

SEISMIC HAZARD ALONG A CRUDE OIL PIPELINE IN THE EVENT OF AN 1811-1812 TYPE NEW MADRID EARTHQUAKE

by

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PREFACE

The National Center for Earthquake Engineering Research (NCEER) is devoted to the expansion and dissemination of knowledge about earthquakes, the improvement of earthquake-resistant design, and the implementation of seismic hazard mitigation procedures to minimize loss of lives and property. The emphasis is on structures and lifelines that are found in zones of moderate to high seismicity throughout the United States.

NCEER's research is being carried out in an integrated and coordinated manner following a structured program. The current research program comprises four main areas:

- Existing and New Structures
- Secondary and Protective Systems
- Lifeline Systems
- Disaster Research and Planning

This technical report pertains to Program 3, Lifeline Systems, and more specifically to crude oil transmission systems.

The safe and serviceable operation of lifeline systems such as gas, electricity, oil, water, communication and transportation networks, immediately after a severe earthquake, is of crucial importance to the welfare of the general public, and to the mitigation of seismic hazards upon society at large. The long-term goals of the lifeline study are to evaluate the seismic performance of lifeline systems in general, and to recommend measures for mitigating the societal risk arising from their failures.

From this point of view, Center researchers are concentrating on the study of specific existing lifeline systems, such as water delivery and crude oil transmission systems. A seismic performance analysis of crude oil transmission systems in the New Madrid area is underway. The study focuses on the vulnerability of these systems to seismic events. Technical and societal issues arising from disruption of the supply of crude oil from the south to the northeast following an earthquake are addressed, as is potential envircnmental pollution due to seismically induced failure of pipelines.

The research activities comprising the crude oil transmission system study are shown in the accompanying figure.

This report provides an assessment of the seismic hazard that exists along a major crude oil pipeline traversing the New Madrid seismic zone. An 1811-1812 *type New Madrid earthquake with moment magnitude* 8.2 *is assumed to occur at three locations where large historical earthquakes have occurred. Six pipeline crossings of the major rivers in West Tennessee are chosen as the sites for hazard evaluation because of the liquefaction potential at these sites. A seismologically-based model is used to predict the bedrock accelerations. Uncertainties in three model parameters, i.e., stress parameter, cutoff frequency, and strong-motion duration are included in the analysis. Each parameter is represented by three typical values. From the combination of these typical values, a total of* 27 *earthquake time histories are generated for each selected site due to an* 1811-1812 *type New Madrid earthquake occurring at a postulated seismic source.*

ABSTRACT

An assessment of the seismic hazard that exists along the major crude oil pipeline (pipeline 22) is presented in this report. An 1811- 1812 type New Madrid earthquake with moment magnitude 8.2 is assumed to occur at three locations where large historical earthquakes have occurred. Six pipeline crossings of the major rivers in West Tennessee are chosen as the sites for hazard evaluation because of the liquefaction potential at these sites. A seismologicallybased model is used to predict the bedrock accekrations. Uncertainties in three model parameters, i.e., stress parameter, cutoff frequency, and strong-motion duration are included in the analysis. Each parameter is represented by three typical values. From the combination of these typical values, a total of 27 earthquake time histories can be generated for each selected site due to an 1811- 1812 type New Madrid earthquake occurring at a postulated seismic source.

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INTRODUCTION

A major crude oil pipeline, Capline (pipeline 22), passes through the recharge area of underground water supplies in the West Tennessee region (figure 1-1). Because of the age of this pipeline, it is doubtful that any seismic design was included in the pipeline design. The vulnerability of the pipeline subject to an 1811-1812 type New Madrid earthquake is a major concern. In a pilot study conducted by the NCEER lifeline group [1], the buried pipeline was considered to fail due to soil liquefaction induced by earthquakes. The areas of high probability of liquefaction in Tennessee include mosl of the stream beds in West Tennessee (figure 1-1). If a pipeline were ruptured during an earthquake, the spilled material could contaminate the recharge area, thus impacting the water supply of almost all West Tennessee. Such an impact would be enormous since there is no immediately alternative source of water supply in this region. In this study, the seismic hazards along pipelin \approx 22 is assessed, assuming that a scenario earthquake with magnitude similar to the 1811-1812 New Madrid earthquakes will oceur at a postulated source in the New Madrid seismic zone.

Pipeline 22 and Liquefiable Soil Areas in Tennessee FIGURE 1-1

SECTION 2

PIPELINE AND SEISMIC SOURCES

The Capline system is a 40-inch pipeline operated by Shell Pipeline Corporation of Houston, Texas. The pipeline transports crude oil from southeastern Louisiana to Patoka, Illinois [1], traveling through soft sediments of the Mississippi Valley and crossing major rivers in West Tennessee: Wolf River, Loosahatchie River, Hatchie River, Forked Deer-South Fork, Forked Deer-North Fork, and Obion River. These river crossings are the potential locations of soil liquefaction during a large earthquake, and therefore are selected as the sites for seismic hazard evaluation. The locations of these six river crossings are listed in table 2-1 and also shown in figure 2-1.

The New Madrid seismic zone (NMSZ) is clearly delineated by the concentration of epicenters (figure 2-2). From the pattern of epicenters, at least three distinct linear trends suggesting the orientations of three subsurface fault segments in the New Madrid region are observed $[2,3]$. These three fault segments are (1) a southern segment extending from Marked Tree, Arkansas, to Caruthersville, Missouri, roughly along the axis of the Reelfoot rift complex; (2) a middle segment trending northwest and extending from Ridgely, Tennessee, to west of New Madrid, Missouri; and (3) a relatively shorter northern segment extending from west of New Madrid, Missouri, to southern Illinois. From the focal mechanism studies [3,4], both the northern and southern segments exhibit a predominantly right-lateral strike-slip fault-plane motion. The middle segment, on the other hand, exhibits a thrus:-faulting mechanism.

During the winter of 1811-1812, three great earthquakes occurred in the Mississippi Valley. Because of low attenuation in the eastern and central United States, the intensity of the 1811-1812 New Madrid earthquakes (according to Nuttli [5]) produced far greater damage than that produced by any other historical earthquake in the North American continent. The shocks were felt over a wide area of the central and eastern United States. Nuttli constructed a generalized isoseismal map of the first principal shock that occurred on December 16, 1811 (figure 2-3). The area estimated to be 600,000 k m 2, experienced damaging earth motions (modified Mercalli intensity \geq VII). Figure 2-4 illustrates the ground disruptions such as

 $\sim 10^7$

TABLE 2-1 Site Locations

FIGURE 2-1

Source and Site Locations

FIGURE 2-2 Epicenters of New Madrid Earthquakes

FIGURE 2-3 Isoseismal Map of the 1811 New Madrid Earthquake (after Nuttli, O.W., 1973)

FIGURE 2-4 Map of Ground Effects of the 1811-12 New Madrid Earthquakes (adapted from Fuller, 1912; taken from Jibson and Keefer, 1988)

uplift and sand extrusion due to the 1811-1812 New Madrid earthquakes [6].

In addition to the three 1811-1812 New Madrid earthquakes, there were two other damaging earthquakes that occurred in the NMSZ (figure 2-5). One occurred in 1843 near Marked Tree, Arkan ;as, and the other occurred in 1895 near Charleston, Missouri. Table 2 -II lists the locations and moment magnitudes M of these five historical events [7].

In this study, a scenario earthquake similar to an 1811-1812 type New Madrid earthquake is postulated to occur in the NMSZ. Considering the concentration of instrumental epicenters, the historical events, and the site locations selected for this study, three earthquake sources $(A, B, and C)$ are selected as shown in table 2-III. Since the scenario earthquake is assumed to have a magnitude similar to the 1811-1812 New Madrid earthquakes $(M = 8.2, 8.09, ...)$ and 8.3), the moment magnitude $M = 8.2$ is assigned to the postulated scenario earthquake. The epicentral distances from these three seismic sources to six selected sites are listed in table 2 -IV.

FIGURE 2-5 Epicenters of Five Historical New Madrid Earthquakes

Event	Epicenter	M	
	Latitude	Longitude	
1811 Dec 16	36.0N	90.0W	
1812 Jan 23	36.3N	89.6W	8.20 8.09
1812 Feb 07	36.5N	89.6W	8.30
1843 Jan 05	35.5N	90.5W	6.47
1895 Oct 31	37.0N	89.4W	6.81

TABLE 2-11 Historical New **Madrid Earthquake5**

 $\overline{}$ and $\overline{}$

Source	Latitude	Longitude	Historical Reference
\mathbf{A}	35.5N	90.5W	1843 0105 Marked Tree event
B	36.0N	90.0W	1811 1216 New Madrid event
\mathcal{C}	36.3N	89.6W	1812 0123 New Madrid event

TABLE 2-111 Source Locations

SECTION 3

SEISMIC HAZARD ESTIMATION

Estimating the characteristics of ground motions induced by a large New Madrid earthquake is quite challenging because of the lack of strong-motion data in the New Madrid region. In this study, a seismologically-based model is used to predict the horizontal bedrock motions primarily due to shear waves generated from a seismic source [8]. The model is centered on a power spectrum which in turn is developed from a Fourier amplitude spectrum. The Fourier amplitude spectrum is constructed based on the physical characteristics of an earthquake including source mechanism and path attenuation. From the power spectrum, earthquake time histories and probability-based response spectra can be generated directly. The power spectrum can also be used to estimate the peak earthquake accelerations.

3.1 Fourier Amplitude Spectrum

The Fourier amplitude spectrum used in this study essentially follows the Boore and Atkinson approach [9].

$$
A(f) = C \times S(f) \times D(f) \times I(f)
$$
 (3.1)

where C is a scaling factor; $S(f)$ is a source spectral function; $D(f)$ is a diminution function; and $I(f)$ is a shape filter.

The source spectral function $S(f)$ is a frequency-domain representation of the seismic energy released by an earthquake. In this study, we use an ω^2 source spectrum proposed by Brune [10]. This source spectrum is a function of a single corner frequency f_0 and a seismic moment M_0 :

$$
S(f) = \frac{M_0}{1 + (\frac{f}{f_0})^2}
$$
 (3.2)

The corner frequency f_0 is determined by the seismic moment M₀, crustal shear wave velocity β , and stress parameter $\Delta \sigma$:

$$
f_0 = 4.9 \times 10^6 \beta \left(\frac{\Delta \sigma}{M_0}\right)^{1/3} \tag{3.3}
$$

The seismic moment M_0 is related to average fault displacement \tilde{D} , fault rupture area A, and modulus of rigidity μ in the source zone:

$$
M_0 = \mu A \widetilde{D} \tag{3.4}
$$

Stress parameter $\Delta \sigma$ describes the level of the source spectral function above the corner frequency f_0 . The stress parameter $\Delta \sigma$ proposed by Atkinson [11], Boore and Atkinson [9], Brune [10], and Joyner [12] is a constant and is independent of earthquake magnitude, while the stress parameter suggested by Nuttli [13] increases with magnitude. In this study, a constant $\Delta \sigma$ is used.

According to the scaling law suggested by Boore [14] and Joyner [12], the scaling factor C is

$$
C = \frac{\langle R_{\theta\phi} \rangle \ F \ V}{4 \ \pi \ \rho \ \beta^3} \cdot \frac{1}{r}
$$
 (3.5)

where:

$$
\langle R_{\theta\phi} \rangle = \text{radiation pattern}
$$
\n
$$
F = \text{free-surface effect}
$$
\n
$$
V = \text{partition of a vector into horizontal components}
$$
\n
$$
\rho = \text{crustal density}
$$
\n
$$
r = \text{hypocentral distance}
$$

 $\langle R_{\theta_0} \rangle$ is the radiation pattern corresponding to different types of seismic waves over a range of azimuths (θ) and takeoff angles (ϕ) . For ϕ and θ averaged over the whole focal sphere, the shear-wave radiation pattern $\langle R_{\theta\phi} \rangle$ is 0.55 [15]. F accounts for the amplification due to the free surface and is taken as 2 [14]. V is the factor that accounts for the partition of a vector into horizontal components and

is chosen as $1/\sqrt{2}$. Based on the hypocentral locations of instrumentally recorded microearthquakes in the NMSZ, the focal depths normally range from 5 to 20 km below the surface where the granitic basement rock of the upper and middle continental crusts are found. **In** this study, an average focal depth is taken as 10 km. The density ρ of continental crust at this focal depth is taken as 2.7 $g m / c m^3$, and the shear-wave velocity β is 3.5 km/sec. The hypocentral distance r is the distance measured from the focus of an earthquake to the site. The term $1/r$ in equation (3.5) accounts for the diminishing of source spectrum as a result of body-wave geometric spreading.

The diminution function *D(f)* represents the anelastic attenuation that accounts for the damping of the earth's crust and a sharp decrease of acceleration spectra abc ve a cutoff frequency f_m .

$$
D(f) = \exp\left[\frac{-\pi f r}{Q(f) \beta}\right] P(f, f_m)
$$
 (3.6)

where:

 $Q(f)$ = frequency-dependent quality factor $P(f, f_m)$ = high-cut filter

The exponential term in equation (3.6) accounts for path attenuation. The quality factor $Q(f)$ describes the attenuation of seismic waves and is frequency dependent. The great distance at which the 1811-1812 New Madrid earthquakes were felt has been attributed to the low attenuation of seismic waves in central and eastern United States [5]. Several frequency-dependent quality factors *Q(f)* have been proposed [16-18]. Based on the attenuation study conducted by Dwyer and others [16] in central United States, the quality factor of shear and Lg waves is

$$
Q(f) = 1500 f \, 0.40 \tag{3.7}
$$

Since this study is related to the NMSZ, we feel that the quality factor proposed by Dwyer and others is more appropriate and so is used in this study.

The high-cut filter $P(f, f_m)$ accounts for the observation that the acceleration spectra often show a sharp decrease above a cutoff frequency f_m . In this study, a Butterworth filter is used as the highcut filter:

$$
P(f, f_m) = \left[1 + \left(\frac{f}{f_m}\right)^8\right]^{-1/2} \tag{3.8}
$$

The shape filter $I(f)$ is to shape the source spectral function for a particular type of earthquake motion of interest. For acceleration, the shape filter is given as

$$
I(f) = (2\pi f)^2
$$
 (3.9)

3.2 Power Spectral Density Function

An earthquake accelerogram generally shows a build-up segment followed by a strong-motion segment which is in turn followed by a decay segment. The frequency content of earthquake accelerograms is approximately constant during the strong-motion segment. Thus, the strong-motion segment of an acceleration time history is considered as a stationary random process. The one-sided power spectral density function (power spectrum) $S_a(f)$ can be derived from the Fourier amplitude spectrum.

$$
S_a(f) = \frac{2}{T} |A(f)|^2
$$
 (3.10)

where $A(f)$ is the Fourier amplitude spectrum as in equation (3.1) , and T is the strong-motion duration.

3.3 Acceleration Time Histories

In this study, synthetic earthquake time histories are generated using the method proposed by Shinozuka [19]. Given the power spectrum, the stationary acceleration time histories $a_s(t)$ can be generated as follows:

$$
a_{s}(t) = \sqrt{2} \sum_{k=1}^{N} \sqrt{S_{a}(\omega_{k}) \Delta \omega} \cos(2\pi \omega_{k} t + \phi_{k})
$$
 (3.11)

where:

$$
S_a(\omega_k) = one-sided earthquake power spectrum\nN = number of frequency intervals\n\Delta\omega = frequency increment\n\omega_k = k \Delta\omega\n\phi_k = random phase angles uniformly distributed\nbetween 0 and 2\pi
$$

The nonstationary acceleration time histories $a(t)$ can be obtained from the multiplication of an envelope function $w(t)$:

$$
a(t)=a_s(t) w(t) \qquad (3.12)
$$

The envelope function $w(t)$ proposed by Hwang et al. [2] and used in this study is comprised of three segments: (l) a parabolically increased segment simulating the initial rise part of the accelerogram and its duration is chosen as one fifth of T, (2) a constant segment representing the strong-motion portion of an earthquake excitation and has a duration equal to T, and (3) a linearly decayed segment extending four fifths of T. Thus, the total duration is 2T. It is noted that real seismograms are commonly observed with long coda durations; however, the coda durations are considered unimportant in most engineering applications.

3.4 Peak Values of Earthquake Motions

The peak value A_p of an acceleration time history $a(t)$ over a duration T is defined as

$$
A_p = \max |a(t)| \tag{3.13}
$$

The statistical distribution of A_p can be approximated by the asymptotic distribution function of the extreme values [20]. In this study, the cumulative distribution function of A_p proposed by Vanmarcke [21] is used.

$$
F_{Ap}(r) = [1 - exp(-\frac{\alpha^2}{2})] exp[-v_0 T \frac{1 - exp(-\sqrt{\pi/2} \alpha \delta_e)}{exp(\alpha^2/2) - 1}; r > 0 \qquad (3.14)
$$

in which v_0 is the mean zero-crossing rate of the displacement response; $\delta_e = \delta^{1.2}$ in which δ is a shape factor [21]; and $\alpha = r/\sigma_a$ is a normalized barrier level. The standard deviation σ_a can be computed from the power spectrum $S_a(\omega)$.

The mean and standard deviation of A_p are particularly useful for engineering applications. The mean value \overline{A}_{p} and standard deviation σ_{Ap} can be expressed as

$$
\overline{A}_p = p_m \sigma_a \tag{3.15}
$$

$$
\sigma_{\text{Ap}} = q \sigma_a \tag{3.16}
$$

From equation (3.14), Der Kiureghian [22] obtained the following empirical equations for p_m and q:

$$
p_{m} = \sqrt{2 \ln(v_{e}T) + 0.5772 / \sqrt{2 \ln(v_{e}T)}}
$$
 (3.17)

$$
q = 1.2 / \sqrt{2 \ln(v_e T)} - 5.4 / (13 + [2 \ln(v_e T)]^{3.2})
$$
 (3.18)

in which

$$
\mathbf{v}_{e} \mathbf{T} = \begin{cases} \max(2.1, 2\delta \mathbf{v}_{0} \mathbf{T}); & 0.00 < \delta \le 0.10 \\ (1.63 \delta^{0.45} - 0.38) \mathbf{v}_{0} \mathbf{T}; & 0.10 < \delta < 0.69 \\ \mathbf{v}_{0} \mathbf{T} & 0.69 \le \delta < 1.00 \end{cases}
$$
 (3.19)

3.5 Parameter Uncertainties

The seismologically-based model for the horizontal bedrock accelerations is defined by several parameters as summarized in table 3-1. In this study, the moment magnitude M is chosen as 8.2 to simulate the $1811-1812$ type New Madrid earthquak?. The epicentral distance R is dependent on the site locations. Some parameters such as the radiation pattern is determined from the consideration of earthquake seismology. The focal depth and the corresponding crustal density ρ and shear-wave velocity β are chosen from the data related to the New Madrid region. Since these parameters appear to have less influence on the resulting bedrock accelerations, a deterministic value is assigned to each cf these parameters. A deterministic quality factor $Q(f)$ is also used in this study due to the lack of pertinent information to quantify the variation. On the other hand, the stress parameter $\Delta \sigma$, cutoff frequency f_m , and strong-motion duration T have significant effects on the resulting bedrock accelerations. Thus, the uncertainties in these parameters are included in the analysis.

In this study, a constant stress parameter is used to predict the horizontal bedrock time histories. Boore and Atkinson [9], McGuire and others [23], and Somerville and others [24] suggested an average $\Delta \sigma$ around 100 bars for central and eastern North America. Using the 1988 Saguenay earthquake, Atkinson and Boore [25] determined that the stress parameter is about 200 bars. Thus, three values, $\Delta \sigma = 100$, 150, and 200 bars, are selected for this study.

The cutoff frequency f_m is a parameter to model the decay of the Fourier amplitude spectrum beyond certain frequency. The selection of *fm* also affects the peak value of an earthquake. According to Boore and Atkinson [9], *fm* is uncertain in eastern North America. McGuire and others [23] suggested f_m = 40 bars for rock sites in this region. From the study of the 1988 Saguenay earthquake, Chen and Hwang [26] found that the value of f_m seems to decrease as the epicentral distance increases and $fm = 20$ to 30 Hz are obtained for rock sites at epicentral distances less than 100 km. In this study, we select 20, 30, and 40 Hz to cover the uncertainty in cutoff frequencies.

The strong-motion duration T is ar other important factor that affects the outcome of the time histories. Hanks [27] suggested tha: strong

TABLE 3-1 Earthquake Parameters

motion duration T is equal to the source duration, which is the reciprocal of the corner frequency *fo:*

$$
T = \frac{1}{f_0} \tag{3.20}
$$

On the basis of this equation, the strong-motion duration T for an $M = 8.2$ earthquake with a stress parameter of 150 bars is about 32 sec. According to studies conducted by Johnston [28], Krinitzsky and others [29], Lai [30], and Sues and others [31], strong-motion duration has significant variation. Thus, the coefficient of variation is taken as 50%. Based on this consideration, three values of 16, 32, and 48 sec are chosen for an earthquake with M of 8.2 and $\Delta\sigma$ of 150 bars. For earthquakes with different stress parameters, the strongmotion durations are computed according to the same procedure.

3.6 Numerical Results

In this study, the uncertainties in model parameters $\Delta \sigma$, f_m , and T are represented by three values. From the combination of these typical values, a total of 27 samples of earthquakes can be generated and are listed in table 3-11. For the purpose of illustration, the seismic source is assumed at Marked Tree, Arkansas (source A) and the site is taken as the Wolf River crossing (Site 1). $\Delta\sigma$, f_m , and T are selected as 150 bars, 30 Hz and 32 sec, respectively. These parameters, together with those listed in table 3-1, are used to generate synthetic earthquakes. The Fourier amplitude spectrum, power spectrum and a synthetic time history corresponding to these parameters are shown in figures 3-], 3-2, and 3-3, respectively. The 27 peak bedrock accelerations (PBA) computed directly from the power spectrum are listed in table 3-111. Appendix A shows the peak bedrock accelerations of all six sites from three different seismic sources. A statistical analysis is carried out to determine the maximum value A_{max} , minimum value A_{min} , mean value A_{mean} , and coefficient of variation (COV). Table 3-IV shows these statistics for all six sites from three seismic sources. Notice that the COV for all cases are about 0.33 and the mean value is very close to the so-called best-estimate value, i.e. the peak value determined using the mean parameter values.

Sample	Δσ (bars)	$f_{\rm m}$ (Hz)	T (sec)
$\mathbf{1}$	100.00	20.00	18.00
$\overline{2}$	100.00	20.00	36.00
\mathfrak{Z}	100.00	20.00	54.00
4	100.00	30.00	18.00
5	100.00	30.00	36.00
6	100.00	30.00	54.00
$\overline{\mathcal{I}}$	100.00	40.00	18.00
8	100.00	40.00	36.00
9	100.00	40.00	54.00
10	150.00	20.00	16.00
11	150.00	20.00	32.00
12	150.00	20.00	48.00
13	150.00	30.00	16.00
14	150.00	30.00	32.00
1 ₅	150.00	30.00	48.00
16	150.00	40.00	16.00
17	150.00	40.00	32.00
18	150.00	40.00	48.00
19	200.00	20.00	15.00
20	200.00	20.00	30.00
21	200.00	20.00	45.00
22	200.00	30.00	15.00
23	200.00	30.00	30.00
24	200.00	30.00	45.00
25	200.00	40.00	15.00
26	200.00	40.00	30.00
27	200.00	40.00	45.00

TABLE 3-11 Samples of Earthquake

Fourier Amplitude Spectrum (M=8.2, R=85 km) FIGURE 3-1

POWER SPECTRUM

Power Spectrum (M=8.2, R=85 km) FIGURE 3-2

 0.0

LOG OF FREQUENCY (Hz)

 $1.0\,$

 2.0

INTENSITY (cm**2/sec**3)

 0.0

 -2.0

 -1.0

TABLE **3-111** Peak Accelerations at Wolf River Crossing due to Seismic Source at Marked Tree

Source	Site	A_{max} (g)	A_{min} (g)	A_{mean} (g)	COV
$\mathbf A$	$\mathbf{1}$	0.49	0.13	0.27	0.33
	\overline{c}	0.46	0.12	0.25	0.33
	3	0.42	0.11	0.23	0.33
	$\overline{4}$	0.38	0.10	0.21	0.33
	5	0.36	0.10	0.20	0.32
	6	0.31	0.09	0.17	0.32
B	$\mathbf{1}$	0.42	0.11	0.23	0.33
	$\boldsymbol{2}$	0.44	0.12	0.24	0.33
	3	0.65	0.17	0.36	0.33
	$\overline{4}$	0.95	0.25	0.51	0.33
	5	1.09	0.28	0.59	0.34
	6	0.80	0.21	0.43	0.33
C	$\mathbf{1}$	0.36	0.10	0.20	0.32
	\overline{c}	0.37	0.10	0.20	0.32
	$\overline{3}$	0.49	0.13	0.27	0.33
	$\overline{4}$	0.76	0.20	0.41	0.33
	5	1.03	0.27	0.56	0.34
	6	1.21	0.31	0.65	0.34

TABLE 3·IV Statistics of Peak Bedrock Accelerations

SECTION 4

SUMMARY AND CONCLUSIONS

Estimation of seismic hazard is an essential task in carrying out seismic risk assessment studies. In this study, we present the seismic hazard assessment along pipeline 22 (Capline). An 1811-1812 type New Madrid earthquake with $M = 8.2$ is assumed to occur at three locations, where large historical earthquakes have occurred. Six pipeline crossings of major rivers in West Tennessee are chosen as the sites for hazard evaluation because of the liquefaction potential at these locations. A seismologically-based model is used to predict the bedrock accelerations. Uncertainties in three model parameters, i.e., stress parameter $\Delta \sigma$, cutoff frequency f_m , and strong-motion duration T are included in the analysis. Each parameter is represented by three values. From the combination of these parameters, 27 samples of earthquake are generated for eact pair of seismic sources and selected sites. The results of the seismic hazard analysis can be used to evaluate the liquefaction potential at these sites, assess the vulnerability of pipeline facilities, investigate the risk from a postulated pipeline break, and prepare an emergency response plan.

SECTION 5

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Appendix A Peak Bedrock Accelerations

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