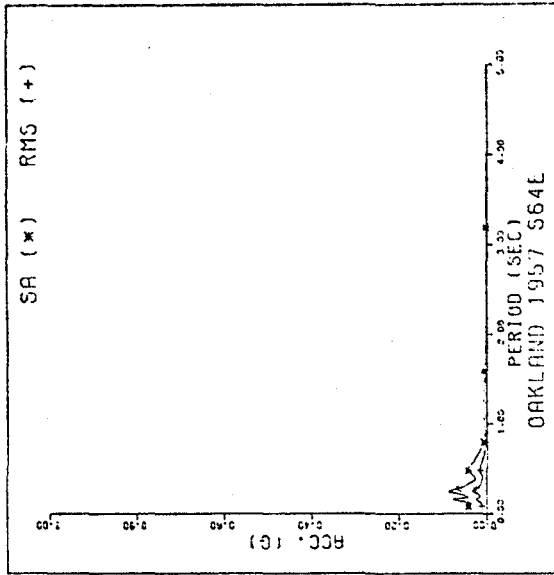


(A011) (A013) FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)

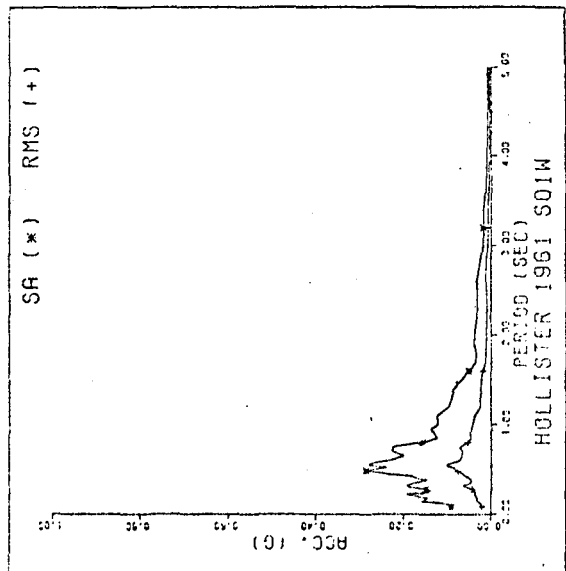
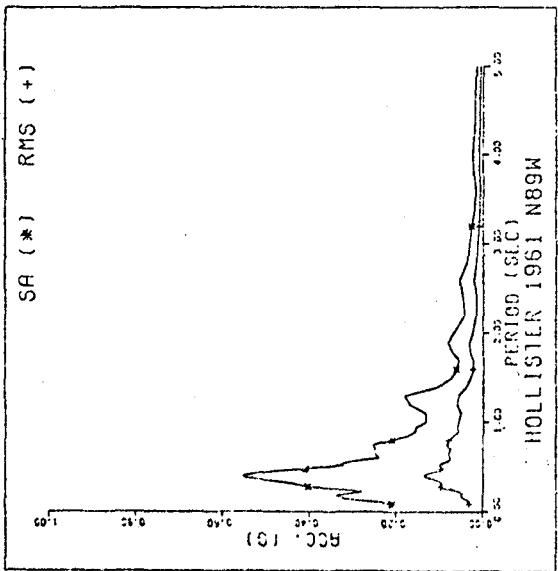
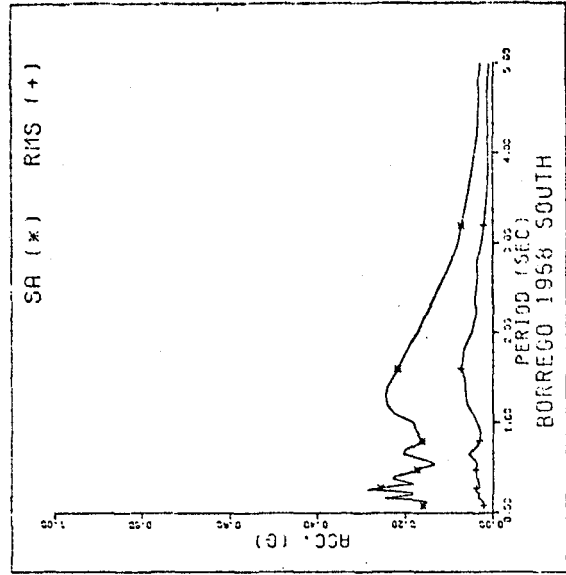
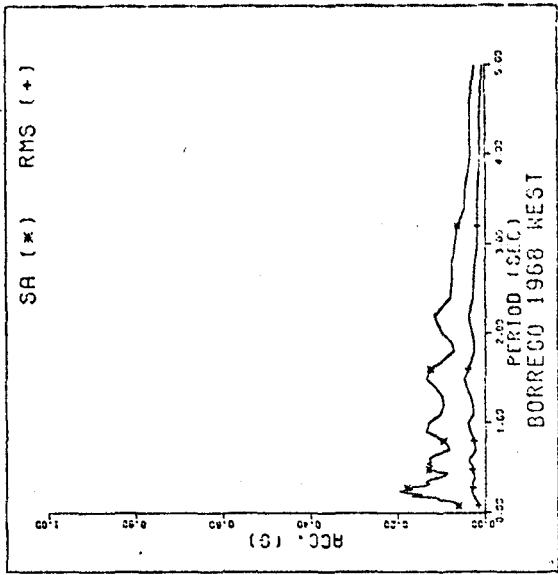


(A015)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)

(A014)



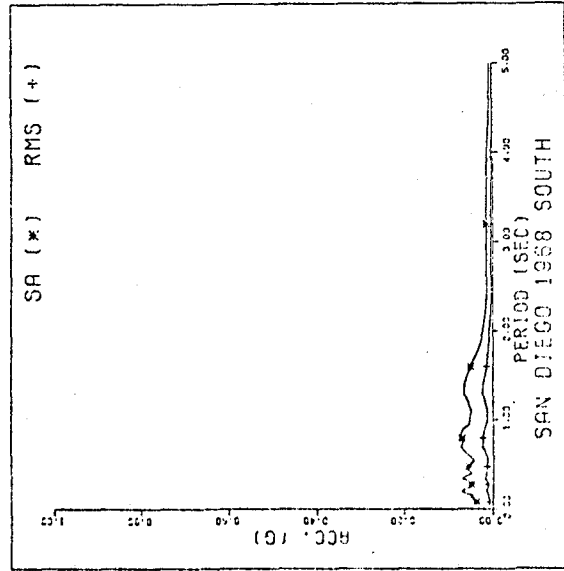
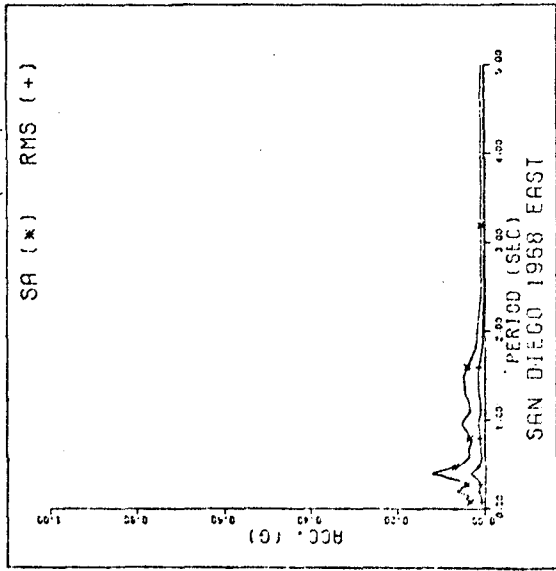
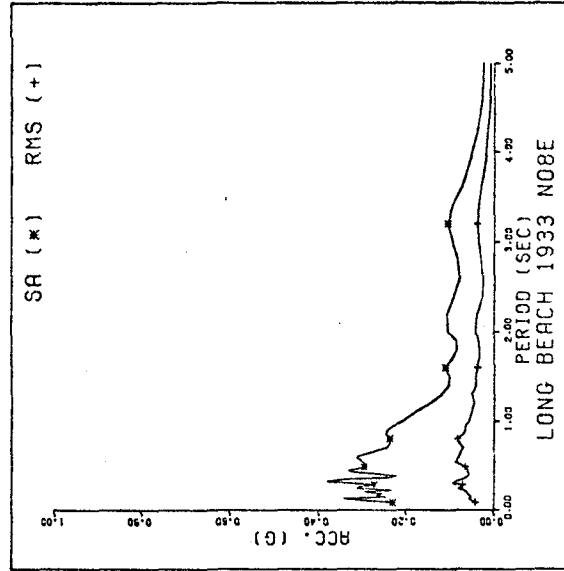
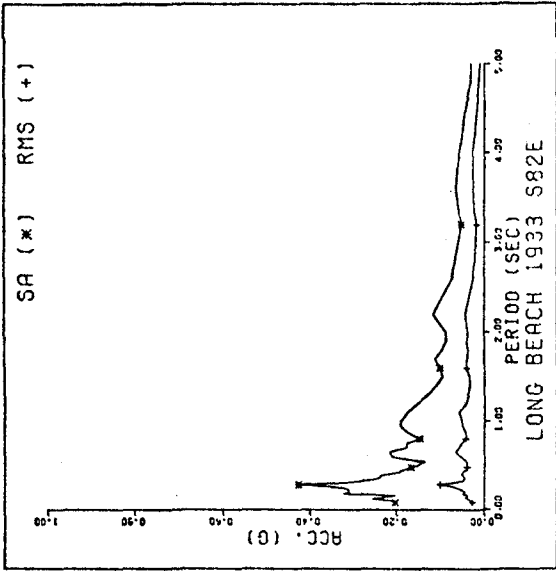
(A016) (A017)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



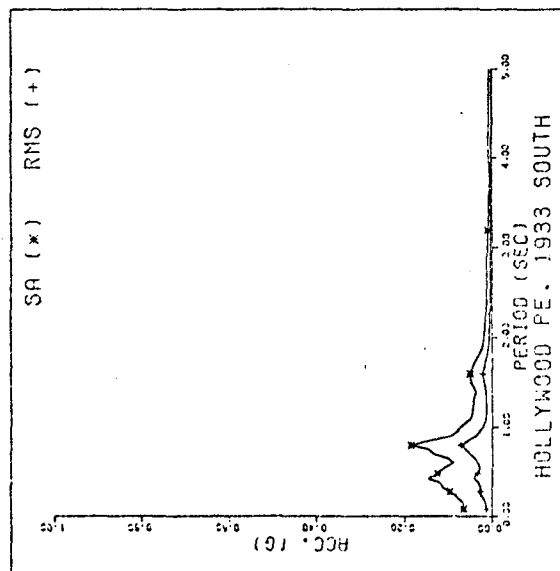
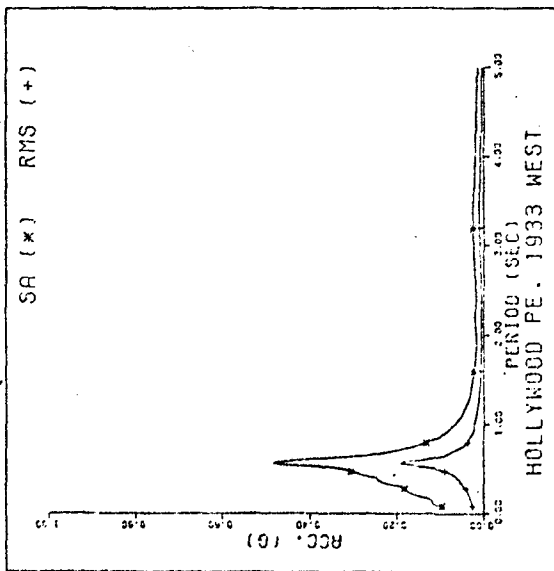
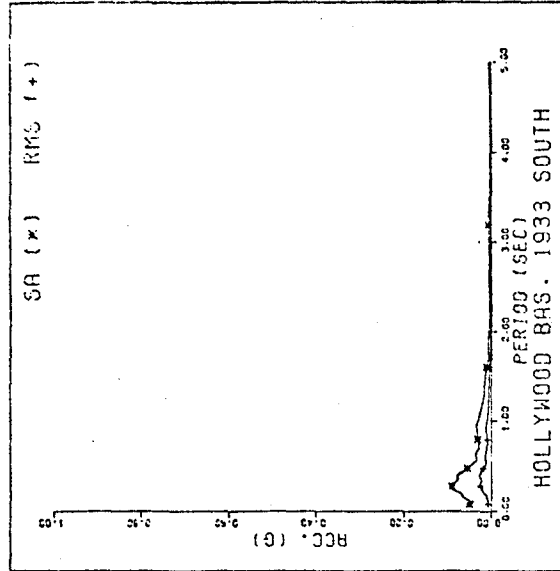
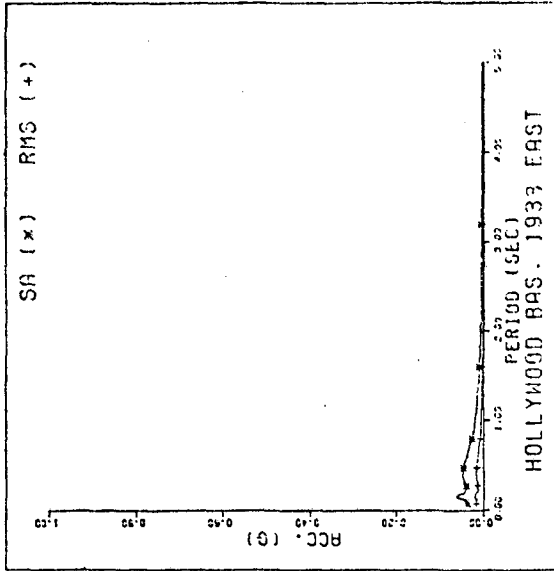
(A019)

(A018)

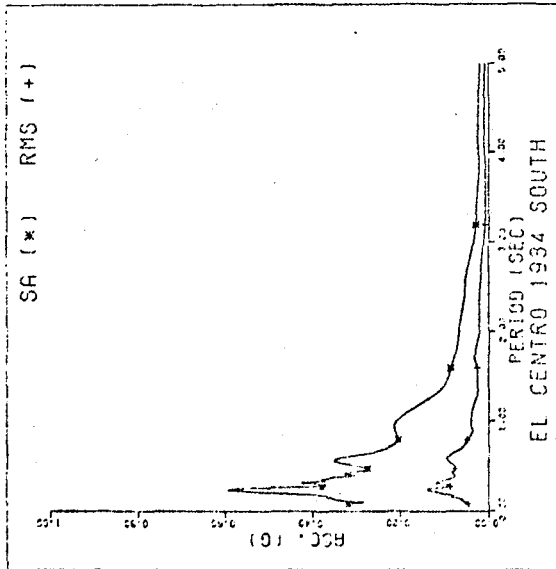
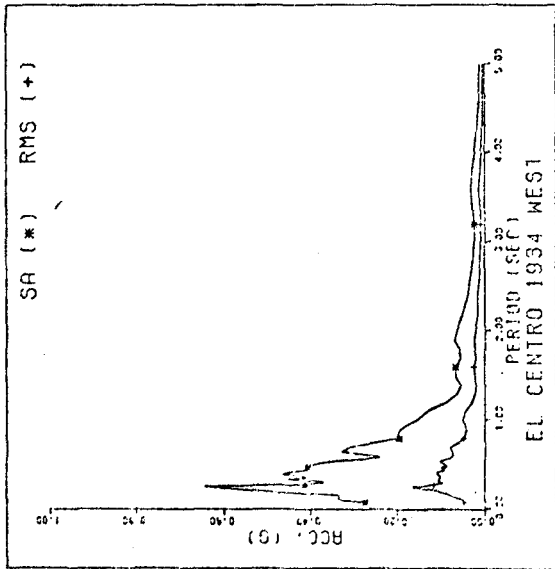
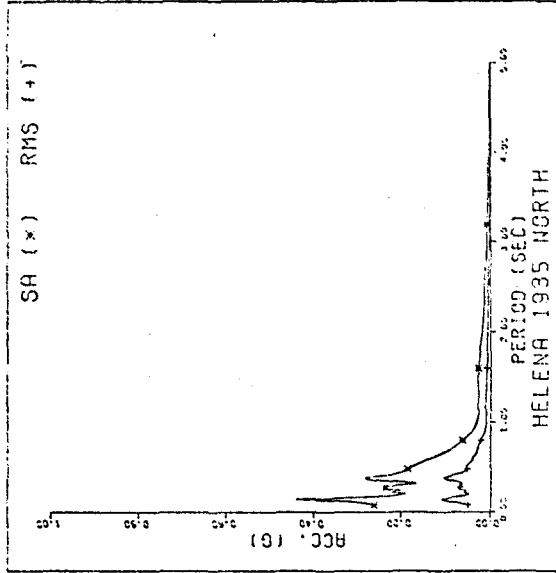
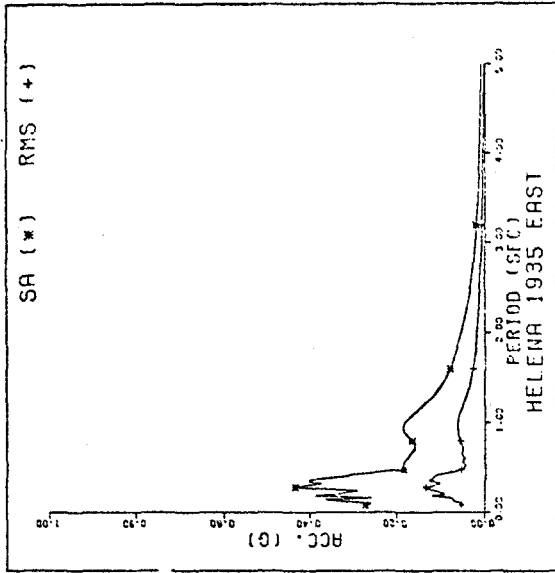
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



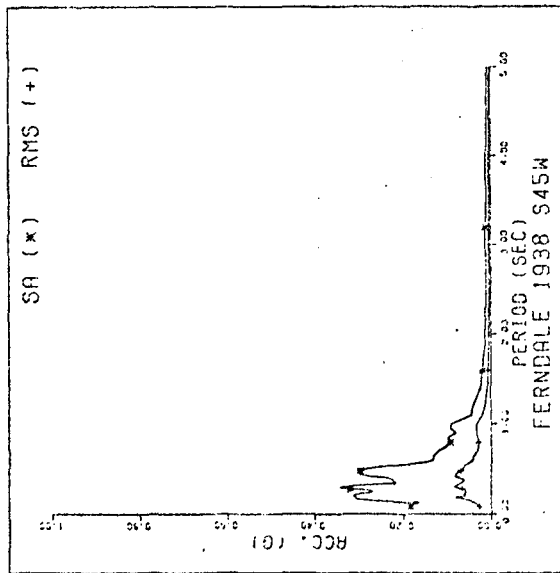
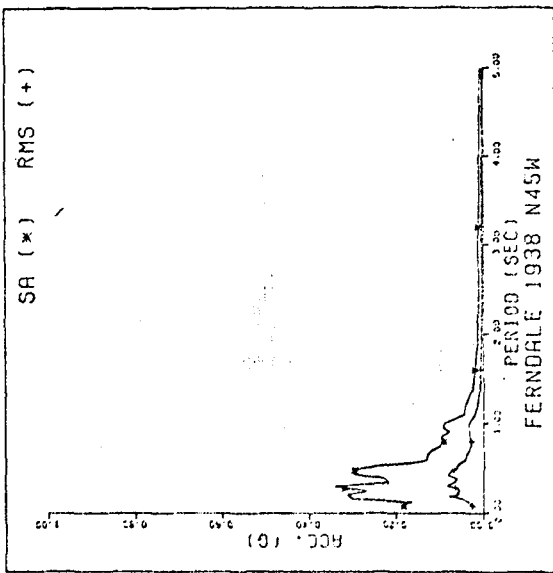
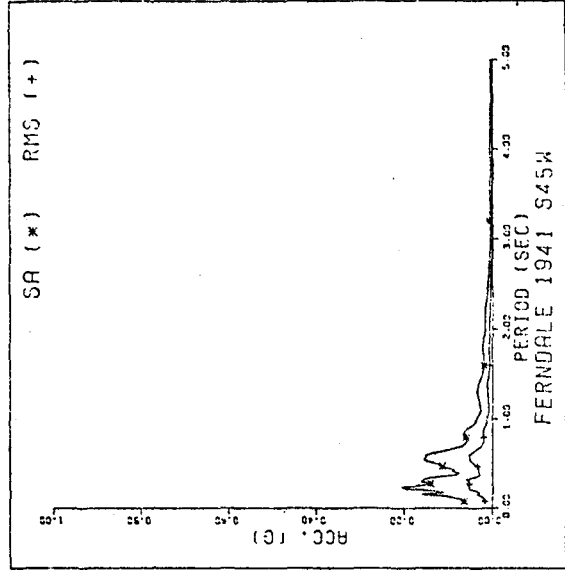
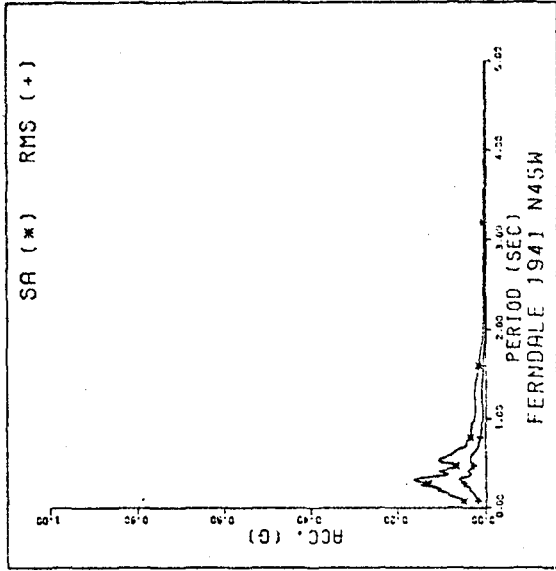
(A020) (B021)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



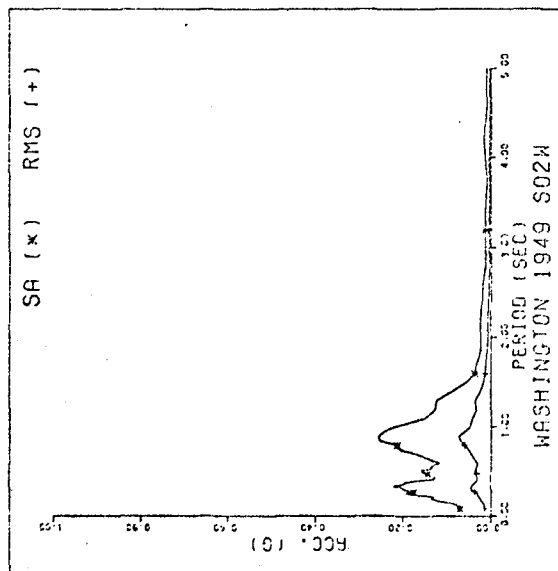
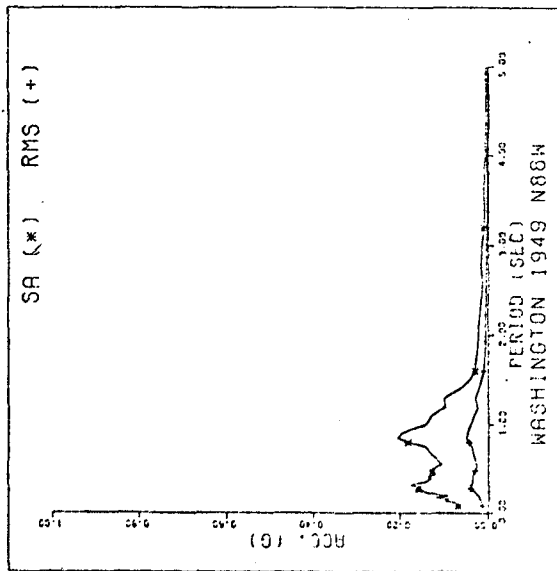
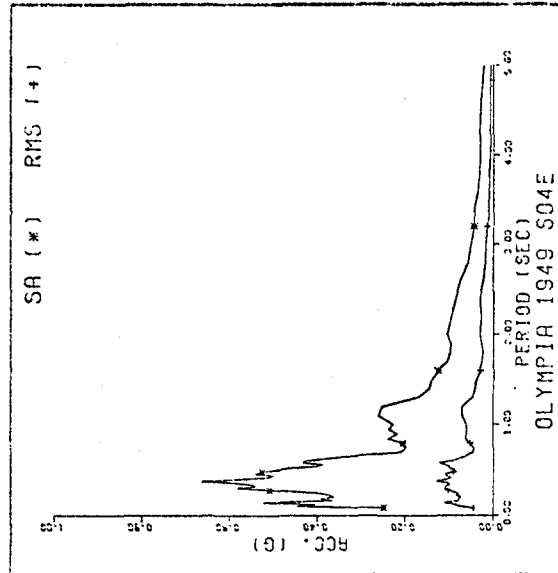
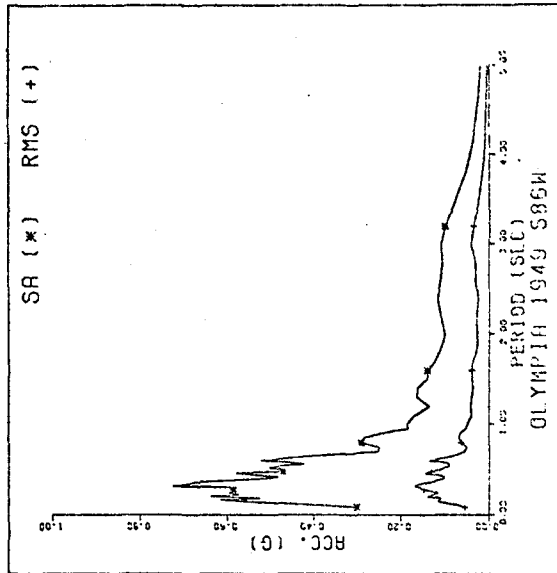
(B023)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)
(B022)



(B024)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)
(B025)



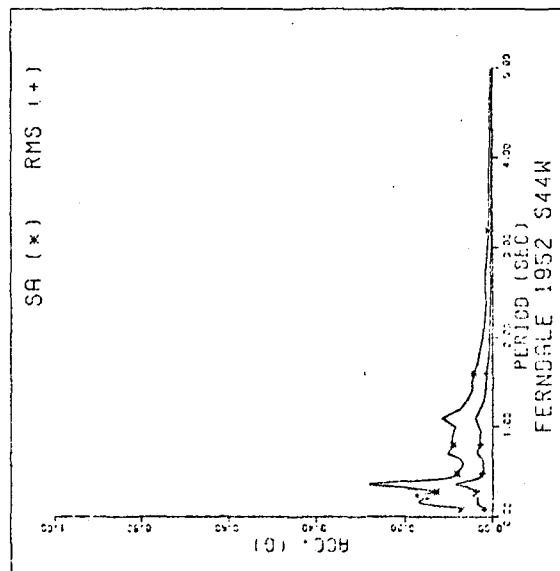
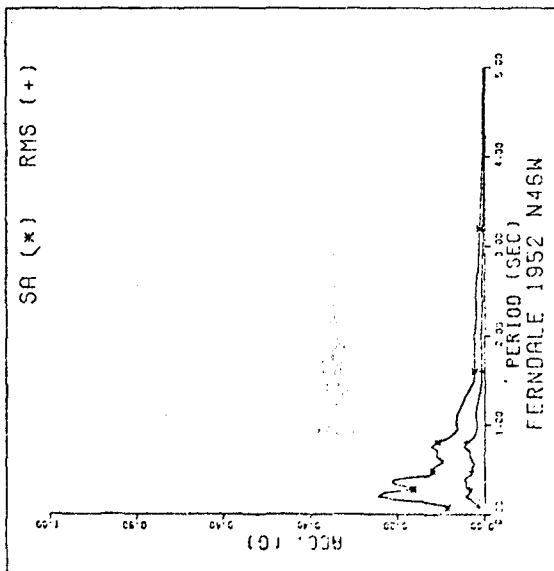
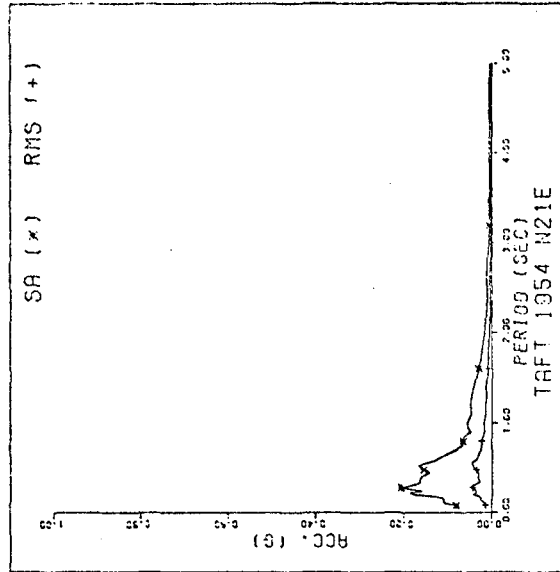
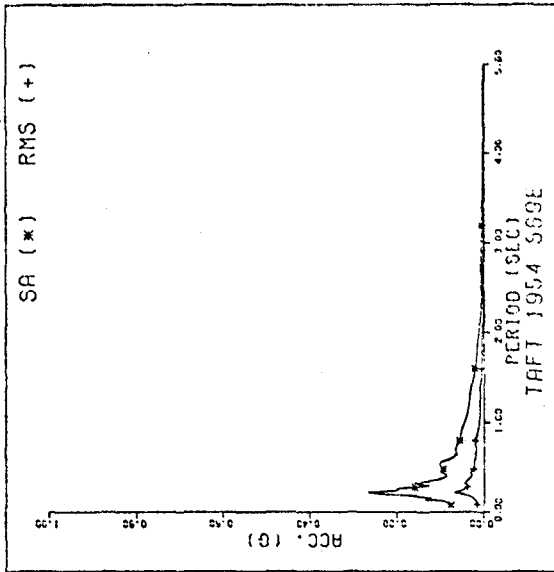
(B026) (B027)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



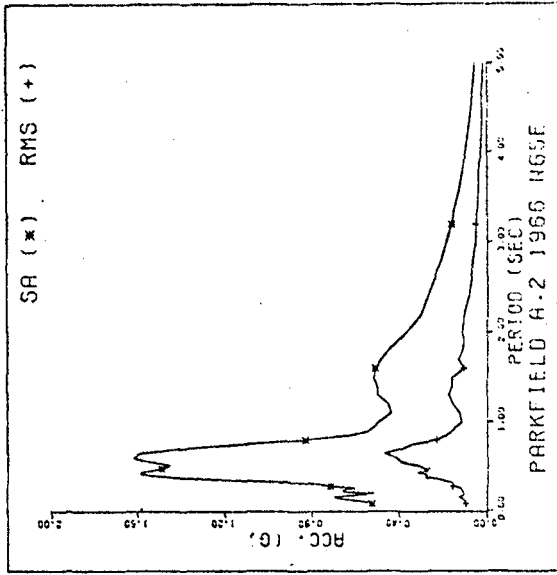
(B029)

FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)

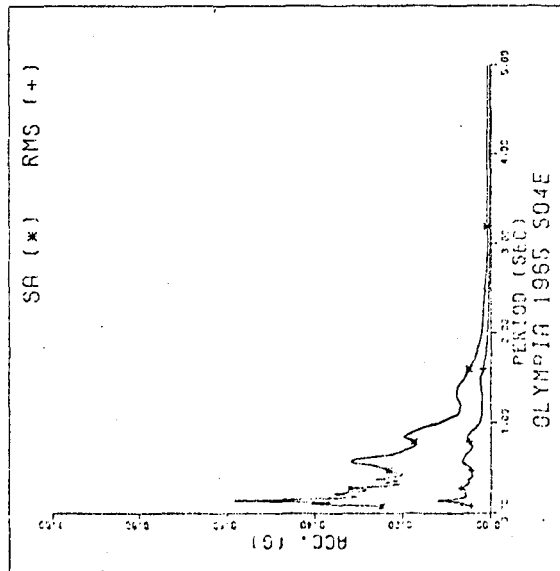
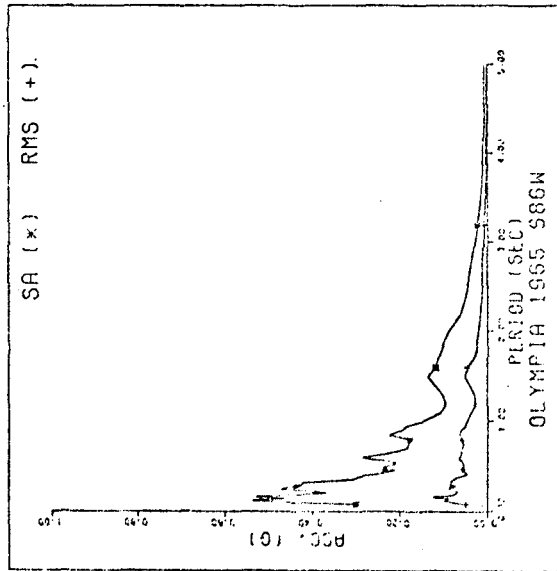
(B028)



(B030) (B031)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



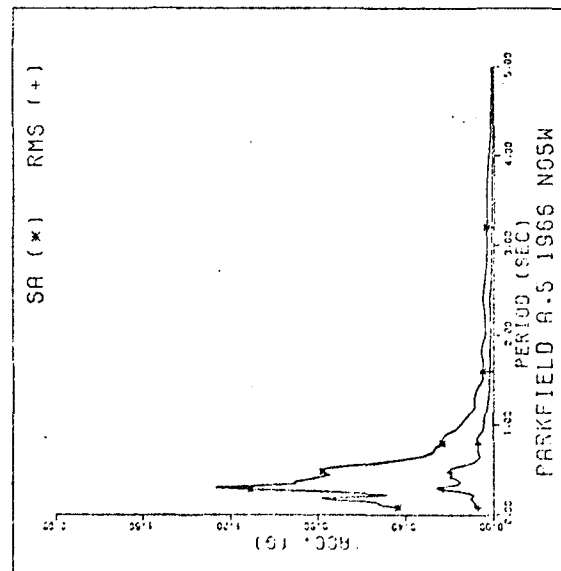
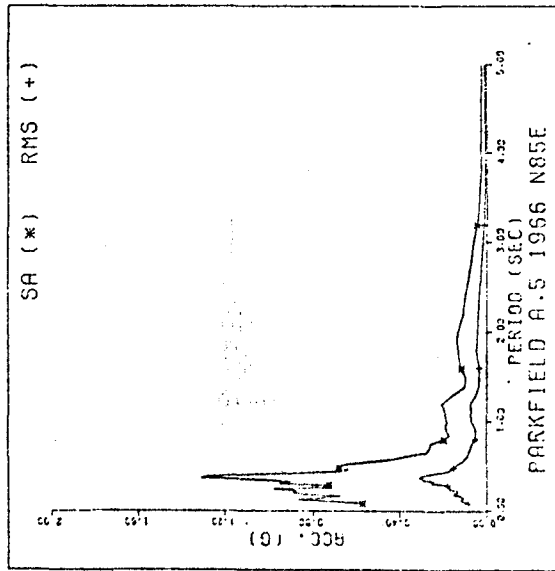
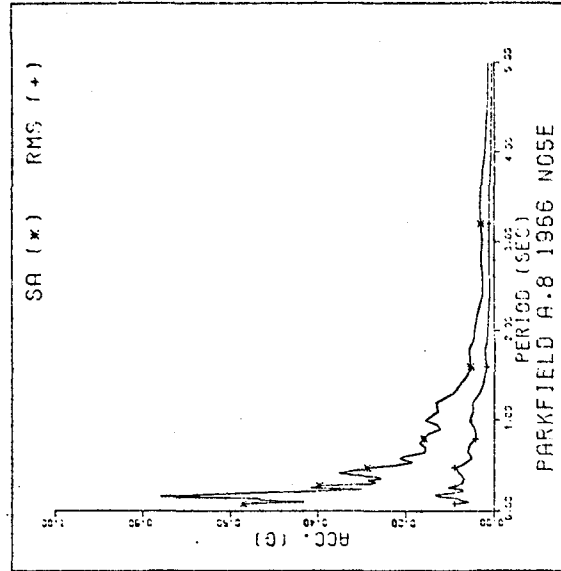
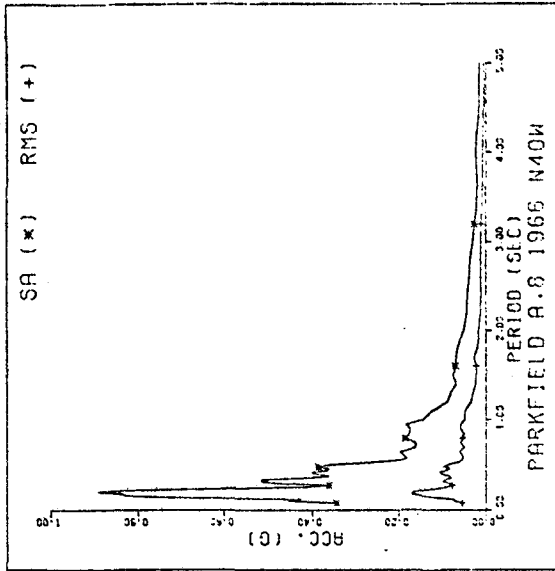
(B033)



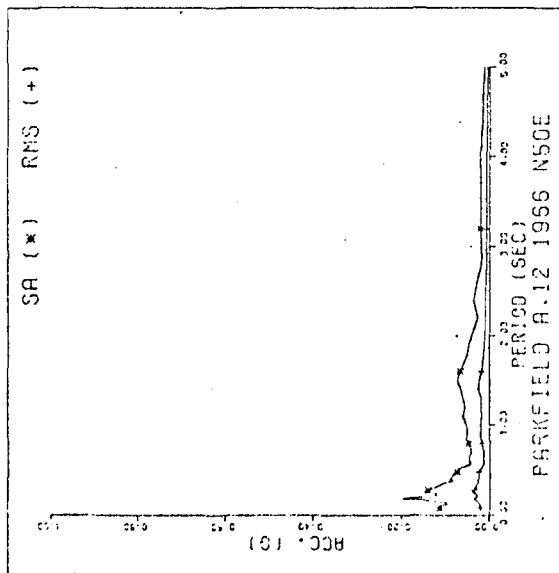
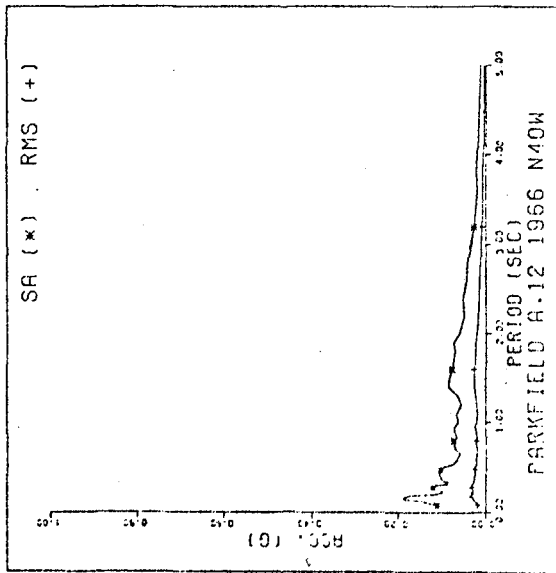
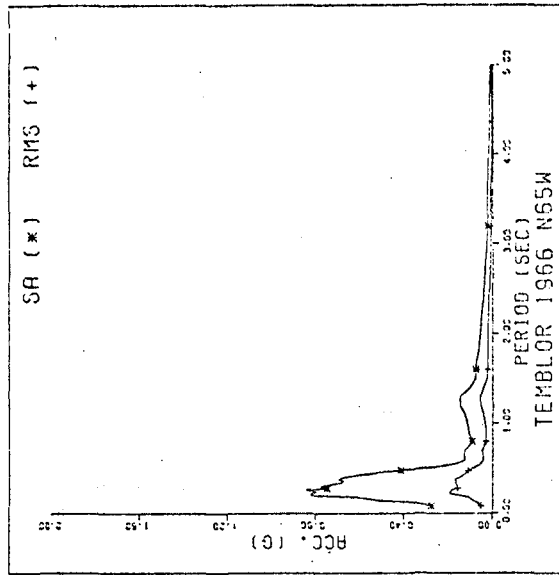
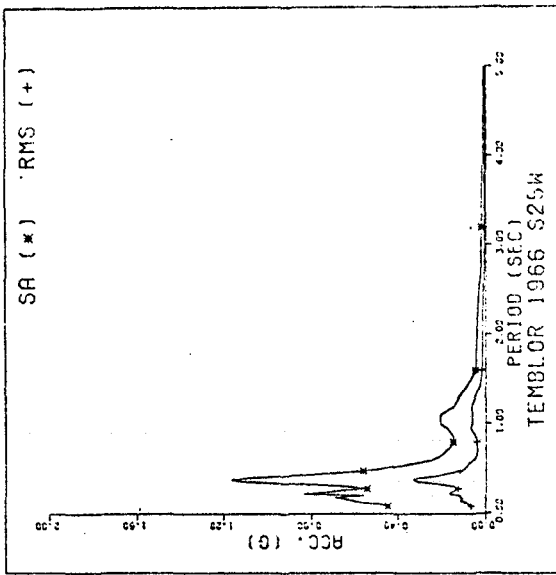
(B032)

FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)

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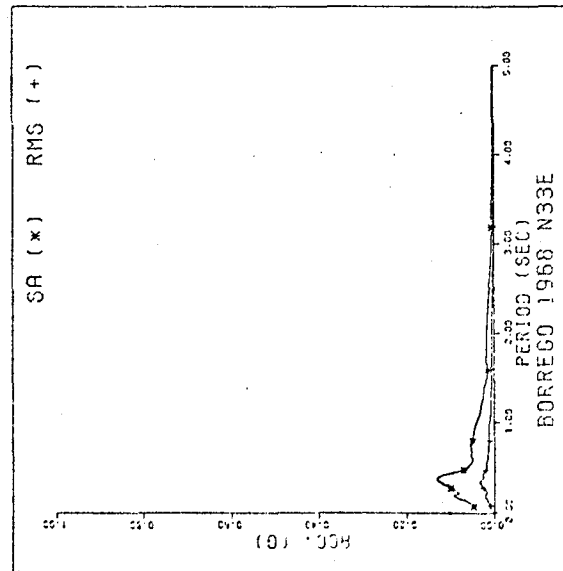
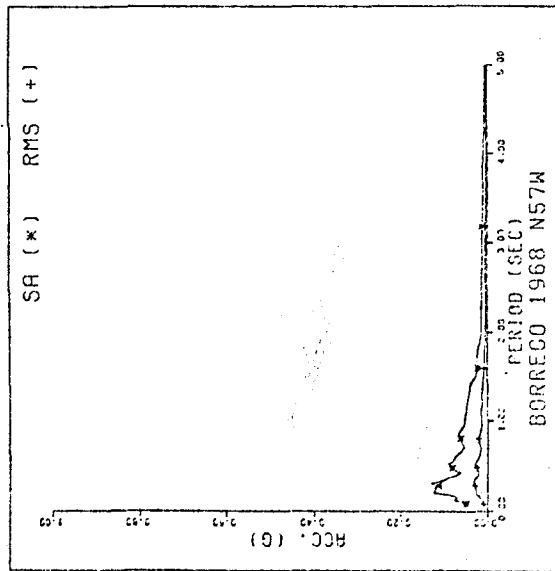
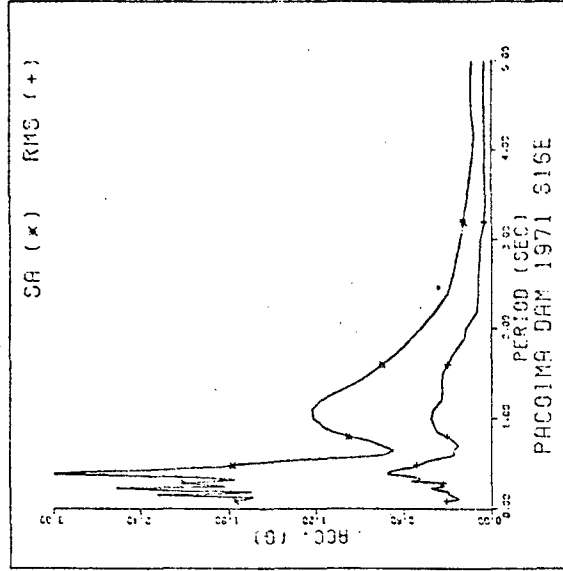
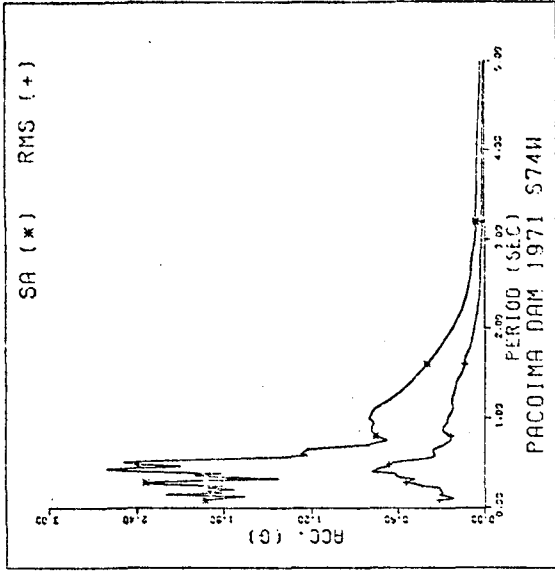
(B034)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)
(B035)



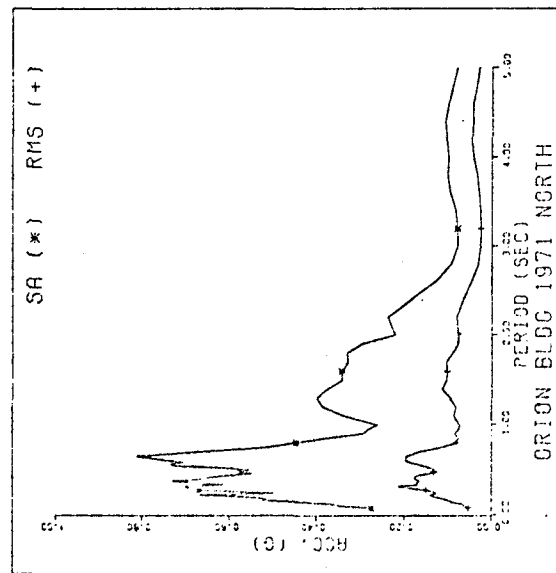
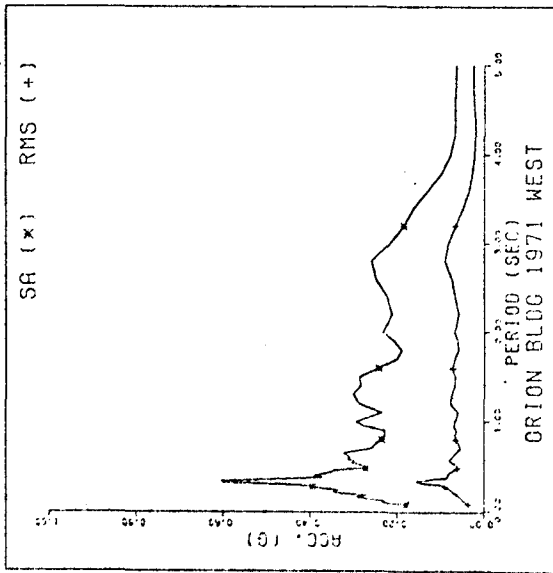
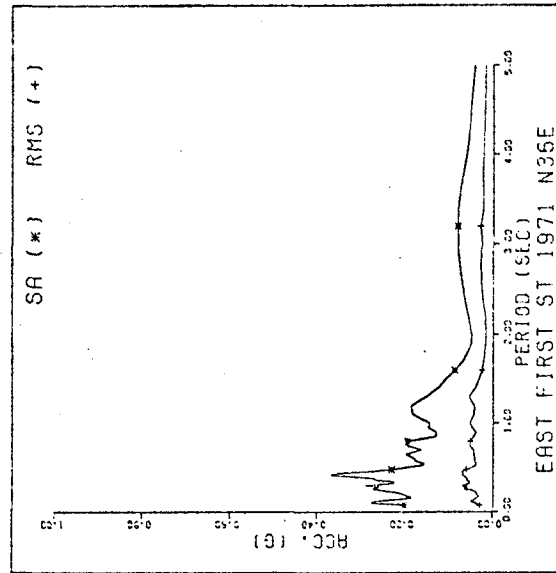
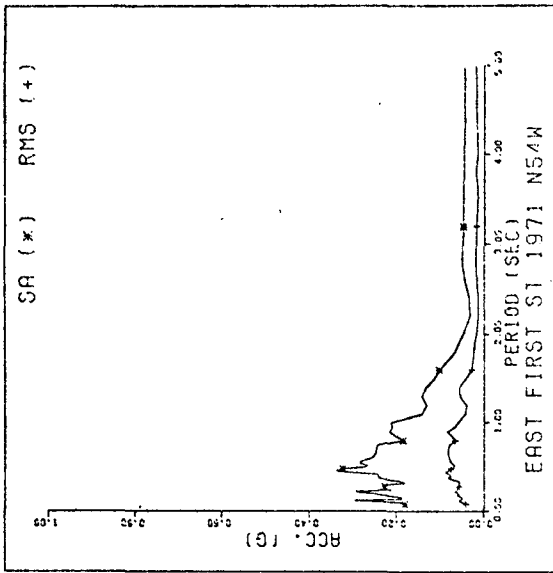
(B037)

FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)

(B036)



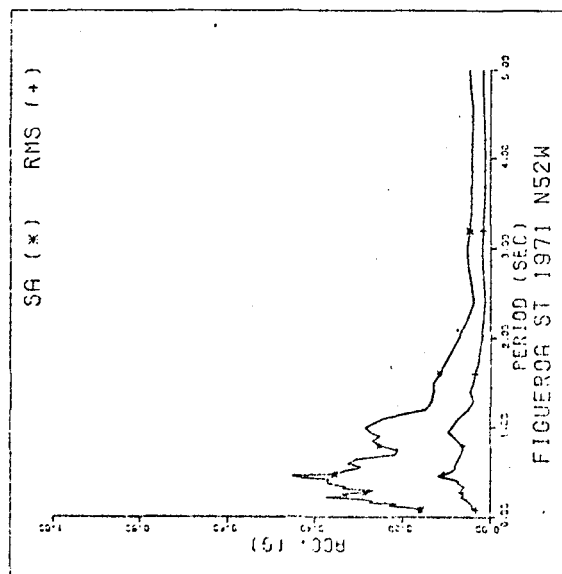
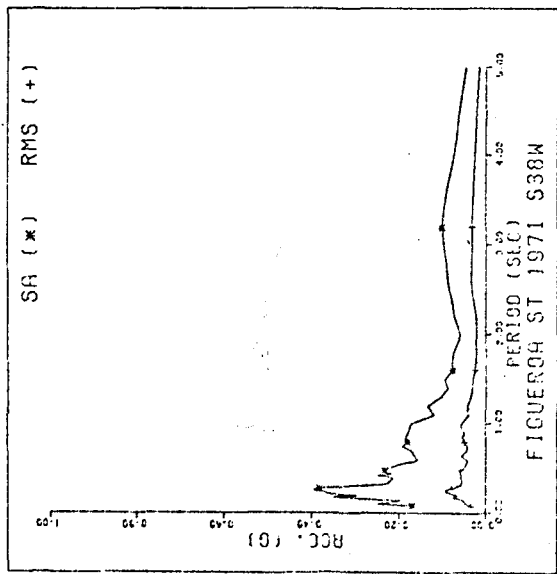
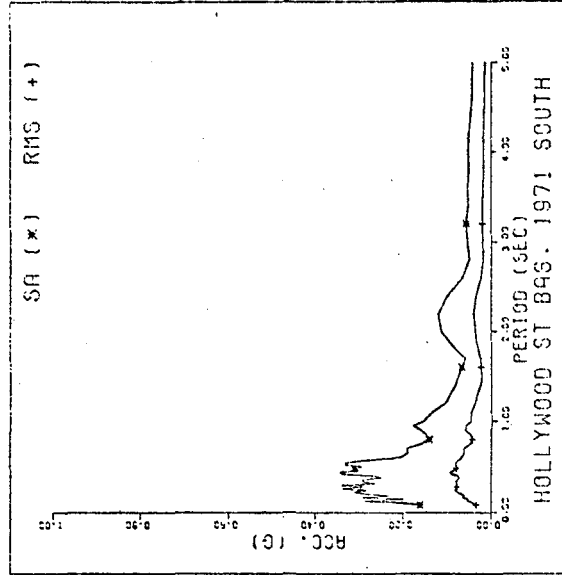
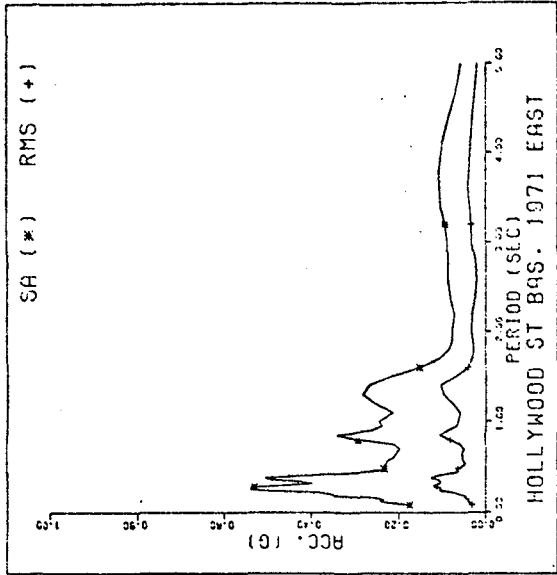
(B040) (C041) FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



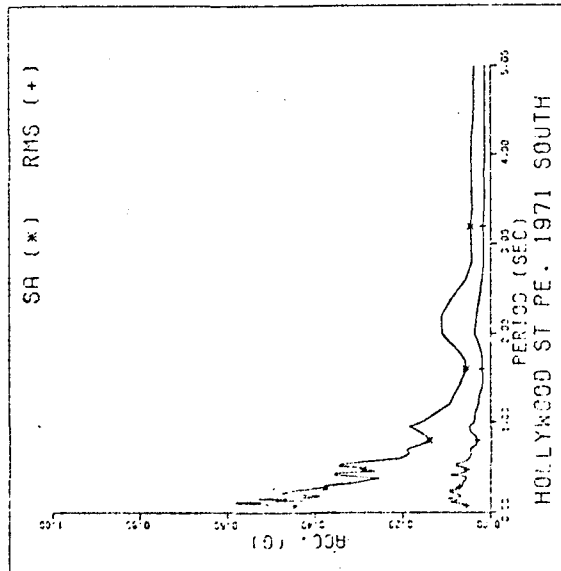
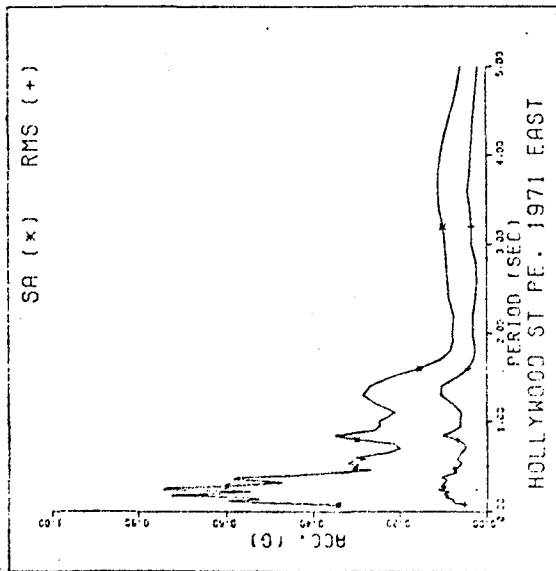
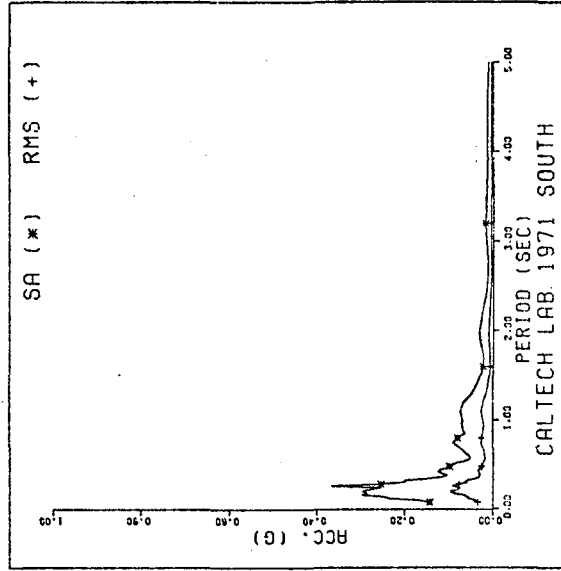
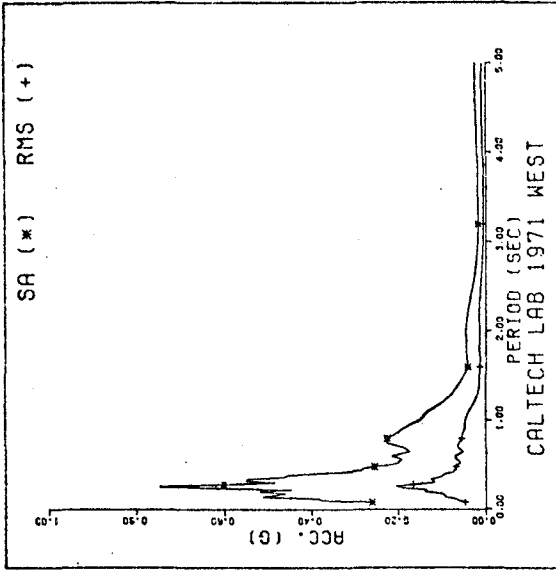
(C051)

FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)

(C048)



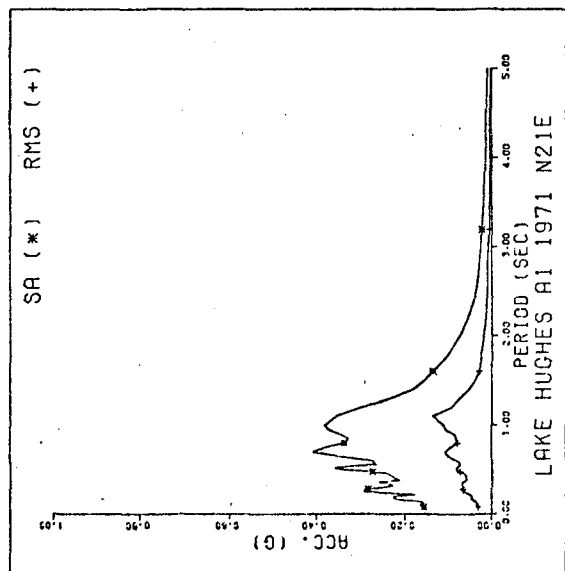
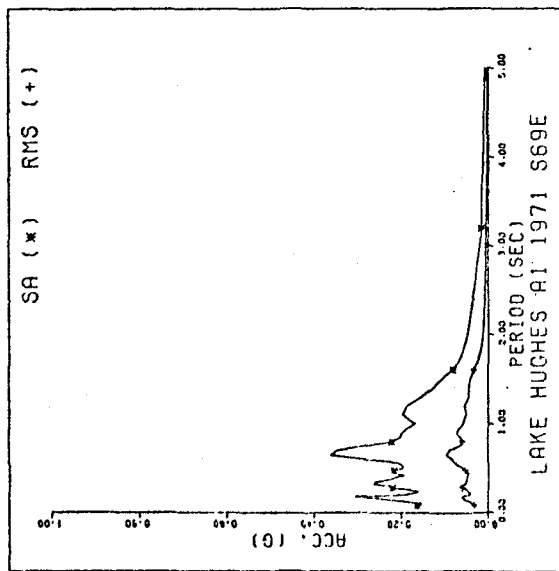
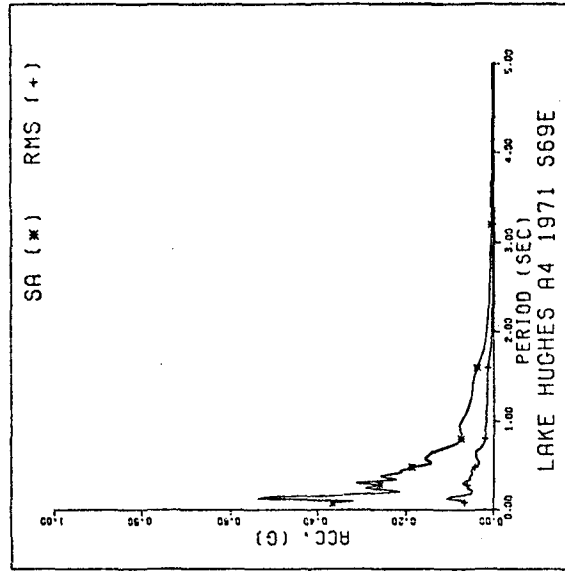
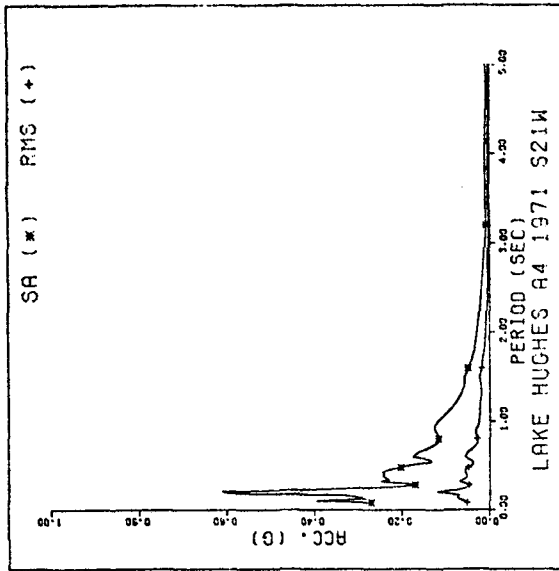
(D057)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)
(C054)



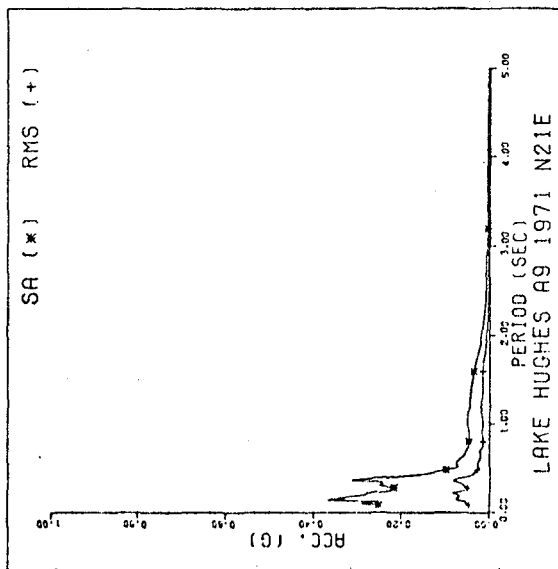
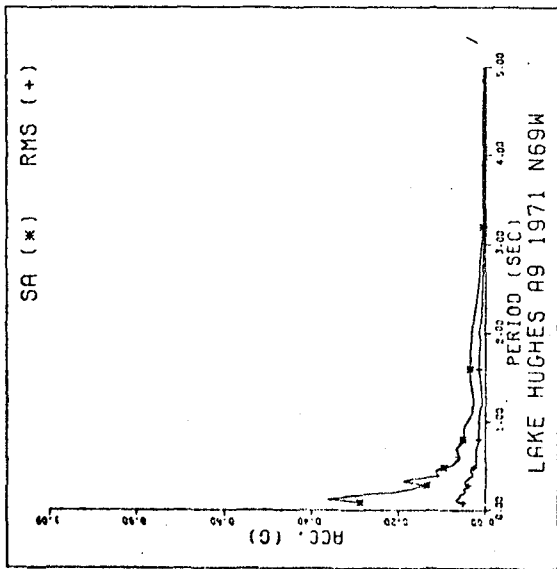
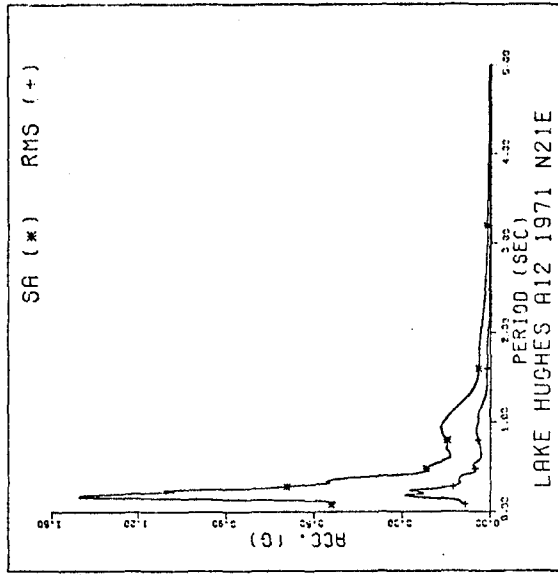
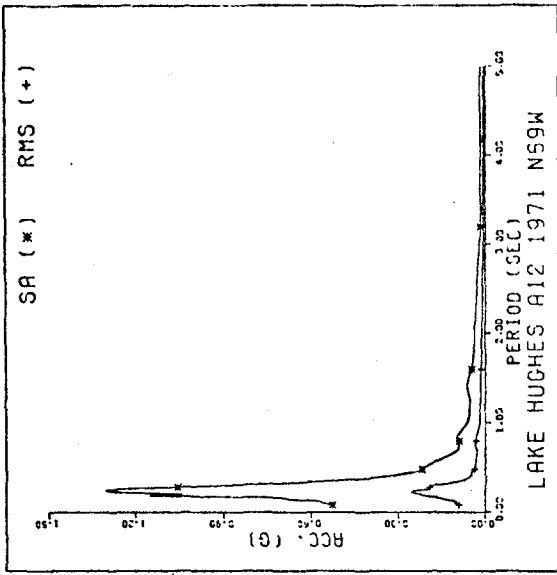
(G106)

FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)

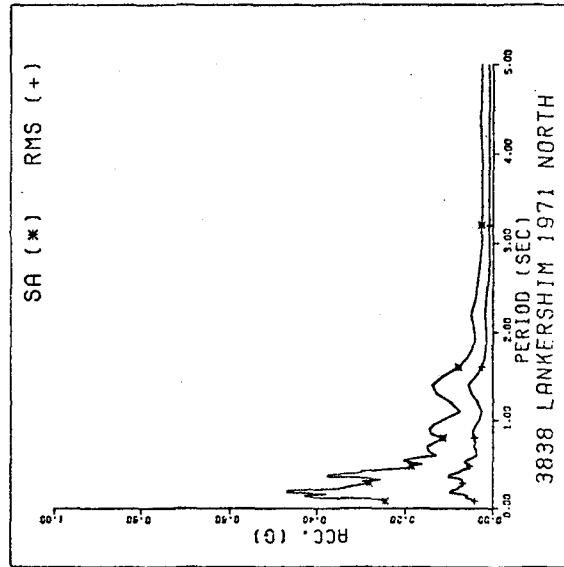
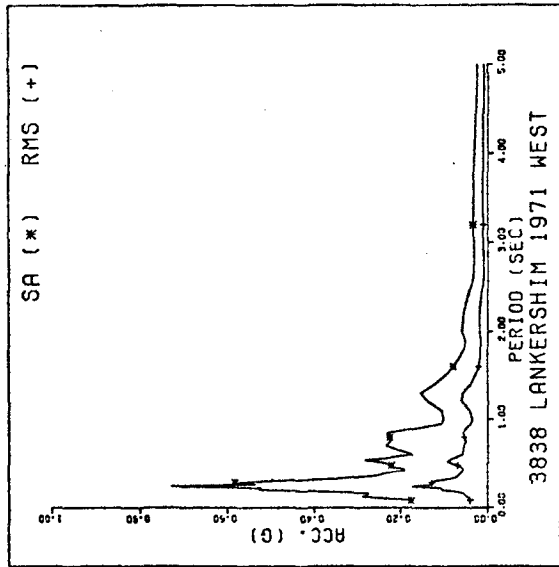
(D058)



(J141) (J142)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



(J143) (J144)
FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)



(L166)

FIGURE 4.4. Acceleration and RMS Response Spectra (Cont.)

the exponential distribution is used in this study. The probability distribution function for exponential shape is given by

$$f_A(a) = \lambda e^{-\lambda a} \quad (4.5)$$

where A = random variable defining the acceleration amplitude

a = the values A takes

λ = constant of the exponential distribution

It should be pointed out that once the distribution of the amplitudes is known, any probability statement regarding the response acceleration can be made. As an example

$$\begin{aligned} P(\text{peak amplitude } A \geq a) &= \int_a^{\infty} \lambda e^{-\lambda A} dA \\ &= e^{-\lambda a} \end{aligned} \quad (4.6)$$

This information is extensively used in determining a_p .

- Mean response acceleration (Column 6: λ)

This parameter is the average of the acceleration throughout the record. It shows the same stability as the RMS.

It is equal to $1/\lambda$ where λ is the parameter of the exponential distribution.

- a_p (Column 7: $p = 5\%$; Column 8: $p = 10\%$)

This variable represents the response acceleration which has a probability p of exceedence for a given acceleration response (Fig. 4.2).

- ENGY (Column 9)

The cumulative potential energy stored at any time in the spring of the single degree of freedom system is computed

for the duration of the response. Consider the single degree of freedom system shown in Figure 4.5. The motion of the support is known in terms of its acceleration \ddot{x}_0 (earthquake record) and the quantity of interest is the relative displacement $y = x_1 - x_0$ (Biggs, 1964). The total energy in the system is equal to the kinetic energy of the mass plus the potential energy stored in the spring. During a free motion of this system, the potential energy is at times stored in the spring. While this energy is stored, the kinetic energy of the vibrating mass decreases. Thus, when there is maximum potential energy in the spring, the kinetic energy of the mass is zero. Conversely, when the kinetic energy is maximum the potential energy in the spring

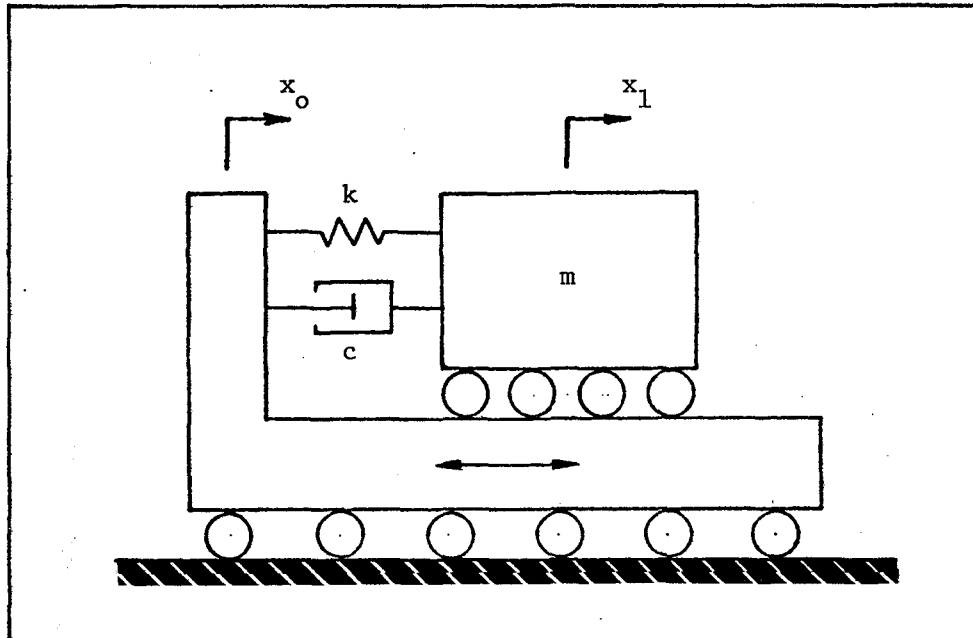


FIGURE 4.5. Single Degree of Freedom System. Mass m Elastically Supported from a Moving Foundation with Viscous Damping

is zero. At any time one could calculate the kinetic energy of the mass and the potential energy in the spring. In addition to the above energy transfer, for a forced vibration (earthquake motion), there is a continuous addition of energy from the ground to the vibratory system. In this process, part of the energy is dissipated in the dashpot. The amount of dissipation is a function of the damping of the system. For any time increment Δt the energy per unit mass stored in the spring is equal to the relative displacement (Δy) multiplied by the absolute acceleration felt in the spring (\ddot{x}_1). The cumulative potential energy per unit mass (ENGY) stored in the spring is therefore

$$\text{ENGY} = \sum_{\text{all } \Delta t} \Delta y \cdot \ddot{x}_1, \quad \text{for } \Delta y \cdot \ddot{x}_1 > 0 \quad (4.7)$$

It seems that this parameter should be highly correlated to the damage of the vibratory system. It is also reasonable to assume that a design based on energy criteria is more rational than the current peak amplitude based procedure. The energy parameter is based explicitly on the acceleration level at each time increment and also on the duration of the duration of the response.

- Ratio $K_1 = a_p/\text{RMS}$ (Column 10: $p = 5\%$; Column 11: $p = 10\%$)

This ratio is one of the parameters investigated in this study for stable behavior. It can be seen that for all the records that are analyzed, its value is almost constant.

Further discussions on this parameter are presented in Chapter V.

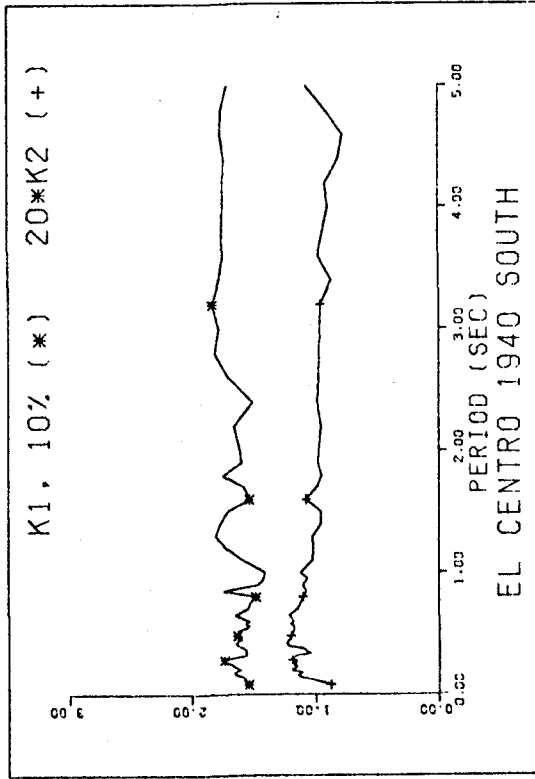
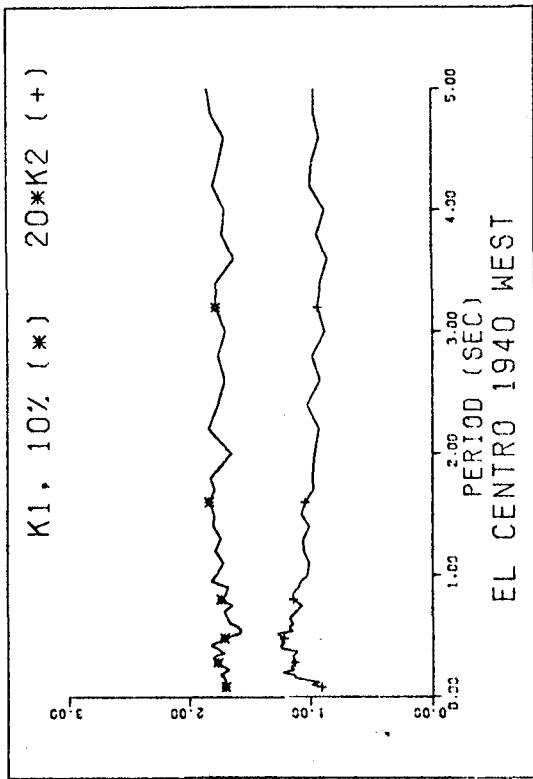
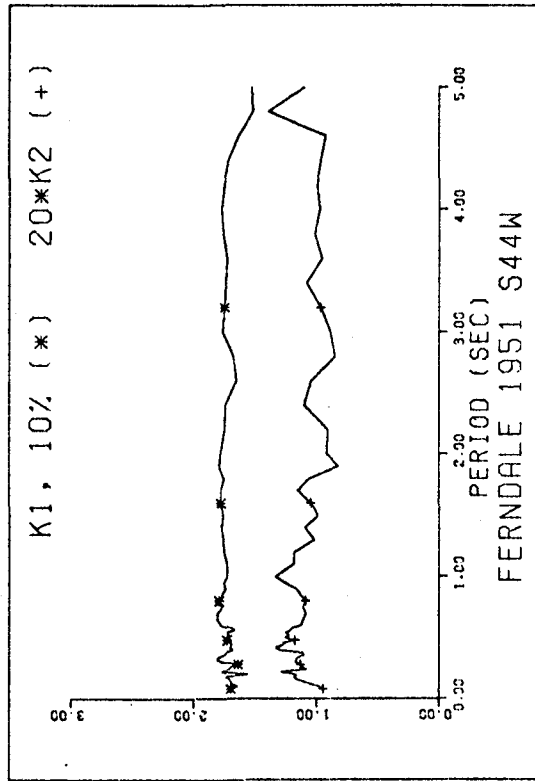
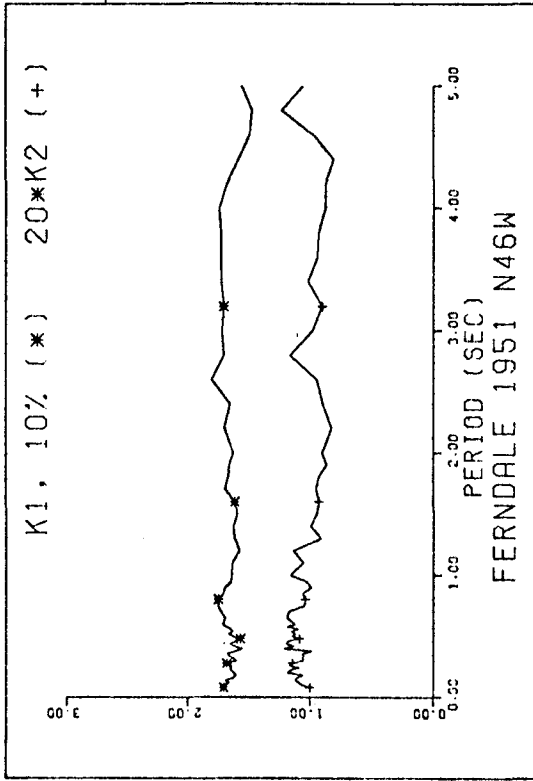
- Ratio $K_2 = \text{RMS}^2 \cdot T^2 / (\text{ENGY}/\text{NBPK})$ (Column 12)

Again, the stable behavior of this parameter prompted this study to evaluate it for all the available earthquake records. A detailed discussion will follow in Chapter V.

Both parameters K_1 ($p = 10\%$) and K_2 are plotted in Fig. 4.6.

The reason for developing the parameters presented in Appendix B are:

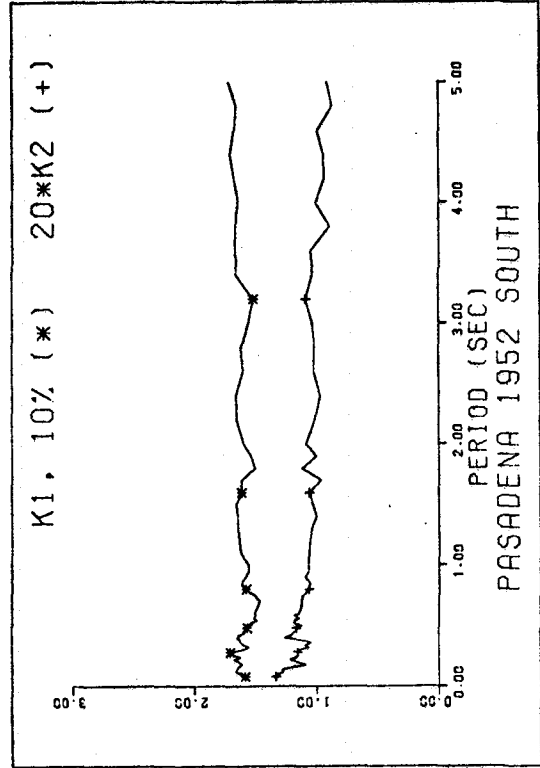
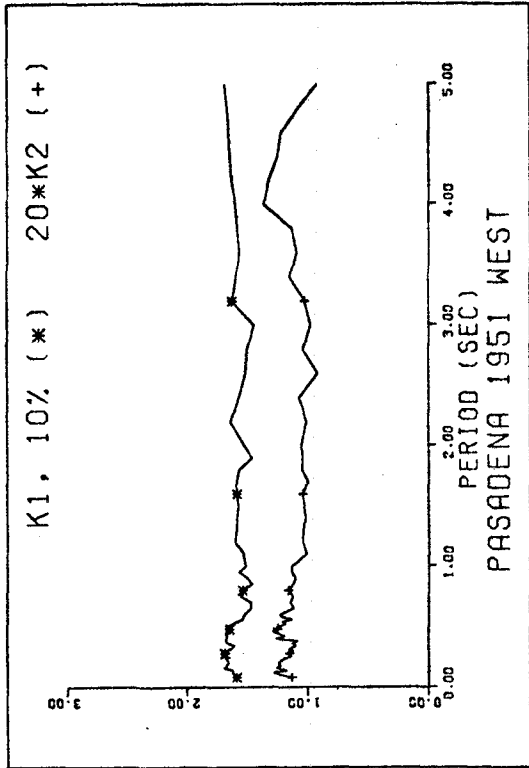
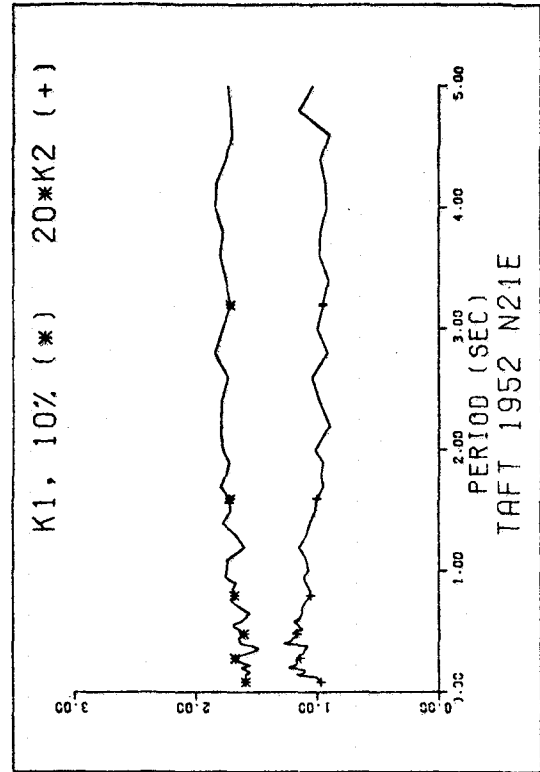
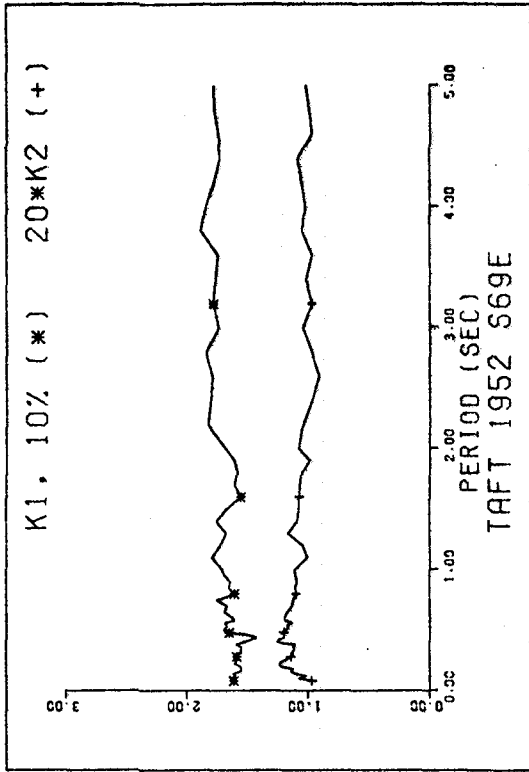
- To numerically summarize the relevant parameters describing the response history
- To look at the behavior of the summary parameters for trends and similarities
- To obtain a stable input parameter which could be used for a better description of the response history.



(A002)

FIGURE 4.6. Factors K_1 and K_2

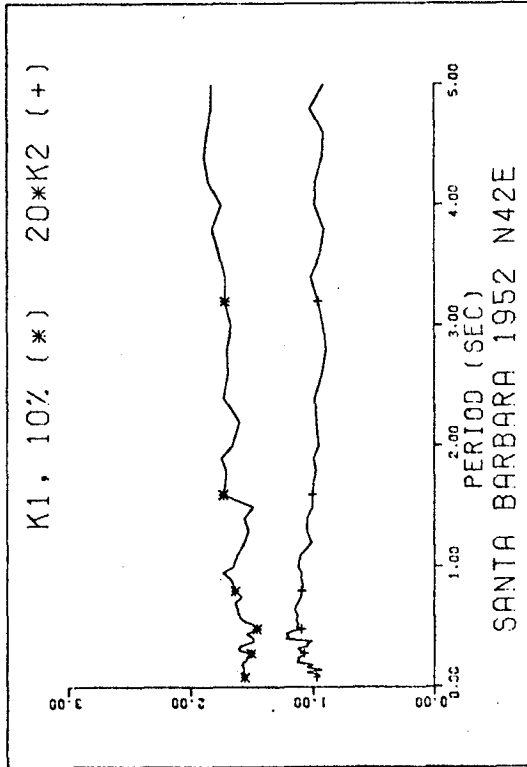
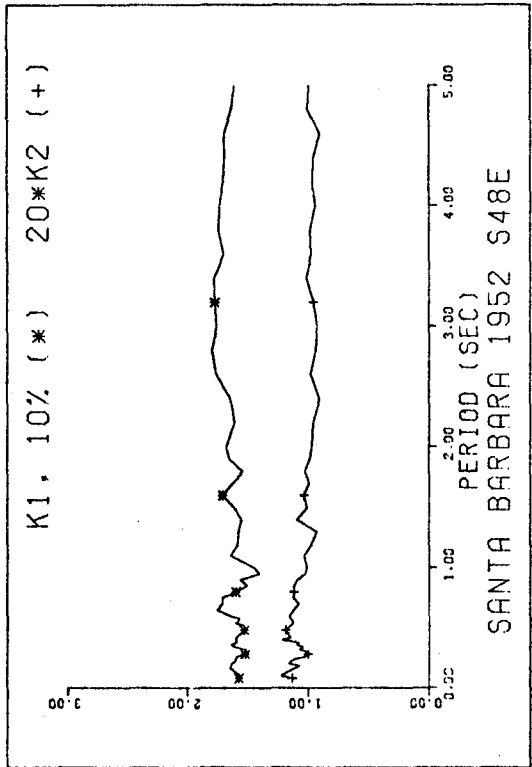
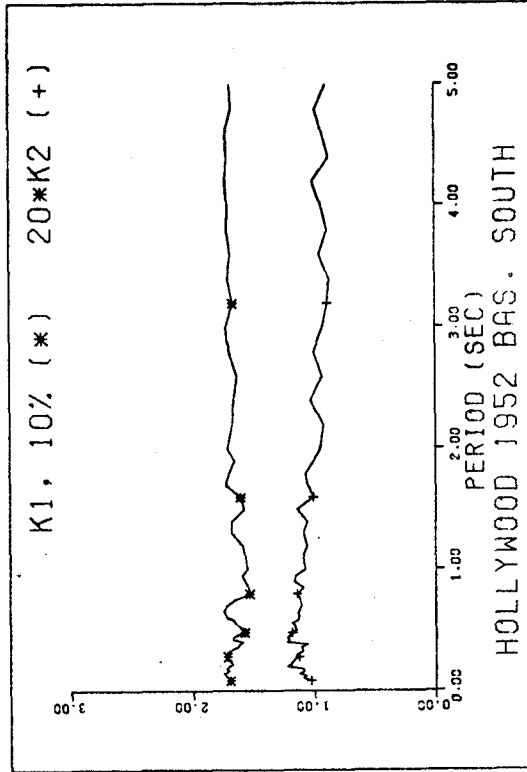
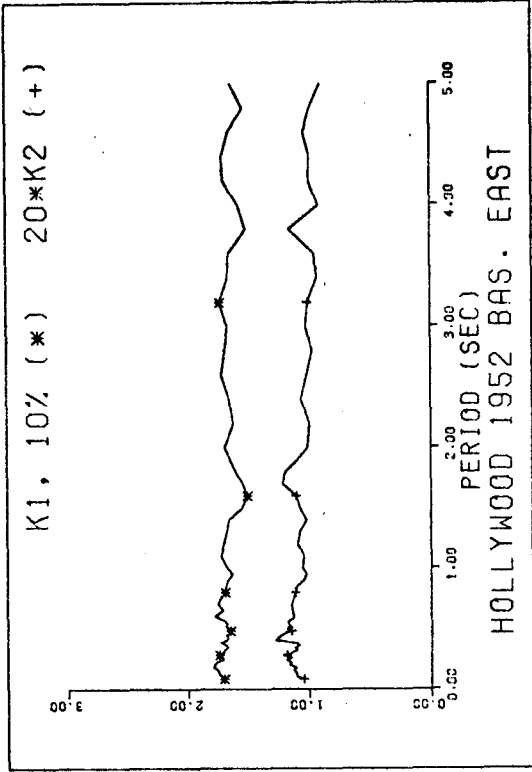
(A001)



(A004)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

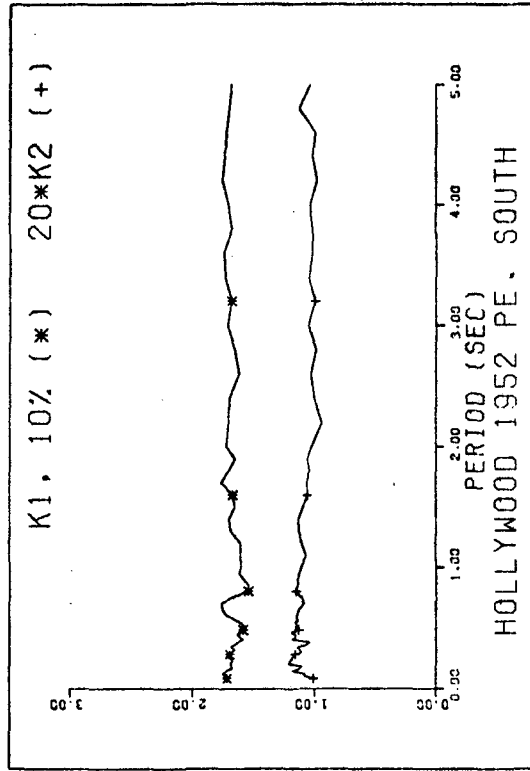
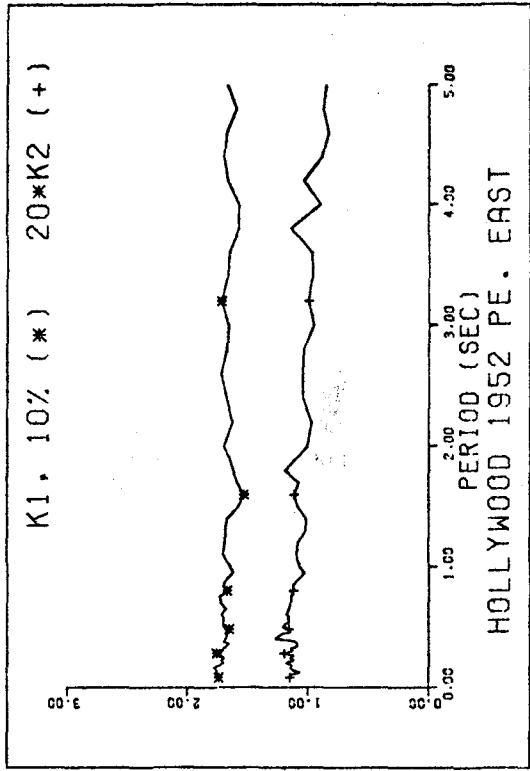
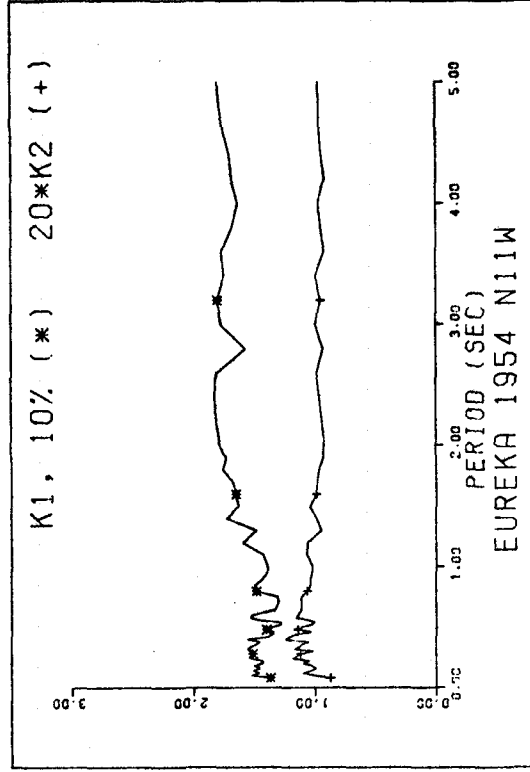
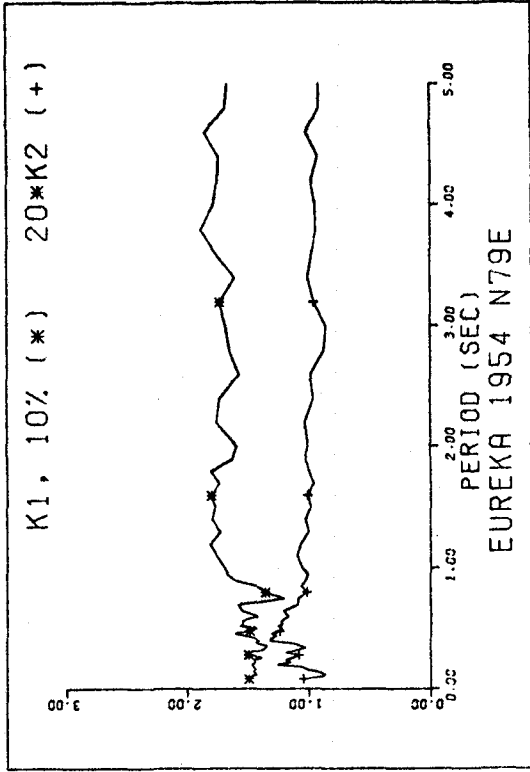
(A003)



(A006)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

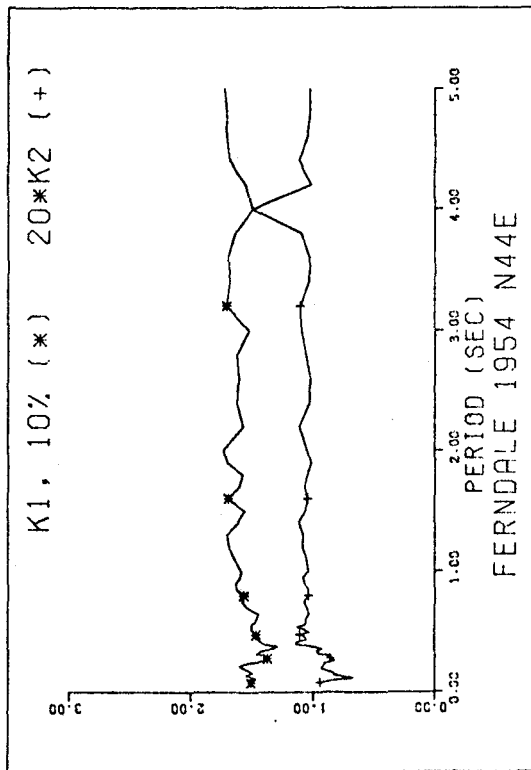
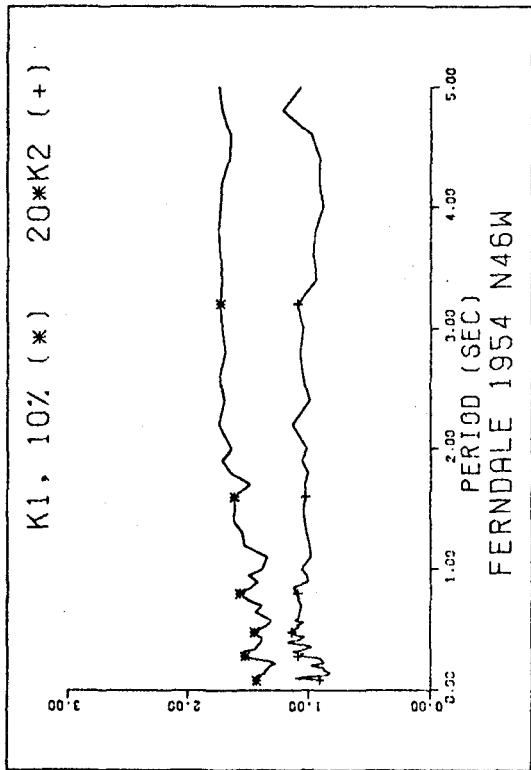
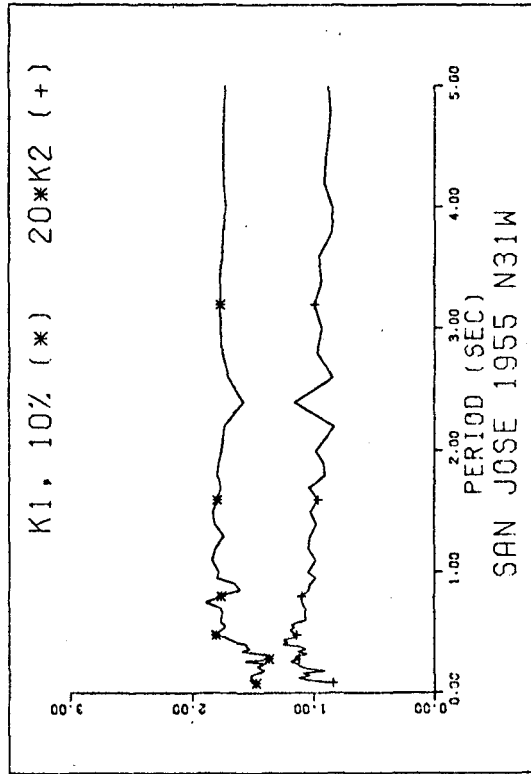
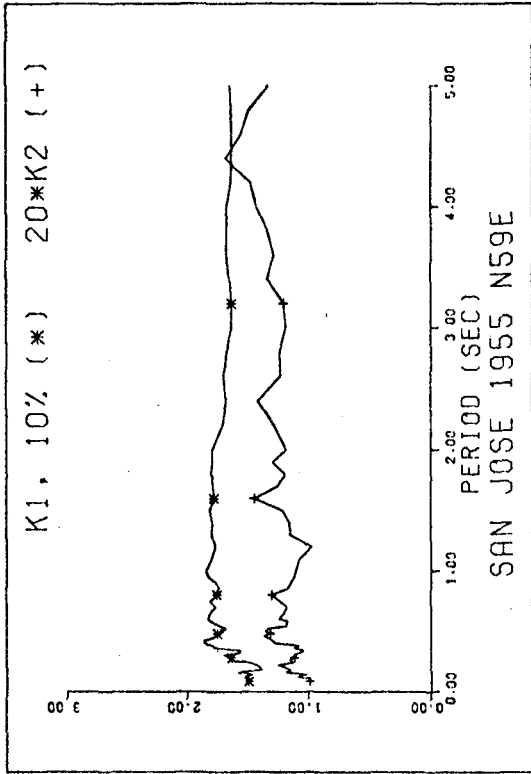
(A005)



(A008)

(A007)

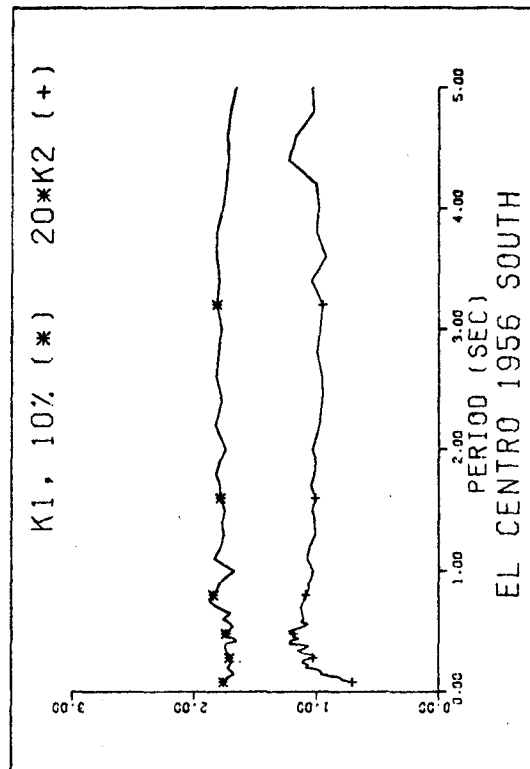
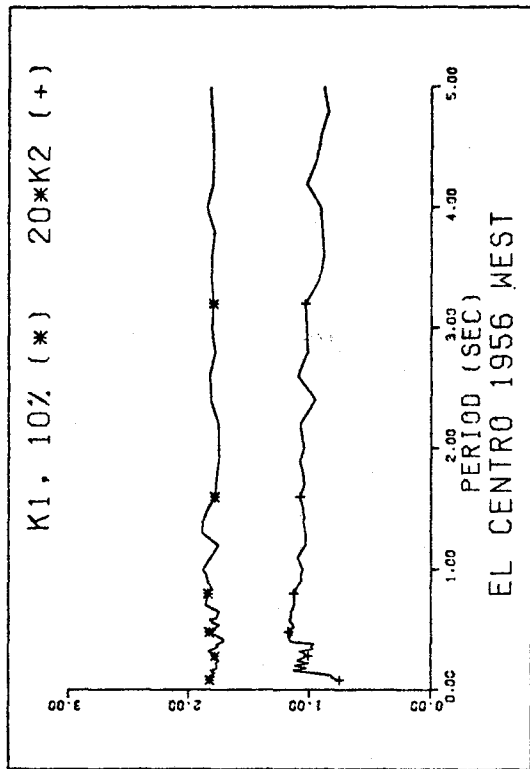
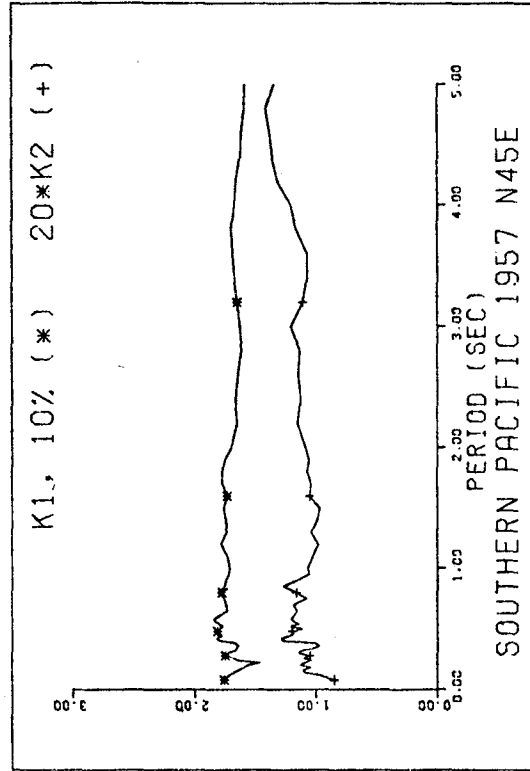
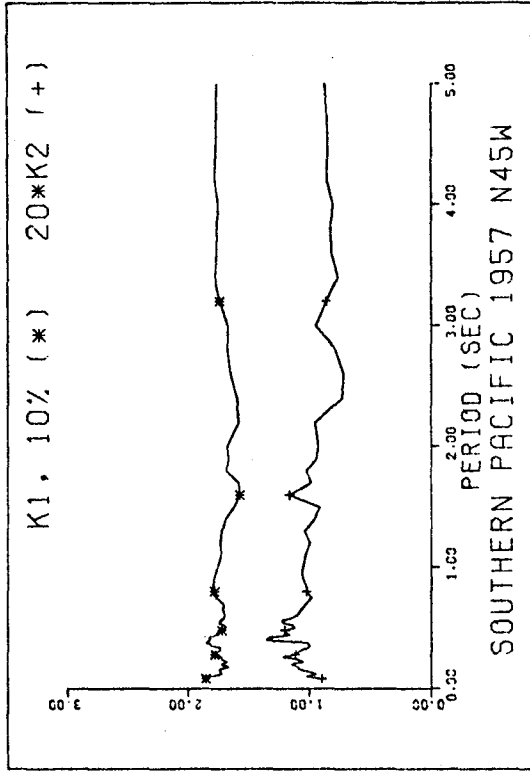
FIGURE 4.6. Factors K_1 and K_2 (Cont.)



(A010)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

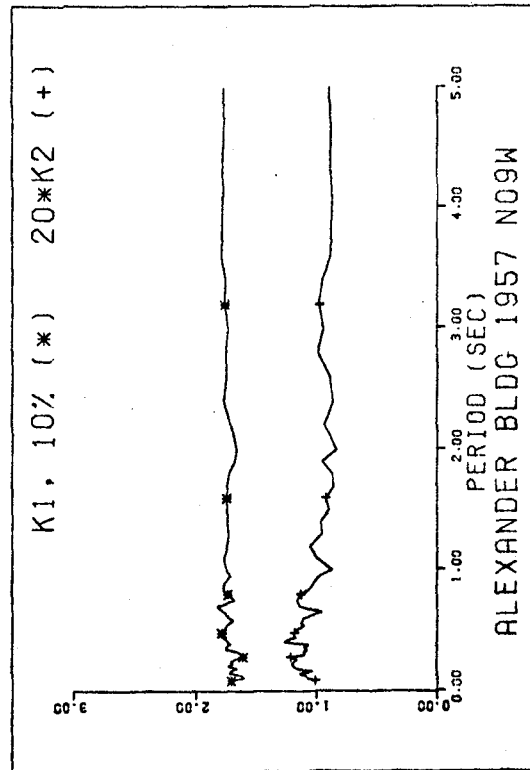
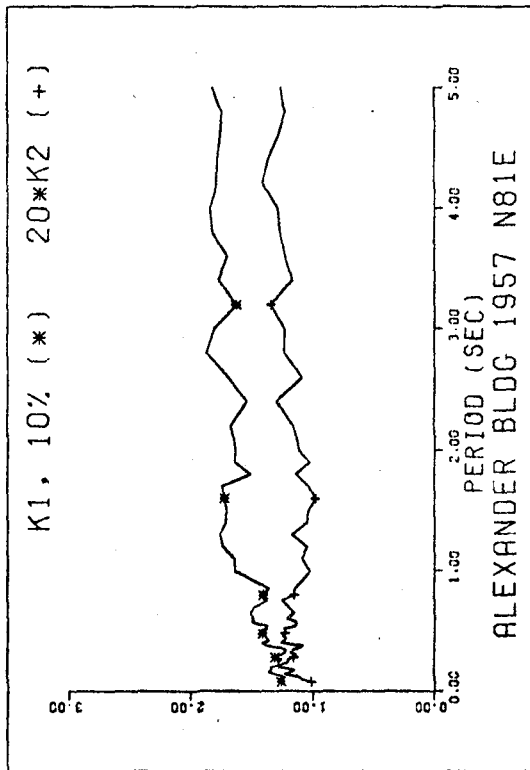
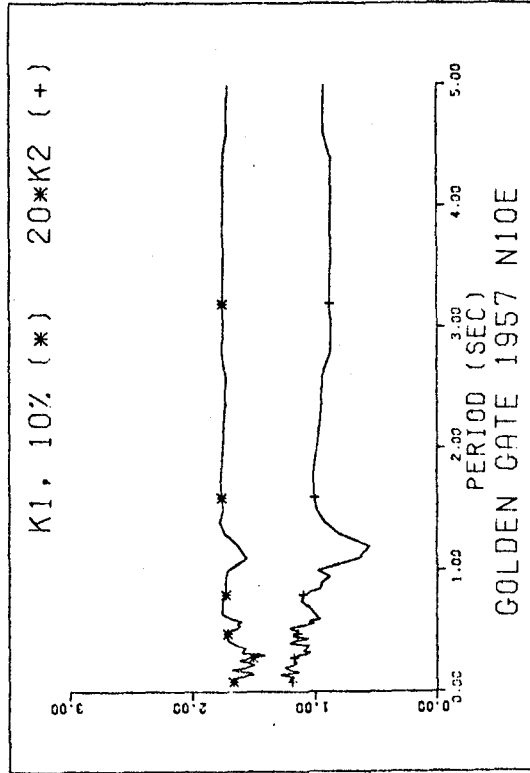
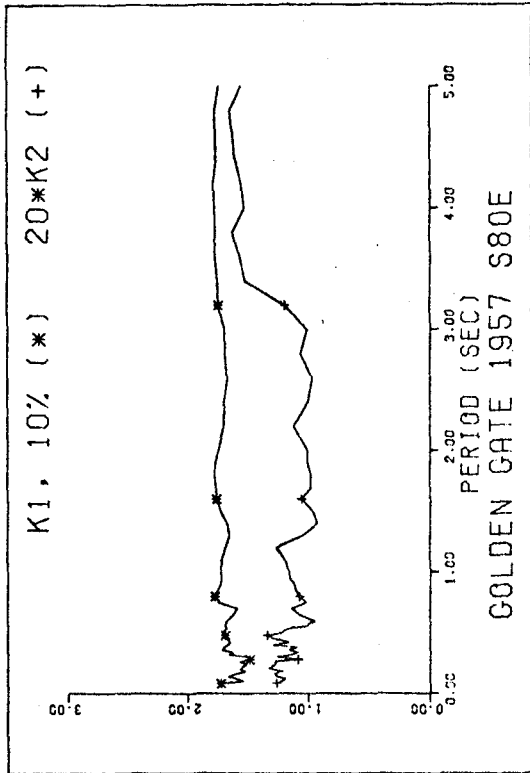
(A009)



(A013)

(A011)

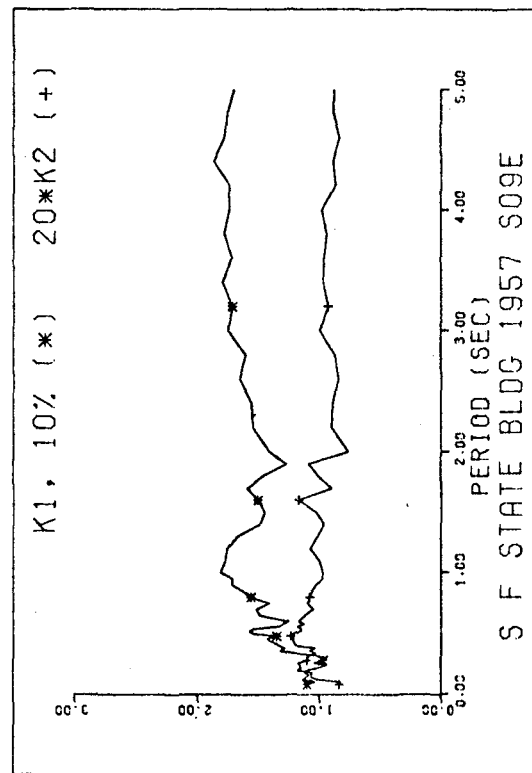
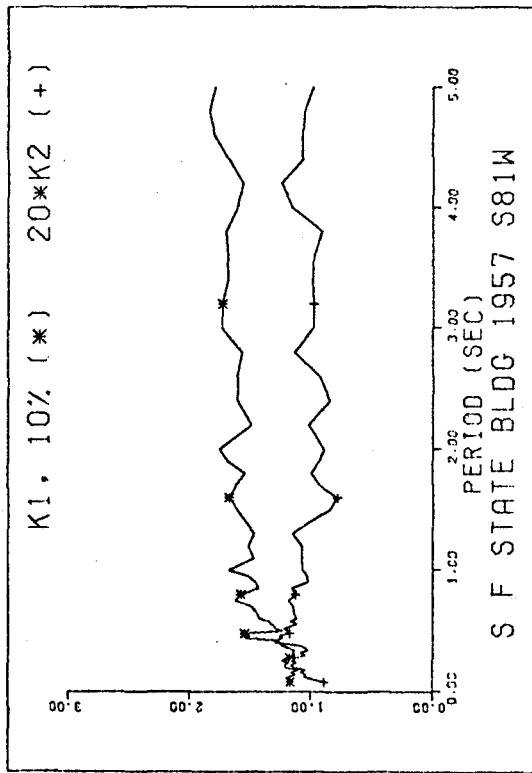
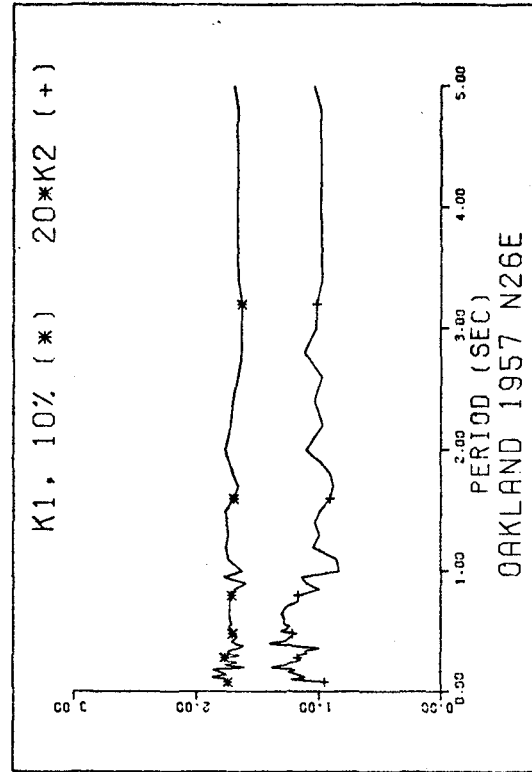
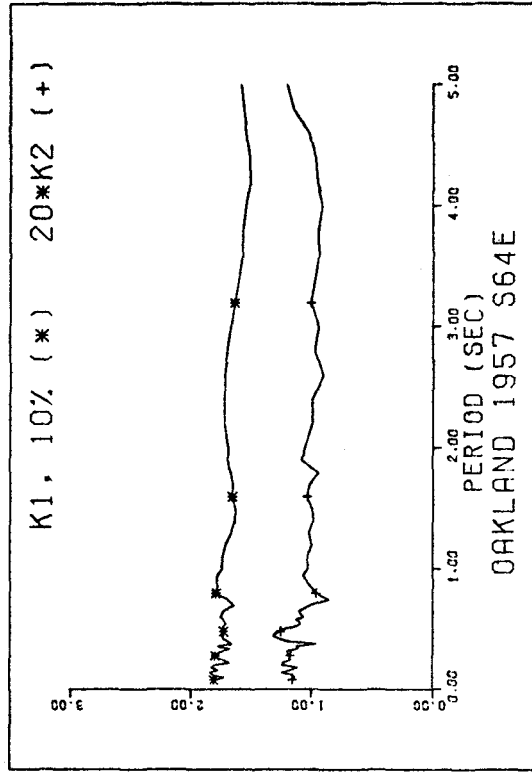
FIGURE 4.6. Factors K_1 and K_2 (Cont.)



(A015)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

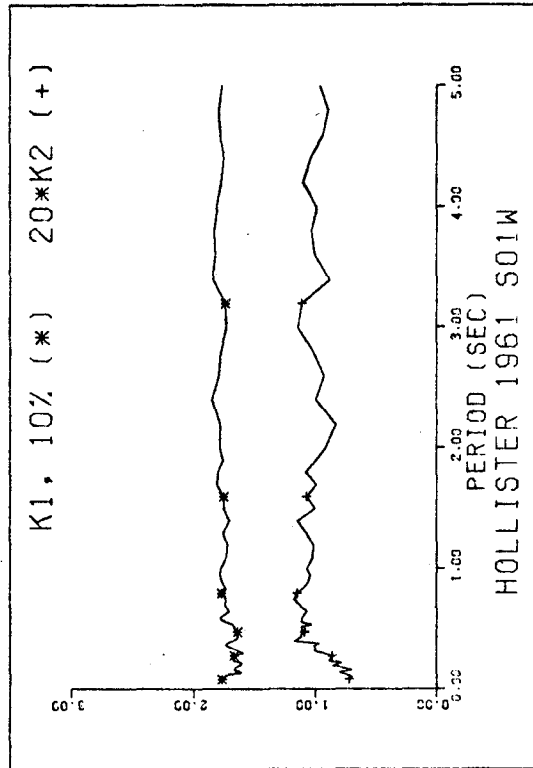
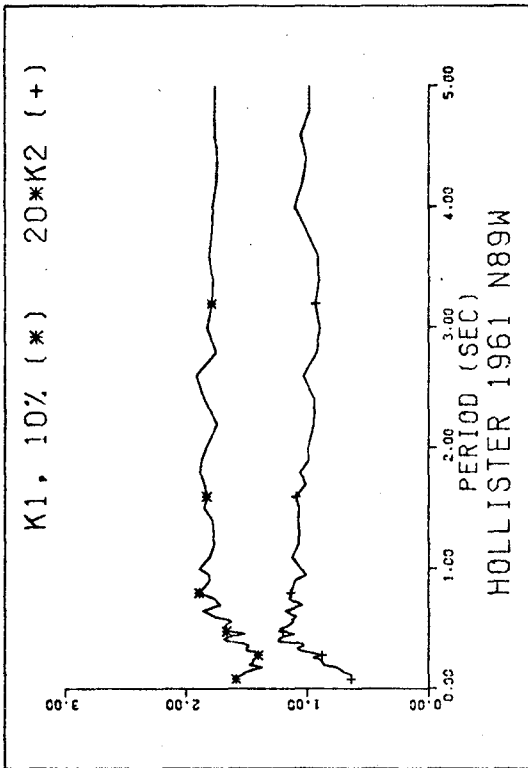
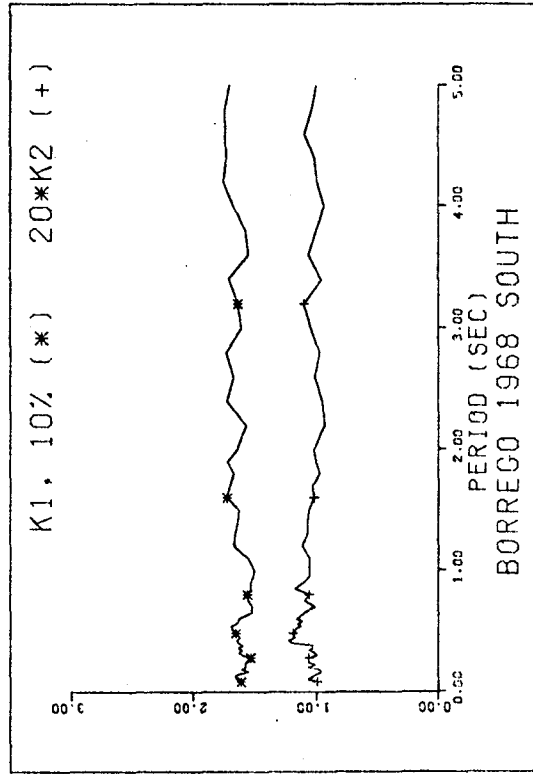
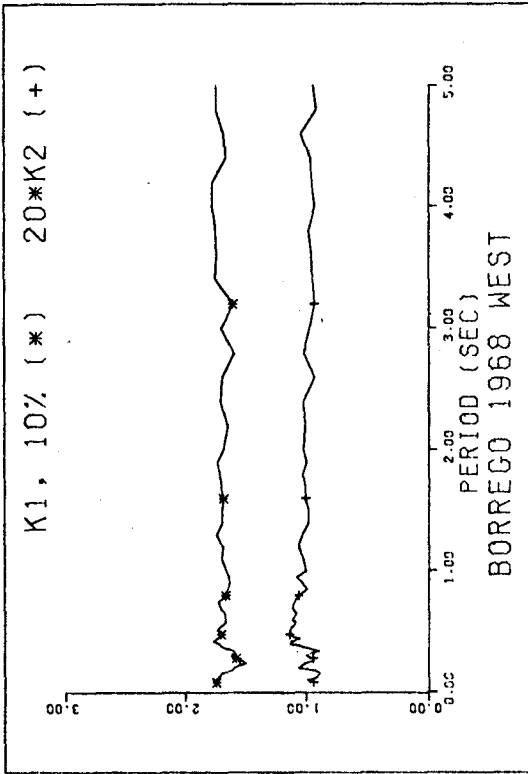
(A014)



(A017)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

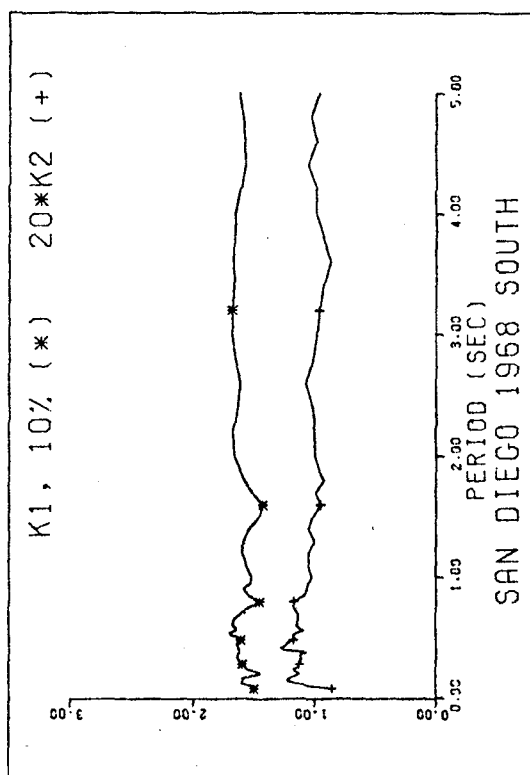
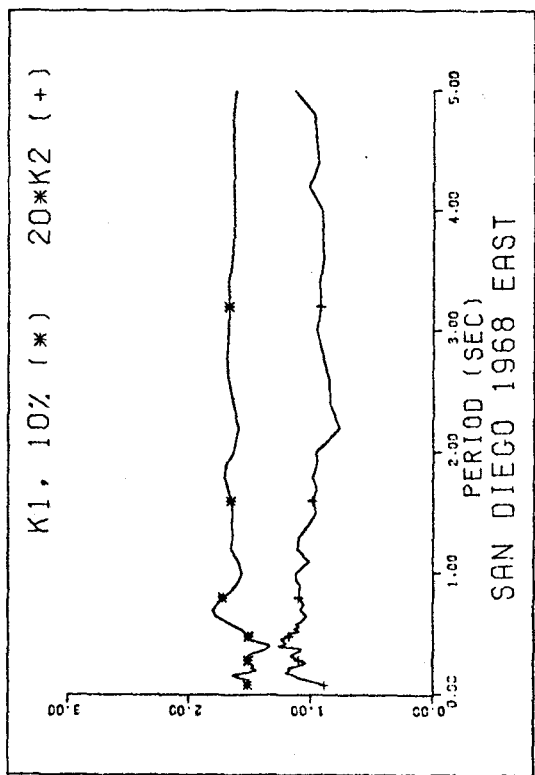
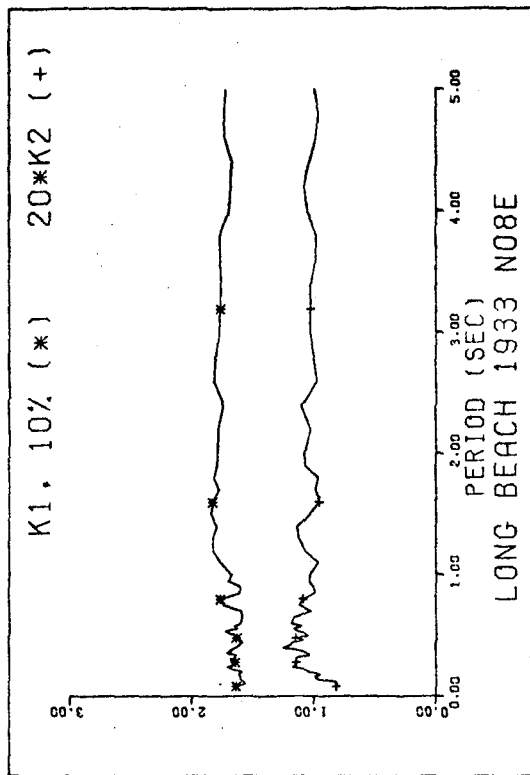
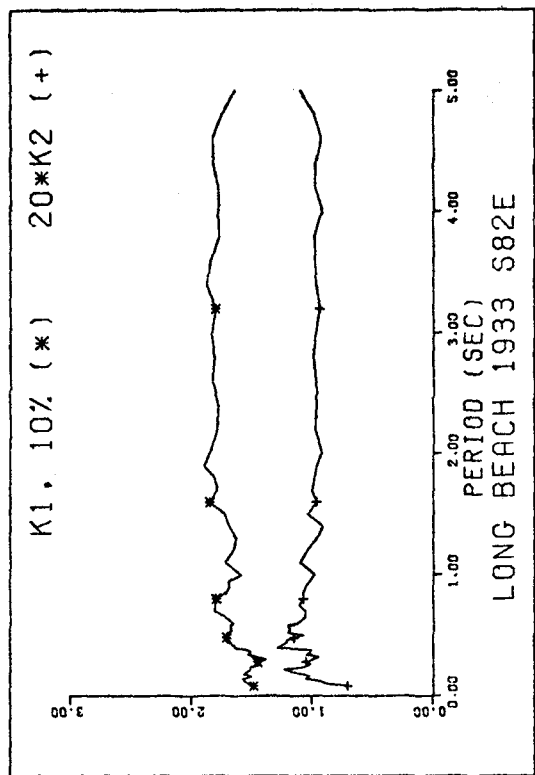
(A016)



(A019)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

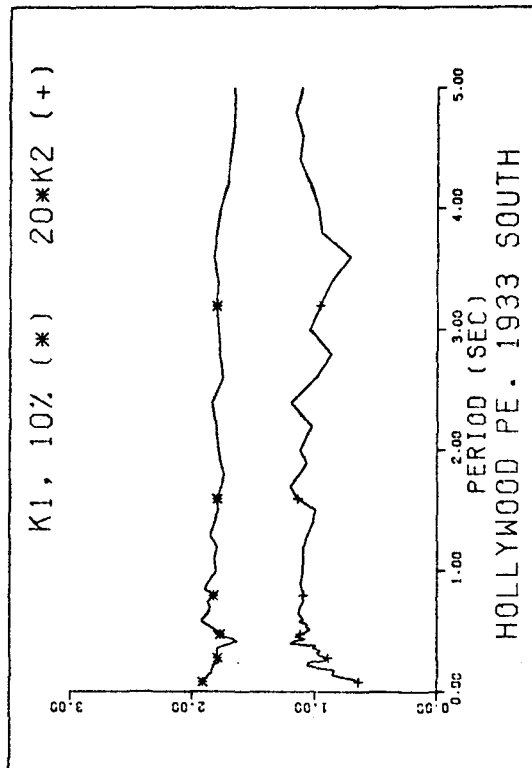
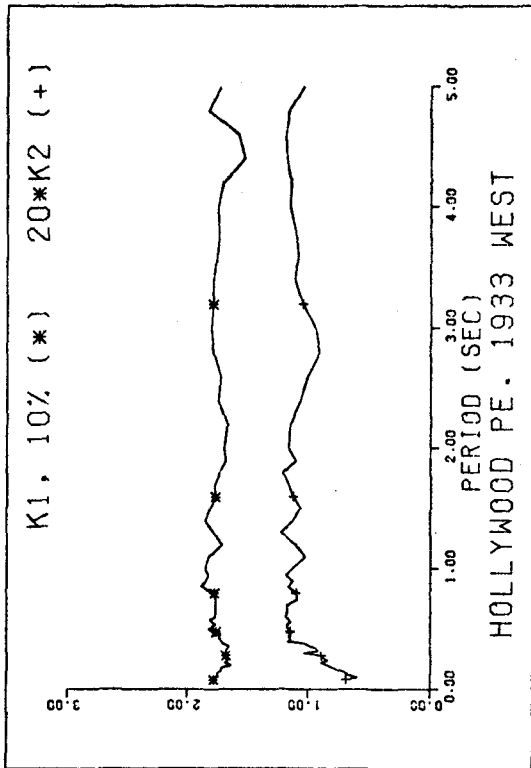
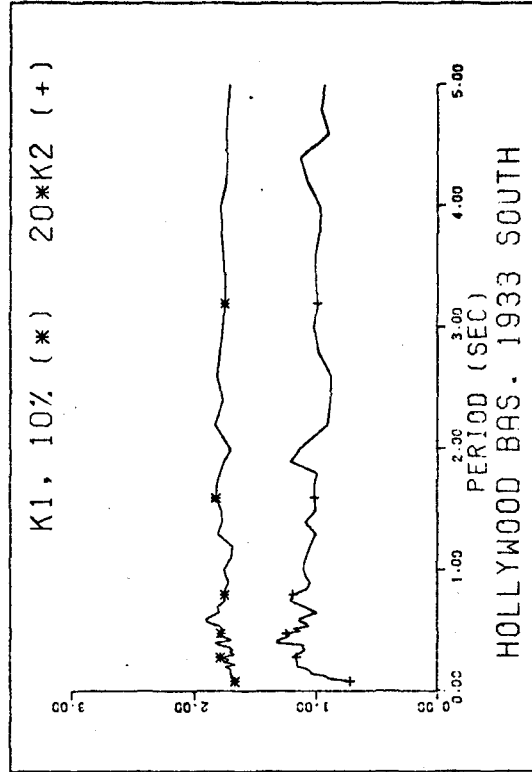
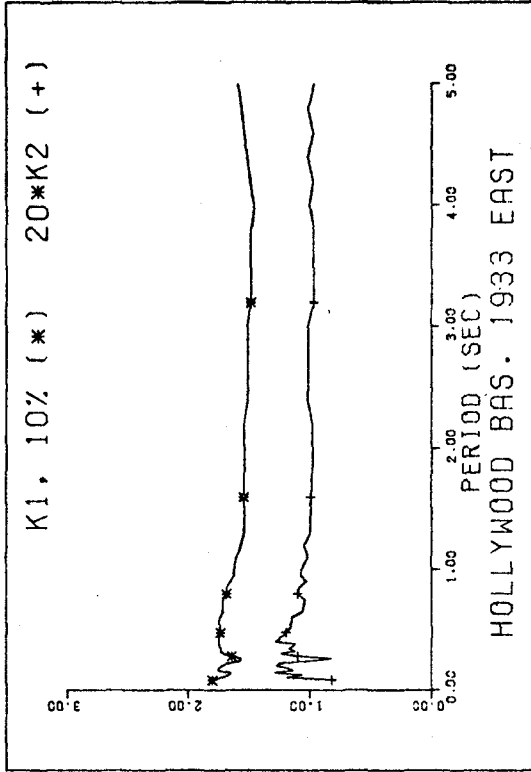
(A018)



(B021)

(A020)

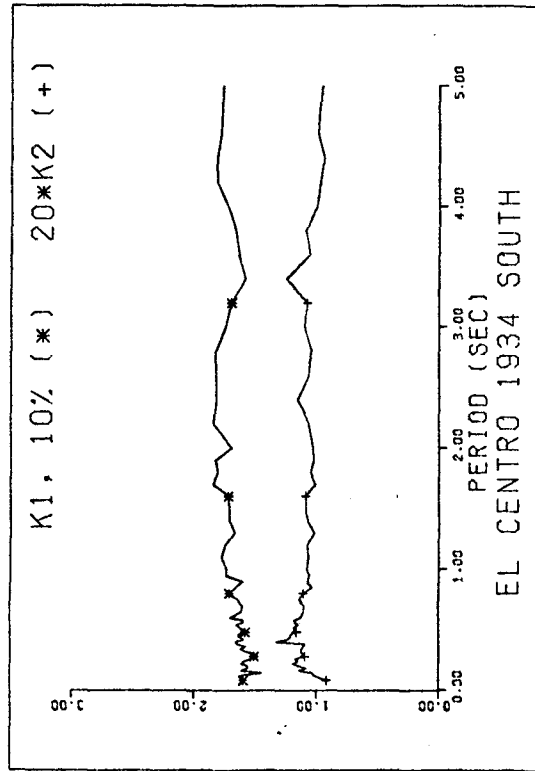
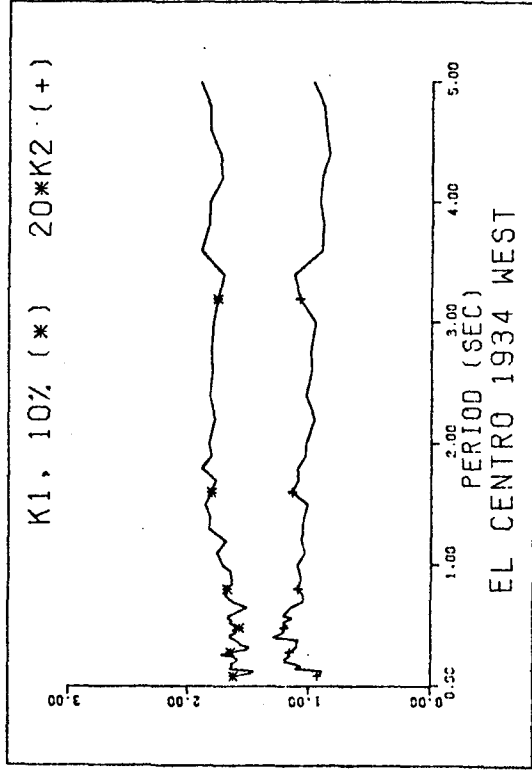
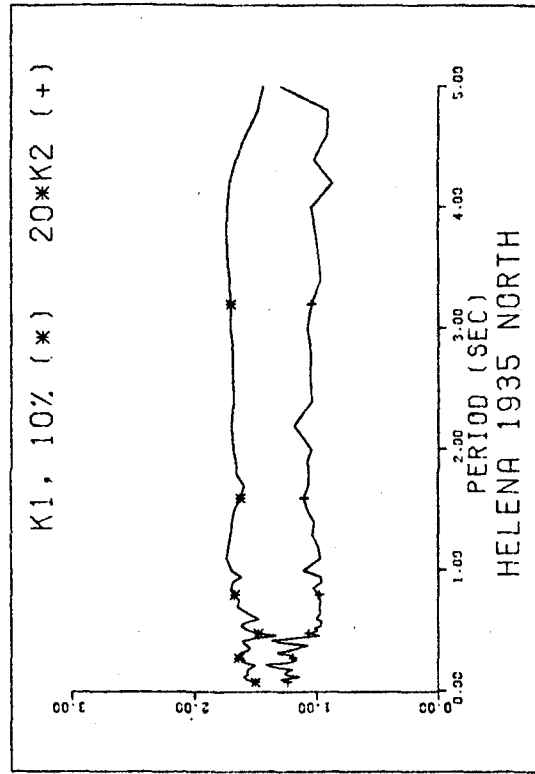
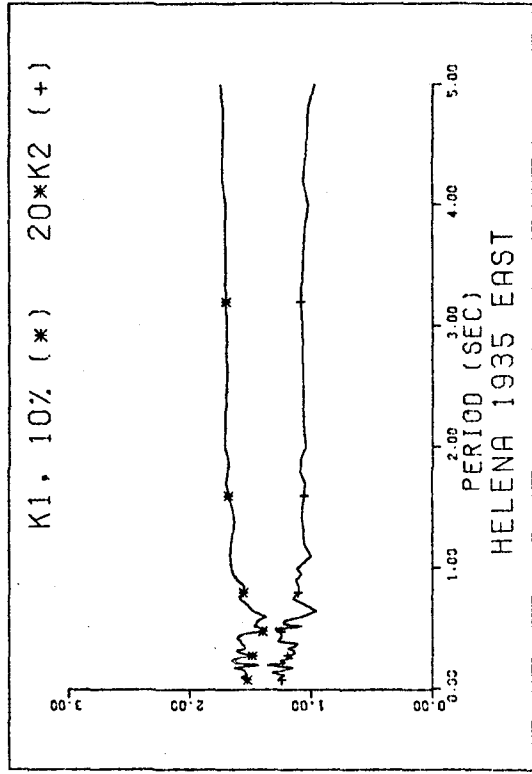
FIGURE 4.6. Factors K_1 and K_2 (Cont.)



(B023)

(B022)

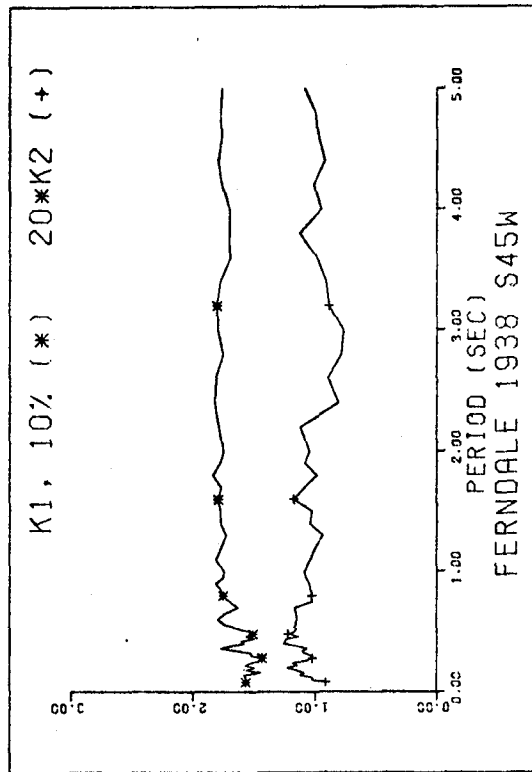
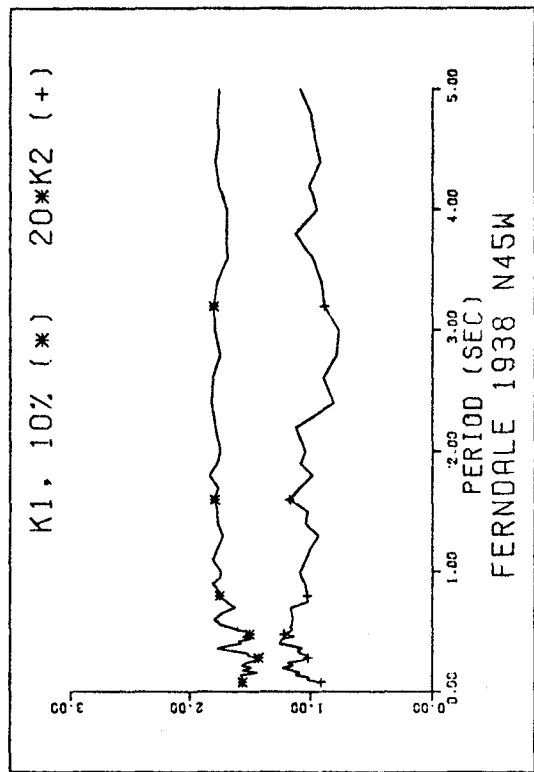
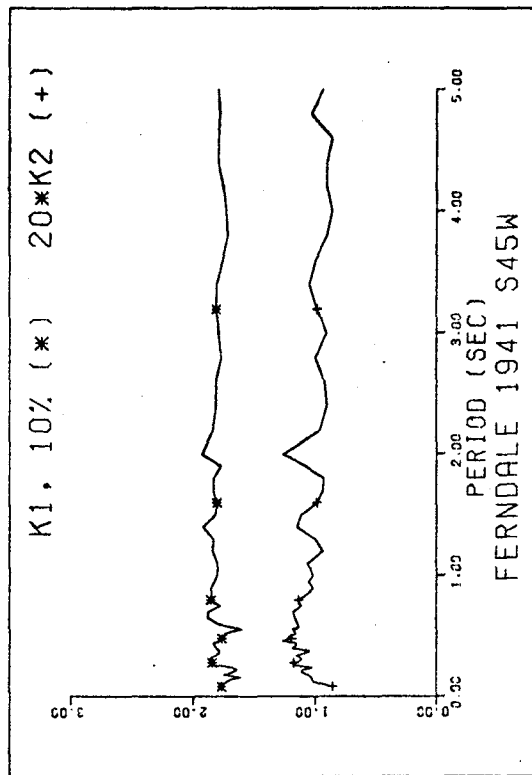
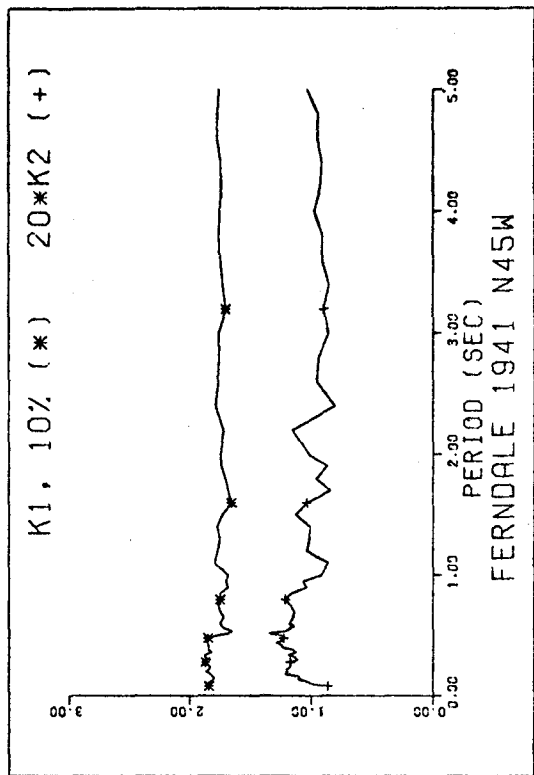
FIGURE 4.6. Factors K_1 and K_2 (Cont.)



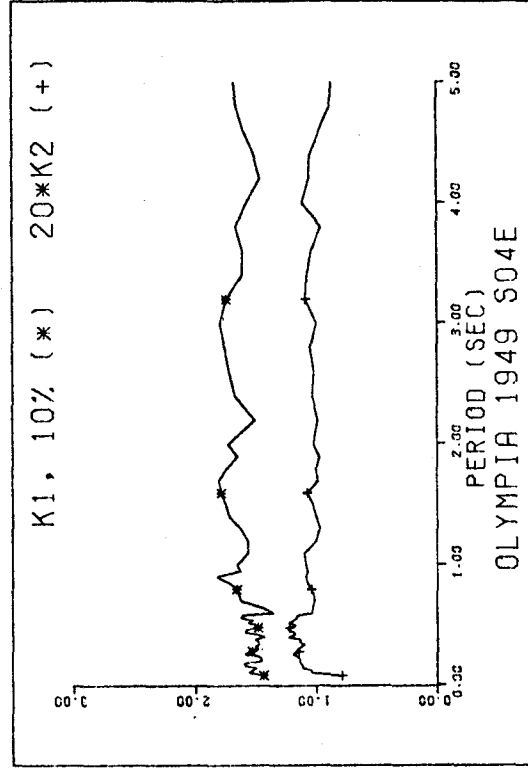
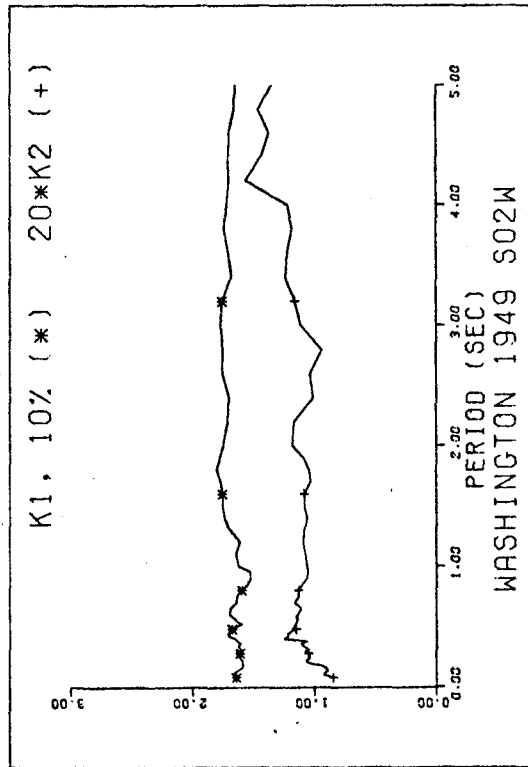
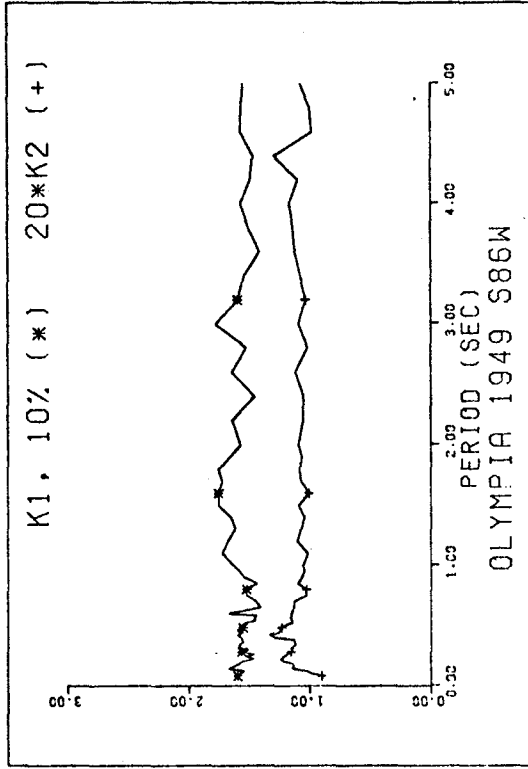
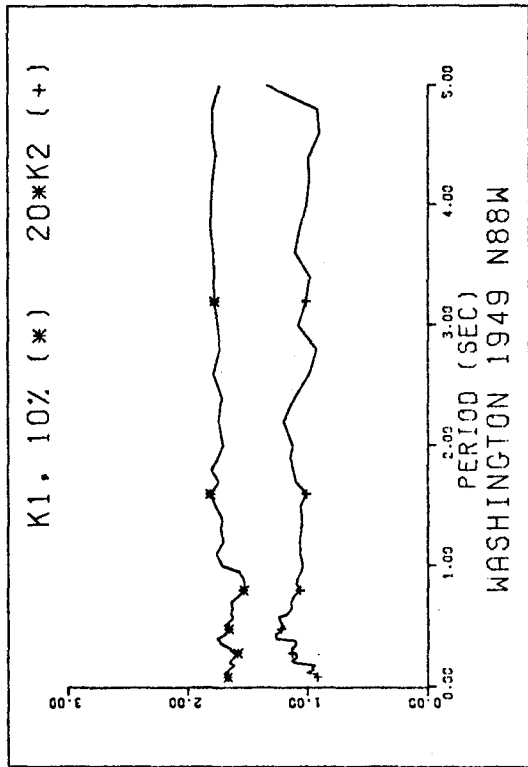
(B025)

(B024)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)



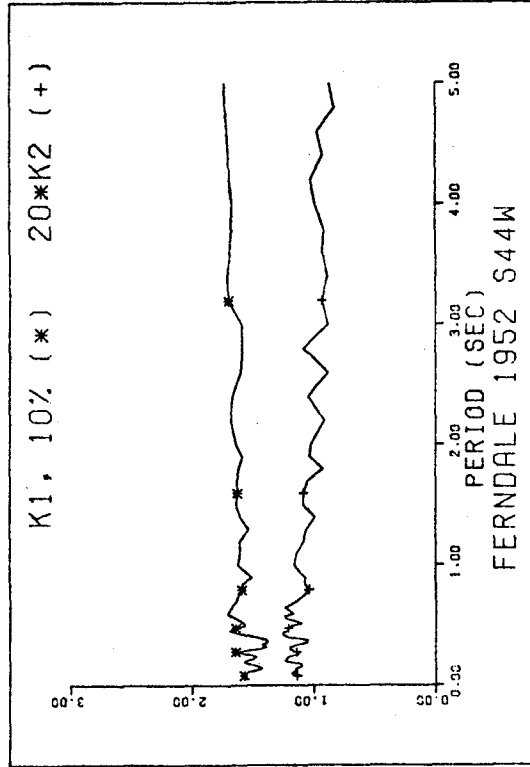
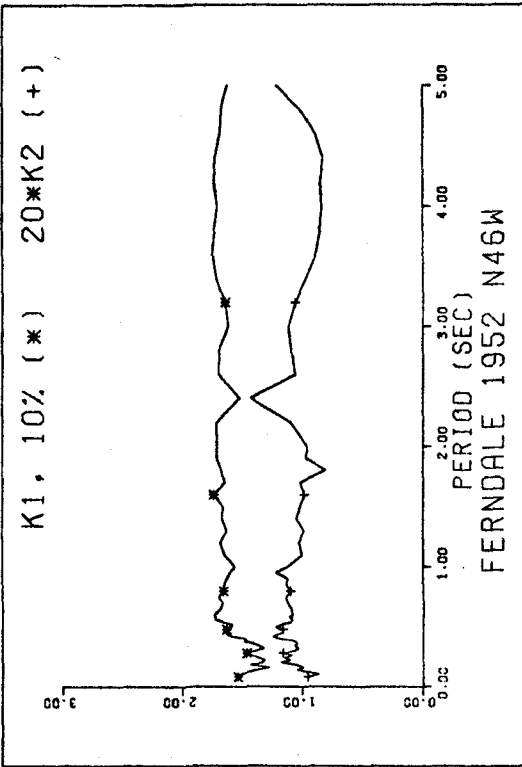
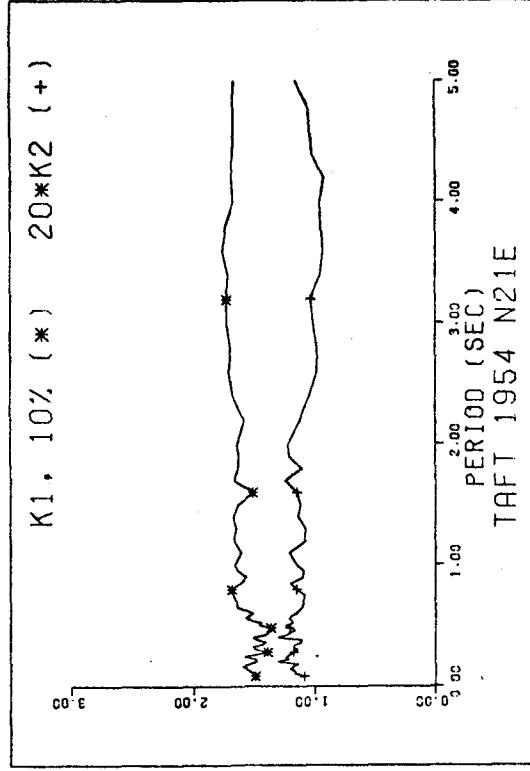
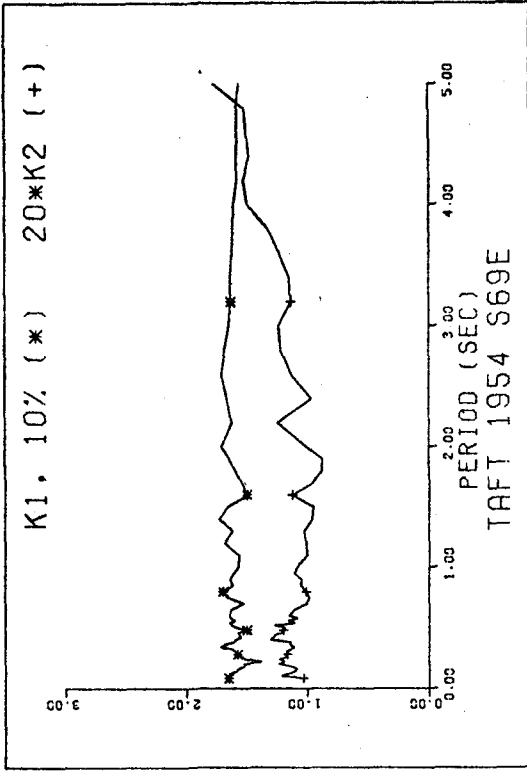
(B026) FIGURE 4.6. Factors K_1 and K_2 (Cont.) (B027)



(B028)

(B029)

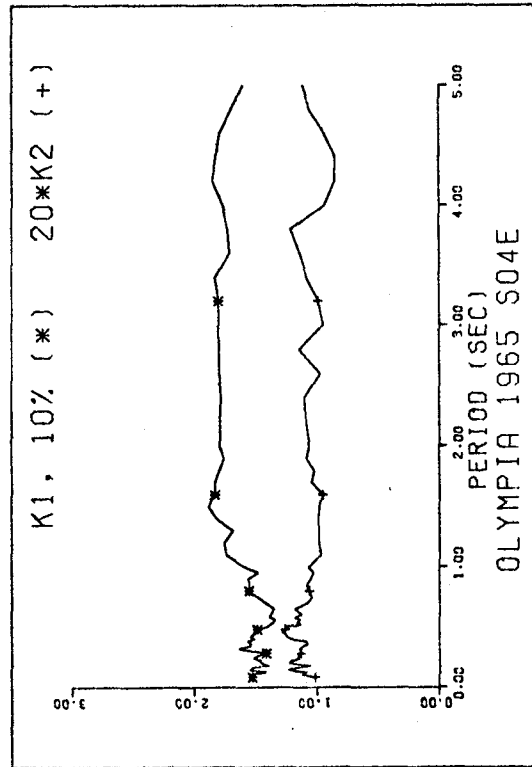
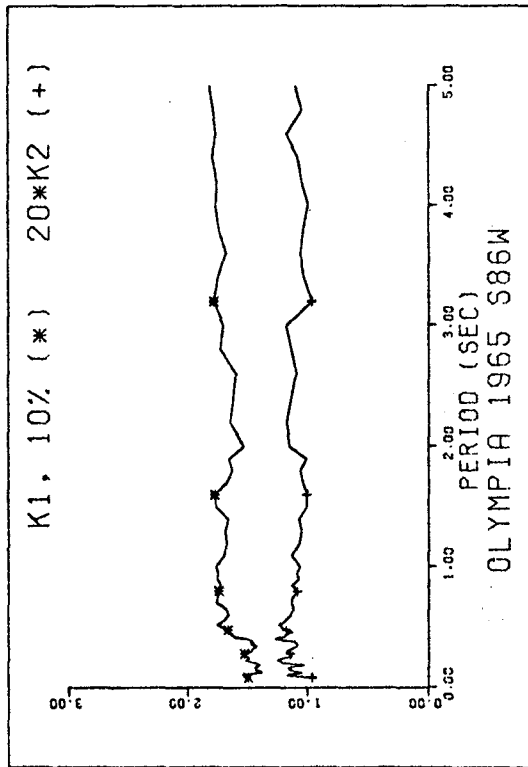
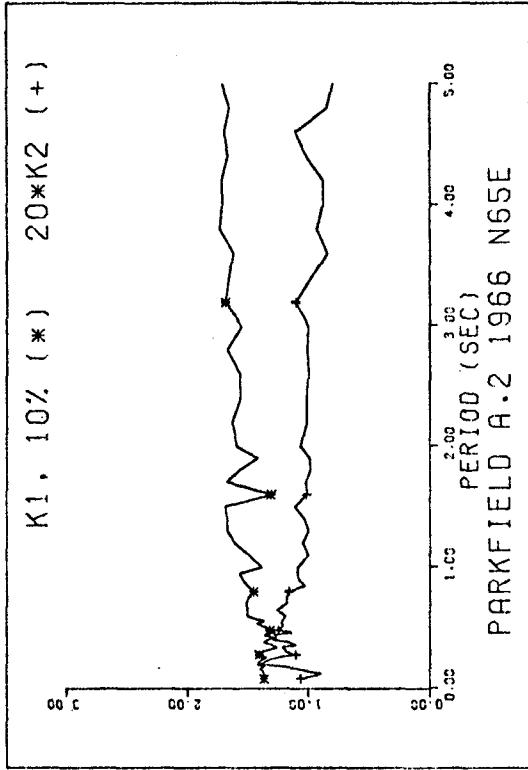
FIGURE 4.6. Factors K_1 and K_2 (Cont.)



(B031)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

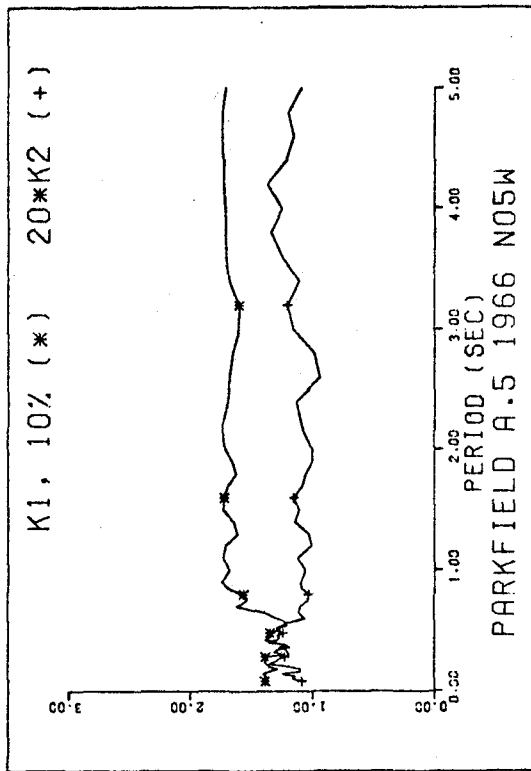
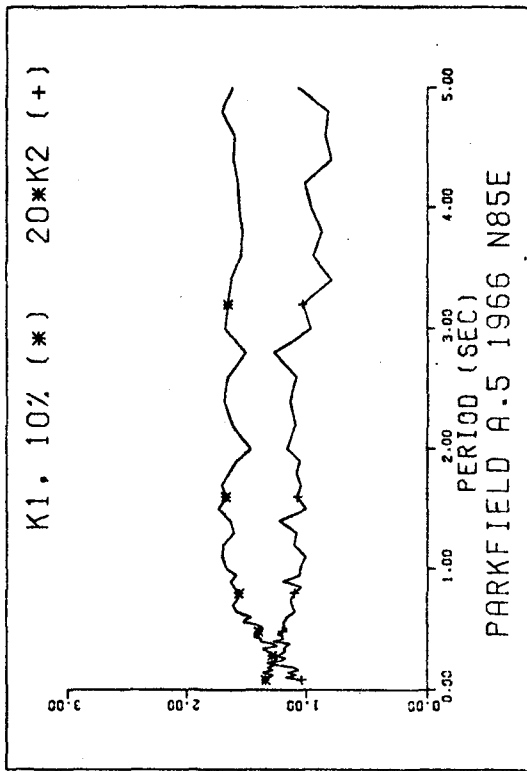
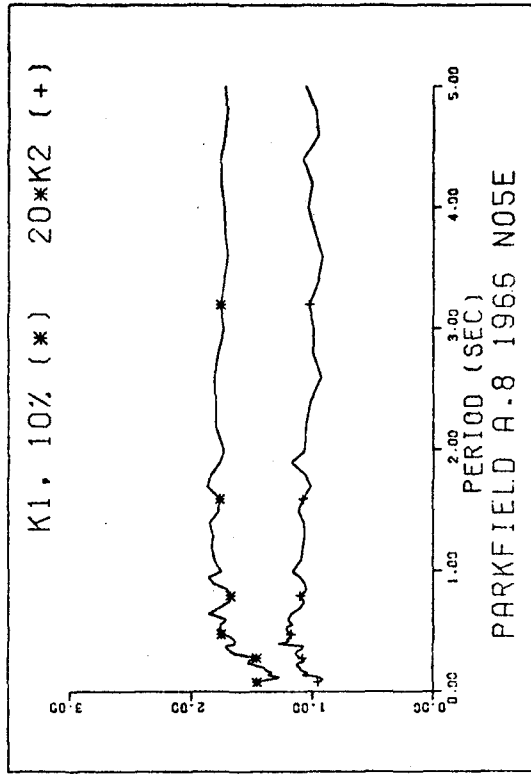
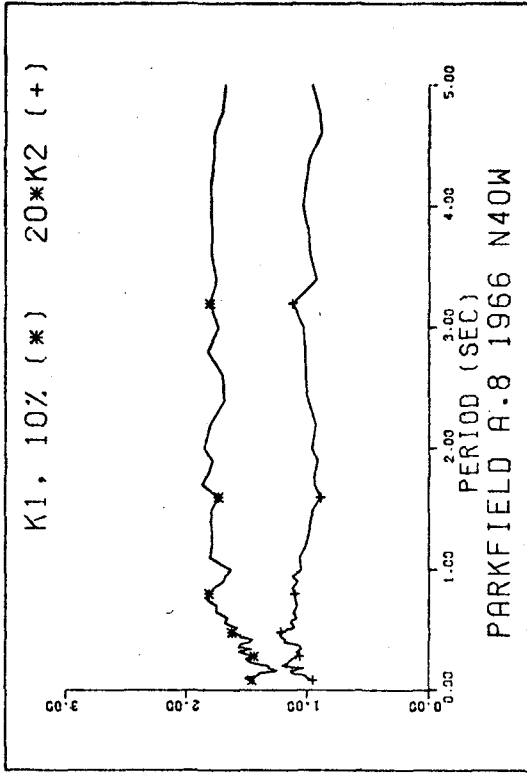
(B030)



(B033)

(B032)

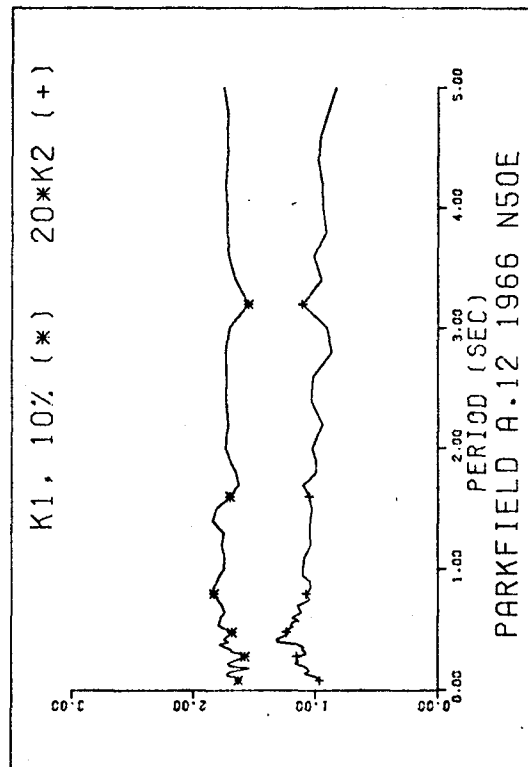
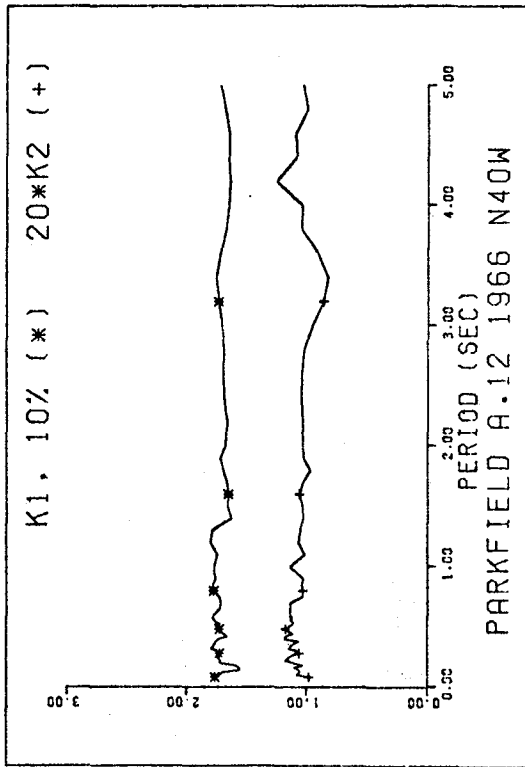
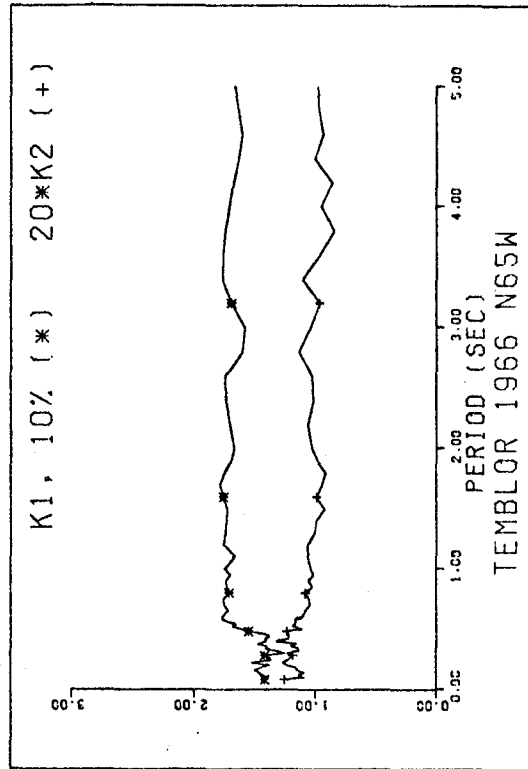
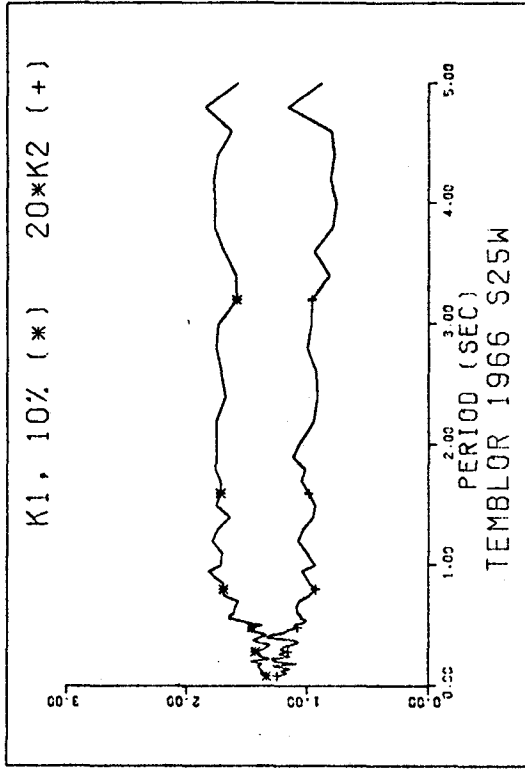
FIGURE 4.6. Factors K_1 and K_2 (Cont.)



(B035)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

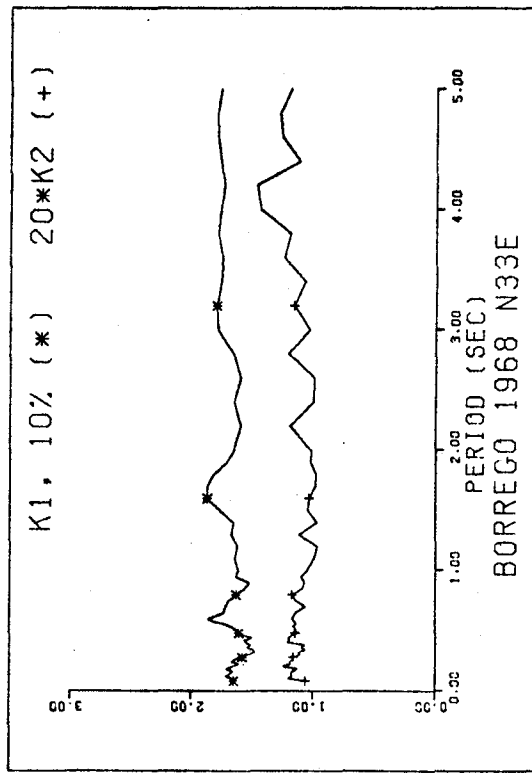
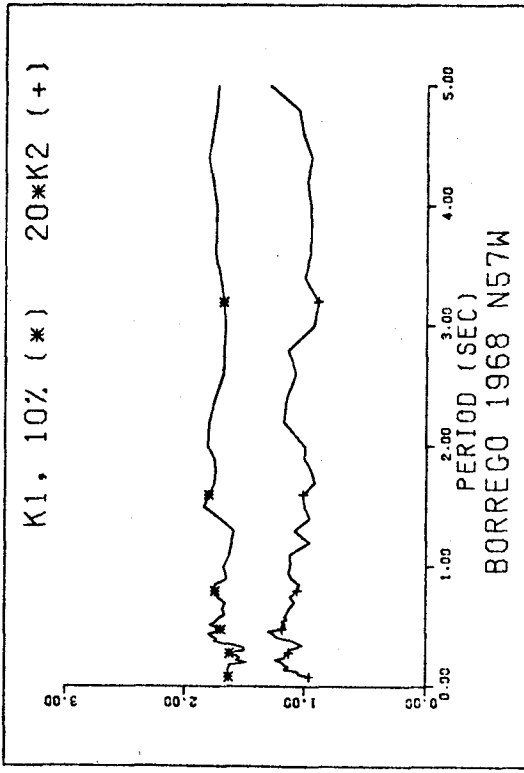
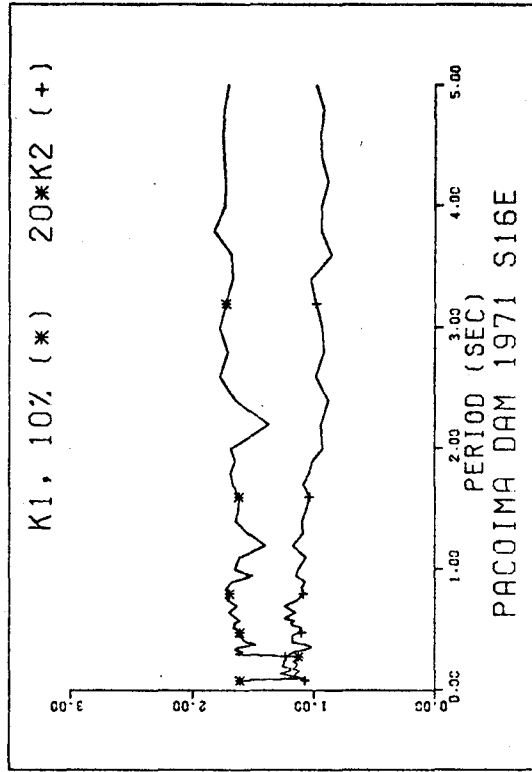
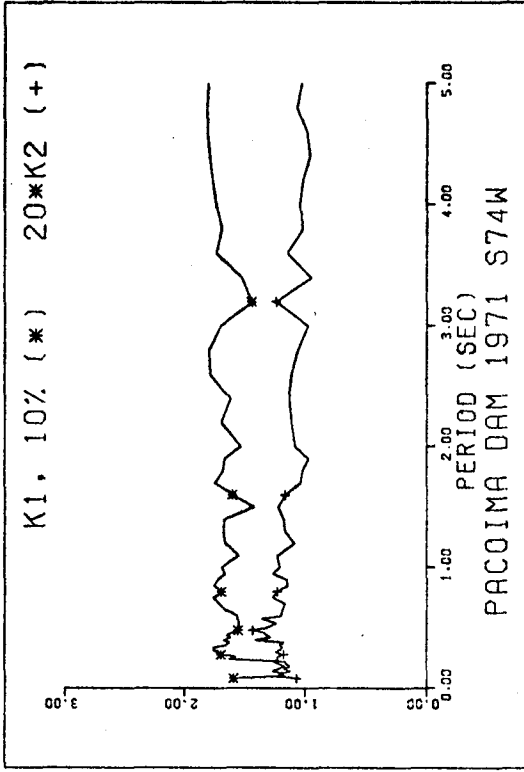
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(B037)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

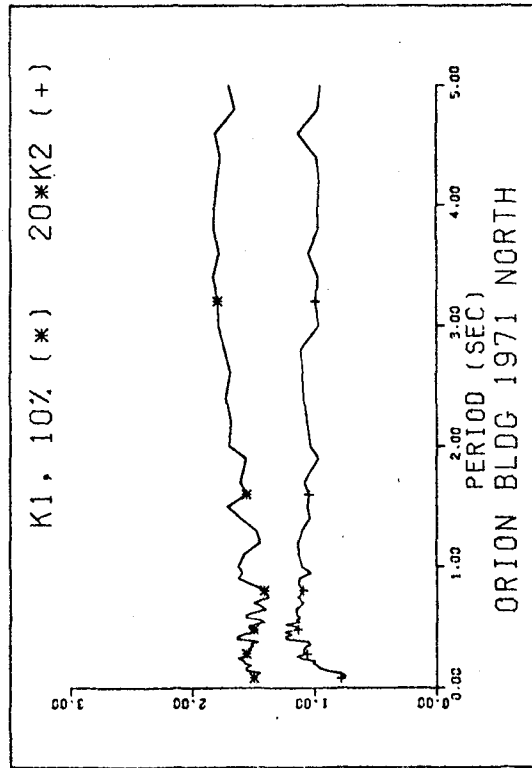
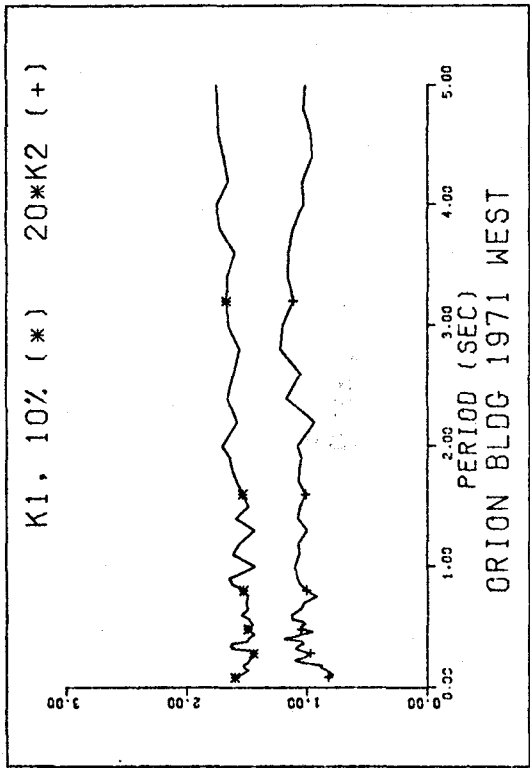
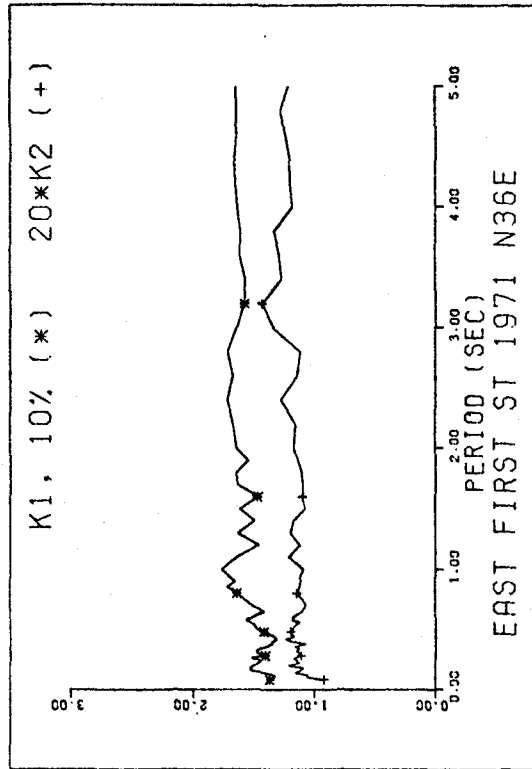
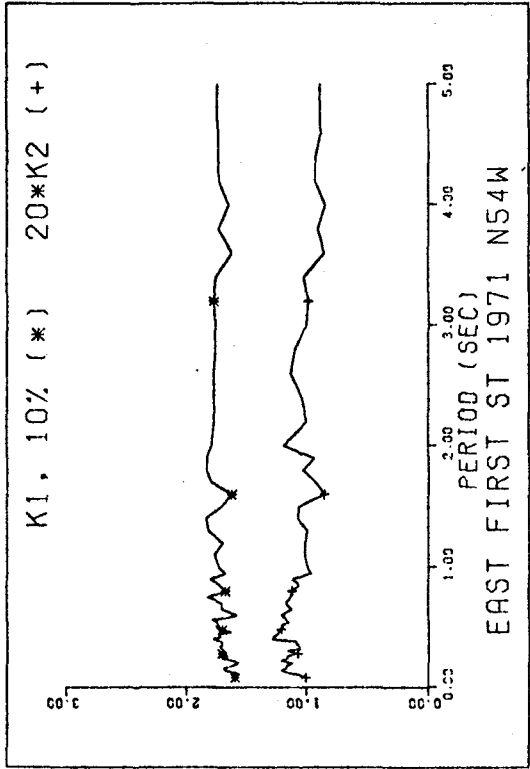
(B036)



(C041)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

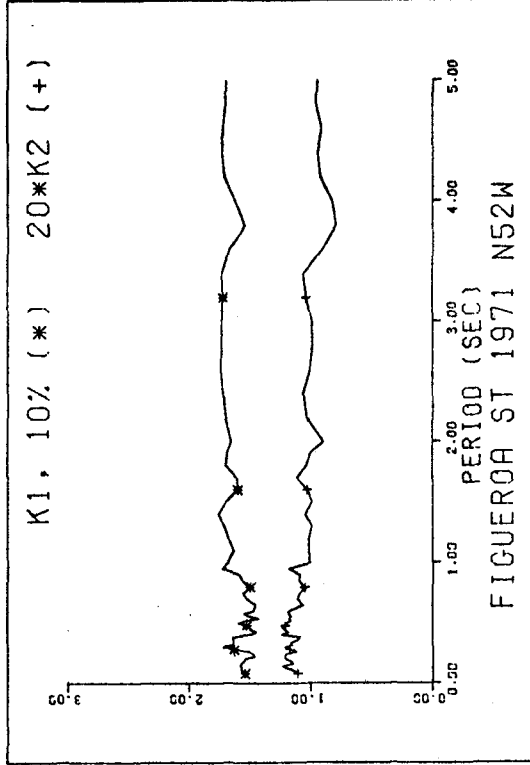
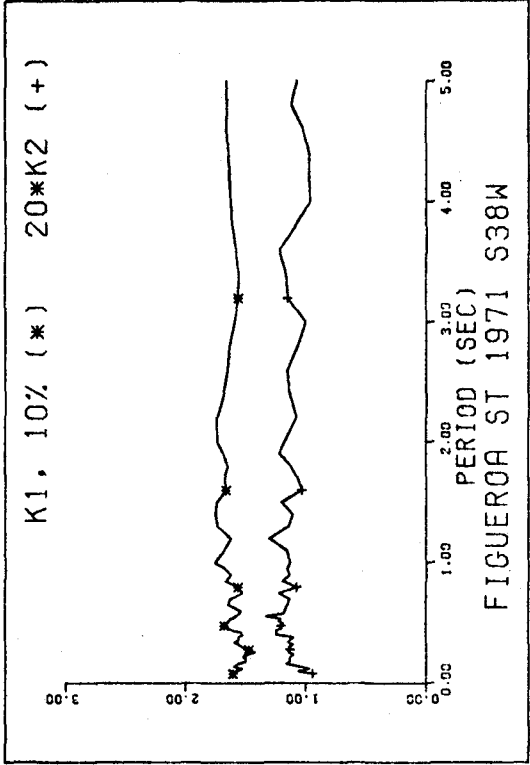
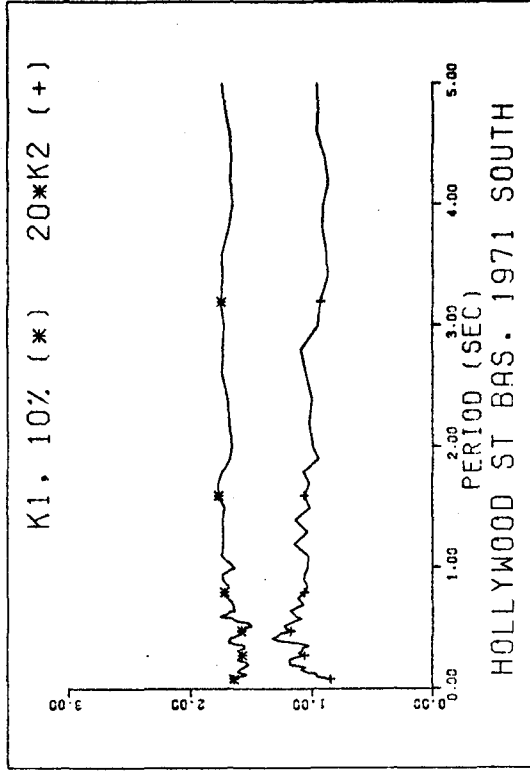
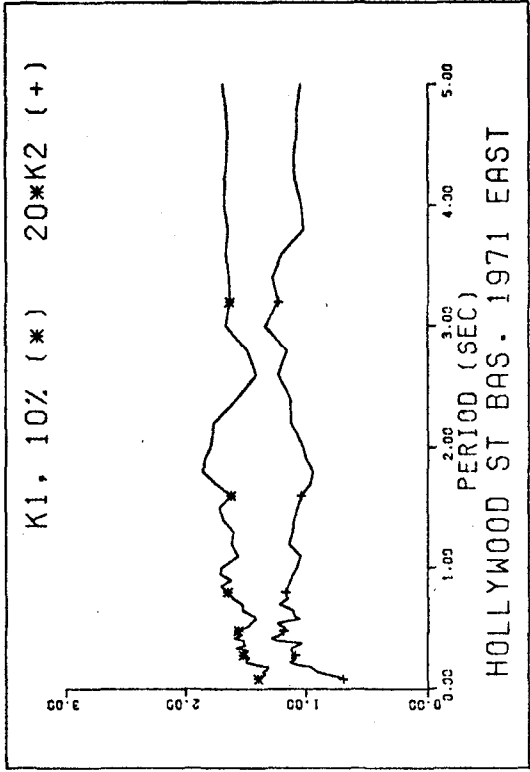
(B040)



(C051)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

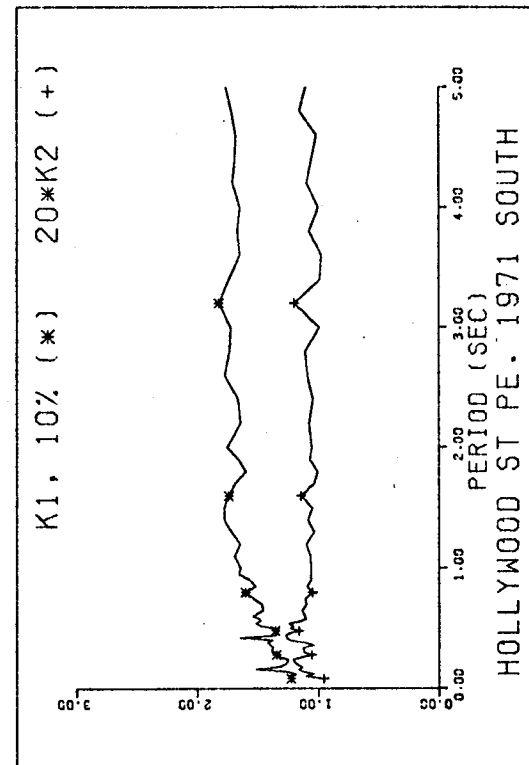
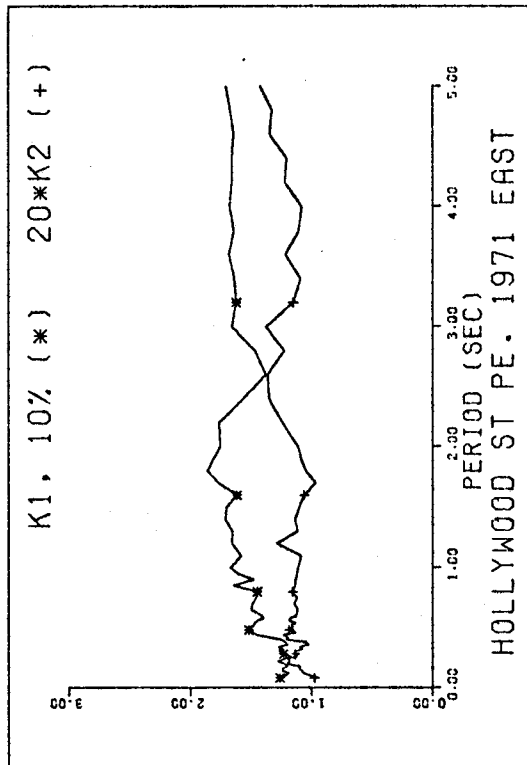
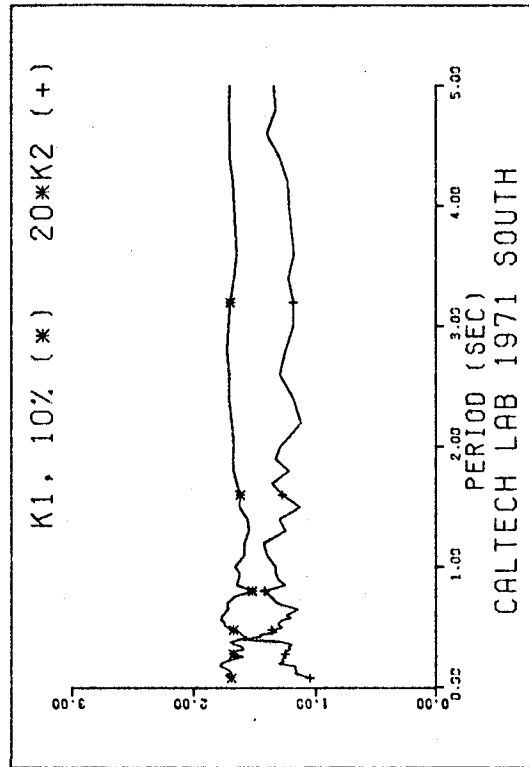
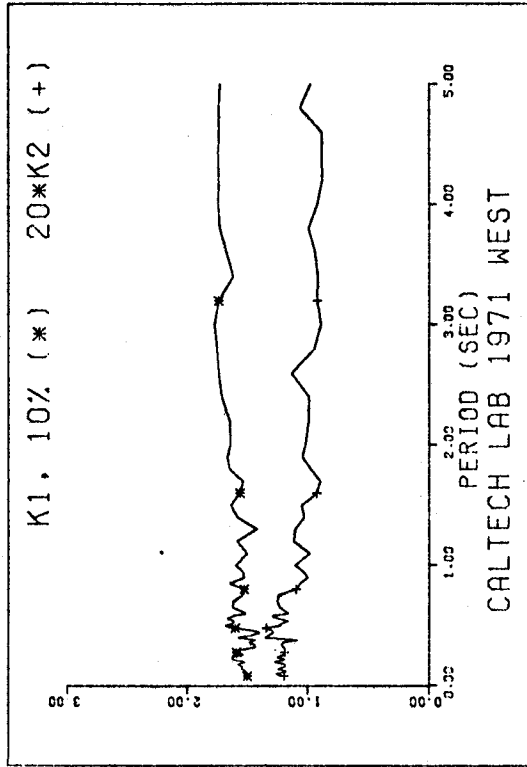
(C048)



(D057)

(C054)

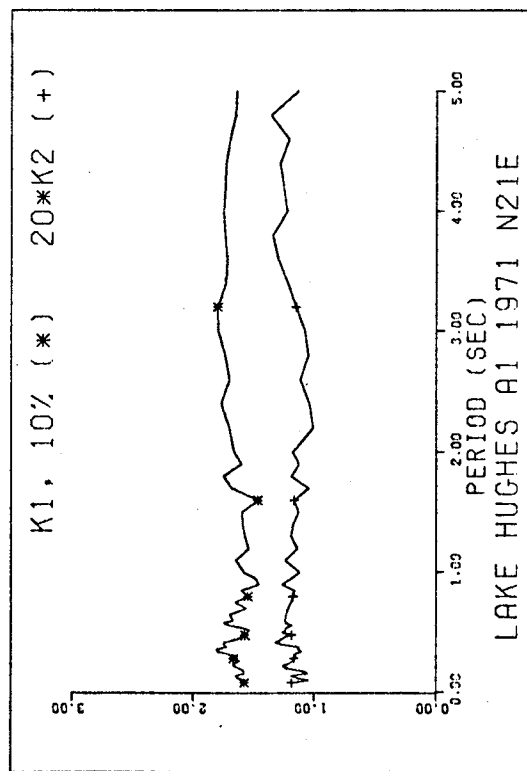
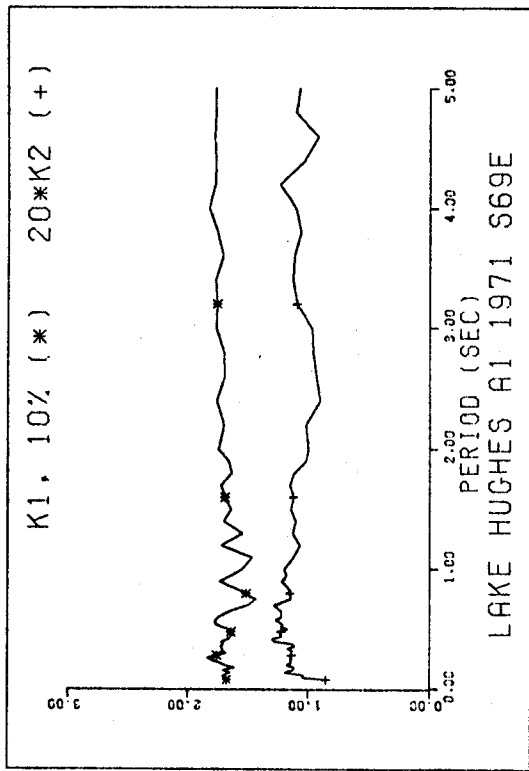
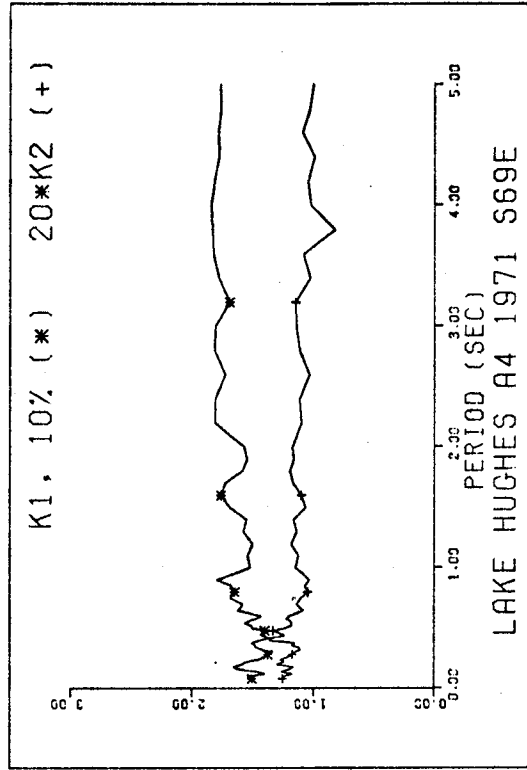
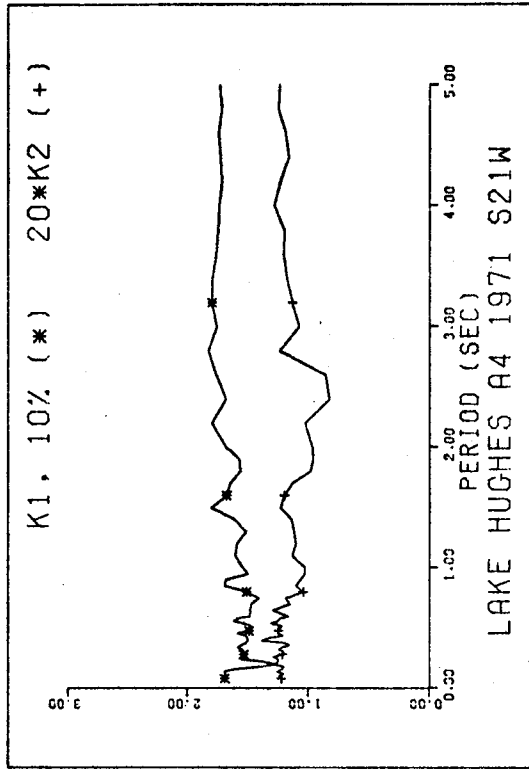
FIGURE 4.6. Factors K_1 and K_2 (Cont.)



(G106)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

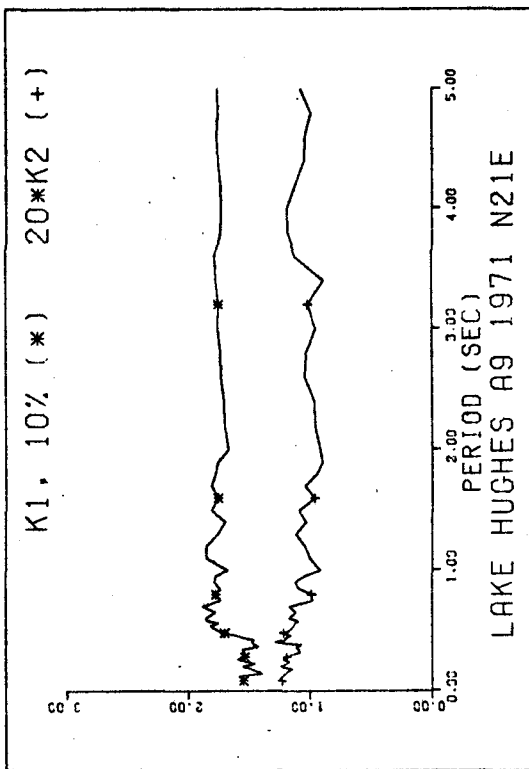
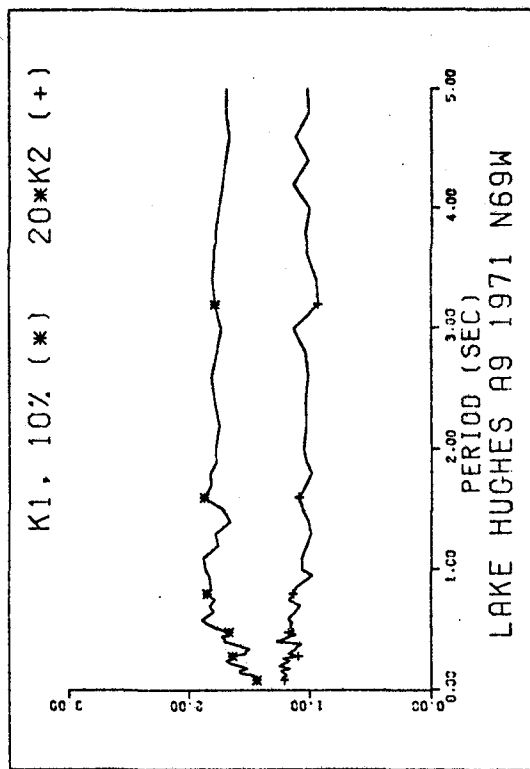
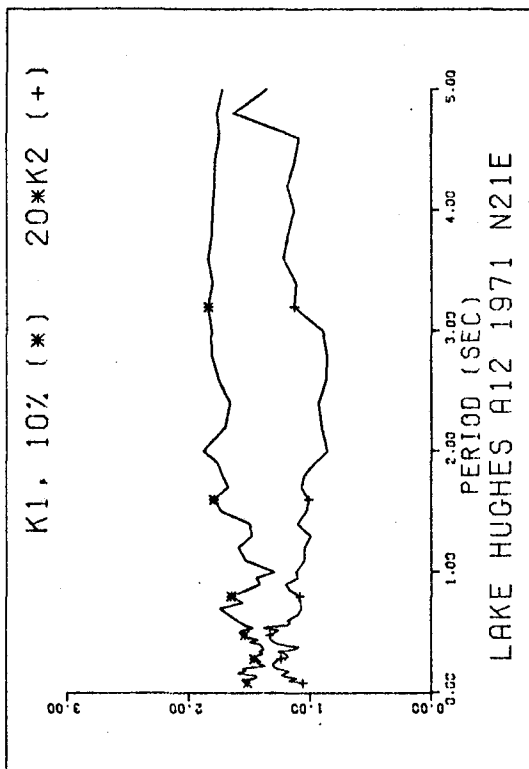
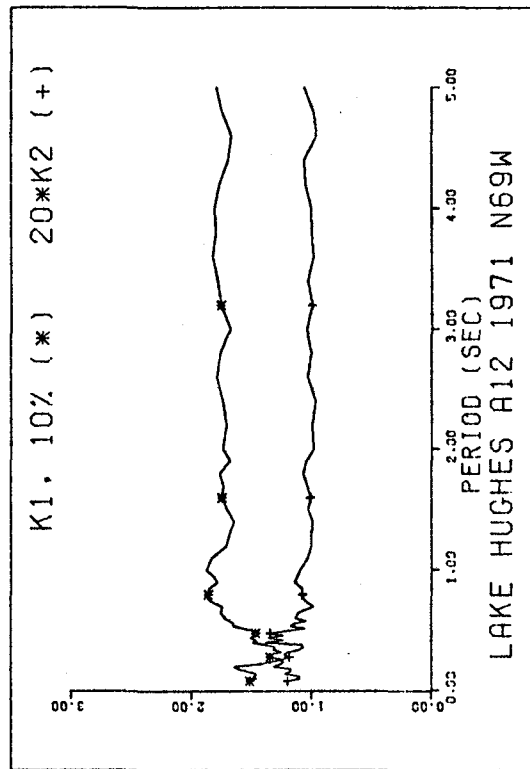
(D058)



(J142)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

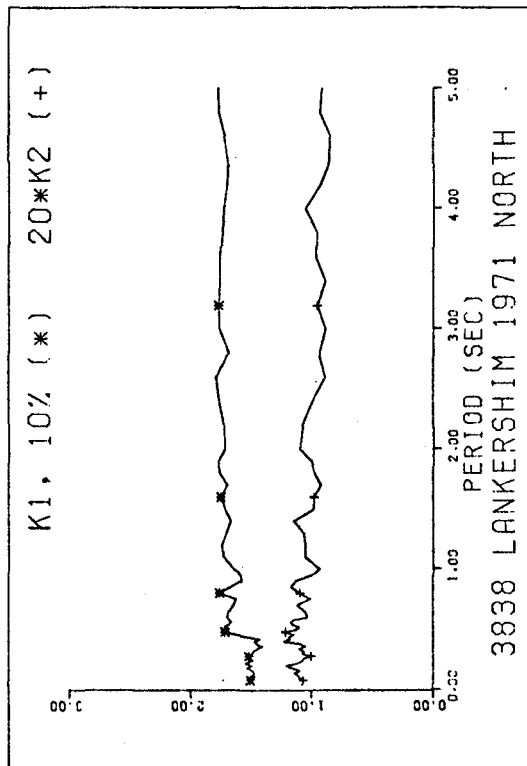
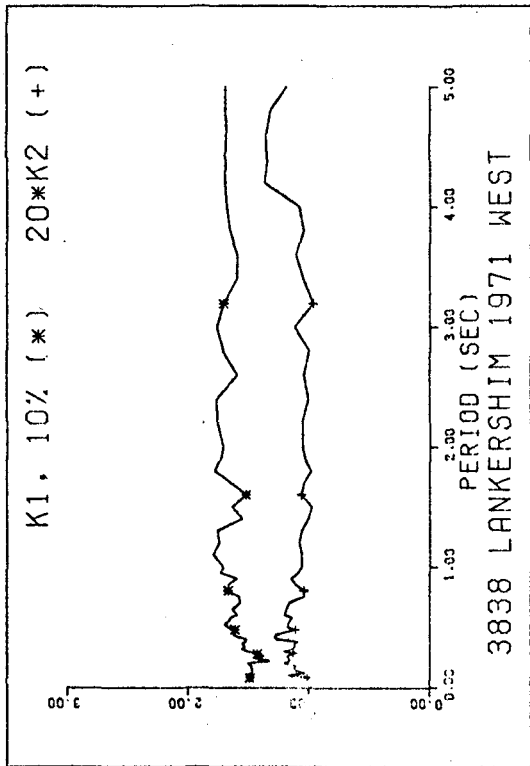
(J141)



(J144)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

(J143)



(L166)

FIGURE 4.6. Factors K_1 and K_2 (Cont.)

CHAPTER V
STABLE DESIGN PARAMETERS



5.1 Introduction

In Chapter IV, 97 earthquake accelerograms were analyzed for various parameters of possible engineering interest. The analysis of these accelerograms was conducted in two steps:

- Various parameters associated with the recorded ground motion were first analyzed. This step was necessary to understand the input characteristic of ground motion for any structural analysis. This step was also useful in determining the "zero period" frequency domain response analysis. Numerical values of the input parameters computed in this study are presented in Table 4.2.
- Parameters associated with the response of a single degree of freedom system for different periods and dampings were also obtained. This step was necessary to determine the variation in response parameters for various recorded ground motions. Appendix B gives the numerical values of all the response parameters of interest in this study.

It was mentioned in Section 1.4.2 that the above analysis was conducted with the following goals in mind:

- Develop stable parameters describing the earthquake ground motion and the response time histories. Currently used design parameters such as peak ground motion values (acceleration, velocity, displacement) or spectral response values do not seem to be good predictors of structural performance or damage potential. This is understandable because, by definition, the peak values are extremes and hence,

statistically, they are not good summaries of the overall behavior.

- Relate the developed stable parameters with the currently used design parameters. This is an important and desirable goal. The current design practice is totally entrenched in the use of peak values such as PGA for recorded ground motion and response spectrum for frequency content. Unless the newly developed stable parameters can be related to these widely used quantities, their usefulness and engineering applicability will be very limited.
- Incorporate new stable parameters in the development of seismic hazard maps. If the potential seismic hazard for a region can be described in terms of parameters more stable than the peak values currently used, a better and more reliable design strategy can be developed to reduce the potential seismic risk.

In the following sections, various possible applications and justifications are given for the parameters presented in Chapter IV.

5.2 Duration of the Input Accelerogram and of the Response Time History

As mentioned previously, more research is needed to develop a standard and meaningful (in engineering terms) definition of duration both for recorded ground motion and response time histories. In this work, the record is truncated at the time when the acceleration becomes and remains smaller than 0.02 G (Table 4.1). Obviously, this is an arbitrary way of defining the earthquake duration. A better method would be to look at the rate at which, for a given record, the input

energy increases with time. Duration could then be defined as the time during which 90% or 95% of the total energy is transferred. (Trifunac and Brady, 1975; Dobry, et al., 1977). This definition has the advantage of disregarding both the beginning and the end of the record. It should be noted that the value of duration selected affects the RMS and the total energy. The probabilistic duration maps developed in Chapter III can be revised and updated if a better definition of the input duration is developed. In Table 4.2 (column 4), the number of peaks (NEPK) is directly related to the input duration.

Even though most earthquake engineers are of the opinion that the damage potential of a given structure increases with an increase in response duration (for a given response amplitude), no design parameter explicitly dependent on duration is currently used. In this work, the response is arbitrarily terminated when the response acceleration peak reaches 10% of the highest response peak and does not exceed that value thereafter. The numerical values are listed in seconds in Appendix B (column 3). They are used to calculate the RMS and the cumulative potential energy of each response. No detailed analysis of the definition of duration is conducted in this study. However, it can be seen that the duration of the response is a function of the input duration, input frequency content, the oscillator natural frequency and the damping (Fig. 4.3). It is interesting to note that a small spectral acceleration value could be associated with a long duration and hence large potential energy. Figure 5.1 shows two response parameters, S_a and ENGY, on a normalized scale such that their shapes can be compared. Contrarily to S_a , the value of ENGY is a function of duration. It can

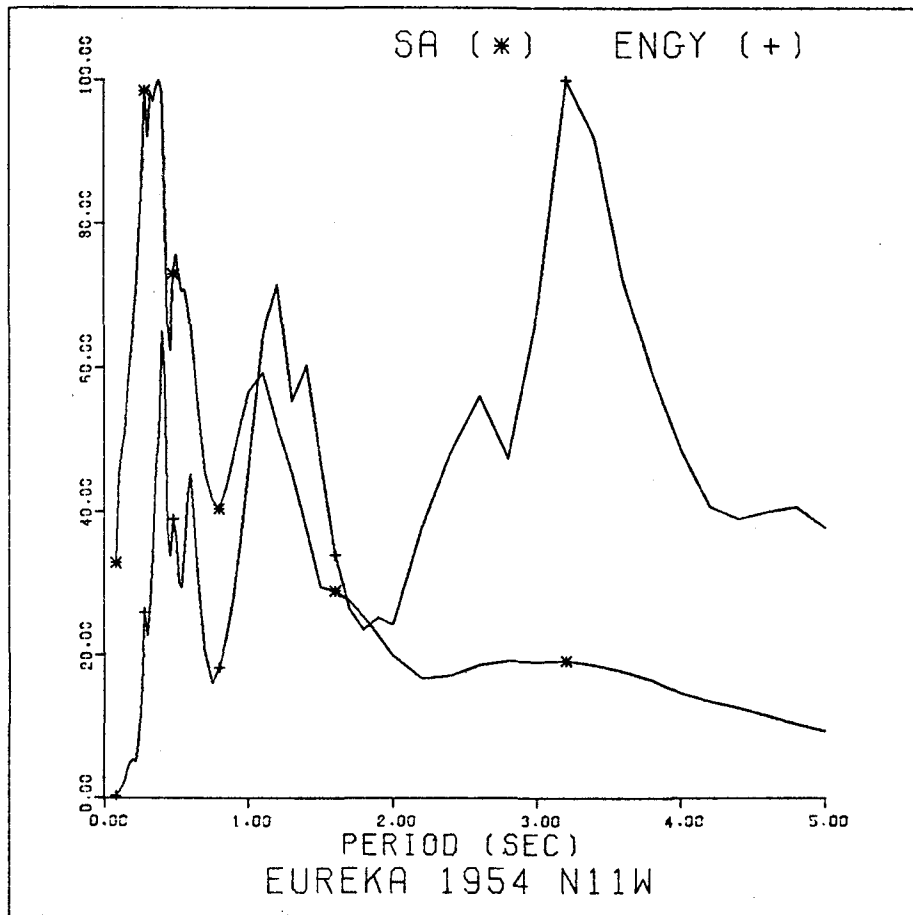


FIGURE 5.1. S_a and ENGY Plotted on Normalized Scale

be seen that in this case for longer period oscillators, the relative value of ENGY is much larger than the corresponding S_a . It is hoped that in the future the damage potential could be better correlated with the duration dependent parameter ENGY than it is today with S_a . More will be said about this parameter in a later section.

5.3 RMS of the Input Accelerogram and of the Response Time History

As mentioned previously in Chapter IV, peak amplitudes by themselves are poor indicators of seismic loading. Probabilistically, they show great scatter. A better parameter defining the amplitude of the input ground motion and of the response spectral value is the root mean square (RMS). Since the RMS is a statistical summary of many peaks, the scatter of this parameter between two similar events (same magnitude, same distance) is small compared to the scatter of peak values. Also, the attenuation relationship for the RMS amplitude with distance should be more stable (with less scatter) than the attenuation relationship for peak values. The values given in Table 4.2 (column 5) are the RMS for the various input accelerograms. They are also the zero period value of the "RMS response spectrum". Values of the "RMS response spectrum" for the other periods are given in Appendix B (column 5). Figure 4.4 shows the RMS response spectrum for all the 97 accelerograms considered. On the same plot, the conventional acceleration spectrum (S_a) is also shown. From these plots, the following observations can be made.

- The general shape of the RMS acceleration spectrum is similar to the general shape of the acceleration response spectrum.
- The RMS acceleration spectrum is always smaller than the acceleration response spectrum
- The shape of the RMS acceleration spectrum is "smoother" than the shape of the acceleration response spectrum. The RMS response spectrum represents a statistical summary of all the response peaks and hence has a lower probabilistic scatter.

From the above observations, it can be said that a design spectrum obtained by scaling upward the RMS spectrum (depending on the level of the acceptable probability of exceedence) will be smoother and will contain more probabilistic information than the acceleration spectrum used currently.

It is also worthwhile to note that a total design methodology based on RMS input and response may not be too far-fetched. Thus, hazard maps could be developed where probabilistic information on various RMS amplitude levels could be presented. This could be achieved by developing RMS amplitude attenuation curves and using them in estimating ground motion information at the site. Instead of using acceleration, velocity or displacement response spectra for design, probabilistic response spectra developed by using RMS response spectra (Fig. 4.4) and a multiplying factor could be developed. This aspect of developing probabilistic spectra will be described in the next section.

5.4 Study of the Stable Parameter K_1

In Chapter IV, it was shown that for all the input accelerograms and the response time histories, the shape of the acceleration amplitude distribution was gamma or exponential. Assuming, for simplicity, that the exponential distribution is acceptable, the parameter λ of the distribution is given for all cases considered in Appendix II (column 6). If the assumption of the exponential distribution for the input or response peaks is valid, the following derivation can be made.

Figure 5.2 represents schematically the shape of the peak distribution whether they belong to the input accelerogram or to the response of an oscillator of period T and damping β .

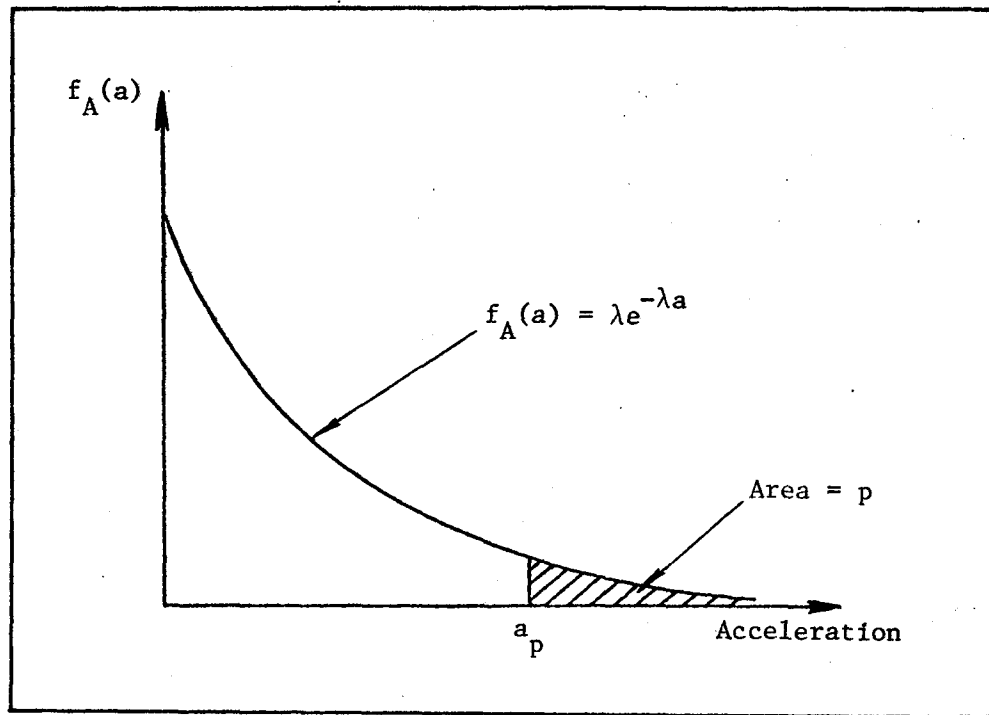


FIGURE 5.2. Exponential Distribution

The RMS value for this distribution is given by

$$\text{RMS} = \frac{\sqrt{2}}{\lambda} \quad (5.1)$$

Considering the acceleration a_p which has $p\%$ chance of being exceeded

$$e^{-\lambda a_p} = p$$

or

$$\lambda a_p = \ln\left(\frac{1}{p}\right)$$

Hence

$$a_p = \frac{1}{\lambda} \ln\left(\frac{1}{p}\right) \quad (5.2)$$

and the ratio

$$\begin{aligned} K_1 &= \frac{a_p}{\text{RMS}} \\ &= \frac{\ln\left(\frac{1}{p}\right)}{\sqrt{2}} \end{aligned} \quad (5.3)$$

From equation 5.3 it can be seen that the ratio K_1 depends only on p and is independent of λ or the individual peaks. Thus, for $p = .05$, K_1 should be 2.12 and for $p = 0.10$, K_1 should be 1.63.

The values of K_1 ($p = 0.05$ and 0.10) for the input accelerograms are given in Table 4.2 (columns 9 and 10). It is worth noting that the values of K_1 for $p = 0.05$ range between 1.33 and 2.51 with a mean of 2.00 and a standard deviation of 0.24. For $p = 0.10$, the values of K_1 range between 1.02 and 1.93 with a mean of 1.54 and a standard deviation of 0.19. A detailed study of the ratio K_1 as a function of the input accelerogram, soil condition, distance, magnitude, etc., is needed. The values of K_1 ($p = 0.05$ and 0.10) for the response time histories, are given in Appendix B (columns 10 and 11). It is remarkable to note that for all earthquakes, all periods and all dampings, this ratio is a constant and very close to the theoretical value obtained using the exponential distribution (eq. 5.3). Figure 5.3 shows the mean and the standard deviation of K_1 versus period for all the accelerogram responses considered.

If the value of the damping varies from 5% to any other value such as 10% or 20%, the value of K_1 does not change appreciably. Figure 5.4 shows a typical behavior of K_1 ($p = 0.10$) as a function of damping. This type of behavior is typical for all the accelerograms and response time histories. From the numerical results presented in Chapter IV, in Appendix B and from the above discussions, the following statements can be made about the ratio K_1 :

- For a given probability of exceedence, the value of K_1 is very stable.

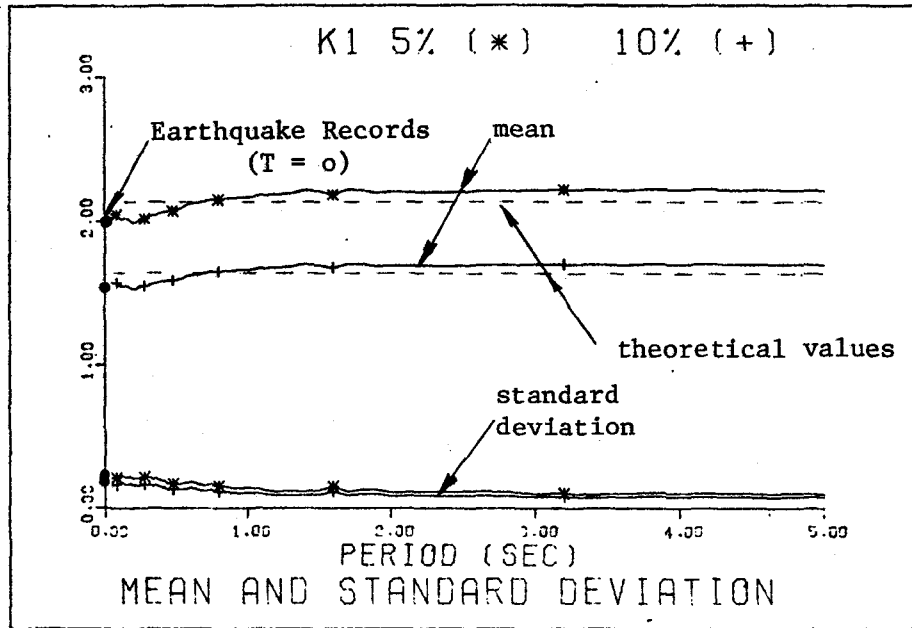


FIGURE 5.3. Factor K_1 (Statistics for 97 Responses)

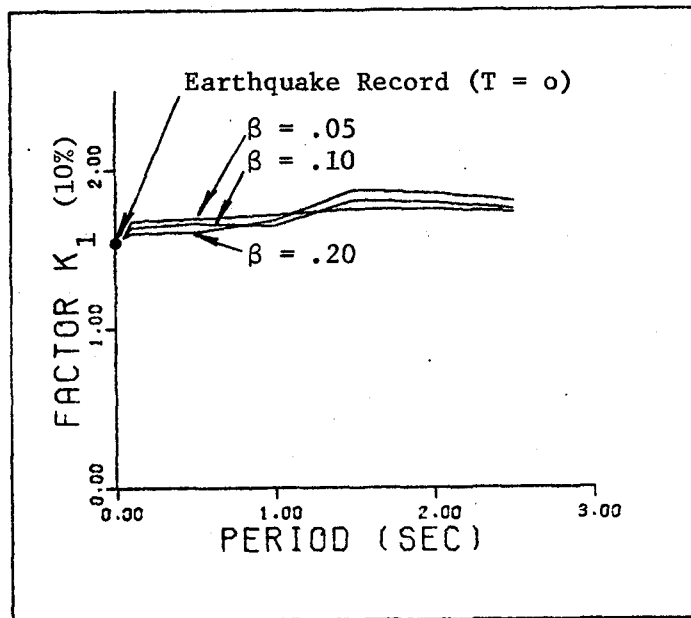


FIGURE 5.4. Typical Behavior of Factor K_1 for Different Dampings β .

- The numerical value of this constant varies with the parameter p but is independent of
 - A) the period of the oscillator
 - B) the damping
 - C) the input accelerogram
- Using the exponential assumption, the theoretical value of K_1 for $p = 0.05$ and 0.10 should be 2.12 and 1.63 respectively. Theoretical and actual values are plotted for comparison in Figure 5.3.
- The remarkable stability of K_1 and the relative stability of RMS can be utilized for deriving response spectra that can be considered as an "equal probability spectra". The values of the equal probability spectra have the same probability of being exceeded for all period bands. In Figures 4.4, the response spectra for all the accelerograms considered in this study were presented. These spectra do not have any probability statement associated with them. Thus development of response spectra which are associated with a predetermined level of non-exceedence essentially constitutes a refinement and an improvement in the current procedure. If for a given region and for a given damping, the shape of RMS versus period can be estimated, then using the stable parameter K_1 any probabilistic spectrum can be generated.

Using

$$a_{p_T} = K_1 \cdot \text{RMS}_T$$

where K_1 depends on p and is constant for all periods and dampings

RMS_T is the RMS corresponding to period T

a_{p_T} is the acceleration corresponding to period T and having a probability p of being exceeded.

A response acceleration spectrum can be developed for a given probability of exceedence P . Figure 5.5 shows the probabilistic spectra developed from the RMS spectra. The multiplication factor K_1 for two values of p is given in Figure 5.3. Thus, the use of K_1 can help in elegantly developing a probabilistic response spectra if the RMS spectrum is given. It is generally felt by seismologists and engineering geologists that a "design RMS spectrum" would be easier to obtain for a given region than the response spectrum. It should be emphasized that much more work is needed on this particular aspect before the stable parameter K_1 can be used for practical development of probabilistic spectrum.

5.5 Study of the Stable Parameter K_2

In Chapter IV, the cumulative potential energy per unit mass was defined and numerically calculated for all the response time histories. Appendix B (column 9) gives the numerical values of this parameter.

In the engineering profession, it is often mentioned that the damage potential of a structure increases with an increase in seismic

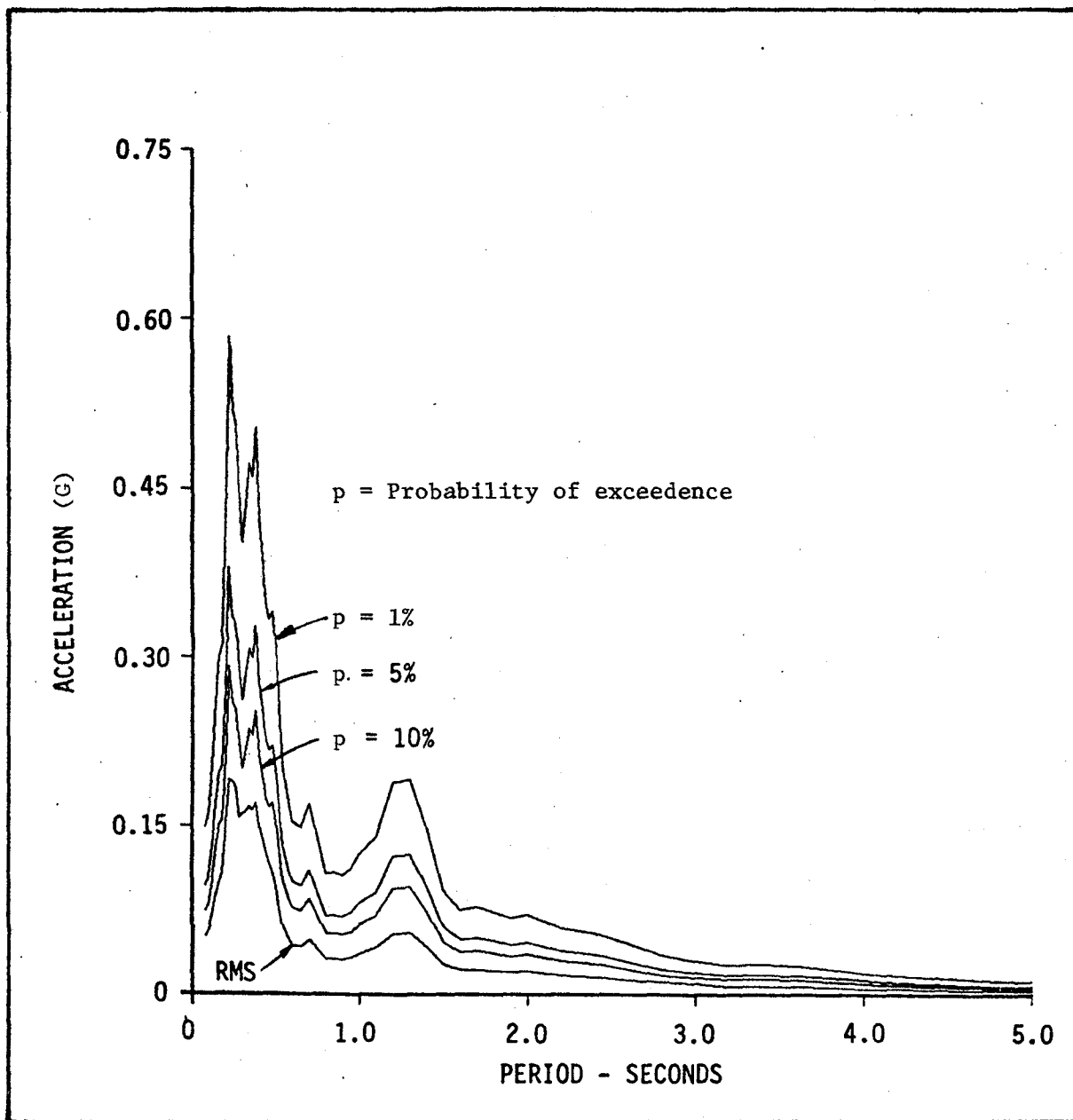


FIGURE 5.5. RMS and Equal Probability Spectra

energy input. Thus, it is quite conceivable that in the future, an energy based method of design could be developed. At the time of this writing, no such comprehensive method is available. Also, if one can develop "energy spectra", it would be desirable to relate the "energy spectra" to the conventional response spectra. For this purpose, another stable parameter (K_2) is evaluated for all the response time histories.

$$K_2 = \frac{\text{RMS}^2 \cdot T^2}{(\text{ENGY}/\text{NEPK})}$$

where T is the period of the oscillator

RMS is the RMS of the response

ENGY is the potential energy of the response

NEPK is the number of peak in the response

It is found that the ratio K_2 is also almost constant for all the periods of the oscillator, Figures 4.6 show the behavior of this ratio for all the records considered. Figure 5.6 shows the mean and the standard deviation of K_2 versus period for all the accelerogram responses considered. It can be seen that this parameter is quite stable. An in-depth study of this ratio and its practical application is beyond the scope of this study. However, one possible use of this parameter could be as follows:

For a given "energy spectrum" and the ratio K_2 , the RMS spectrum could be obtained. Then using the ratio K_1 and the RMS spectrum, a probabilistic response spectrum can be derived. Thus, with the use of K_1 and K_2 , a "mapping" procedure between the currently used parameter (response spectrum) and a future design parameter (RMS or ENGY) can be developed.

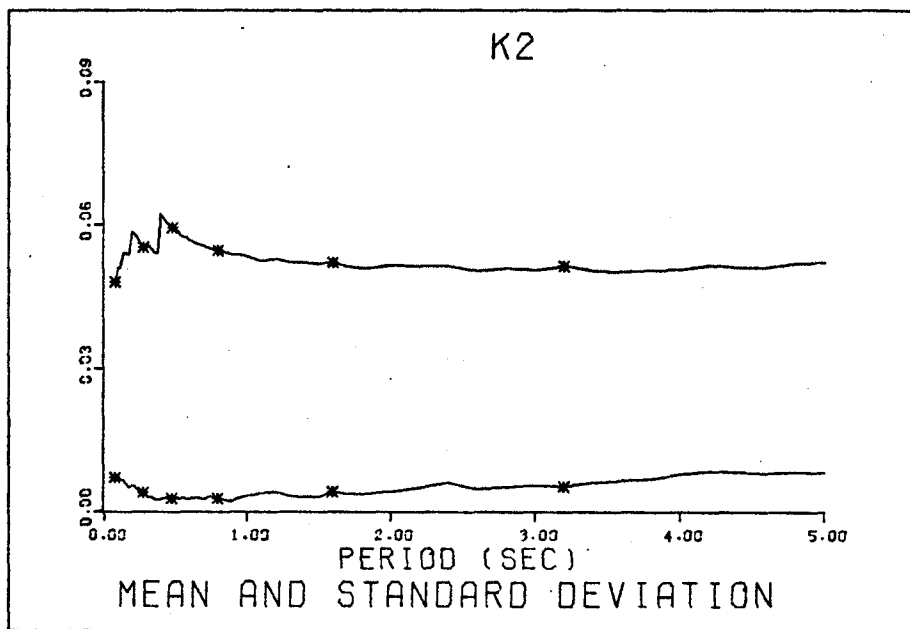


FIGURE 5.6. Factor K_2 (Statistics for 97 Responses)

CHAPTER VI

SUMMARY, CONCLUSIONS AND RECOMMENDATIONS

6.1 Summary of Present Work

In this report, a Bayesian procedure for seismic hazard mapping is developed. As a case study, it is applied to Nicaragua. The method developed has the advantage of combining, in a rational and consistent way, all the available historical, seismological and geological information. This increases the versatility and reliability of the proposed method especially in cases where the historical data is scarce or the time period for which historical data is available is short.

The mapping parameters considered for this study are the peak ground acceleration (PGA) and the duration of the strong ground motion. In order to be consistent with the probabilistic approach of the model and to take into consideration the large scatter for attenuation data (for the peak ground acceleration and the duration), probabilistic treatment of the attenuation relationships in the form of coefficient of variation is presented. A uniform probability distribution function is assumed for the PGA and the duration at a given distance and for a given magnitude.

Comparison between the iso-acceleration maps developed in this study with previous results (Shah, et al., 1975) is made. It is shown that for the same data base, the shape of the iso-acceleration lines is very similar. However, for this study the numerical values of the expected peak ground accelerations change with the degree of uncertainty in attenuation equations. No comparative results are available in the literature. The iso-duration maps are also unique and again no comparable results are available in the literature.

developed in this work may help in dissipating some of those doubts. There are various major faults around the world where not a single major seismic event has occurred in historical time. However, these faults are considered potentially very active. With the previous methods of developing probabilistic hazard maps, the fault would be considered inactive leading to incorrect estimate of the potential hazard. With the Bayesian method presented in this work, the seismicity of the fault could be realistically evaluated.

There may be cases where there is no subjective geological information. In that case, using a diffuse prior and the historical data, the seismicity of the region can be estimated.

- The explicit use of the probability attenuation relationship is another new development in the hazard mapping methodology. Past data on peak ground acceleration or duration, when plotted against distance, have shown great scatter. Using a single curve through these scattered points is not realistic. In the method developed here, the scatter in the data is explicitly incorporated. Thus, the final numerical value of the peak ground acceleration or duration on the hazard map depends not only on the seismicity of the region, the probability of exceedence and the time period for which the maps are drawn, but also on the uncertainty in the attenuation laws used.

- With the available information of the ground motion duration a hazard map for duration can be developed. Such an "iso-duration" map could be useful in determining the damage potential for a given region. Currently, engineers do not use this parameter explicitly in designing their structures. However, it is felt that the duration of the strong motion is an important parameter. Some combination of the peak ground acceleration information from the iso-acceleration map with the duration information from the iso-duration map may increase the correlation of the estimated damage potential with the earthquake severity.

From the second part of this study, where a parameter study of the 97 accelerograms is conducted, the following conclusions can be drawn:

- For the input accelerograms as well as the response acceleration time histories, the probability distribution function for the amplitudes is either gamma or exponential. If one assumes that the exponential shape is adequate, various useful and practical observations can be made.
- There is no standard definition of a duration of an earthquake or of its response even though the importance of this parameter is widely recognized in earthquake engineering. The definition presented in this dissertation is somewhat arbitrary. However it is an essential step to evaluate parameters such as RMS, ENGY, K_1 and K_2 .

- It is shown that the RMS is a better and more stable measure of earthquake ground motion than the currently used peak values. This should be expected since the RMS is a statistical summary and takes into account all the peaks and the duration of the motion. This is true for either input accelerograms or for the response time histories. Also, the uncertainty in RMS is much smaller than the uncertainty in the peak values. Hence the RMS response spectrum is much "smoother" than the conventional response spectrum. This observation has many practical implications.
- Since the peaks of the input or the response time histories have an exponential distribution, the ratio $K_1 = a_p / \text{RMS}$ is a constant and only depends on the value of p . This is proved conclusively by analyzing 97 accelerograms. This observation can be very effectively used for practical applications. As an example, assume that one wishes to obtain a design acceleration spectrum which has probability p of being exceeded. Then, the first step would be to obtain, for the region, an RMS response spectrum for a given damping. As mentioned previously, this step should be more reliable and easier than the current method of obtaining a design response spectrum shape. Having obtained the RMS response spectrum, the constant K_1 for a given p can be used to map the RMS response spectrum to the design response spectrum with probability p of exceedence. This is a simple and elegant way of obtaining a probabilistic design spectrum.

- Since it is also observed that the parameter K_2 is a constant for the 97 accelerograms considered, it can be used to a great advantage. Any future energy related design methodology can be correlated to the current response spectra related design method by the use of K_1 and K_2 . However, it should be confessed at this time that no detailed practical applications of K_2 are presented in this dissertation. The purpose of this overall study is to explore various different ways of hazard mapping and data analysis. The practical applications and possible usage can only be inferred after further research.

6.3 Recommendations for Future Work

During the development of this work, various shortcomings in definitions and in current usage were observed. In spite of those shortcomings, some parameters are used based on current practice. For a better utilization of the methodology presented in this study, the following parameters and ideas need further research.

- Sensitivity analysis of the Bayesian model to source location, seismicity description, type of subjective information and uncertainty in attenuation relations.
- Use of Bayesian model with no historical data
- Development of hazard maps for RMS acceleration, velocity and displacement. This would be a truly useful step towards a better overall design methodology.
- A detailed study of the two stable parameters K_1 and K_2 before their practical applications can be assessed.

- A better definition of the input and response duration is a must. Without a standard and rational definition of duration one cannot obtain a reliable description of the earthquake or its damage potential.
- A better understanding of the energy in input and response time histories. This may help in developing energy based design methods.

Before the ideas presented in this dissertation can be used or implemented, the above suggested work must be completed. Hopefully, at that time, a better way of developing design load criteria will be made available through the application of the work such as the one presented here.

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APPENDIX A

LISTING OF EARTHQUAKES

AJ-13

EARTHQUAKES IN CHRONOLOGICAL ORDER

TIME FIRST DIGIT OF MAGNITUDES IN MS OR RICHTER (CORRESPOND TO UNIT)

Table with columns: D S, T U, A R, A C, E, D M Y H M S, L A T I T U D E, L O N G I T U D E, M B M A G, P S M A G, R I C H T E R M A G, R I C H T E R M A G. Contains earthquake data for stations like DEM 56, DEM 57, etc.

AJ-12

EARTHQUAKES IN CHRONOLOGICAL ORDER

TIME FIRST DIGIT OF MAGNITUDES IN MS OR RICHTER (CORRESPOND TO UNIT)

Table with columns: D S, T U, A R, A C, E, D M Y H M S, L A T I T U D E, L O N G I T U D E, M B M A G, P S M A G, R I C H T E R M A G, R I C H T E R M A G. Contains earthquake data for stations like DEM 56, DEM 57, etc.

APPENDIX B
PARTIAL LISTING OF
RESPONSE PARAMETERS

(A complete listing can be obtained by writing the John A. Blume Earthquake Engineering Center, Department of Civil Engineering, Stanford University, Stanford, CA 94305).

EL CENTRO 1940 SOUTH DAMPING π 5% (STATISTICS BASED ON INCREMENTS)

PERIOD (SEC)	NBPK	DURATION (SEC)	SA	RMS	LAMDA	A* (P=5%)	A* (P=10%)	ENGY	(P=5%)	K1	K2
.08	365	30.31	.578	.0795	18.80	.159	.122	.336	2.003	1.540	.0439
.10	361	30.31	.566	.0900	16.40	.183	.140	.615	2.029	1.560	.0476
.12	359	30.30	.636	.1036	14.21	.211	.162	1.086	2.034	1.564	.0511
.14	347	30.30	.704	.1209	12.04	.249	.191	1.771	2.058	1.582	.0561
.16	349	32.16	.543	.1250	11.18	.268	.206	2.439	2.142	1.647	.0573
.18	279	28.54	.731	.1306	10.67	.281	.216	2.783	2.151	1.653	.0554
.20	279	31.63	.648	.1443	9.95	.301	.232	3.887	2.088	1.605	.0598
.22	253	31.78	.686	.1454	9.59	.312	.240	4.503	2.149	1.651	.0575
.24	243	31.68	.860	.1534	8.96	.334	.257	5.616	2.178	1.674	.0587
.26	223	30.68	.905	.1719	7.64	.392	.301	7.701	2.281	1.753	.0579
.28	223	32.01	.759	.1524	8.72	.344	.264	6.834	2.255	1.713	.0594
.30	207	32.07	.708	.1460	9.60	.312	.240	6.721	2.137	1.642	.0591
.32	211	34.69	.702	.1481	10.52	.285	.219	7.300	2.033	1.562	.0581
.34	183	35.33	.649	.1277	11.55	.259	.199	6.844	2.031	1.561	.0519
.36	181	35.56	.658	.1342	10.99	.233	.210	7.860	2.032	1.562	.0537
.38	177	35.59	.677	.1444	10.09	.237	.228	9.684	2.055	1.560	.0550
.40	163	36.02	.615	.1471	9.60	.312	.240	9.237	2.121	1.630	.0611
.42	165	36.98	.600	.1502	4.33	.321	.242	10.634	2.137	1.645	.0617
.44	159	37.04	.730	.1690	8.46	.354	.272	14.374	2.096	1.611	.0611
.46	155	37.08	.864	.1858	7.77	.366	.296	18.706	2.075	1.595	.0605
.48	151	37.16	.796	.1740	8.13	.369	.283	16.216	2.118	1.628	.0600
.50	143	36.54	.855	.1672	8.59	.319	.268	16.716	2.087	1.604	.0598
.52	135	35.80	.880	.1730	8.22	.364	.280	18.312	2.106	1.619	.0590
.54	131	35.86	.907	.1766	8.37	.358	.275	20.283	2.027	1.558	.0586
.56	125	35.12	.913	.1771	8.44	.355	.273	20.890	2.004	1.540	.0588
.58	115	32.42	.893	.1743	8.46	.354	.272	19.882	2.033	1.563	.0603
.60	109	31.88	.858	.1628	9.21	.325	.250	17.606	1.998	1.536	.0591
.65	101	30.58	.737	.1778	7.85	.362	.293	22.215	2.146	1.649	.0607
.70	109	37.18	.621	.1612	9.25	.334	.249	24.215	2.009	1.544	.0573
.75	101	37.22	.583	.1261	11.98	.250	.192	16.146	1.983	1.524	.0560
.80	93	37.24	.549	.1336	11.63	.258	.198	19.306	1.927	1.481	.0550
.85	67	28.64	.598	.1768	7.42	.404	.310	28.855	2.257	1.735	.0536
.90	85	37.30	.536	.1410	11.20	.267	.206	24.743	1.897	1.458	.0553
.95	79	37.36	.541	.1190	13.61	.250	.169	18.926	1.850	1.422	.0534
1.00	71	32.90	.517	.1092	15.03	.199	.156	15.115	1.825	1.403	.0560
1.10	69	37.98	.386	.0966	14.76	.203	.156	15.823	2.058	1.582	.0513
1.20	63	38.12	.331	.0952	14.11	.212	.163	16.138	2.230	1.714	.0510
1.30	65	41.36	.237	.0700	18.23	.164	.126	18.505	2.349	1.806	.0512
1.40	57	41.16	.182	.0595	21.65	.137	.105	8.297	2.305	1.772	.0476
1.50	53	41.38	.190	.0557	24.16	.124	.095	7.750	2.226	1.711	.0477
1.60	53	40.20	.195	.0528	28.47	.105	.081	7.052	1.993	1.532	.0537
1.70	47	40.32	.187	.0558	26.20	.114	.088	8.555	2.048	1.574	.0495
1.80	43	40.68	.173	.0604	21.87	.117	.105	10.668	2.269	1.744	.0476
1.90	41	36.64	.173	.0559	25.84	.116	.089	9.462	2.072	1.593	.0489
2.00	39	40.02	.179	.0538	26.64	.112	.086	9.294	2.090	1.606	.0486
2.20	27	28.12	.189	.0564	24.68	.121	.093	8.676	2.153	1.655	.0479
2.40	27	30.36	.190	.0497	30.75	.097	.075	7.622	1.960	1.506	.0491
2.60	31	40.98	.162	.0993	27.44	.109	.084	10.444	2.213	1.701	.0489
2.80	27	37.70	.140	.0561	22.67	.132	.102	13.744	2.355	1.610	.0485
3.00	27	40.94	.115	.0661	27.76	.108	.083	10.946	2.313	1.778	.0483
3.20	33	51.32	.100	.0334	37.76	.079	.061	7.831	2.379	1.628	.0480
3.40	29	50.82	.084	.0226	57.45	.052	.040	3.914	2.311	1.776	.0436
3.60	33	51.94	.067	.0174	75.76	.040	.030	2.663	2.272	1.747	.0486
3.80	31	51.26	.053	.0147	89.39	.034	.026	2.072	2.280	1.752	.0467
4.00	25	44.74	.046	.0145	90.97	.033	.025	1.884	2.266	1.742	.0488
4.20	23	42.50	.041	.0130	101.46	.030	.023	1.509	2.266	1.742	.0456
4.40	23	49.46	.036	.0104	127.68	.023	.018	1.202	2.254	1.732	.0402
4.60	23	55.32	.032	.0098	135.84	.023	.017	1.206	2.293	1.762	.0385
4.80	27	59.76	.030	.0100	131.33	.022	.018	1.256	2.286	1.757	.0457
5.00	33	65.26	.030	.0097	139.75	.021	.016	1.444	2.217	1.704	.0534

(A001)

EL CENTRO 1940 WEST DAMPING B 5% (STATISTICS BASED ON INCREMENTS)

PERIOD (SEC)	NSPK	DURATION (SEC)	SA	RMS	LAMDA	A* (P=5%)	A* (P=10%)	ENGY	K1 (P=5%)	K1 (P=10%)	K2
.08	336	32.04	.292	.0587	23.07	.130	.100	.163	2.210	1.699	.0458
.10	348	32.00	.330	.0648	20.85	.144	.110	.293	2.218	1.705	.0498
.12	328	32.00	.329	.0749	18.12	.165	.127	.570	2.209	1.698	.0464
.14	326	32.09	.384	.0851	15.83	.189	.145	.884	2.223	1.709	.0524
.16	338	31.93	.416	.1067	12.53	.239	.184	1.748	2.240	1.721	.0564
.18	306	31.07	.450	.1181	11.18	.268	.206	2.488	2.270	1.745	.0565
.20	284	31.06	.510	.1255	10.74	.279	.214	2.488	2.223	1.709	.0614
.22	246	31.09	.584	.1240	9.72	.270	.207	3.212	2.177	1.673	.0570
.24	242	31.61	.470	.1380	8.98	.308	.237	4.555	2.235	1.718	.0582
.26	230	32.30	.485	.1424	8.98	.333	.256	5.536	2.333	1.793	.0572
.28	216	32.21	.415	.1334	9.77	.307	.236	5.332	2.298	1.766	.0565
.30	206	32.27	.427	.1279	10.03	.299	.230	5.275	2.336	1.795	.0575
.32	190	31.77	.452	.1329	9.59	.312	.240	6.038	2.351	1.807	.0569
.34	184	32.02	.396	.1326	9.74	.308	.237	6.490	2.321	1.784	.0576
.36	174	32.42	.445	.1355	9.91	.302	.232	7.350	2.232	1.715	.0563
.38	166	32.68	.481	.1394	9.48	.316	.243	8.370	2.268	1.743	.0557
.40	152	32.38	.566	.1569	8.17	.367	.282	9.938	2.309	1.775	.0618
.42	150	32.26	.502	.1684	7.49	.400	.307	12.009	2.375	1.825	.0625
.44	142	32.16	.544	.1645	7.79	.384	.295	12.157	2.337	1.796	.0612
.46	142	32.70	.590	.1665	8.47	.354	.272	11.881	2.259	1.737	.0620
.48	136	32.80	.622	.1450	9.29	.323	.268	10.816	2.225	1.710	.0609
.50	134	32.86	.638	.1456	9.58	.314	.240	11.365	2.147	1.650	.0625
.52	134	32.68	.638	.1524	9.53	.324	.242	13.193	2.063	1.586	.0638
.54	112	30.98	.625	.1569	9.36	.320	.246	13.988	2.040	1.568	.0575
.56	110	30.76	.604	.1531	9.48	.316	.243	13.766	2.063	1.586	.0587
.58	108	30.80	.585	.1491	9.61	.312	.240	13.686	2.091	1.607	.0590
.60	102	30.78	.569	.1438	9.62	.311	.239	13.211	2.166	1.665	.0575
.62	108	33.58	.452	.1196	11.40	.263	.202	11.114	2.190	1.689	.0587
.70	96	33.02	.430	.1065	12.65	.237	.182	9.533	2.224	1.710	.0561
.75	86	32.82	.433	.0979	14.29	.210	.161	8.670	2.255	1.733	.0569
.80	78	30.10	.435	.1112	11.94	.251	.193	10.857	2.141	1.646	.0535
.85	78	31.50	.337	.0964	14.10	.212	.163	9.157	2.205	1.695	.0571
.90	70	31.22	.342	.0911	15.06	.199	.153	8.588	2.185	1.679	.0550
.95	68	32.52	.284	.0832	15.21	.197	.151	7.846	2.367	1.819	.0540
1.00	58	30.20	.278	.0791	16.39	.185	.141	7.049	2.312	1.771	.0513
1.10	50	26.90	.270	.0847	15.62	.189	.146	8.575	2.236	1.718	.0506
1.20	50	29.86	.348	.1098	11.74	.255	.196	16.463	2.325	1.787	.0527
1.30	50	31.44	.285	.0940	14.08	.213	.164	14.073	2.263	1.740	.0531
1.40	52	36.68	.214	.0827	15.48	.194	.149	13.830	2.341	1.800	.0504
1.50	54	38.82	.190	.0614	20.88	.143	.110	8.515	2.338	1.797	.0537
1.60	50	39.58	.157	.0561	22.35	.134	.103	7.713	2.389	1.836	.0522
1.70	40	36.14	.145	.0501	25.64	.117	.090	5.957	2.332	1.792	.0487
1.80	34	32.56	.147	.0570	22.10	.136	.104	7.310	2.377	1.827	.0490
1.90	34	33.16	.187	.0670	19.78	.151	.116	11.317	2.260	1.737	.0487
2.00	32	34.26	.218	.0758	18.42	.163	.125	15.260	2.146	1.650	.0482
2.20	34	39.10	.189	.0768	16.29	.184	.141	20.835	2.393	1.840	.0466
2.40	32	36.86	.156	.0519	25.22	.119	.091	9.702	2.288	1.759	.0512
2.60	24	33.06	.147	.0512	26.36	.114	.087	9.255	2.220	1.706	.0459
2.80	22	29.96	.148	.0561	23.52	.128	.099	11.006	2.290	1.760	.0493
3.00	22	36.66	.142	.0485	27.94	.107	.082	10.574	2.209	1.698	.0441
3.20	26	41.60	.133	.0432	30.00	.100	.077	10.575	2.314	1.778	.0469
3.40	26	44.52	.123	.0351	37.03	.081	.062	8.069	2.307	1.773	.0458
3.60	22	43.82	.113	.0266	49.52	.060	.046	5.451	2.117	1.627	.0427
3.80	24	48.50	.103	.0264	50.75	.059	.045	5.112	2.240	1.622	.0471
4.00	20	43.40	.094	.0253	53.60	.056	.043	4.678	2.211	1.600	.0438
4.20	24	50.22	.085	.0248	51.86	.058	.044	5.210	2.333	1.793	.0498
4.40	26	57.68	.078	.0241	54.70	.055	.042	5.457	2.273	1.747	.0491
4.60	24	57.24	.071	.0241	58.91	.051	.039	5.788	2.221	1.707	.0460
4.80	24	58.38	.067	.0250	50.80	.059	.045	7.150	2.363	1.816	.0483
5.00	26	64.74	.066	.0265	43.89	.068	.052	10.911	2.398	1.843	.0483

(A001)

FERNDALE 1951 S44W DAMPING = 5% (STATISTICS BASED ON INCREMENTS)

PERIOD (SEC)	NBPK	DURATION (SEC)	SA	RMS	LAMDA	A* (P=5%)	A* (P=10%)	ENGY	(P=5%)	M1 (P=10%)	K2
.08	125	9.46	.138	.0293	46.39	.065	.050	.015	2.202	1.693	.0473
.10	123	9.59	.165	.0375	37.51	.080	.061	.045	2.132	1.639	.0487
.12	121	9.69	.172	.0412	35.09	.091	.070	.058	2.197	1.669	.0522
.14	115	9.87	.240	.0570	23.84	.126	.097	.133	2.204	1.694	.0550
.16	113	10.03	.239	.0582	23.15	.129	.099	.169	2.223	1.709	.0580
.18	107	9.95	.243	.0752	17.71	.169	.130	.333	2.248	1.728	.0590
.20	87	9.69	.296	.0729	20.30	.148	.113	.317	2.025	1.257	.0582
.22	89	9.64	.191	.0672	19.35	.155	.119	.302	2.306	1.772	.0642
.24	69	9.42	.194	.0622	21.67	.138	.106	.295	2.224	1.709	.0538
.26	71	9.96	.221	.0598	22.69	.132	.101	.307	2.207	1.696	.0559
.28	69	10.17	.236	.0634	22.23	.135	.104	.388	2.126	1.634	.0564
.30	71	10.09	.183	.0601	21.66	.136	.106	.394	2.302	1.770	.0565
.32	65	10.34	.169	.0567	22.53	.133	.102	.368	2.345	1.803	.0582
.34	63	10.67	.145	.0524	24.45	.123	.094	.347	2.339	1.798	.0576
.36	54	10.46	.146	.0508	21.74	.137	.108	.510	2.269	1.744	.0551
.38	57	10.88	.239	.0693	19.16	.156	.120	.709	2.256	1.734	.0558
.40	53	10.60	.241	.0638	21.49	.139	.107	.511	2.185	1.680	.0650
.42	49	9.90	.238	.0676	19.96	.150	.115	.595	2.219	1.705	.0665
.44	54	10.74	.231	.0641	21.23	.141	.108	.637	2.202	1.692	.0637
.46	51	11.10	.201	.0517	26.33	.114	.087	.455	2.202	1.692	.0633
.48	49	11.92	.165	.0462	30.15	.099	.076	.435	2.246	1.727	.0589
.50	51	12.28	.143	.0424	31.57	.095	.073	.365	2.239	1.721	.0628
.52	49	12.68	.143	.0464	29.00	.103	.079	.468	2.226	1.711	.0610
.54	49	12.56	.152	.0501	26.62	.115	.087	.515	2.245	1.725	.0624
.56	45	12.16	.135	.0483	28.63	.105	.080	.543	2.165	1.664	.0607
.58	39	11.02	.136	.0467	29.21	.103	.079	.476	2.195	1.687	.0601
.60	41	12.56	.150	.0417	31.13	.096	.074	.455	2.308	1.774	.0564
.62	43	14.20	.112	.0403	31.64	.095	.073	.533	2.348	1.805	.0554
.70	41	14.84	.125	.0492	25.93	.116	.089	.694	2.349	1.806	.0544
.75	35	12.86	.120	.0438	27.83	.100	.077	.619	2.295	1.764	.0555
.80	35	13.34	.084	.0279	46.06	.065	.050	.321	2.329	1.790	.0545
.85	39	14.92	.062	.0206	65.37	.047	.036	.211	2.299	1.767	.0565
.90	37	16.54	.053	.0172	77.12	.039	.030	.152	2.261	1.758	.0582
.95	41	16.00	.044	.0159	82.51	.036	.028	.150	2.277	1.750	.0620
1.00	45	17.40	.047	.0156	85.31	.035	.027	.165	2.246	1.726	.0667
1.10	41	18.82	.035	.0112	119.02	.025	.019	.105	2.246	1.726	.0593
1.20	39	19.08	.033	.0103	127.59	.024	.018	.101	2.278	1.751	.0590
1.30	31	19.50	.030	.0113	115.73	.026	.020	.132	2.285	1.756	.0509
1.40	31	19.50	.030	.0106	122.21	.025	.019	.126	2.302	1.769	.0547
1.50	25	17.36	.025	.0081	161.77	.019	.014	.074	2.290	1.760	.0495
1.60	17	11.42	.016	.0068	189.70	.016	.012	.059	2.311	1.776	.0523
1.70	25	16.26	.015	.0048	268.55	.011	.009	.029	2.310	1.776	.0575
1.80	27	19.48	.014	.0043	504.21	.010	.008	.031	2.279	1.752	.0529
1.90	19	16.30	.014	.0043	297.25	.010	.008	.029	2.330	1.791	.0409
2.00	15	14.02	.014	.0047	274.93	.011	.008	.029	2.315	1.779	.0457
2.20	19	20.06	.012	.0041	321.48	.009	.007	.034	2.272	1.747	.0456
2.40	23	22.60	.013	.0038	535.50	.009	.007	.030	2.264	1.740	.0551
2.60	15	16.22	.014	.0039	557.16	.006	.006	.035	2.149	1.652	.0523
2.80	13	17.00	.013	.0040	340.29	.009	.007	.038	2.181	1.676	.0424
3.00	19	20.32	.011	.0037	551.80	.009	.007	.033	2.292	1.762	.0441
3.20	23	34.64	.010	.0036	349.19	.007	.007	.070	2.269	1.744	.0463
3.40	25	36.44	.010	.0037	354.48	.008	.006	.075	2.256	1.734	.0541
3.60	23	40.04	.009	.0034	397.09	.008	.006	.070	2.246	1.726	.0477
3.80	23	40.02	.008	.0030	443.44	.007	.005	.057	2.284	1.756	.0507
4.00	21	39.80	.006	.0025	521.94	.006	.004	.043	2.301	1.769	.0487
4.20	21	39.56	.005	.0020	647.73	.005	.004	.031	2.274	1.748	.0498
4.40	19	36.48	.005	.0017	802.19	.004	.004	.021	2.230	1.714	.0483
4.60	15	30.62	.005	.0014	973.35	.003	.002	.014	2.124	1.632	.0463
4.80	17	24.30	.004	.0014	1104.15	.003	.002	.011	1.969	1.513	.0695
5.00	15	24.18	.004	.0013	1162.48	.003	.002	.010	1.982	1.523	.0551

DAMPING = 5% (STATISTICS BASED ON INCREMENTS)

FERNOALE 1951 N46H

PERIOD (SEC)	MBPK	DURATION (SEC)	SA	RMS	LAMDA	A* (P=5%) A* (P=10%)	ENG	(P=5%)	K1 (P=10%)	K2
.08	116	10.31	.126	.0291	46.39	.065	.013	2.216	1.704	.0900
.10	134	10.24	.146	.0361	37.15	.081	.033	2.231	1.715	.0532
.12	128	10.34	.160	.0401	34.63	.087	.054	2.159	1.659	.0544
.14	116	10.67	.163	.0440	30.93	.097	.074	2.202	1.642	.0556
.16	110	10.19	.233	.0634	22.20	.135	.104	2.111	1.623	.0556
.18	96	10.15	.362	.0753	19.03	.157	.121	2.090	1.606	.0528
.20	94	10.49	.273	.0678	21.05	.142	.109	2.100	1.614	.0591
.22	86	10.47	.259	.0661	21.15	.142	.109	2.141	1.614	.0543
.24	74	10.30	.237	.0683	20.39	.147	.113	2.152	1.654	.0542
.26	74	10.33	.207	.0659	21.31	.141	.106	2.134	1.640	.0571
.30	64	10.65	.337	.0779	17.57	.170	.131	2.188	1.682	.0571
.32	66	11.36	.272	.0989	14.57	.206	.158	2.079	1.682	.0528
.34	64	11.54	.220	.0743	17.90	.167	.129	2.162	1.622	.0531
.36	58	11.58	.202	.0566	24.49	.122	.094	2.162	1.622	.0529
.38	48	10.70	.235	.0534	25.74	.116	.089	2.179	1.675	.0528
.40	50	10.90	.269	.0648	22.05	.136	.104	2.097	1.612	.0495
.42	46	10.94	.313	.0789	18.64	.159	.122	2.015	1.549	.0603
.44	46	10.52	.283	.0879	16.28	.140	.111	2.092	1.608	.0565
.46	50	11.70	.255	.0886	15.73	.190	.146	2.051	1.608	.0565
.48	44	11.56	.226	.0752	19.03	.157	.121	2.091	1.607	.0594
.50	46	11.64	.193	.0630	23.37	.128	.099	2.033	1.563	.0543
.52	38	10.36	.164	.0547	25.80	.116	.089	2.123	1.632	.0573
.54	44	12.16	.147	.0515	26.68	.112	.086	2.165	1.664	.0536
.56	44	13.36	.137	.0446	31.66	.095	.073	2.145	1.629	.0574
.58	48	14.10	.136	.0441	31.65	.095	.073	2.192	1.649	.0549
.60	46	14.56	.136	.0466	29.32	.102	.079	2.192	1.665	.0592
.65	48	14.98	.172	.0480	27.02	.107	.082	2.234	1.717	.0560
.70	46	15.44	.134	.0577	23.23	.127	.098	2.207	1.696	.0592
.75	38	15.44	.130	.0484	27.99	.109	.084	2.252	1.731	.0589
.80	38	15.88	.150	.0494	26.55	.113	.087	2.286	1.757	.0523
.85	38	15.92	.132	.0571	23.04	.130	.100	2.277	1.750	.0520
.90	32	15.00	.115	.0449	29.97	.100	.077	2.227	1.712	.0530
.95	30	14.50	.116	.0377	35.97	.083	.064	2.206	1.696	.0495
1.00	32	14.46	.113	.0371	37.75	.079	.061	2.141	1.645	.0516
1.10	30	15.78	.090	.0366	38.11	.074	.060	2.141	1.641	.0581
1.20	22	11.28	.072	.0305	46.05	.065	.050	2.130	1.637	.0528
1.30	22	15.28	.061	.0247	59.16	.051	.039	2.051	1.576	.0568
1.40	26	16.64	.053	.0180	78.98	.038	.029	2.104	1.617	.0455
1.50	18	12.22	.044	.0155	91.46	.033	.025	2.115	1.626	.0496
1.60	22	16.26	.037	.0148	97.28	.031	.024	2.077	1.596	.0475
1.70	24	19.44	.032	.0099	124.61	.024	.018	2.094	1.610	.0461
1.80	22	19.28	.031	.0087	138.03	.022	.017	2.203	1.693	.0474
1.90	16	15.22	.030	.0082	158.09	.019	.015	2.170	1.668	.0465
2.00	12	10.76	.029	.0082	166.74	.018	.014	2.163	1.662	.0433
2.20	18	21.04	.025	.0064	162.02	.018	.014	2.113	1.624	.0451
2.40	18	22.38	.019	.0055	212.70	.014	.011	2.210	1.698	.0414
2.60	8	10.10	.014	.0070	252.68	.016	.013	2.156	1.657	.0447
2.80	20	24.68	.013	.0048	282.12	.011	.008	2.347	1.804	.0474
3.00	22	32.20	.012	.0047	263.88	.011	.008	2.216	1.793	.0580
3.20	22	35.52	.013	.0047	289.52	.010	.006	2.217	1.719	.0491
3.40	24	35.60	.012	.0045	295.69	.010	.008	2.217	1.704	.0450
3.60	22	37.28	.011	.0041	295.69	.009	.007	2.246	1.726	.0508
3.80	20	37.08	.011	.0041	326.47	.009	.007	2.251	1.730	.0470
4.00	18	34.68	.010	.0032	371.12	.008	.006	2.251	1.730	.0464
4.20	18	34.02	.009	.0027	415.75	.006	.005	2.267	1.743	.0438
4.40	14	28.60	.009	.0025	505.24	.006	.005	2.185	1.674	.0437
4.60	12	26.92	.008	.0025	579.96	.005	.004	2.056	1.561	.0404
4.80	14	20.94	.007	.0026	589.51	.005	.004	1.936	1.488	.0486
5.00	16	28.56	.007	.0022	609.48	.005	.004	1.914	1.471	.0615
					674.41	.004	.003	2.021	1.553	.0530

PASADENA 1952 SOUTH DAMPING = 5% (STATISTICS BASED ON INCREMENTS)

PERIOD (SEC)	NBPK	DURATION (SEC)	SA	RMS	LAMDA	A* (P=5%) A* (P=10%)	ENG	K1 (P=5%)	K2 (P=10%)
.08	250	30.62	.048	.0106	137.67	.022	.003	2.059	.0667
.10	294	30.74	.054	.0111	128.20	.023	.005	2.111	1.583
.12	296	30.97	.053	.0117	118.19	.025	.009	2.159	.0664
.14	296	30.86	.050	.0127	109.99	.027	.014	2.151	.0638
.16	246	30.54	.059	.0128	109.88	.028	.017	2.131	.0642
.18	226	30.58	.063	.0134	106.50	.028	.024	2.092	.0598
.20	222	30.55	.073	.0149	94.16	.032	.034	2.129	.0545
.22	232	30.63	.080	.0172	79.85	.038	.054	2.177	.0576
.24	220	31.39	.094	.0191	74.14	.040	.079	2.114	.0610
.26	208	31.49	.093	.0198	68.80	.044	.096	2.199	.0576
.28	202	31.66	.090	.0214	63.19	.047	.126	2.215	.0517
.30	194	31.83	.092	.0233	61.53	.049	.166	2.088	.0572
.32	172	31.13	.111	.0253	58.62	.051	.211	2.023	.0533
.34	168	31.15	.096	.0237	60.65	.049	.200	2.080	.0547
.36	152	31.73	.101	.0212	67.16	.045	.169	2.107	.0523
.38	150	31.81	.093	.0209	67.65	.044	.172	2.118	.0549
.40	146	31.68	.106	.0230	60.65	.049	.196	2.144	.0631
.42	138	31.72	.104	.0265	54.13	.056	.282	2.128	.0606
.44	134	32.46	.117	.0284	54.88	.055	.304	2.069	.0595
.46	126	32.58	.119	.0253	57.60	.052	.293	2.059	.0581
.48	124	32.96	.113	.0270	54.62	.055	.356	2.033	.0583
.50	118	32.84	.134	.0317	46.70	.064	.527	2.024	.0563
.52	122	32.72	.155	.0351	42.48	.071	.527	2.008	.0594
.54	118	32.58	.155	.0350	44.11	.068	.712	1.941	.0592
.56	112	32.70	.155	.0333	45.64	.066	.681	1.968	.0573
.58	112	32.52	.154	.0310	49.41	.061	.612	1.953	.0593
.60	104	32.30	.156	.0321	49.41	.061	.674	1.962	.0574
.65	98	32.54	.177	.0406	38.21	.078	1.210	1.932	.0564
.70	92	31.96	.131	.0372	42.28	.071	1.112	1.906	.0560
.75	86	32.10	.107	.0289	52.71	.057	.728	1.966	.0553
.80	84	34.94	.082	.0246	59.23	.051	.612	2.054	.0553
.85	82	36.14	.090	.0261	54.60	.055	.766	2.103	.0526
.90	62	36.54	.098	.0292	49.92	.060	1.037	2.055	.0546
.95	76	36.48	.094	.0246	60.22	.050	.778	2.025	.0532
1.00	72	35.80	.089	.0231	64.09	.047	.720	2.025	.0533
1.10	68	36.82	.062	.0207	68.43	.044	.665	2.111	.0532
1.20	66	39.30	.070	.0218	64.59	.046	.857	2.124	.0528
1.30	62	38.92	.064	.0200	69.87	.043	.810	2.139	.0520
1.40	50	35.26	.046	.0150	93.23	.032	.442	2.139	.0500
1.50	54	39.92	.048	.0162	85.58	.035	.622	2.156	.0515
1.60	54	41.52	.056	.0170	83.93	.036	.762	2.095	.0527
1.70	48	41.48	.052	.0156	91.11	.033	.701	2.104	.0484
1.80	44	35.06	.065	.0181	84.44	.035	.835	1.959	.0560
1.90	40	38.62	.065	.0179	83.90	.036	.924	1.997	.0500
2.00	48	44.24	.055	.0165	87.33	.034	.965	2.077	.0542
2.20	46	48.16	.039	.0129	108.05	.028	.737	2.145	.0505
2.40	40	47.56	.028	.0104	133.97	.022	.509	2.158	.0486
2.60	28	33.94	.035	.0124	115.72	.026	.575	2.080	.0509
2.80	34	45.60	.031	.0108	131.92	.023	.615	2.102	.0506
3.00	34	47.14	.026	.0065	174.98	.017	.425	2.020	.0510
3.20	30	39.68	.024	.0070	217.19	.014	.278	1.971	.0541
3.40	36	50.70	.021	.0060	232.98	.013	.288	2.156	.0517
3.60	38	55.78	.017	.0050	276.04	.011	.235	2.165	.0526
3.80	34	55.46	.013	.0037	370.34	.008	.146	2.161	.0485
4.00	34	54.70	.010	.0028	498.41	.006	.086	2.133	.0501
4.20	32	57.94	.008	.0024	575.56	.005	.069	2.179	.0469
4.40	32	59.32	.007	.0021	634.75	.005	.060	2.219	.0470
4.60	34	58.28	.006	.0019	738.28	.004	.051	2.170	.0496
4.80	28	54.30	.007	.0018	775.98	.004	.048	2.152	.0435
5.00	30	60.02	.006	.0018	757.68	.004	.052	2.230	.0456

PASADENA 1952 WEST DAMPING = 5% (STATISTICS BASED ON INCREMENTS)

PERIOD (SEC)	NBPK	DURATION (SEC)	SA	RMS	LAMDA	A* (P=5%)	A* (P=10%)	ENDY	K1 (P=5%)	K1 (P=10%)	K2
.08	164	30.26	.056	.0146	99.18	.030	.023	.008	2.064	1.587	.0565
.10	168	30.24	.058	.0149	96.42	.031	.024	.007	2.080	1.599	.0604
.12	212	30.26	.059	.0154	92.57	.032	.025	.011	2.102	1.615	.0641
.14	224	30.23	.075	.0163	86.94	.034	.026	.020	2.110	1.622	.0580
.16	216	30.49	.073	.0169	81.06	.037	.028	.025	2.192	1.665	.0628
.18	232	30.25	.079	.0186	74.15	.040	.031	.042	2.176	1.672	.0612
.20	232	30.19	.089	.0205	67.66	.044	.034	.063	2.156	1.657	.0620
.22	228	31.18	.085	.0213	65.24	.046	.035	.081	2.158	1.659	.0613
.24	218	31.42	.099	.0236	58.32	.051	.039	.117	2.181	1.676	.0594
.26	200	31.27	.110	.0245	56.59	.053	.041	.142	2.163	1.663	.0569
.28	188	31.50	.098	.0250	54.83	.055	.042	.161	2.183	1.678	.0574
.30	186	31.90	.110	.0286	48.16	.062	.048	.245	2.172	1.669	.0561
.32	182	31.56	.138	.0311	45.50	.066	.051	.316	2.116	1.626	.0572
.34	166	31.14	.140	.0337	42.66	.070	.054	.388	2.085	1.603	.0547
.36	164	31.71	.130	.0296	46.70	.064	.049	.336	2.151	1.653	.0563
.38	148	32.01	.111	.0271	51.14	.059	.045	.289	2.179	1.675	.0543
.40	142	31.90	.112	.0270	51.14	.059	.045	.263	2.173	1.670	.0629
.42	132	31.78	.131	.0266	48.58	.062	.048	.321	2.164	1.663	.0544
.44	132	31.64	.138	.0329	41.92	.071	.056	.450	2.175	1.672	.0613
.46	136	31.72	.127	.0378	37.57	.080	.062	.681	2.123	1.632	.0642
.48	128	31.56	.139	.0374	37.57	.080	.061	.661	2.133	1.639	.0624
.50	120	31.16	.122	.0338	41.86	.073	.056	.535	2.139	1.659	.0617
.52	108	30.12	.156	.0346	40.79	.073	.056	.589	2.124	1.632	.0593
.54	108	30.10	.156	.0381	38.12	.079	.060	.756	2.061	1.584	.0606
.56	104	31.52	.156	.0398	37.74	.079	.061	.913	1.995	1.534	.0546
.58	112	32.50	.153	.0392	38.31	.078	.060	.924	1.994	1.532	.0607
.60	110	32.64	.168	.0380	39.92	.075	.058	.929	1.976	1.514	.0614
.62	94	32.58	.215	.0459	34.25	.087	.067	1.500	1.906	1.465	.0558
.64	94	32.80	.199	.0598	26.28	.114	.088	2.882	1.905	1.464	.0570
.66	90	35.76	.183	.0476	31.17	.096	.074	2.054	2.019	1.551	.0559
.68	84	32.46	.168	.0492	30.45	.098	.076	2.246	2.007	1.537	.0579
.70	80	33.40	.171	.0505	31.26	.096	.074	2.625	1.899	1.459	.0561
.72	74	33.12	.199	.0532	28.94	.104	.080	3.083	1.946	1.496	.0550
.74	78	35.28	.185	.0571	25.72	.116	.090	4.067	2.039	1.564	.0564
.76	74	34.66	.168	.0489	31.21	.096	.074	3.131	1.964	1.509	.0565
.78	62	34.16	.122	.0375	40.04	.075	.057	2.088	1.991	1.531	.0504
.80	66	38.92	.090	.0287	50.11	.060	.046	1.507	2.083	1.601	.0519
.82	64	40.46	.093	.0263	55.21	.054	.042	1.446	2.063	1.586	.0510
.84	58	39.28	.075	.0222	65.71	.046	.035	1.100	2.055	1.579	.0509
.86	54	39.08	.072	.0195	75.26	.040	.031	.901	2.042	1.570	.0513
.88	54	41.30	.079	.0231	63.18	.047	.036	1.421	2.055	1.580	.0518
.90	50	42.16	.070	.0206	70.49	.042	.033	1.236	2.062	1.585	.0497
.92	48	41.22	.064	.0181	81.18	.037	.028	.977	2.037	1.566	.0522
.94	40	36.12	.054	.0168	94.45	.032	.024	.781	1.893	1.455	.0519
.96	42	40.10	.045	.0152	99.83	.030	.023	.737	1.976	1.519	.0526
.98	42	45.70	.056	.0206	68.41	.044	.034	1.716	2.124	1.633	.0504
1.00	36	40.44	.057	.0202	72.84	.041	.032	1.582	2.034	1.564	.0536
1.02	32	40.78	.047	.0152	100.00	.030	.023	1.098	2.034	1.564	.0536
1.04	34	40.68	.040	.0111	138.00	.022	.017	.640	1.948	1.497	.0517
1.06	30	38.02	.030	.0087	183.54	.016	.013	.420	1.875	1.441	.0488
1.08	32	46.92	.021	.0074	191.22	.016	.012	.353	2.111	1.622	.0511
1.10	34	46.94	.019	.0066	201.85	.014	.011	.302	2.064	1.586	.0573
1.12	32	48.62	.018	.0054	272.89	.011	.008	.221	2.036	1.565	.0545
1.14	34	51.98	.014	.0044	326.13	.009	.007	.171	2.057	1.581	.0566
1.16	42	57.56	.013	.0041	354.49	.008	.006	.164	2.076	1.595	.0681
1.18	40	58.92	.013	.0042	334.55	.009	.007	.191	2.122	1.631	.0660
1.20	40	62.72	.013	.0042	332.50	.009	.007	.221	2.136	1.642	.0622
1.22	40	64.44	.012	.0040	351.08	.009	.007	.250	2.148	1.651	.0608
1.24	36	66.12	.011	.0035	399.08	.008	.006	.185	2.168	1.666	.0539
1.26	30	65.24	.010	.0029	469.67	.006	.005	.139	2.190	1.683	.0459